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**The IRI-2007 model overestimates electron density during the 23/24 solar minimum**

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**Abstract.** We compare electron density predictions of the International Reference Ionosphere (IRI-2007) model with in-situ measurements of the satellites CHAMP and GRACE for the years 2000-2009. Orbital-averages of the electron density are considered. During the first half of the period (2000-2004) measurements and collocated model predictions track each other reasonably well at both sampling heights. From 2005 onward the overestimation of the electron density by the model is progressively increasing. Annual averages show that IRI-2007 values are too high by 50% for 2008 and by more than 60% by 2009. An inspection of the latitudinal and local time distributions reveals that the too high predictions primarily occur at low latitudes during daytime hours. From comparison with observations it becomes obvious that IRI-2007 is strongly overestimating the equatorial ion fountain effect during the last deep solar minimum.

**1. Introduction**

The ionospheric electron density is a highly variable quantity. Significant changes are observed on various time scales. Among these the solar cycle is an important variation. At fixed height the electron density is modulated by the expansion of the ionosphere (change in F2 peak height) but also by the variation of the electron concentration. Both these effects

30 vary with the irradiance of solar extreme ultraviolet (EUV). For instance, the solar cycle-  
31 dependence of the maximum F region electron density (NmF2) and the related critical  
32 frequency (foF2) have been studied among others by L. Liu et al. [2006]. They report an  
33 increase of the average NmF2 value from  $3 \cdot 10^{11}$  to  $2 \cdot 10^{12} \text{ m}^{-3}$  for solar fluxes  $F_{10.7} = 70$  to  
34 200 sfu at low latitudes. Similarly, the F region peak height (hmF2) is changing with the solar  
35 cycle. According to L. Liu et al. [2006] hmF2 is increasing from about 260 to 330 km for  
36  $F_{10.7} = 70$  to 200 sfu at the station Wuhan ( $19.5^\circ$  MLat) during noon time. The total electron  
37 content (TEC) variation over the course of a solar cycle has been investigated, for example,  
38 by Huang and Cheng [1995]. They find a close relation between electron content and sunspot  
39 number, R. Around the ionisation anomaly crest TEC values increase from 30 to 110 TECU  
40 for  $R = 20$  to 140 at 16:00 local time in the East Asian region.

41 Ionospheric models, such as the International Reference Ionosphere (IRI) [e.g., Bilitza,  
42 1992, 2003], are expected to reflect these effects. However, there is evidence that some  
43 important parameters (NmF2 and hmF2) deviate significantly, in particular at low and  
44 equatorial latitudes, from observations [e.g., Adeniyi et al., 2003; Obrou et al., 2003; Souza et  
45 al., 2003]. This can strongly influence the electron density,  $N_e$ , distribution at a certain height  
46 level. H. Liu et al [2007] have used CHAMP electron density data at 400 km altitude over the  
47 period Aug. 2000 to Aug. 2006 for validating the IRI-2001 model [Bilitza, 2003]. Particular  
48 emphasis was put on the equatorial ionization anomaly (EIA) during daytime and post-sunset  
49 hours. They conclude that IRI reproduces the EIA around noon rather well for all levels of  
50 solar flux. Conversely, the ion fountain effect around 20:00 LT is too weak in the model, in  
51 particular for moderate to strong solar activity. As can be seen in Figure 6 of H. Liu et al.  
52 [2007], the observed  $N_e$  profiles across the EIA exhibit higher crest values, which appear at  
53 larger distances to the dip-equator, and  $N_e$  is more depleted above the equator. All this  
54 indicates that IRI-2001 underestimates the effect of the vertical plasma drift during post-  
55 sunset hours.

56           The IRI is a purely empirical model based on a large collection of satellite and ground-  
57 based observations, and it is expected to give a reasonably accurate description of the  
58 ionosphere under quiet geomagnetic conditions. However, the minimum of the solar cycle  
59 23/24 was quite special since it had a record number of days without sun spots [Livingston  
60 and Penn, 2009]. There is no representative dataset included in IRI that comes from a  
61 comparable solar minimum. Therefore the model may not be able to predict the ionospheric  
62 conditions correctly during the years 2008 and 2009.

