

Originally published as:

Park, J., Lühr, H., Min, K. W., Lee, J.-J. (2010): Plasma density undulations in the nighttime mid-latitude F-region as observed by CHAMP, KOMPSAT-1, and DMSP F15. - Journal of Atmospheric and Solar-Terrestrial Physics, 72, 2-3, 183-192

DOI: 10.1016/j.jastp.2009.11.007

Plasma density undulations in the nighttime mid-latitude F-region as observed by CHAMP, KOMPSAT-1, and DMSP F15

Jaeheung Park^{a,b}, Hermann Lühr^a, Kyoung Wook Min^b, Jae-Jin Lee^c

^aHelmholtz Center Potsdam (GFZ), Sect. 2.3, Telegrafenberg, D-14473 Potsdam, Germany.

^bDepartment of Physics, Korea Advanced Institute of Science and Technology (KAIST), Yuseong-gu, Daejeon 305-701, Republic of Korea.

^cSolar and Space Weather Research Group, Korea Astronomy and Space Science Institute (KASI), Yuseong-gu, Daejeon 305-348, Republic of Korea.

11 Abstract

1

2

3

5

6

7

8

9

10

We investigate plasma density undulations in the nighttime mid-latitude 12 topside F-region. During solar maximum years the undulations are found 13 at CHAMP, KOMPSAT-1, and DMSP F15 altitudes. The occurrence rate 14 is higher at KOMPSAT-1 than at DMSP F15 altitude. The undulations 15 occur infrequently during equinoxes, and the occurrence peaks are in the 16 Asian/Oceanian (eastern Pacific/American) region during June (December) 17 solstice. At CHAMP altitude the undulations are observed all through the 18 night, and the occurrence rate is anti-correlated with the solar cycle. As all 19 these results are in general agreement with known climatology of MSTIDs, 20 we suggest that the undulations are a topside signature of MSTIDs. The 21 undulations are often but not always accompanied by magnetic signatures 22 indicating the presence of field-aligned current (FAC). The partial lack in 23 correspondence might be due to the ionospheric conductivity variation. The 24 similar distribution is, however, in support of a connection between density 25

Preprint submitted to Journal of Atmospheric and Solar-Terrestrial PhysicsNovember 4, 2009

- ²⁶ undulations and FACs.
- 27 Key words: Ionospheric irregularities, Mid-latitude ionosphere, Topside
- ²⁸ ionosphere, Ionospheric current

²⁹ 1. Introduction

The mid-latitude ionospheric F-region is known to be relatively calm com-30 pared to the equatorial region where equatorial plasma bubbles (EPBs) fre-31 quently occur or the high-latitude region where auroral precipitation and 32 polar cap convection generate complex current and plasma structures. How-33 ever, though relatively weak, the mid-latitude F-region also possesses plasma 34 structures. Plasma density irregularities in the nighttime mid-latitude F-35 region were first reported by *Peterson et al.* (1955). They conducted scatter-36 sounding experiments at the west coast of the United States, and observed 37 echoes which were notably different from ground scattering. Those echoes 38 were aligned with the geomagnetic field lines at the ionospheric E- and F-39 layer. They were generally accompanied by sporadic E (E_s) patches, but 40 the occurrence rate had little connection with the geomagnetic activity. Pe-41 terson et al. (1955) suggested that the echoes originate from equatorward 42 extension of the auroral oval. Afterwards, a number of papers were pub-43 lished on the topic. Using radio scintillation data above Australia, Munro 44 (1963) found discrete patches of scintillation regions in the nighttime mid-45 latitude ionosphere. Dyson (1968) investigated ionograms recorded by the 46 Alouette-I satellite, and found that the occurrence rate of topside spread F 47 above Australia reached its maximum around June solstice. 48

Behnke (1979) reported band structures in the nighttime mid-latitude 49 F region using the Arecibo incoherent scatter radar. They had horizon-50 tal widths of about 100 km, and drifted southwestward. Afterwards, such 51 band structures were called medium-scale traveling ionospheric disturbances 52 (MSTIDs). Using airglow images and total electron content (TEC) data 53 at Arecibo, Garcia et al. (2000) found that most MSTIDs are aligned from 54 northwest to southeast and drift southwestward. Shiokawa et al. (2003a) 55 showed that the MSTID occurrence rate in the eastern Asian longitude sec-56 tor maximizes during June solstice. Otsuka et al. (2004) and Shiokawa et al. 57 (2005) found that airglow images of MSTIDs taken at geomagnetic conju-58 gate points exactly mirror each other. Saito et al. (1995) observed electric 59 and magnetic fluctuations in the nighttime mid-latitude ionospheric F region, 60 and attributed them to mid-latitude plasma irregularities. Shiokawa et al. 61 (2003b) supported the close relationship between MSTIDs and electromag-62 netic fluctuations by conjugate observations of airglow intensity and in-situ 63 E-field measurements by a DMSP satellite. Recently, Park et al. (2009a) 64 reported linearly-polarized magnetic field fluctuations in the nighttime mid-65 latitude F region, and related them to MSTIDs. 66

There is a consensus that MSTIDs are connected with the Perkins instability (*Perkins*, 1973) whose growth rate, γ_P , is approximately given by (*Tsunoda*, 2006):

$$\gamma_P = \frac{E_0 \sin(\theta - \alpha) \sin\alpha \cos I}{BH} \tag{1}$$

⁷⁰ where E_0 is the ambient electric field, *B* magnetic field strength, *H* neutral ⁷¹ density scale height, θ the polarization angle of the ambient E-field measured ⁷² from magnetic east, and α the polarization angle of the wave normal of the

perturbed band structure. But, this equation cannot explain all the MSTID 73 features to full satisfaction. For example, Kelley and Makela (2001) had to 74 assume a finite MSTID size along the wavefront in order to reconcile the 75 observed MSTID drift (southwestward in the northern hemisphere) with the 76 Perkins instability theory. The linear growth rate of the Perkins instability 77 is too small to explain the plasma density structure (e.g. Saito et al. (1998) 78 and Shiokawa et al. (2003b)). Shiokawa et al. (2003b) argued that the effect 79 of conjugate F-region and E-regions has to be considered for reconciling the 80 disagreement. Tsunoda and Cosgrove (2001) and Tsunoda (2006) suggested 81 that E_s and MSTIDs give positive feedback to each other through electro-82 dynamic coupling. Recently, Yokoyama et al. (2009) conducted the first 3D 83 computer simulation of the coupled E_s /Perkins instability. The directional 84 preference of MSTID wavefronts and phase velocities was explained in terms 85 of E_s -MSTID coupling. 86

