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1 **Wood anatomical analysis of Swiss willow (*Salix helvetica*) shrubs growing on**  
2 **creeping mountain permafrost**

3

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19

20 **Abstract**

21 Permafrost and related landforms (rockglaciers) are widespread phenomena in alpine  
22 geosystems. In the context of changing environments due to the significant warming,  
23 there is a need for thoroughly monitoring and analyzing the complex responses of these  
24 cryospheric geosystems. Here, the first-time application of wood anatomical methods in  
25 this context is presented in order to investigate whether rockglacier movement is  
26 reflected in varying cell structures of plants growing on top of the rockglaciers.

27 In order to determine the influence of ground movements (by permafrost creep) and  
28 their influence on the conductive elements within roots of plants, wood-samples were  
29 taken from active and inactive rockglaciers in the Turtmann Valley, southern Swiss  
30 Alps. Since the occurrence of trees is limited altitudinally, the investigation was  
31 restricted to Swiss willow shrubs (*Salix helvetica*) frequently growing in permafrost  
32 areas above the timberline in the European Alps. This rather new approach concentrates  
33 on general vessel size differences as a result of mechanical stresses. The comparison of  
34 vessel sizes in roots of Swiss willow shrubs growing on active and inactive permafrost  
35 bodies depicts differences within the roots, which are related to the activity status of the  
36 respective rockglacier creep.

37

38 Keywords: Dendroecology, rockglacier activity, vessel size, roots

39

40

41 **Introduction**

42 Permafrost and related landforms (e.g., rockglaciers) are widespread phenomena in  
43 many high mountain geosystems (Barsch, 1992). In the context of changing  
44 environments due to the significant warming, there is a need for thoroughly monitoring  
45 and analyzing the sensitivity and the complex response of the considered geosystems.

46 Various methods like geomorphological mapping, terrestrial geodetic survey as well as  
47 photogrammetric analyses have been applied in order to assess the activity of  
48 rockglaciers and to quantify their movements (Roer et al., 2005a; Roer and Nyenhuis,  
49 2007). In general, the activity of rockglaciers is classified by their ice content and flow  
50 behavior (Haeberli, 1985; Barsch, 1996). Hence, active rockglaciers contain ice and  
51 move with at least 0.1 m/a, whereas inactive landforms still contain ice but actually do  
52 not move. Relict rockglaciers indicate former permafrost conditions; they lack the ice  
53 and definitely stopped moving (often several hundred to thousand years ago)  
54 (Frauenfelder et al., 2001).

55 In addition to these techniques, a new approach has recently been designed which  
56 assesses rockglacier activities by means of dendrogeomorphological methods combined  
57 with wood anatomical analyses (Roer, 2007). In recent years, annual rings of several  
58 shrub species in high arctic and tundra areas proved to be feasible indicators for  
59 environmental change (Hallinger et al., 2010), but comparable studies have hardly ever  
60 been conducted in alpine areas. Particularly with regard to rockglacier studies,  
61 dendrogeomorphological methods have rarely been applied. This is mainly due to the  
62 fact that the vegetation cover on rockglacier surfaces is generally sparse, since the  
63 surfaces mostly consist of coarse and blocky material (e.g., Giardino et al., 1984; Roer  
64 and Nyenhuis, 2007). Thus, on active rockglaciers herbs and small shrubs are mostly  
65 restricted to areas with fine material usually found at the lobe fronts and on ridges.

