Neutron - and X-ray diffraction analysis of residual stress in rocks of the KTB

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X-Ray diffraction:

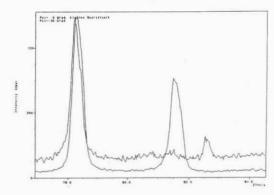
X-ray and neutron diffraction analysis of residual stress is based on the measurement of the interplanar spacings of familiar planes (hkl) of rock forming minerals which have been deformed.

The strains, caused by the stress, is given by Bragg's law (Maeder et al., 1980),

$$\varepsilon = \frac{1}{2} \cdot \cot g\Theta \cdot 2\Delta\Theta$$

where θ_0 is the Bragg angle in a deformed lattice and $\Delta\Theta$ is the shift of a diffraction line of the (hkl) family caused by stress.

Fig. 1 shows the result of the lattice-deformation for specimen KTB-607Ala resulting from the position of the quartz 311-reflection. A Bragg-diffraction angle shift of $\Delta 2\Theta = 0.06^{\circ}$ can be recognised.



Pig. 1: Diffraction angle and reflection profile for the Bragg reflection (311) of quartz at various stress levels ($\Delta 2\Theta = 0.06$)

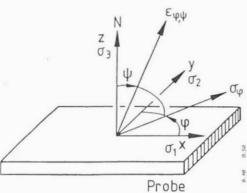


Fig. 2: Coordinate system for fixing a triaxial stress condition

For the stress analysis a polycrystalline plane specimen is assigned to an orthogonal coordination lattice with the orientation of κ_1 -axes, as shown in Figure 2.

These mentioned relations are a function of the orientation of planes (hkl) with respect to the sample surface.

This condition can be realised by adding in addition to the 20 circle two further independent circles, w and ϕ , to a diffractometer.

Their geometrical relation to the orientation of the strain σ_i -axes of the sample is shown in Figure 2.

For stress analysis it is postulated that the strain $\epsilon_{\phi,\psi}$ is uniform to the measured lattice spacing of the same direction (Glocker, 1958):

$$\varepsilon_{\theta, \psi} = \frac{d_{\theta, \psi} - d_0}{d_0}$$

By determination of the lattice spacings of the same family in several w-directions at a given azimuth o the following regularities are established:

> 1. The lattice spacings $\epsilon_{\phi,\psi}$ are independent azimuth ϕ and linear to $\text{min}^2\psi.$

2. The slope of the sin'y straight line

$$m\varphi = \frac{\delta\varepsilon_{\varphi, \psi}}{\sin^2 \psi} = \frac{1}{2}s_2 * \sigma_{\varphi}$$

is proportional to the stress component σ_{ip} at a given azimuth ϕ .

3. Finally, by employing the Voigt constant e, the intercept of the strain plot on the axis of ordinates, which is the same for all azimuths w. gives the sum of the principal stresses according to the equation:

$$\varepsilon_{\varphi, \psi} = 0 = \varepsilon_3 = s_1(\sigma_1 + \sigma_2)$$
.

Using X-rays the residual elastic strain locked in quartz grains of naturally deformed rocks can be monitored.

The interatomic spacings were determined as a function of sin's on KTB cores. The X-ray diffraction datas of specimen 607Ala are documented in figure 3.

From the diagram in Fig. 3 a residual strees

-280 MPa could be determined.

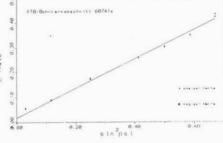


Fig. 3: Lattice spacing distribution measured on 311 quartz planes of KTB-core 607A1a versus various vi-angles

Neutron diffraction:

In the case of silicates the penetration depth of neutron radiation has a magnitude of centimetres. Therefore, and with the aid of disphragms and he xyz-translation device it was possible to choose any crystal of the specimen for determination of the residual stress analysis by neutron diffraction. Even stress gradients in smallest regions of a sample could be determined.

For investigating coarse grained rocks, the single crystal diffraction technique with neutron radiation ie an indispensable tool.

These experiments were carried out at the Berlin Neutron Scattering Centre - BENSC, using a powder diffractometer which was installed at the beam tube

For the single crystal technique an Eulerian cradle with an xyz-translation device is installed in the instrument E3 (Fig.4).

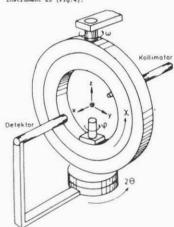


Fig 4 Euler cradle with integrated xyz-translation device

On a sample of hexagonal quartz crystals different compounds of the deformation tensor \mathbf{e}_{ij} could be measured from several deformed do.w-spacings Hence the deformation tensor & could be established from the following datas as shown below:

d-value (calc.)	d-value (obs.)	20 (calc.)	20 (obs.)	Intensity	hkl	d _{cate} - d _{obe}
2.237	2.227	48.844	49.066	37	111	9.4 x 10 ⁻³
2.127	2.125	51.55	51.615	76	200	2.5 x 10 ⁻³
1.9792	1.9732	55,719	55,901	94	201	6.0 x 10 -3
1.5418	1,5391	73.723	73.871	27	211	2.7 x 10 ⁻³
1.3718	1.3702	84.787	84.910	54	301	1.6 x 10 ⁻³
1.2558	1.2535	94.8686	95.098	61	302	2.3 x 10 ⁻³
1.1804	1.1799	103.173	103.233	43	310	2.3 x 10 ⁻³
1.1405	1.1399	108.379	108.464	57	204	6.0 x 10-4
1.1143	1.1137	112.203	112.295	61	303	6.0 x 10-4

with c = 5.3972

The experiments were carried out on KTB-sample H027.

It must be mentioned that the direction of the principal stresses do not directly ensue out of the principle deformation direction.

Our investigation are being continued.

Crostack HA, Reimers W (1983) Analysis of Residual Stresses in Coarse Grained Materials. In: Macherauch, E. & Hauk, V. (eds) Residual Stresses in Science and Technology, OGM, Oberussel Crostack HA, Reimers W (1983) Eminituog von Materialspanningen nit Hilfe von Ronigen- und Neutroneobeuguingsmessungen an einzelnen Kristallen. In: Macherauch E & Hauk V (eds) Residual Stresses in Science and Technology, DGM, Oberussel Glocker (1988) Materialpriifung mit Röntgenstrahlung. Berlin: Springer Verlag.

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