Various plasma undulation phenomena in the mid-latitude nighttime F 87 region, i.e. field-aligned irregularities (FAIs), mid-latitude spread F (MSF) 88 and MSTIDs are closely connected with each other. Fukao et al. (1991) 89 reported that FAIs are generally collocated with strong MSFs. According 90 to Bowman (1992) mid-latitude spread F (MSF) is generated by tilted iso-91 density contours which in turn originate from MSTIDs. Saito et al. (2001) 92 showed that FAIs do not always occur together with MSTIDs. Shiokawa et 93 al. (2003a) reported that only 10-15% of MSTIDs were accompanied by MSFs 94 while their statistical properties are similar. Kotake et al. (2006) argued that 95 different scale sizes of the two phenomena might lead to such discrepancy. 96 Lately, Otsuka et al. (2009) showed that intense FAIs occur in the plasma 97

⁹⁸ depletion region of MSTIDs, where enhanced gradient drift instabilities are
⁹⁹ suspected to generate FAIs. In this study we will mainly deal with MSTIDs
¹⁰⁰ whose scale size is several hundred kilometers.

Though there have been a lot of studies on the mid-latitude plasma struc-101 tures, their full global/seasonal occurrence pattern is not known yet. Most of 102 the previous investigations were constrained to ground observations: radar 103 experiments (Saito et al., 1998), 630.0 nm airglow imaging (Garcia et al., 104 2000; Otsuka et al., 2004; Candido et al., 2008), and GPS/TEC measure-105 ments (Kotake et al., 2006, 2007; Lee et al., 2008). Kotake et al. (2006) 106 explored the global MSTID climatology using worldwide GPS networks, but 107 their results are restricted to the landmasses. This lack of global data cov-108 erage is because MSTID is basically a phenomenon at the bottomside iono-109 sphere (Saito et al., 2001) and there are only a few satellites orbiting at such 110 a low altitude (below 350 km). However, some plasma density undulations do 11 exist also in the topside ionosphere. For example, Seker et al. (2009) showed 112 that MSTIDs can affect the topside as well as the bottomside ionosphere by 113 combining airglow imager and incoherent scatter radar observations. Livneh 114 et al. (2009) reported that the Arecibo incoherent scatter radar found plasma 115 density undulations from 160 km to altitudes higher than 500 km. Also, a few 116 groups investigated the midlatitude plasma structures by means of topside 117 satellite observations. Hanson and Johnson (1992) found mid-latitude iono-118 spheric disturbances below 300 km, which were morphologically similar to 119 EPBs, using Atmosphere Explorer E satellite measurements. Shiokawa et al. 120 (2003b) gave one example of plasma density undulation observed by DMSP 12 F15 at 840 km altitude. Recently Onishi et al. (2009) observed a daytime 122

MSTID with DEMETER satellite (650 km altitude). The MSTID showed 123 quasi-periodic modulation of plasma density and field-aligned flow, and were 124 collocated with similar GPS/TEC variations. Earle et al. (2006) studied top-125 side ionograms from the ISS-b satellite, and constructed for the first time a 126 global MSF statistics. However, the observation period was only from August 12 1978 to December 1980, i.e. near solar maximum, and the global data cover-128 age was poor during June solstice. Su et al. (2006) investigated mid-latitude 129 irregularities of in-situ plasma density as observed by ROCSAT-1 (at 600 130 km altitude). They found that the irregularities generally appear in different 13 longitude sectors in the two solstices. The occurrence rate showed a broad 132 maximum centered at midnight, and was anti-correlated with the solar cycle. 133 Yet, as the orbit inclination angle of ROCSAT-1 was low (40°) , the satellite 134 did not reach the latitude of the irregularities at all longitudes. Saito et al. 135 (1995) and Park et al. (2009a) presented the global distributions of electric 136 and magnetic fluctuations, respectively, associated with MSTIDs. However, 13 their results might be different from true MSTID climatology because other 138 factors than the plasma distribution, such as ionospheric conductivity levels, 139 can affect the electromagnetic signals. 140

In this study we will investigate the plasma density undulations in the topside nighttime midlatitude F-region as observed by three satellites: CHAMP, KOMPSAT-1, and DMSP F15. As each satellite has global coverage at different altitudes, we can give the global climatology of the plasma density undulations at several altitudes as well as the solar-cycle dependence. In Section 2 we briefly describe the instruments and event detection approach. Section 3 shows the global climatology of mid-latitude plasma density undulations, which are discussed in detail in Section 4. Finally, the results aresummarized in Section 5.

150 2. Observations

The Challenging Mini-Satellite Payload (CHAMP) was launched in July 151 2000 with the orbit inclination angle of 87.3° . The precession period of the 152 orbit through 12 hours in local time is about 131 days. Right after the 153 launch its orbit altitude was about 450 km, but slowly decreased to around 154 300 km as of 2009. Onboard CHAMP a Flux-Gate magnetometer (FGM) and 155 a scalar Overhauser Magnetometer (OVM) measure the geomagnetic field, 156 and a Planar Langmuir Probe (PLP) monitors the plasma density. After 157 preprocessing the data rate of FGM and OVM (PLP) is 1s (15s). These 158 three payloads are still in full operation as of June 2009, and all the data are 159 open for public use (http://isdc.gfz-potsdam.de/index.php). 160

The Korea Multi-Purpose Satellite-1 (KOMPSAT-1) was launched in 161 1999. Its orbit is circular at 685 km altitude and sun-synchronous at 1050-162 2250 LT. A scientific payload, Ionospheric Measurement Sensor (IMS), has a 163 Langmuir Probe (LP) and an Electron Temperature Probe (ETP) (Lee et al., 164 2002). The LP measures electron density and temperature with a temporal 165 resolution of 4s. The ETP samples the electron temperature each second. 166 The IMS operated with a duty cycle of about 30% from the end of June 2000 16 to the beginning of August 2001. 168

The Defense Meteorological Satellite Program (DMSP) F15 is in a sunsynchronous orbit (0930-2130 LT) at the altitude of 840 km. It was launched in 1999, and has a Special Sensors-Ions, Electrons, and Scintillation (SSIES) to measure the ion density, temperature, drift, and composition. The nominal
data sample period is 4s, and the duty cycle is almost 100%.