66

67 The application of tree-ring analysis in geomorphology (dendrogeomorphology; cf.,  
68 Alestalo, 1971) facilitates the reconstruction and dating of gravitational processes (e.g.,  
69 Schweingruber, 1983, 1996; Shroder, 1980; Gärtner 2007a). The underlying principle of  
70 the ‘process-event-response-chain’ (cf., Shroder, 1978) describes the link between  
71 geomorphological processes, their influences on trees and the corresponding tree growth

72 reactions. These reactions include not only variations in tree-ring width (e.g., sudden  
73 reduction in growth), but also changes in the structure of the cells, such as the  
74 development of compression or tension wood (Fantucci and McCord, 1995; Heinrich  
75 and Gärtner, 2008) caused by the tilting of the stem. So far, various  
76 dendrogeomorphological approaches have been used in reconstructing erosion rates  
77 (e.g., LaMarche, 1968; Gärtner et al., 2001; Gärtner, 2007b), enlargements of cracks at  
78 the Questa-scarp (Sahling et al., 2003), in determining frequency and magnitudes of  
79 debris flow events (e.g., Strunk, 1991; Gärtner et al., 2003; Sorg et al., 2010), in  
80 identifying landslides (e.g., Paolini et al., 2005) or dating flood-plain sedimentation  
81 (Friedman et al., 2005). Annual ring-width variations and related growth variations due  
82 to environmental stress have rarely been studied in shrubs (e.g., Gers et al., 2001). A  
83 comprehensive overview on growth rings in herbs and shrubs has recently been  
84 presented by Schweingruber and Poschlod (2005). However, only a few publications  
85 concentrate on the combination of shrubs and geomorphic processes (e.g., Owczarek  
86 2010), the ability to build tree-ring chronologies and to reconstruct temperature changes  
87 (Bär et al., 2008). In general, recent studies on wood anatomical structures in tree and  
88 shrub rings have demonstrated the great potential of wood anatomical investigations  
89 (Copenheaver et al., 2010; Fonti et al., 2010; Roer et al., 2007). Regarding  
90 dendrogeomorphological studies on periglacial processes, Zoltai (1975) was the first to  
91 report different phases of activity related to gelifluction in the Subarctic. The  
92 conclusions were based on the analysis of reaction wood in spruce (*Picea mariana*,  
93 *Picea glauca*) and larch (*Larix laricina*) trees. Giardino et al. (1984) studied reaction  
94 wood, tree-ring width variations and tilting of 283 trees on a rockglacier complex and  
95 revealed different periods of movement since the 15th century. Jakob (1995) monitored  
96 dwarf-shrubs in the Canadian Arctic, which had been run over by gelifluction lobes. He  
97 was able to quantify movement rates of 1.9 to 3.5 cm/a. Bachrach et al. (2004) used

98 dendrogeomorphological methods, to document the long-term development of a  
99 rockglacier in Alberta, Canada. In this case, trees (*Picea engelmannii* and *Abies*  
100 *lasiocarpa*) were buried by an advancing rockglacier. By comparing the death-dates of  
101 different trees, a front advance of 1.6 cm/a was estimated. More recently, Körner and  
102 Hoch (2006) investigated the dwarfing of trees growing on low elevation permafrost  
103 islands concluding that the temperature of the root zone most dominantly determined  
104 the growth of these trees. Moreover, there is evidence that changes of the direct  
105 environment of roots (e.g., soil pressure, mechanical stress or temperature) result in a  
106 variation of their anatomical structure (Gärtner et al., 2001; Gärtner 2007b).

107 At the site investigated here (Turtmann Valley, Swiss Alps), different techniques have  
108 been used to quantify rockglacier movements (Roer, 2007; Roer et al., 2005a). While  
109 terrestrial geodetic survey and digital photogrammetry are standard techniques for this  
110 purpose (cf., Käab et al., 2005), wood anatomical techniques have never been applied in  
111 this regard, but have the potential to supply additional information (Filion and Gärtner,  
112 2010; Gärtner and Heinrich, 2010). Therefore, the main purpose of this study is to  
113 demonstrate the usefulness of wood anatomical analyses especially of roots of alpine  
114 shrubs for detecting ground movements resulting from permafrost creep. The key  
115 question is whether the movements due to rockglacier activity are reflected in the  
116 anatomical structure of the shrubs, thus allowing a long-term monitoring as well as a  
117 better understanding of the particular conditions in which accelerated movements of  
118 rockglaciers occur.