63           The purpose of this study is thus to test the reliability of IRI, in particular during the  
64 recent deep solar minimum. For a direct comparison we make use of measurements from the  
65 satellites CHAMP and GRACE. They provide electron density readings over the past decade  
66 from two altitudes. This dataset allows for a direct comparison between IRI predictions and  
67 observations over the whole range from solar maximum to minimum. Based on these results  
68 we perform an assessment of the differences.

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## 71           **2. Data**

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73           The CHAMP satellite was launched on 15 July 2000 into a circular, near-polar orbit  
74 (inclination:  $87.3^\circ$ ) with an initial altitude of 456 km. By the end of 2009 it has decayed to  
75 310 km. The local time of the orbit changes by 1 hour in 11 day, requiring about 130 day for  
76 covering all local times [Reigber et al., 2002]. The Planar Langmuir Probe (PLP) on board the  
77 satellite takes in-situ measurements of the electron density every 15 s. The  $N_e$  readings of the  
78 PLP have been verified by comparison against digisonde measurements at Jicamarca  
79 [McNamara et al., 2007]. The authors report an average discrepancy between PLP and  
80 digisonde recordings of only 4% with a standard deviation of 8.8%. This good agreement  
81 adds confidence in the reliability of the CHAMP  $N_e$  measurements.

82 GRACE, comprising two spacecraft GRACE-A and GRACE-B, was launched on 17  
83 March 2002 into a near-circular, polar orbit (inclination: 89°) with an initial altitude of about  
84 490 km. The altitude of these satellites is quite stable over the years. An overview of the  
85 orbital evolution of CHAMP and GRACE is given in Figure 2 of Xiong et al. [2010]. The  
86 local time of the orbital plane precesses by 4.5 minutes every day taking the mission 160.5  
87 days to sample all local times [Tapley et al., 2004].

88 The two GRACE satellites follow each other at a distance of about 220 km. For the  
89 determination of the electron density we make use of the K-band ranging (KBR) system,  
90 which measures the dual one-way range change between the two satellites. The total electron  
91 content (TEC) between the spacecraft can be deduced from the KBR data. When dividing the  
92 horizontal TEC by the distance between the spacecraft we get the average electron density. A  
93 more detailed description of the electron density retrieval is given in section 3 of Xiong et al.  
94 [2010]. TEC measurements based on radio wave modifications do in principle not require  
95 calibration. There is, however, the uncertainty of an arbitrary bias value. By considering long  
96 and continuous data series we think, we have constrained the applied bias quite well.

97 For comparison electron density predictions have been calculated from the latest  
98 model version, IRI-2007 [Bilitza and Reinisch, 2008], for each measurement point of the  
99 satellite readings. The International Reference Ionosphere (IRI) is an international project  
100 sponsored by the Committee on Space Research (COSPAR) and the International Union of  
101 Radio Science (URSI). IRI describe monthly averages of the electron density, electron  
102 temperature, ion composition ( $O^+$ ,  $H^+$ ,  $N^+$ ,  $O_2^+$ ,  $NO^+$ ,  $Cluster^+$ ), ion temperature and ion drift  
103 in the ionospheric altitude range of 50-1500km.

104 Comparing to the former IRI models, there are some important changes in the newest  
105 version, IRI-2007. Relevant for this study are the two new options for the topside electron  
106 density profile. One option is a correction factor for the 2001 model based on over 150,000  
107 topside profiles from Alouette 1, 2, and ISIS 1, 2, and this term varies with altitude, modified

108 dip latitude, and local time [Bilitza, 2004]. The other option is the NeQuick topside model  
109 that was developed by S. Radicella and his collaborators over the last decade [Radicella and  
110 Leitinger, 2001; Coisson et al., 2006].

111 Furthermore there is a much improved model for the topside ion composition. The  
112 present IRI model is largely based on a compilation of Russian high-altitude rocket  
113 measurements [Danilov and Yaichnikov, 1985; Danilov and Smirnova, 1995], and on a  
114 limited amount of incoherent scatter radar data. Working with the satellite in situ  
115 measurements from AE-C, -E, and Intercosmos 24, Triskova et al. [2003] have developed a  
116 new model for the ion composition in the topside ionosphere.

117 The inputs for the IRI-2001 and IRI-2007 models are (1) sunspot number, R, (2) global  
118 ionospheric index, IG, and (3) the 3-hourly magnetic,  $a_p$  index. IG is an ionospheric-effective  
119 solar index that is based on foF2 measurements from selected ionosondes and a correlation  
120 with the Consultative Committee of International Radio Communications (CCIR) maps [R.  
121 Liu et al., 1983]. For the calculations presented here we have used the version updated on 28  
122 May 2010 (<http://nssdcftp.gsfc.nasa.gov/models/ionospheric/iri/>).