Figure 1(a)-(b) gives an example of mid-latitude plasma density undu-174 lations as observed by CHAMP. In panel (a) the thick solid line represents 175 the plasma density, n, measured by CHAMP/PLP, and the thin dashed line 176 is n_0 , the plasma density low-pass filtered by a median filter ($T_c=52$ s). The 17 undulation amplitude is defined as $|n - n_0|$, which is plotted in panel (b) 178 as a thick solid line. Near the geomagnetic equator (from -10° to $+15^{\circ}$) 179 we can see equatorial plasma bubble (EPBs), whose undulation amplitudes 180 reach well beyond $1 \times 10^5 \ cm^{-3}$. On the other hand, in the southern mid-18 latitude region (from -40° to -25°) there exist quasi-periodic undulations 182 with smaller amplitudes (about $3 \times 10^4 \ cm^{-3}$). 183

Figure 1(c)-(e) shows another example encountered by KOMPSAT-1. In 184 panel (c) the original and low-pass-filtered plasma densities are given in the 185 same format as in panel (a), and panel (d) represents electron tempera-186 ture. Again, panel (e) is the undulation amplitude as in panel (b). Between 18 $20^{\circ} \sim 40^{\circ}$ MLAT the plasma density shows quasi-periodic variations, and 188 the electron temperature is nearly anti-correlated with it. The undulation 189 amplitude of the plasma density is about $1.5 \times 10^4 \ cm^{-3}$ as shown in panel 190 (e). 191

In Figure 1(f)-(g) is an example observed by DMSP F15. In panel (f) the original and low-pass-filtered plasma densities are given in the same format as in panel (a), and panel (g) is the undulation amplitude as panel (b). The plasma density shows quasi-periodic variations between $-40^{\circ} \sim -20^{\circ}$ MLAT.

196

In this study we search for such plasma density undulations automatically

based on the following procedure. First, we gather nighttime satellite passes 19 during magnetically quiet times $(K_p < 3.7)$ and apply a low-pass median fil-198 ter (window size ≈ 100 s) to the plasma density. Second, if the undulation with 199 respect to the filtered data is larger than a certain threshold $(8 \times 10^3 \, cm^{-3})$, 200 it is considered as an event. However, an event should be surrounded by 20 calm background (undulation amplitude $< 4 \times 10^3 \ cm^{-3}$) at each end for 202 at least 30s. In this way adjacent events are tied together to be one large 203 event. Third, to exclude EPBs and subauroral density troughs (e.g. Yizen-204 gaw et al. (2005)), CHAMP events are neglected when their center points 205 are outside $\pm 20^{\circ} \sim 40^{\circ}$ MLAT. Considering the field-aligned geometry of the 206 ionospheric F-region features, the latitude window is shifted equatorward to 20 $\pm 17^{\circ} \sim 37^{\circ}$ MLAT for KOMPSAT-1 and DMSP F15. Events lasting shorter 208 than 60s (about 470 km), which might originate from data outliers, are also 209 discarded. Finally, as CHAMP/FGM continuously measures the geomag-210 netic field strength, we imposed an additional criterion on the CHAMP/PLP 21 events. Stolle et al. (2006) reported that the magnetic field strength shows 212 deflections inside EPBs. In order to exclude EPBs we neglected plasma den-213 sity undulations within which an EPB was detected by an approach similar 214 to that in Stolle et al. (2006) (high-pass filtered/rectified $|\Delta B|$ (T_c=30s) ex-215 ceeds 0.2 nT). For details, the readers are referred to Stolle et al. (2006) and 216 Park et al. (2009b). The mid-latitude plasma irregularities detected by the 21 above-mentioned procedure are marked by a horizontal bar in Figures 1(b), 218 1(e), and 1(g). On the other hand, the region marked by a series of 'A' 210 in Figure 1(e), which also shows electron density/temperature undulations 220 above the threshold between $\pm 17^{\circ} \sim 37^{\circ}$ MLAT, was excluded automatically 22

from our statistics due to its short length. In the subsequent sections we will concentrate on the events identified by the automatic procedure described above.

225 3. Results

Figure 2 shows the seasonal occurrence rate of mid-latitude density undu-226 lations as observed by CHAMP during the period common with KOMPSAT-227 1/LP: 28 June 2000 - 31 July 2001 (hereafter 'KOMPSAT-1/LP lifetime' for 228 simplicity). Within each geographic bin $(5^{\circ} \text{ in latitude by } 10^{\circ} \text{ in longitude})$ 229 the occurrence rate is calculated as the ratio between the number of detected 230 events to the number of satellite passes over this bin. The final global im-231 age is filtered by a 3-by-3 2D median filter so that outliers are suppressed. 232 Note that events whose center point lies outside $\pm 20^{\circ} \sim 40^{\circ}$ MLAT were au-233 tomatically neglected. Global occurrence patterns in each season are given 234 in panels (a)-(c), and panel (d) shows the occurrence rate as a function of 235 invariant latitude versus local time. 230

First, it is noted that the occurrence rate of undulations are lowest during equinoxes. Second, the occurrence maximum during June solstice is in the Asian/Oceanian region, while it is in the eastern Pacific/American region during December solstice. In general, the occurrence rate is higher in the summer hemisphere. Third, the undulations are spread over the night, as can be seen in panel (d), with a weak maximum near midnight.

Figure 3 has the same format as Figure 2, but it is based on KOMPSAT-1/LP observations. From panels (a)-(c) we can see again that the undulations are rare during the equinoxes, and that the occurrence peaks in the

Asian/Oceanian (eastern Pacific/American) region during June (December) 246 solstice. The occurrence rate during June solstice is higher in the summer 24 (i.e. northern) hemisphere. While the same trend is expected for December 248 solstice, the observed occurrence rate is not higher in the summer (i.e. south-240 ern) hemisphere. However, we should keep in mind that during December 250 solstice the operation of KOMPSAT-1/LP in the southern hemisphere was 25 generally limited to equatorward of -30° GLAT (see Figure 2 of Kim et al. 252 (2006) for reference). In panel (d) the event occurrence is confined to 22-23 253 LT due to the sun-synchronous orbit characteristic of KOMPSAT-1. 254

The same detection procedure was applied to DMSP F15 data during the KOMPSAT-1/LP lifetime, and the result is shown in Figure 4. Similar to Figures 2 and 3, the occurrence rate of the events is lowest during equinoxes, and has a weak local maximum in the Asian/Oceanian (eastern Pacific/American) region around June (December) solstice, with the peak in the summer hemisphere. Generally, the occurrence rate is lower than at the KOMPSAT-1 altitude.