119 The investigation focuses on wood-anatomical features in the xylem of shrubs, more  
120 precisely the conduction tissue (vessels) in their roots, which have not been examined in  
121 this regard yet neither in the given context nor for the selected species. The resulting  
122 data are interpreted and verified in relation to information on rockglacier movement  
123 derived from digital photogrammetric analyses (Roer et al., 2005a, b). The verification

124 will help to judge whether wood anatomical techniques can be applied to periglacial  
125 investigations and hence established as an innovative approach in this field of research.

126

## 127 **Study area**

128 The sampling of shrubs on rockglaciers was conducted in the Turtmann valley, located  
129 south of the River Rhone in southern Switzerland (7° 38' E, 46° 13' N) (Fig. 3). Due to  
130 its inner-alpine location, the area is characterized by an intramontane climate with an  
131 annual precipitation rate of 600-900 mm at c. 2000 m a.s.l. and a 0°C isotherm of the  
132 mean annual air temperature (MAAT) at c. 2550 m a.s.l. (Van Tatenhove and Dikau,  
133 1990). The timberline is situated at 2200 m a.s.l., while the treeline, which is affected by  
134 grazing during the summer months, runs at 2400 m a.s.l. The geomorphology of the  
135 Turtmann valley is dominated by several hanging valleys and two big glaciers. In  
136 addition, a multitude of periglacial landforms such as gelifluction lobes, ploughing  
137 boulders, and rockglaciers covering 4.2% of the total area of 110 km<sup>2</sup> are found above  
138 the shoulders of the glacial trough. Rockglaciers occur in different states of activity and  
139 with varying degrees of vegetation cover (Roer and Nyenhuis, 2007), thus making them  
140 suitable for investigations by dendroecological methods.

141 The active and inactive rockglaciers from which the samples were taken are situated  
142 next to each other (Fig. 4). Apart from the activity conditions, other factors influencing  
143 plant growth (e.g., altitude, aspect, slope angle, etc.) are similar at the two sites (Roer,  
144 2007). The active rockglacier lobe creeps with extraordinary high velocities and  
145 indicates instability at its terminus (Roer, 2007; Kääh et al., 2005, 2007). At the location  
146 of the sampled shrubs (Fig. 4), horizontal velocities between 0.5 and 2 m/a have been  
147 measured (Roer, 2007).

148

## 149 **Material and Methods**

150 The study focuses on analyzing vessel lumen area (VLA) within the diffuse to semi  
151 ring-porous annual rings of Swiss willow (*Salix helvetica*) shrubs. Due to the fact that  
152 we focus on activity stages of rockglaciers and with this also on permafrost related  
153 movements of the ground, we concentrate on the roots of these shrubs, which are  
154 directly influenced by low ground temperatures as well as by mechanical stress exerted  
155 to the roots. The vessel sizes within the roots of these shrubs were quantified and  
156 analyzed ecophysiologicaly. In doing so we were not concentrating on single rings but  
157 on the overall vessel differences of individual shrub roots.

158

### 159 *Measurement design*

160 During various field campaigns, Swiss willow shrubs growing on an active and an  
161 adjacent inactive rockglacier in the Turtmann valley were sampled (cf., Roer, 2007).  
162 The inactive rockglacier was not affected by creep during the last 30 years and is known  
163 from an outcrop to contain a frozen core (Broccard, 1998; Roer, 2007). This strategy  
164 ensured that all shrubs sampled for the study were affected by the same climatic  
165 conditions and were growing on comparable substrates, both having ice in the  
166 subsurface. Consequently, the main ecological difference at the two sampling sites was  
167 the mechanical stress due to ground movements influencing only the shrubs growing on  
168 the active rockglacier. For that reason, the inactive rockglacier was regarded as the  
169 reference site and respective samples from the inactive rockglacier were treated as a  
170 reference to demonstrate differences to samples taken from the active rockglacier.