123 Selected options:

124 foF2: URSI

125 Ni: DS-95 & TTS-03

126 Ion drift: not computed

127 Te: Te topside (Intercosmos)

128 ESF: not computed

129 Topside model: NeQuick

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### 132 **3. Comparison between model and observation**

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134 For our comparison we made use of all available electron density data from the two  
 135 considered satellites up to the beginning of 2010. Orbit averages of  $N_e$  are calculated for each  
 136 orbit of CHAMP and GRACE. The orbital averages are the basis for further analyses. Each  
 137 satellite circles the Earth more than 5500 times per year. Collocated IRI  $N_e$  predictions are  
 138 generated at the same rate and by considering the actual geophysical conditions. Figure 1  
 139 shows the temporal evolution of the four data sets. Here orbital averages have been smooth by  
 140 a moving boxcar filter over 31 days. This filter length is chosen since IRI predictions are  
 141 representing monthly averages. Both at CHAMP and GRACE heights we find large intra-  
 142 annual variations. They reflect primarily the low-latitude seasonal variation of the ionospheric  
 143 electron density [e.g. L. Liu et al., 2009]. Rather prominent are the minima around June  
 144 solstice. Peak densities are obtained during equinox seasons. These seasonal variations are  
 145 further modulated by the orbit-dependent local time changes of the satellite measurements.

146 During the solar maximum years and the declining phase of the cycle the modeled  $N_e$   
 147 values track the observations at the two altitudes rather well. However, IRI underestimates  
 148 somewhat the amplitude of seasonal excursions at CHAMP height. GRACE returns  
 149 systematically higher  $N_e$  values during fall seasons than the model. After the epoch 2005,  
 150 however, IRI systematically overestimates the electron density at both altitudes. The  
 151 differences between the model and measurement curves are increasing with time up to 2010.

152 In order to obtain a more quantitative estimate of the difference between IRI and satellite  
 153 data, we calculated annual averages of the relative deviations,  $\Delta$ , in percent

$$154 \quad \Delta = \frac{N_{e,IRI} - N_{e,Sat}}{N_{e,Sat}} 100 \quad (1)$$

155 where the subscript  $Sat$  stands either for CHAMP or GRACE.

156 As can be seen in Figure 2, during the solar maximum years the deviation is small. In  
 157 case of CHAMP it is in the range of the measurement uncertainty of a few percent  
 158 [McNamara et al., 2007]. From 2004 onward, however, we observe an increasing difference

159 between the data sets. For 2009 the ratio between  $N_e$  from IRI and from CHAMP is larger  
160 than 1.5. For the altitude of GRACE the model first underestimates the measurements (too  
161 small fall peaks), but from 2004 on the deviation rises even faster than at CHAMP location.  
162 For the average of 2009 the ratio IRI over GRACE surmounts 1.7. Otherwise the differences  
163 at the two altitudes show similar evolutions.

164 We are interested to find out to which regions or local times these large discrepancies  
165 can be attributed. For that reason we went back to the original measurements and binned the  
166 data in magnetic latitude versus local time frames. Figure 3 shows in the upper row the  
167 electron density distribution as observed by CHAMP during the years 2008 and 2009  
168 separately for the three Lloyd seasons (equinoxes: Mar., Apr., Sep., Oct.; June solstice: May-  
169 Aug.; Dec. solstice: Nov.-Feb.). Below collocated IRI densities are plotted. Shown  $N_e$  values  
170 are averages over all longitudes.

171 It is quite obvious that largest differences between modeled and measured  $N_e$  appear at  
172 low latitudes during daytime. During the post-midnight hours until 06 LT differences are  
173 small in all season. The main contribution to the  $N_e$  overestimation by IRI comes from the  
174 daylight hours, in particular from the latitude range  $\pm 30^\circ$ MLat. During post-sunset hours the  
175 predicted  $N_e$  is also somewhat too high at low latitudes. A very similar picture emerges from  
176 the model-to-measurement comparison at GRACE altitude (not shown). At this height the  
177 relative difference is even higher than at CHAMP. Here the post-sunset hours provide a larger  
178 contribution.