Figure 5 shows the event distribution obtained from CHAMP data, but 262 this time during the two-year solar minimum period, 2006 and 2007. During 263 that time much more events are detected. Again, the undulation occurrence 264 rate attains its minimum during equinoxes. During June solstice the occur-265 rence maximum is in the Asian/Oceanian region and the minimum in the 266 eastern Pacific, southern Atlantic, and southern African region as in Figure 26 2(b). The occurrence peak during December solstice is spread across the 268 southern Pacific Ocean with the maximum above the American continent. 269 Just like for solar maximum (Figure 2), the undulations occur through all the 270

night, as seen in Figure 4(d), with a broad maximum shortly before midnight.

272 4. Discussions

273 4.1. statistical properties of nighttime topside mid-latitude plasma density 274 undulations

In Section 3 we described the statistical distribution of mid-latitude plasma 275 density undulations. During solar maximum years, CHAMP (around 400 276 km altitude), KOMPSAT-1 (around 680 km), and DMSP F15 (around 840 277 km) found similar longitudinal distribution patterns: an occurrence maxi-278 mum in Asia/Oceania (eastern Pacific/American) during June (December) 279 solstice and globally low occurrence rate during equinoxes. Unfortunately, a 280 height-dependent relationship between the occurrence rates at CHAMP and 28 KOMPSAT-1 altitudes is not clear from Figures 2 and 3, possibly due to 282 different data rates: CHAMP/PLP (15s) and KOMPSAT-1/LP (4s). DMSP 283 F15 encountered mid-latitude plasma density undulations less frequently 284 than KOMPSAT-1. As the data rate of DMSP/SSIES is also 4s, we can 28 conclude that the occurrence rate is higher at 680 km than at 840 km. The 286 result is compatible with *Livneh et al.* (2009). They found plasma density un-28 dulations up to 750 km altitude using mid-latitude incoherent scatter radars, 288 while no undulation was observed by the DMSP ion drift meter. 289

The undulations occur through all night at CHAMP altitude (around 400 km), and the occurrence rate is anti-correlated with the solar cycle. Also, the occurrence rate shows strong hemispheric asymmetry at CHAMP and KOMPSAT-1 altitude, i.e. preference for the summer hemisphere. The reason might be as follows. MSTIDs can be described as alternating bands which show upward/downward displacement of the F-layer (e.g. see *Tsunoda* (2006)). Let us assume a plasma instability with growth rate γ , and resultant altitude change of ionospheric F-layer, Δz , is given from Eq. (32) in *Tsunoda* (2006) as:

$$\Delta z = z - z_0 = z_0 (e^{\gamma \Delta t} - 1) \tag{2}$$

where $z(z_0)$ is perturbed (unperturbed) altitude, and Δt is elapsed time from the instability initiation. This altitude change leads to density undulation:

$$\Delta n = n - n_0 = n_0 \left(e^{\Delta z / H_{plasma}} - 1 \right) = n_0 \left(e^{\frac{z_0}{H_{plasma}} \left(e^{\gamma \Delta t} - 1 \right)} - 1 \right)$$
(3)

where n_0 is the ambient density and H_{plasma} the plasma scale height. The 30 growth rate γ is expected to be the same in both hemispheres because the 302 effect of an E_s -layer instability, mapped along the geomagnetic field, (rather 303 than the Perkins instability) dominates the initial-stage growth of F-layer 304 undulations (Yokoyama et al., 2009). Stankov and Jakowski (2006) found 305 from CHAMP occultation data that the plasma scale height at the topside 306 (near CHAMP altitude) nighttime mid-latitude ionosphere is slightly lower 307 in the summer than in the winter hemisphere. Therefore, in the topside 308 summer hemisphere a larger background density, n_0 , together with a slightly 309 lower plasma scale height can produce larger-amplitude undulations. 310

Eq. (3) can also interpret the solar cycle dependence of plasma density undulations. First, scale height does not vary much with solar activity in topside (near CHAMP altitude) nighttime mid-latitude F-region (*Stankov* and Jakowski, 2006). Second, though n_0 decreases and γ increases with decreasing solar cycle (e.g. *Tsunoda* (2006)), the effect of γ , which is doubly exponential, dominates Δn . Hence, the resultant Δn is expected to become larger during solar-minimum years than during solar maximum, as confirmed
by our observations shown in Figures 2 and 4.

In this paper we used a fixed threshold $(8 \times 10^3 \text{ cm}^{-3})$ to identify mid-319 latitude plasma density undulations. In order to assess the reliability of 320 Figures 2-5 we also tested the relative undulation amplitude $(|n - n_0|/n_0)$ for 321 the event identification (threshold: undulation amplitude 8% of the ambi-322 ent plasma density). The obtained results are generally compatible with the 323 finding of Figures 2-5 (figures not shown). First, the event occurrence rate is 324 generally higher at KOMPSAT-1 (680 km) than at DMSP F15 (840 km) alti-325 tude. Second, for all the three satellites the occurrence rate is lowest during 326 equinoxes, and maximizes in the Asian/Oceanian (eastern Pacific/American) 327 longitude sector during June (December) solstice. Third, at CHAMP altitude 328 the density undulations occur through almost all night, but with a bias to-329 wards the post-midnight sector. Fourth, the occurrence rate is higher during 330 solar minimum years than during solar maximum. However, the hemispheric 331 distribution patterns show notable difference from Figures 2-5. The pref-332 erence for the summer hemisphere shown in Figures 2-5 disappears when 333 relative undulation amplitude is used for the event identification. During 334 solar maximum CHAMP events generally appear in the winter hemisphere; 335 the occurrence rates for KOMPSAT-1 (DMSP F15) are slightly higher in 336 the northern (southern) hemispheres, respectively. During solar minimum 337 the occurrence rate at CHAMP altitude is slightly higher in the southern 338 hemisphere. 339

Eq. (3) predicts that the relative undulation amplitudes $(|n - n_0|/n_0)$ show hemispheric symmetry if the difference in plasma scale height, H_{plasma} , can be neglected (e.g. Stankov and Jakowski (2006)). However, Eq. (3)
cannot explain why relative density variations of CHAMP, KOMPSAT-1,
and DMSP F15 data do not exhibit perfect hemispheric symmetry. It is
probably related to the seasonal (hemispheric) difference of primary seeding,
which originates from the neutral atmosphere. But, further investigation is
needed to prove this suggestion.