171 In general, reaction wood in shrubs is difficult to investigate due to the lack of a main  
172 stem. The single branches (or stems) of shrubs often show reaction wood which might  
173 be caused by different mechanical stress factors, such as snow cover and wind exposure.

174 In contrast, the root collar and especially the roots of shrubs grow under more stable  
175 conditions, i.e., in the soil, and hence they may be suited best for the investigation of

176 possible influences of ground movements envisaged in this study. Changes in the root  
177 environment may lead to growth stress and result in changes of the anatomical structure  
178 (Gärtner et al., 2001).

179 In order to study these potential changes, entire specimen of Swiss willow shrubs were  
180 taken from an active rockglacier lobe (n = 10) as well as from a reference site, an  
181 adjacent inactive rockglacier (n = 10), including their entire root systems (Fig. 1).

182 In the laboratory, several discs were taken from the individual stems and roots of the  
183 shrubs at a distance of about 2 cm from the root collar. The small size of most roots  
184 (diameters of ~1.5 cm) enabled the preparation of micro-sections (thickness ~15  $\mu\text{m}$ ) of  
185 whole cross-sections (Fig. 2) using a sledge microtome GSL1. These micro-sections  
186 were stained with Safranin and Astrablue to distinguish between lignified (Safranin) and  
187 non-lignified (Astrablue) parts of the rings and to obtain a better contrast for the  
188 following image analysis procedure. For dehydration, the samples were rinsed with  
189 alcohol, immersed in Xylol, embedded in Canada-Balsam and oven-dried at 60° C for  
190 about 24 hours (Gärtner et al., 2001).

191 The resulting micro-slides were then placed under a microscope and digital photos, with  
192 400 times magnification, were taken. The anatomical structure of the rings was clearly  
193 visible, which is not always the case in deciduous plants (Fig. 2).

194 The micro photos were then used to measure vessel sizes of stressed (active rockglacier)  
195 and unstressed (inactive rockglacier) reference samples using the image analysis  
196 program WinCELL (Regent Instruments, Canada). This program automatically  
197 measures cell dimensions, that is, radial length, width and area of individual vessel  
198 lumen. To realize an automated measurement of the vessel lumen area (VLA), a filter  
199 which excluded all cells with an area smaller than 200  $\mu\text{m}^2$  was applied. These cells  
200 were classified as fiber cells, that is, cells of the ground tissue of the rings. All cells  
201 larger than 200  $\mu\text{m}^2$  were classified as vessels.

202 For an a posteriori examination of the variation of the VLA-values between the single  
203 roots as well as regarding their site specific variations, box-whisker plots were created  
204 (Heinrich et al., 2007). Box-whisker plots show some important features, that is, the box  
205 in the center spans the quartiles, limited by the 25% and 75% quartiles; the horizontal  
206 line in the center marks the median; the vertical lines extending out from the box are  
207 capped by horizontal lines often called whiskers indicating certain minima and maxima.  
208 In general, box-plots allow an easy comparison of the mean values and the spread of the  
209 data. Two data sets are considered significantly different if the boundary of one box  
210 does not overlap the median of the other box (Cleveland, 1994).

211 An additional focus was set on analyzing the variation of the biggest vessels occurring  
212 in the roots of the single shrubs. This was done because it is known for several tree  
213 species, that the variation of the biggest vessels in single annual rings show high  
214 correlations to summer temperatures (Garcia-Gonzalez and Fonti, 2006) or other  
215 environmental factors as floodings (George et al., 2002) or forest fires (Cames et al.,  
216 2011).

217

## 218 **Results**

219 A simple comparison of ring-width variations did not show any obvious differences  
220 between unstressed and stressed samples from the inactive and active rockglacier lobe,  
221 respectively. This also applies to variations in the structure of fiber cells formed mainly  
222 to stabilize the plant. However, a basic visual comparison of the vessel sizes between  
223 individual roots of stressed and unstressed samples revealed clear differences (Fig. 5).  
224 The vessels in roots from the active rockglacier tend to be smaller than those from the  
225 inactive site.

