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## Mare2DEM on land:

### MT Data from the Cape Fold Belt (South Africa) revisited

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#### INTRODUCTION

Mare2DEM is a parallel adaptive finite element code for 2D forward and inverse modelling for electromagnetic data (Key & Ovall, 2011), which is now being made freely available. Mare2DEM was originally designed with marine controlled-source electromagnetic (CSEM) and marine magnetotelluric (MT) applications in mind, but it can also be applied to onshore data. Important features of Mare2DEM are:

- se automatic mesh generation and refinement se triaxial (intrinsic) anisotropy
- s topography

To test this inversion code with onshore MT data we use the data set measured in 2005 in the Cape Fold Belt, South Africa. The stations are aligned along a ~ 100km long profile with significant topography and a close-by ocean. Several stations indicate the existence of electrical anisotropy in the subsurface through phases > 90°. Commonly, these phases out of quadrant are explained by an anisotropy strike oblique to the mean conductivity extructions.

While in earlier inversion studies only data without 3D/anisotropy effects were inverted using WinGLink, we can now compare inversion results of this study, in particular how Mare2DEM deals the anisotropy effects in the MT data.



### Figure 1 User interface Mamba2D for Mare2DEM to specify the following isotropic / anisotropic

conductivities import a priori information, e.q.

2D Models

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## topography • create and modify nodes segments • set conductivities / lock cells generate mesh

♦ target rms ♦ conductivity bounds
♦ regularization parameters



Figure 2

Figure 2 Map of the study area in South Africa. The blue dots show the location of the MT-Stations, which were projected onto the . dashed profile line. The MT profile consists of 54 stations and was measured in 2005. It crosses the Cape Fold Belt in South Africa between Prince Albert and Mosselbay.





## [ImD)<sup>01</sup> Boj] <sup>e</sup>

# Figure 3 Isotropic conductivity model calculated with Mare2DEM [Key and Ovall, 2011]. White triangles show the location of the MT stations. The topography was included with 50m resolution together with the Ocean (rough bathymetry) at the southern end of the profile. Total rms: 2.04

#### Figure 4

Figure 4 Anisotropic conductivity model allowing for triaxial anisotropy. Here we present the y component of the conductivity which shows similar features compared to the isotropic inversion. The anisotropy introduced did not exceed a ratio of 2, but phases >90° cannot be modelled. Total rms: 1.87

Figure 5 Comparison between (a) surface geology, (b) 2D isotropic inversion inded (Mare2DEM) and (c) 2D isotropic inversion model (RLM2D)) (Weckmann et al., 2012). Prominent conductors e.g. under the Oudsthoorn Basin are included in both inversions, but have different shapes. FE inversion shows more structures (overfitting?), but vertical anomalies seem to correlate with positions of syn- and anticlines and might image fluid pathways. Including topography in EE inversion was vital to improve data fit. While the FD inversion did not coverage, 2 stations north of the profile had to be included an a-prioric conductivity structure beneath fixed. Figure 5



Figure 6 Pseudo-sections of TE and TM apparent resistivity and phase. Phases >90° (circled) are observed in the middle of the profile at periods >1s in both modes [Chen, 2012].



Mare2DEM works with land MT data.

topography has improved data fit

♦ Only triaxial anisotropy --> phase values over 90° cannot be reproduced.

Anisotropic inversion absorbes 3D effects



Figure 9 Pseudosections of observed and modelled TE and TM mode phases

## Data fit (lines) of the inversion after omiting data with phases >90°. The total rms improved, but phases approaching or higher than 90° usually correlate with decrasing resistivities. The apparent resistivities can be fitted, but not the phases approaching 90°.

MT Data & fit

aching 90°. along the profile. White areas at long periods show omitted data due to phases >90°. Because of static shift, apparent resistivities were downwaited in the inversion and not shown here. Therefore only pha-ses are presented, which could be explained well by the inversion model.

#### Data fit (lines) of the inversion Figure 10

Figure 10 Development of total rms and roughness during the inversion run. While the rms does not change a lot after the 20th iteration the model includes more complex structures. The final iteration and "reasonable" inversion results have to be compared, to prevent overfitting of the data.

#### Figure 11

Figure 11 rms for each ♦ station (in depen-dance of data type) ● period (in dependance of data type) ● data type





Chen, X.: Two-dimensional anisotropic inversion of magnetotelluric data, 2012 Key, K.: MARZEDM: Tutorial and Training, Scripps Institution of Oceanography, UC San Diego, 2014 Key, K., Ovali, J.: A parallel goal-oriented adaptive finite diemeit method for 2.5-D electromagnetic modelling, Geophysical Journal International, 2011 ckie, R.: Nonlinear conjugate ithm for 2-D magnetotelluric

inversion, Geophysics, 2001, volume 66, No.1, S. 174-187 174-187 Rullf, P.: Vergleich isotroper und anisotroper 2-D Inversion (MARE2DEN) von magnetotelluirschen Daten aus dem Cape Fold Belt, Südafrika, Bsc. Thesis 2015 Weckmann et al.: Magnetotelluric image linked to surface geology across the Cape Fold Belt, South Africa, Terra Nova, 2012



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