348 4.2. comparison with other satellite observations

We want to compare the climatology given in Figures 2-4 with previ-349 ous satellite observations of topside mid-latitude plasma density structures. 350 In Su et al. (2006) the mid-latitude plasma density irregularities observed 35 by ROCSAT-1 in-situ measurements occur mainly in the Asian/Oceanian 352 (American) longitude region during June (December) solstice. The occur-353 rence rate showed a broad maximum around midnight, and was anti-correlated 354 with the solar cycle. All these results are consistent with ours. Yet, the num-35! ber of mid-latitude plasma density irregularities encountered during equinoxes 356 by ROCSAT-1 was comparable to that during solstices, which is slightly dif-357 ferent from the trends shown in this paper. Mid-latitude undulations ob-358 served by ROCSAT-1 show no preference of the summer hemisphere, as seen 359 in our Figure 2. The differences might be due to the different detection ap-360 proach (i.e. relative plasma density variation) applied to ROCSAT-1 data. 36

According to *Earle et al.* (2006) the occurrence rate of topside MSFs, which originate from topside plasma density structures, was minimum during the spring equinox (February-April) as observed by the ISS-b satellite. The longitudinal occurrence distribution peaked in the Asian (north American) region from August to October (from November to January). Although the seasonal division is slightly different from that in this study, the results generally agree with those given in Section 3. It is noteworthy in their study that both the August-October and November-January periods show the longitudinal peaks in the northern hemisphere.

4.3. comparison with other ground observations

Next, let us compare our results with previous ground-based observa-372 tions, which measure MSTIDs (i.e. plasma density structures) seen at the 373 bottomside (by airglow imaging) or around F-region peak (by GPS/TEC 374 measurements) in mid-latitude region. From airglow observation Garcia et 375 al. (2000) reported that MSTID occurrence at Arecibo, Puerto Rico peaks 376 during December solstice, which supports our results. The MSTID occur-377 rence rate in the Asian/Oceanian longitude sector, estimated from airglow 378 images, is maximum during June solstice (*Shiokawa et al.*, 2003a), which is 379 also consistent with our results. By interpreting GPS/TEC data Kotake et 380 al. (2006) also confirmed that relative nighttime MSTID amplitudes in the 381 Asian/Oceanian and the western U. S. longitude sector are largest during 382 June solstice. Only, in South America MSTID amplitudes were largest dur-383 ing June solstice, which is inconsistent with our result. From the similarities 384 in the distribution patterns, as shown above, we suggest that topside mid-385 latitude plasma density undulations addressed in this paper are a topside 386 signature (or extension) of MSTIDs. 38

$_{388}$ 4.4. comparison with E_s layers

The relationship between mid-latitude density undulations and E_s layers might be important because MSTID generation is promoted by E_s layers

(see the theoretical works of *Tsunoda and Cosgrove* (2001); *Tsunoda* (2006); 391 Yokoyama et al. (2009)). Otsuka et al. (2008) showed a good correlation 392 between the observational climatologies of MSTIDs and E_s in East Asia, 393 but there has been no such investigation on a global scale. Using CHAMP, 394 GRACE-A, and FORMOSAT-3/COSMIC Arras et al. (2008) presented a 395 global E_s climatology for the years 2006 and 2007 (see their Figures 3 and 396 4). Their Figure 3 shows that (1) E_s occurrence is lowest during equinoxes, 39 (2) the occurrence prefers the summer hemisphere, (3) during June solstice 398 the occurrence is lowest in the American and eastern Pacific region, (4) during 399 December solstice the occurrence is lowest between the Atlantic and western 400 Indian Ocean, and (5) there are practically no E_s layers within the South 40 Atlantic Anomaly (SAA). All of these properties are in remarkable agreement 402 with our results at the topside ionospheric F-layer, corroborating the close 403 relationship between the density undulations (or MSTIDs) and E_s layers as 404 suggested by the theoretical work of Yokoyama et al. (2009). 405

406 4.5. relationship to midlatitude magnetic fluctuations (MMFs)

Now, we want to discuss our results in relation with mid-latitude mag-407 netic fluctuations (MMFs). Park et al. (2009a) showed MMF statistics 408 obtained by CHAMP/FGM measurements, and concluded that MMFs are 409 magnetic signatures of MSTIDs. The argument was supported by the sea-410 sonal/longitudinal distribution and linear polarization of MMFs. The fea-41 tures in Figure 2 and 4 are generally compatible with Figure 2 and 3 of 412 Park et al. (2009a): few events in equinoxes and occurrence maxima over 413 Asia/Oceania (eastern Pacific Ocean/Americas) during June solstice (De-414 cember solstice). However, several differences are notable. First, the mid-415

latitude plasma density undulations occur throughout the night while MMFs 416 predominantly occur in the premidnight sector. Second, the occurrence rate 41 of the plasma density undulations is anti-correlated with the solar activity 418 while MMF occurrence has a clear positive correlation. Park et al. (2009a) 419 argued that not only MSTIDs but also ambient conditions such as ionospheric 420 conductivity seem to influence the MMF generation. As the ionospheric con-42 ductivity constantly decreases through the night, MMF generation peaks in 422 the premidnight sector. Also, the conductivity increases with the solar cycle, 423 and promotes MMF generation during the solar activity. 424

We will investigate the relationship between the plasma density undula-425 tions and MMFs in more details. Figure 5 presents an early-night example 426 of the mid-latitude plasma density undulations. Panel (a) shows original 427 and filtered plasma density as in Figure 1(a), panel (b) fluctuations of the 428 zonal magnetic field measured by CHAMP/FGM, panel (c) fluctuations of 429 the meridional component, and panel (d) presents field-aligned current den-430 sity calculated from the magnetic field deviations (*Park et al.*, 2009b). The 431 plasma undulation event detected by our automatic procedure is marked 432 by a series of circles in panel (a). Within the event periods we can see 433 quasi-periodic plasma density undulations collocated with linearly polarized 434 magnetic fluctuations in the plane perpendicular to the main geomagnetic 435 field. The correlation coefficient between zonal and meridional fluctuations, 436 as well as the polarization angle (measured counter-clockwise from the mag-43 netic east direction), is noted in panel (d). Because of the good linear polar-438 ization we can calculate reliably the field-aligned current density by Ampère's 439 law, which is plotted in panel (d) (see *Park et al.* (2009a,b) for details). We 440

note that the plasma density and magnetic signatures exhibit hemispheric
conjugacy. In panel (d) large positive (field-aligned) peaks of FAC density
roughly correspond to the plasma density depletion (enhancement) regions
in the northern (southern) hemisphere.