226 To support this visual impression, the vessel sizes in all micro-sections were measured.

227 Although there is a certain variability of the vessel lumen area (VLA) within and

228 between the single roots, there is a distinct difference when comparing the VLA of  
229 single roots as well as of the single shrubs growing on the active and inactive  
230 rockglacier (Fig. 6). The overall average VLA values for roots were 2.6 times higher in  
231 shrubs growing on the inactive rockglacier ( $1441 \mu\text{m}^2$ ) than on the active rockglacier  
232 ( $554 \mu\text{m}^2$ ).

233 This rather simple analysis does show an obvious difference in the VLA values of the  
234 two sites. On average, 659 vessels were measured per micro-section (average numbers:  
235 active - 602; inactive – 716).

236 The box-plot results confirm the visual impressions (Fig. 7); although vessels sizes are  
237 reduced in stressed roots on the active rockglacier compared to undisturbed roots on the  
238 inactive lobe, these results are barely significant (lower diagram in Figure 7).  
239 Nevertheless, the median for vessel sizes in roots of all plants from the active part is on  
240 average 60% lower than the value in roots of the inactive rockglacier. Moreover, the  
241 spread of the data around the median values, the maximum values as well as the outliers  
242 are lower in all plants growing on the active lobe (boxes upper left diagram in Figure 7).  
243 Consequently, further analysis concentrated on the measurements closer to the median,  
244 that is, all data within the range of the upper and lower quartile around the median  
245 values (Fig. 8).

246 Within this more confined data range, significant differences exist between the VLA-  
247 values of shrub roots growing on active and inactive rockglaciers. As it was shown for  
248 the entire dataset, the spread of the data in the confined dataset around the respective  
249 median values is also lower for all plants growing on the active lobe (Fig. 8).

250 When limiting the analysis of the entire data set on the largest vessels only, the  
251 differences between shrubs growing on the active and inactive rockglaciers are  
252 significant, no matter if concentrating on the 10% or 35% biggest vessels occurring in

253 individual roots (Fig. 9). In contrast to the analysis before, the differences in the spread  
254 of the data around the median values are less pronounced.

255

256

## 257 **Discussion**

258 When comparing the overall VLA data of the single roots of shrubs growing on the  
259 inactive and active rockglacier, the difference of the average values is significant,  
260 although the distance between the two boxes is not large. This problem might be related  
261 to the fact, that the lower threshold for lumen sizes to be defined as vessels was set to  
262  $200 \mu\text{m}^2$ . This threshold value is based on various tests using the images of the sample  
263 roots which identified no vessels with VLA smaller than  $200 \mu\text{m}^2$ . While it is important  
264 to state that this value is helpful for the current analysis, it is difficult to define a  
265 universally valid threshold for vessel sizes. The general size as well as the variability of  
266 vessels is species-specific, but there is also a huge variation depending on the position  
267 of the vessels within the plant (Aloni and Zimmermann, 1983) and depending on  
268 ecological conditions during the vegetation period (Sass and Eckstein, 1995). For this  
269 reason, focusing on the main data spreading around the median values is reliable. The  
270 additional concentration on the 10% and 35% biggest vessels showed highly significant  
271 results which are comparable to the findings of Garcia-Gonzalez and Fonti (2006). They  
272 found significant correlations when comparing the biggest vessels of single annual rings  
273 in ring-porous species to temperature variations during the vegetation period.