179 Dominant density features in Figure 3 are the two latitude bands representing the  
180 equatorial ionization anomaly. The EIA electron density peaks are clearly visible for equinox  
181 and December solstice seasons. From the modeled  $N_e$  we deduce a latitude separation of the  
182 peaks of  $30.5^\circ$ , while from CHAMP measurements the peaks appear to be separated only by  
183  $27.5^\circ$  in latitude. Both these facts imply that IRI is strongly overestimating the equatorial ion



184 fountain effect during the considered solar minimum. This means, more plasma is moved up  
185 and it reaches larger heights.

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#### 188 **4. Discussion**

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190 Based on comparison with CHAMP and GRACE measurements we found that the IRI-  
191 2007 model overestimates the electron density by a factor of more than 1.5 in the height range  
192 300 to 500 km during the deep solar minimum of solar cycle 23/24. Conversely, during solar  
193 maximum years up to 2004 on average a reasonable agreement of the model at the two  
194 altitudes is found. During the minimum years 2008 and 2009 CHAMP was cruising at  
195 altitudes of 320-340 km, and GRACE stayed between 450 and 470 km. According to  
196 observations by L. Liu et al. [2006] the F2 peak height during solar minimum stays below the  
197 altitudes of the two spacecraft. One of the new features of IRI-2007 is the improvement in  
198 topside electron density. Previous versions tend to overestimate  $N_e$  above the F2 peak [Bilitza  
199 and Reinisch, 2007]. We also tried an earlier version of IRI-2007 (updated on 9 Feb. 2009).  
200 There we found overestimations of more than a factor of 2 for the year 2009. Obviously the  
201 improvements of IRI go in the right direction, but they are not sufficient to account for the  
202 special conditions during the recent solar minimum.

203 By inspecting the latitude and local time distribution of the differences we find that the  
204 discrepancy between model and measurements stems primarily from a too strong equatorial  
205 ionization anomaly. Furthermore, the distribution of  $N_e$  suggests that the EIA from IRI peaks  
206 at a too large height above the magnetic equator. These two facts can be caused by a too  
207 strong vertical plasma drift.

208 Stolle et al. [2008] have studied the L-value dependence of the EIA crests at CHAMP  
209 altitude on the prevailing vertical plasma velocity. Their result for  $L_{crest}$  obtained over  
210 Jicamarca reveals a linear correlation of the form

$$211 \quad L_{crest} = 1.083 + 1.97 \cdot 10^{-3} v_Z \quad (2)$$

212 where  $v_Z$  is the vertical plasma velocity measured in m/s and  $L_{crest}$  in Earth radius. We have  
213 used the definition of the L-value for dipole geometry

$$214 \quad L = \frac{r}{R_E} \frac{1}{\cos^2 \beta} \quad (3)$$

215 where  $r$  is the radial distance of the measurement point from the Earth center,  $R_E = 6371$  km is  
216 the Earth's radius, and  $\beta$  the magnetic latitude of the EIA crest. From Figure 3 we deduce  
217 latitudinal separations of the crests of  $30.5^\circ$  and  $27.5^\circ$  for the IRI and CHAMP results around  
218 12:00 LT. This corresponds to  $L_{crest}$ -values of 1.130 and 1.115 for the model and observation,  
219 respectively. With the help of Eq. (2), we obtain vertical velocities of 25 m/s for IRI and 17  
220 m/s for CHAMP results. For comparison, Fejer et al. [2008] obtain from their statistical  
221 analysis of ROCSAT-1 plasma drift measurements an average drift velocity of 20.5 m/s over  
222 the F10.7 range 100-160 sfu at 12:00 LT. They find almost no solar flux dependence of  
223 vertical drift around noon. The reported plasma drift velocity is half way between the values  
224 determined in this study. There have to be other causes than just low solar flux level for the  
225 small vertical plasma drift deduced from CHAMP observations.

226 During the recent solar minimum some properties of the ionosphere seem to have  
227 changed that cannot simply be parameterized by F10.7 or Ap. In that case IRI, an empirical  
228 model, by design cannot reproduce the ionospheric properties correctly. Related problems  
229 have been identified by Emmert et al. [2010] for the neutral atmosphere when comparing  
230 thermospheric mass density measurements with predictions from MSISNRL-00. They report a  
231 record low of the thermospheric density during this solar minimum. Largest deviations from  
232 the model are observed just below 500 km altitude amounting to more than 20% depletion.