All the mid-latitude plasma density undulations detected by CHAMP/PLP 445 are reexamined to check whether they are accompanied with MMFs. An 446 event is identified as collocated with MMFs if filtered/rectified ($T_c=10s$) per-44 pendicular B-field deflections exceed 0.45 nT within the event location (1° 448 margin on both sides) or at the conjugate point in the opposite hemisphere 449 (see Park et al. (2009a)). Altogether 683 events were identified in CHAMP 450 data as mid-latitude plasma density undulations during KOMPSAT-1/LP 451 lifetime. Among them 425 events are collocated with MMFs, 222 events 452 were not, and for 36 events appropriate CHAMP/FGM data are missing. 453 On average about 66% of the plasma undulations are collocated with MMFs. 454 MSTIDs, whose scale sizes are several hundred kilometers, are not always 455 collocated with other irregularities of smaller scale sizes, such as FAI (Saito 456 et al., 2001) or MSF (Shiokawa et al., 2003a). Note that a high-pass filter 45 was used here to identify MMFs ($T_c=10$ s, i.e. scale length less than 80 km) 458 based on the studies of Saito et al. (1995) and Park et al. (2009a). Moreover, 459 MMF amplitude is affected by ionospheric conductivity Park et al. (2009a). 460 Hence, (1) the lack of secondary instabilities generating smaller-scale irregu-46 larities on MSTIDs and (2) day-to-day variability of the conductivity might 462 have compromised a one-to-one correspondence between plasma density un-463 dulations and MMFs. This point is worth further investigation with the 464 upcoming Swarm mission, which will be equipped with higher-resolution 46!

⁴⁶⁶ plasma instrument as well as an ion drift meter by which the E-field can be⁴⁶⁷ estimated.

468 5. Summary

In this study we investigated plasma density undulations in the nighttime mid-latitude F-region as observed by CHAMP, KOMPSAT-1, and DMSP F15. In this way the first 'full global' climatology of topside mid-latitude undulations is given at different altitudes and different solar phases. From the results obtained we come to the following conclusions.

- In the nighttime mid-latitude ionosphere during the solar maximum,
 plasma density undulations (scale size < 400 km) are found at CHAMP
 (400 km), KOMPSAT-1 (680 km), DMSP F15 (840 km) altitudes.
 The event occurrence rate is higher at KOMPSAT-1 (680 km) than
 at DMSP F15 (840 km) altitude.
- 479
 2. Mid-latitude plasma density undulations occur infrequently during equinoxes.
 480 Their occurrence peaks are in the Asian/Oceanian (eastern Pacific/American)
 481 region during June (December) solstice.
- 482 3. At CHAMP altitude (400 km) the undulations are observed all through 483 the night, and the occurrence rate is anti-correlated with the solar cycle. 484 All these results are in general agreement with known MSTID and E_s 485 climatology.
- 486 4. The undulations are frequently but not always accompanied by magnetic signatures indicating the presence of field-aligned current. The
 487 partial lack in correspondence might be due to the scale-size difference
 489 in detection algorithms for the two phenomena, but more probably due

to the effect of ionospheric conductivity on FAC density. The quite similar distribution is, however, in support of a connection between density undulations and FACs.

Acknowledgements: The CHAMP mission is supported by the German Aerospace Center (DLR) in operation and by the Federal Ministry of Education (BMBF), as part of the Geotechnology Programme, in data processing. The authors gratefully acknowledge the Center for Space Sciences at the University of Texas at Dallas and the US Air Force for providing the DMSP thermal plasma data.

References

490

491

492

- Arras, C., J. Wickert, G. Beyerle, S. Heise, T. Schmidt, and C. Jacobi (2008), A global climatology of ionospheric irregularities derived from GPS radio occultation, *Geophys. Res. Lett.*, 35, L14809, doi:10.1029/2008GL034158.
- Behnke, R. (1979), F layer height Bands in the Nocturnal Ionosphere over Arecibo, J. Geophys. Res., 84, 974.
- Bowman, G. G. (1992), Upper atmosphere neutral-particle density variations compared with spread-F occurrence rates at locations around the world, Ann. Geo., 10, 676.
- Candido, C. M. N., A. A. Pimenta, J. A. Bittencourt, and F. Becker-Guedes (2008), Statistical analysis of the occurrence of medium-scale traveling ionospheric disturbances over Brazilian low latitudes using OI 630.0 nm emission all-sky images, *Geophys. Res. Lett.*, 35, L17105, doi:10.1029/2008GL035043.

- Dyson, P. L (1968), Topside spread F at midlatitudes, J. Geophys. Res., 73, 2441.
- Earle G. D., A. M. Musumba, and J. P. McClure (2006), A global study of nighttime midlatitude topside spread echoes, J. Geophys. Res., 111, A11306, doi:10.1029/2006JA011614.
- Fukao, S., M. C. Kelley, T. Shirakawa, T. Takami, M. Yamamoto, T. Tsuda, and S. Kato (1991), Turbulent Upwelling of the Mid-Latitude Ionosphere,
 1. Observational Results by the MU Radar, J. Geophys. Res., 96(A3), 3725.
- Garcia, F. J., M. C. Kelley, J. J. Makela, and C. S. Huang (2000), Airglow observations of mesoscale low-velocity traveling ionospheric disturbances at midlatitudes, *J. Geophys. Res.*, 105, 18407.
- Hanson, W. B., and F. S. Johnson (1992), Lower midlatitude ionospheric disturbances and the Perkins instability, *Planet. Space Sci.*, 40, 1615.
- Kelley, M. C., and J. J. Makela (2001), Resolution of the discrepancy between experiment and theory of midlatitude Fregion structures, Turbulent upwelling of the mid-latitude ionosphere, 2. Theoretical framework, *Geophys. Res. Lett.*, 28(13), 2589.
- Kim, H, K. Min, J. Park, J. Lee, E. Lee, H. Kil, V. P. Kim, and S. Park (2006), Comparison of satellite measurements of the low-latitude nighttime upper ionosphere with IRI, J. Atmos. Solar Terr. Phys., 68, 2107.
- Kotake N., Y. Otsuka, T. Tsugawa, T. Ogawa, and A. Saito (2006), Climatological study of GPS total electron content variations caused by medium-

scale traveling ionospheric disturbances, J. Geophys. Res., 111, A04306, doi:10.1029/2005JA011418.