274 The comparison and interpretation of Swiss willow shrubs growing on active and  
275 inactive permafrost bodies is complex, although on both landforms, the topographic and  
276 climatic conditions are similar (Fig. 4). However, beside the distinction in their state of  
277 activity, differences may exist in the thickness of the frozen core of the rockglacier. On  
278 the active rockglacier, the upper part (several meters), the so-called active layer, thaws

279 in summer time but is completely frozen during winter. On the inactive lobe, the thawed  
280 layer presumably is much thicker in summer and the freezing during winter time cannot  
281 completely penetrate through this layer; thus an unfrozen layer may exist in several  
282 meters depth. Unfortunately, suitable temperature data do not exist for the individual  
283 rockglaciers investigated here. But, since the roots that were cropped penetrated to a  
284 maximum depth of 10-15 cm only, we conclude that the ground temperature is not the  
285 main driving factor.

286 At least it is known with certainty that the inactive rockglacier contains a frozen core,  
287 which has been documented by photographs taken during construction works at an  
288 avalanche dam nearby (Broccard, 1998). Based on the present observations, the  
289 anatomical differences regarding the size of the vessels in the roots of Swiss willow  
290 shrubs growing on the active and inactive rockglacier surface may be related to the  
291 following factors: (i) the movement of the active rockglacier exerts mechanical stress to  
292 the roots of the respective shrubs and (ii) differences in ground thermal conditions.

293 We regard constant growth stresses in the ground due to mechanical influences by  
294 rockglacier creep as the more dominant reason for the differences in vessel size. Until  
295 now, this behavior has not been described for roots of shrubs. In general, vessels are  
296 formed to support the plant by transporting water and nutrients and they are known to  
297 be variable in size due to environmental stress (Kozłowski, 1979). Since vessels are  
298 formed only for transporting water it may be assumed that their sizes are closely related  
299 to availability of water rather than temperatures. Furthermore, Heinrich and Gärtner  
300 (2008) showed that vessel sizes are correlated to mechanical stimuli due to geomorphic  
301 events, and thus, it seems more likely that the relatively strong reduction in vessel  
302 lumen areas of the shrubs growing on the active rockglacier is due to mechanical rather  
303 than the thermal stress.

304 Nevertheless, these thermal conditions cannot be neglected as possible additional factor.  
305 Although both rockglaciers contain a frozen core and the upper part of the substratum is  
306 constantly frozen during wintertime, there might be an effect of the reduced and most  
307 likely deeper seated ice core of the inactive rockglacier regarding the temperature of the  
308 upper soil layer during the vegetation period of the shrubs. Compared to the thickness of  
309 the “active layer” of the active rockglacier, the higher thickness of the unfrozen  
310 substratum on the inactive rockglacier may result in potentially higher maximum and  
311 minimum temperatures of the soil layer penetrated by the roots on the inactive  
312 landform.  
313 Differences in daily maximum and minimum temperatures are postulated to be a  
314 potential reason for tracheid size differences in exposed and unexposed roots of conifers  
315 (Fayle, 1968; Gärtner et al. 2001; Gärtner, 2003). But in fact, the definite influence of  
316 soil temperature differences on the anatomical structure of plant roots has never been  
317 tested in detail.

318

### 319 **Conclusion**

320 Anatomical variations in roots of 20 shrubs of Swiss willow were investigated and  
321 statistically quantified with digital imagery. The results showed clear differences  
322 between the vessel sizes in roots of shrubs growing on active and inactive rockglaciers,  
323 respectively. The differences in vessel sizes seem to be caused by two factors, first of all  
324 the differences in mechanical stresses exerted to the roots, and secondly, possibly to a  
325 lower extent, by differences in the thermal regime causing differences in soil  
326 temperature close to the surface.  
327 This first time application of wood anatomical techniques appears to be a feasible  
328 method and the findings are promising for more extended on-going research.

329 The combination of geophysical and wood anatomical methods provides a good  
330 opportunity to thoroughly describe and analyze rockglacier and ground movements in  
331 general. Hence, the current study indicates a new way to carefully interpret the complex  
332 responses of mountain geosystems in the context of changing environments along with  
333 significant warming and related process changes.