233 For explaining this discrepancy they had to reduce the abundance of atomic oxygen by 12% at  
234 120 km altitude and reduction the exospheric temperature by 14K. They propose changes of  
235 chemical and dynamical processes in the mesosphere and lower thermosphere as being  
236 responsible for the depletion of O at the base of the thermosphere. We suggest that also the  
237 ionosphere has experienced fundamental modifications during the deep minimum. An  
238 indication for that may be the weak recovery of electron density at the end of 2009 when the  
239 solar flux has already started to rise again. This fact contributes significantly to the huge  
240 difference in *Ne* between IRI and satellite observations in 2009. It will be interesting to watch  
241 the evolution during the coming months and years. This may help to find out what kind of  
242 modification is required to adapt the IRI model to the new conditions.

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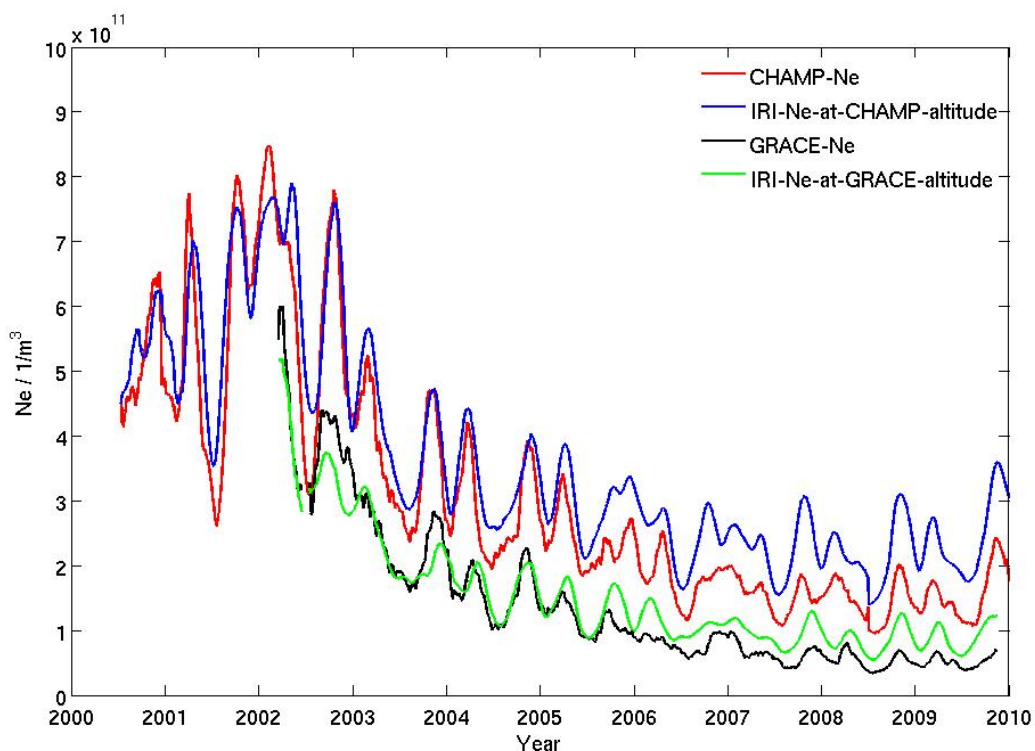
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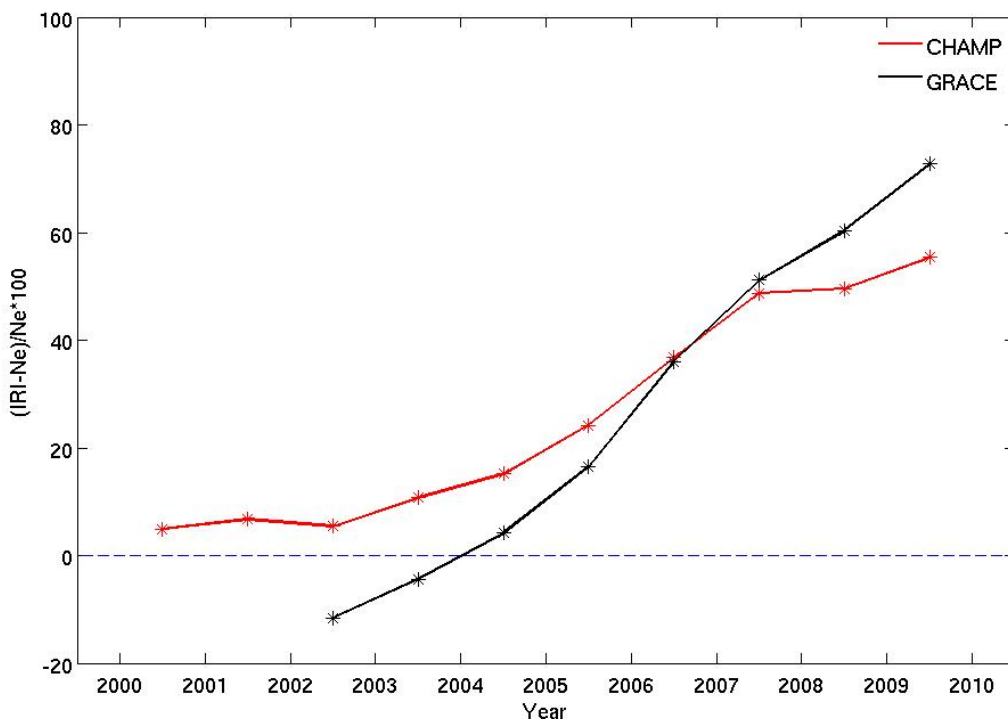
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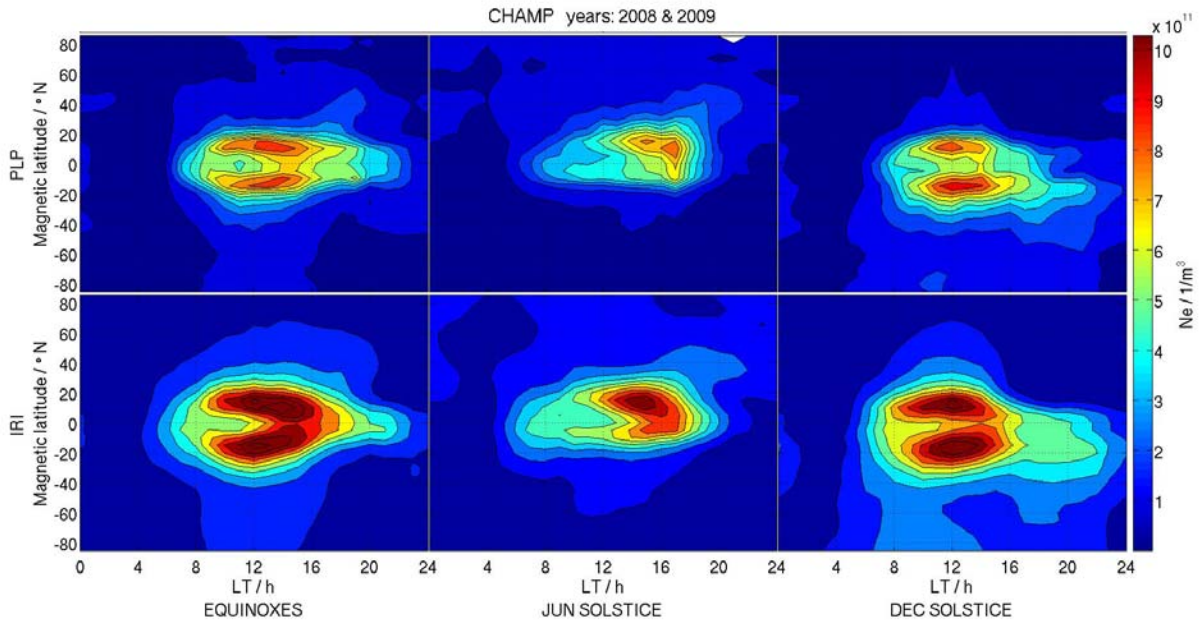
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Fig. 1: Comparison of CHAMP and GRACE electron density measurements with collocated predictions of the IRI model. Orbital averages are averaged over 31 days.



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Fig. 2: Annual average of the deviation in percent of IRI-2007 electron density estimates from collocated measurements by CHAMP and GRACE.



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Fig. 3: Distribution of electron density versus magnetic latitude and local time (averages over all longitudes) separately for the three Lloyd seasons. The top row presents CHAMP observations of the years 2008 and 2009 and the bottom frames show collocated IRI  $Ne$  predictions.