- Kotake, N., Y. Otsuka, T. Ogawa, T. Tsugawa, and A. Saito (2007), Statistical study of medium-scale traveling ionospheric disturbances observed with the GPS networks in Southern California, *Earth Planets Space*, 59, 95.
- Lee, J. J., K. W. Min, V. P. Kim, V. V. Hegai, K.-I. Oyama, F. J. Rich, and J. Kim (2002), Large density depletions in the nighttime upper ionosphere during the magnetic storm of July 15, 2000, *Geophys. Res. Lett.*, 29(3), 1032, doi:10.1029/2001GL013991.
- Lee, C. C., Y. A. Liou, Y. Otsuka, F. D. Chu, T. K. Yeh, K. Hoshinoo, and K. Matunaga (2008), Nighttime medium-scale traveling ionospheric disturbances detected by network GPS receivers in Taiwan, J. Geophys. Res., 113, A12316, doi:10.1029/2008JA013250.
- Livneh, D. J., I. Seker, F. T. Djuth, and J. D. Mathews (2009), Omnipresent vertically coherent fluctuations in the ionosphere with a possible worldwide-midlatitude extent, J. Geophys. Res., 114, A06303, doi:10.1029/2008JA013999.
- Maruyama T., and T. Matuura (1984), Longitudinal variability of annual changes in activity of equatorial spread-F and plasma bubbles, J. Geophys. Res., 89, 10903.
- Munro, G. H. (1963), Scintillation of radio signals from satellites, J. Geophys. Res., 68, 1851.

- Onishi, T., T. Tsugawa, Y. Otsuka, J.-J. Berthelier, and J.-P. Lebreton (2009), First simultaneous observations of daytime MSTIDs over North America using GPS-TEC and DEMETER satellite data, *Geophys. Res. Lett.*, 36, L11808, doi:10.1029/2009GL038156.
- Otsuka, Y., K. Shiokawa, T. Ogawa, and P. Wilkinson (2004), Geomagnetic conjugate observations of medium-scale traveling ionospheric disturbances at midlatitude using all-sky airglow imagers, *Geophys. Res. Lett.*, 31, L15803, doi:10.1029/2004GL020262.
- Otsuka, Y., T. Tani, T. Tsugawa, T. Ogawa, A. Saito (2008), Statistical study of relationship between medium-scale traveling ionospheric disturbance and sporadic E layer activities in summer night over Japan, J. Atmos. Solar Terr. Phys., 70, 2196.
- Otsuka, Y., K. Shiokawa, T. Ogawa, T. Yokoyama, and M. Yamamoto (2009), Spatial relationship of nighttime medium-scale traveling ionospheric disturbances and F region field-aligned irregularities observed with two spaced all-sky airglow imagers and the middle and upper atmosphere radar, J. Geophys. Res., 114, A05302, doi:10.1029/2008JA013902.
- Park, J., H. Lühr, C. Stolle, M. Rother, K. W. Min, J.-K. Chung, Y. H. Kim, I. Michaelis, and M. Noja (2009a), Magnetic signatures of medium-scale traveling ionospheric disturbances as observed by CHAMP, *J. Geophys. Res.*, 114, A03307, doi:10.1029/2008JA013792.
- Park, J., H. Lühr, C. Stolle, M. Rother, K. W. Min, and I. Michaelis, (2009b), The characteristics of field-aligned currents associated with equa-

torial plasma bubbles as observed by the CHAMP satellite, *submitted to* Ann. Geo..

- Perkins, F. (1973), Spread F and ionospheric currents, J. Geophys. Res., 78, 218.
- Peterson, A. M, O. G. Villard, Jr., Rl. L. Leadabrand, and P. B. Gallagher (1955), Regularly-observable aspect-sensitive radio reflections from ionization aligned with the earth's magnetic field and located within the ionospheric layers at middle latitudes, J. Geophys. Res., 60, 497.
- Saito, A., T. Iyemori, M. Sugiura, N. C. Maynard, T. L. Aggson, L. H. Brace, M. Takeda, and M. Yamamoto (1995), Conjugate Occurrence of the Electric Field Fluctuations in the Nighttime Midlatitude Ionosphere, J. Geophys. Res., 100(A11), 21439.
- Saito, A., T. Iyemori, L. G. Blomberg, M. Yamamoto, and M. Takeda (1998), Conjugate observations of the mid-latitude electric field fluctuations with the MU radar and the Freja satellite, J. Atmos. Solar Terr. Phys., 60, 129.
- Saito, A., M. Nishimura, M. Yamamoto, S. Fukao, M. Kubota, K. Shiokawa, Y. Otsuka, T. Tsugawa, T. Ogawa, M. Ishii, T. Sakanoi, and S. Miyazaki (2001), Traveling Ionospheric Disturbances Detected in the FRONT Campaign, *Geophys. Res. Lett.*, 28(4), 689.
- Seker, I., D. J. Livneh, and J. D. Mathews (2009), A 3-D empirical model of F region Medium-Scale Traveling Ionospheric Disturbance bands using incoherent scatter radar and all-sky imaging at Arecibo, J. Geophys. Res., 114, A06302, doi:10.1029/2008JA014019.

- Shiokawa, K., C. Ihara, Y. Otsuka, and T. Ogawa (2003a), Statistical study of nighttime medium-scale traveling ionospheric disturbances using midlatitude airglow images, J. Geophys. Res., 108, 1052, doi:10.1029/2002JA009491.
- Shiokawa, K., Y. Otsuka, C. Ihara, T. Ogawa, and F. J. Rich (2003b), Ground and satellite observations of nighttime medium-scale traveling ionospheric disturbance at midlatitude, *J. Geophys. Res.*, 108, 1145, doi:10.1029/2002JA009639.
- Shiokawa, K., Y. Otsuka, T. Tsugawa, T. Ogawa, A. Saito, K. Ohshima, M. Kubota, T. Maruyama, T. Nakamura, M. Yamamoto, and P. Wilkinson (2005), Geomagnetic conjugate observation of nighttime medium- and large-scale traveling ionospheric disturbances: FRONT3 campaign, J. Geophys. Res., 110(A5), A05303, doi:10.1029/2004JA010845.
- Stankov, S. M., and N. Jakowski (2006), Topside ionospheric scale height analysis and modelling based on radio occultation measurements, J. Atmos. Solar Terr. Phys., 68, 134.
- Stolle, C., H. Lühr, M. Rother, and G. Balasis (2006), Magnetic signatures of equatorial spread F as observed by the CHAMP satellite, J. Geophys. Res., 111, A02304, doi:1029/2005JA011184.
- Su, S.-Y., C. H. Liu, H. H. Ho, and C. K. Chao (2006), Distribution characteristics of topside ionospheric density irregularities: Equatorial versus midlatitude regions, *J. Geophys. Res.*, 111, A06305, doi:10.1029/2005JA011330.

- Tsunoda, R. T., and R. B. Cosgrove (2001), Coupled Electrodynamics in the Nighttime Midlatitude Ionosphere, *Geophys. Res. Lett.*, 28, 4171.
- Tsunoda, Roland T. (2006), On the coupling of layer instabilities in the nighttime midlatitude ionosphere, J. Geophys. Res., 111, A11304, doi:10.1029/2006JA011630.
- Yizengaw E., H. Wei, M. B. Moldwin, D. Galvan, L. Mandrake, A. Mannucci, and X. Pi (2005), The correlation between mid-latitude trough and the plasmapause, *Geophys. Res. Lett.*, 32, L10102, doi:10.1029/2005GL022954.
- Yokoyama, T., D. L. Hysell, Y. Otsuka, and M. Yamamoto (2009), Threedimensional simulation of the coupled Perkins and Es-layer instabilities in the nighttime midlatitude ionosphere, J. Geophys. Res., 114, A03308, doi:10.1029/2008JA013789.

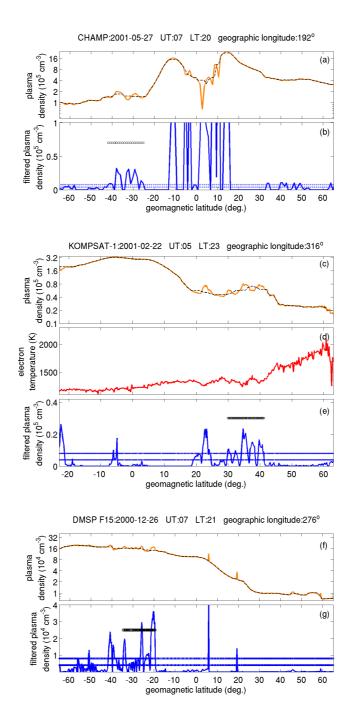


Figure 1: (a) plasma density measured by CHAMP/PLP (thick solid line) and the plasma density low-pass filtered by a median filter (thin dotted line), (b) undulation of the plasma density with respect to the low-pass filtered density, (c) plasma density measured by KOMPSAT-1/LP (thick solid line), and the plasma density low-pass filtered by a median filter (thin dotted line), (d) electron temperature measured by KOMPSAT-1/LP, (e) density undulation calculated as in panel (b), (f) plasma density measured by DMSP/SSIES (thick solid line), and the plasma density low-pass filtered by a median filter (thin dotted line), and the plasma density low-pass filtered by a median filter (thin dotted line), and the plasma density low-pass filtered by a median filter (thin dotted line), and the plasma density low-pass filtered by a median filter (thin dotted line), and the plasma density low-pass filtered by a median filter (thin dotted line), and the plasma density low-pass filtered by a median filter (thin dotted line), and the plasma density low-pass filtered by a median filter (thin dotted line), and (g) density undulation calculated as in panel (b).

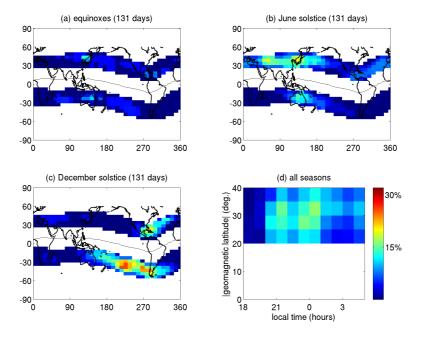


Figure 2: Seasonal/global occurrence rate of mid-latitude density undulations as observed by CHAMP during the KOMPSAT-1/LP lifetime. The distribution is presented for each season centered on: (a) equinoxes, (b) June solstice, and (c) December solstice. Panel (d) shows the occurrence rate as a function of invariant latitude versus local time.

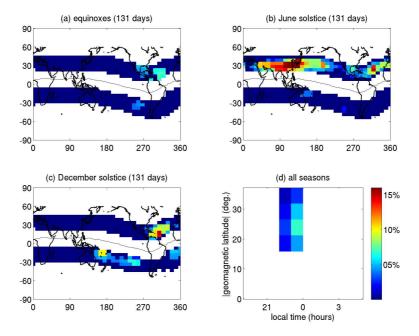


Figure 3: Same as Figure 2, but as observed by KOMPSAT-1/LP. Notice the difference in color scale.

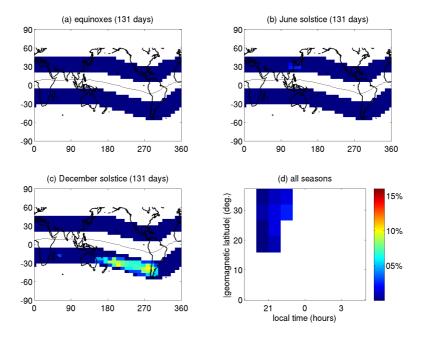


Figure 4: Same as Figure 2, but as observed by DMSP F15. Notice the difference in color scale.

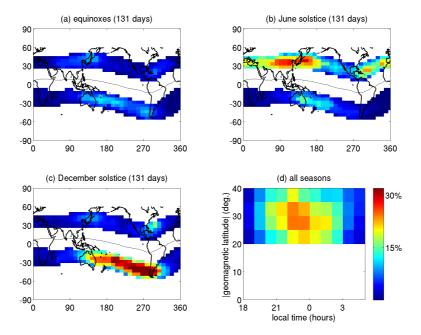


Figure 5: Same as Figure 2, but as observed by CHAMP during the years 2006 and 2007.

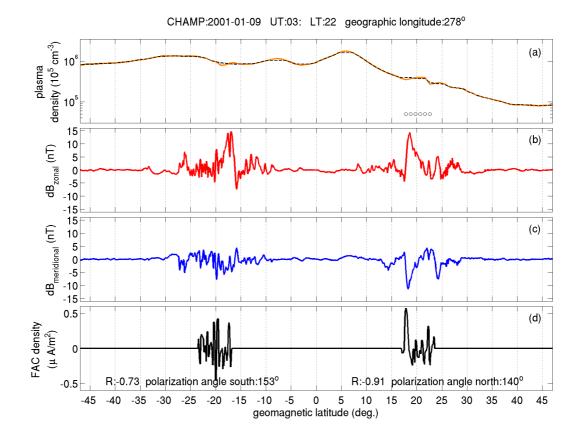


Figure 6: Comparison of density undulations with magnetic signatures. (a) plasma density measured by CHAMP/PLP (thick solid line), and the density low-pass filtered by a median filter (thin dotted line), (b) fluctuation of zonal magnetic field measured by CHAMP/FGM, (c) fluctuation of meridional comoponent, and (d) field-aligned current density calculated from the magnetic fluctuations.