334

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341

### 342 **References**

- 343 Alestalo, J., 1971. Dendrochronological interpretation of geomorphic processes. *Fennia*  
344 105, 140 pp.
- 345 Aloni, R., Zimmermann, M., 1983. The control of vessel size and density along the  
346 plant axis: A new hypothesis. *Differentiation* 24, 203-8.
- 347 Bachrach, T., Jakobsen, K., Kinney, J., Nishimura, P., Reyes, A., Laroque, C.P., Smith,  
348 D.J., 2004. Dendrogeomorphological assessment of movement at Hilda rock glacier,  
349 Banff National Park, Canadian Rocky Mountains. *Geografiska Annaler* 86A, 1-9.
- 350 Bär, A., Pape, R., Bräuning, A., Löffler, J., 2008. Growth-ring variations of dwarf  
351 shrubs reflect regional climate signals in alpine environments rather than  
352 topoclimatic differences. *Journal of Biogeography* 35, 625–36.
- 353 Barsch, D., 1992. Permafrost creep and rockglaciers. *Permafrost and Periglacial*  
354 *Processes* 3, 175–88.

355 Barsch, D., 1996. Rockglaciers: Indicators for the present and former geocology in  
356 high mountain environments. Springer Series in Physical Environment 16, Berlin,  
357 331 pp.

358 Broccard, A., 1998. Géomorphologie du Turtmantal (Valais). Unpublished Diploma  
359 thesis, Department of Geography, University of Lausanne.

360 Comes, S., Tardif, J.C., Bergeron, Y., 2011. Anomalous earlywood vessel lumen area in  
361 black ash (*Fraxinus nigra* Marsh.) tree rings as a potential indicator of forest fires.  
362 *Dendrochronologia* 29, 109-14.

363 Cleveland, W.S., 1994. The elements of graphing data. Hobart Press, Summit, NJ, 297  
364 pp.

365 Copenheaver, C.A., Gärtner, H., Schäfer, I., Vaccari, F.P., Cherubini, P., 2010. The  
366 potential of wood anatomy of a Mediterranean shrub to reflect site-specific responses  
367 to increased drought frequency. *Canadian Journal of Botany* 88, 545–55.

368 Fantucci, R., McCord, A., 1995. Reconstruction of landslide dynamic with  
369 dendrochronological methods. *Dendrochronologia* 13, 1-22.

370 Fayle, D.F.C., 1968. Radial growth in tree roots - distribution, timing, anatomy. Faculty  
371 of Forestry, University of Toronto, Technical Report No. 9, Toronto, 183 pp.

372 Filion, L., Gärtner, H., 2010. Dendrogéomorphologie, in: Payette, S., Filion, L. (Eds.),  
373 *La Dendroécologie: Principes, méthodes et applications*. Les Presses de l'Université  
374 Laval, Québec, Québec, pp. 537-72.

375 Fonti, P., von Arx, G., García-González, I., Eilmann, B., Sass-Klaassen, U., Gärtner, H.,  
376 Eckstein, D., 2010. Studying global change through plastic responses of xylem  
377 anatomy in tree rings. *New Phytologist* 185, 42-53.

378 Frauenfelder, R., Haeberli, W., Hoelzle, M., Maisch, M., 2001. Using relict rockglaciers  
379 in GIS-based modelling to reconstruct Younger Dryas permafrost distribution

380 patterns in the Err-Julier area, Swiss Alps. Norwegian Journal of Geography 55, 195-  
381 202.

382 Friedman, J.M., Vicent, K.R., Shafroth, P.B., 2005. Dating floodplain sediments using  
383 tree-ring response to burial. Earth Surface Processes and Landforms 30, 1077-91.

384 Gärtner, H., 2003. Holzanatomische Analyse diagnostischer Merkmale einer  
385 Freilegungsreaktion in Jahrringen von Koniferenwurzeln zur Rekonstruktion  
386 geomorphologischer Prozesse. Dissertationes Botanicae 378, 118 pp.

387 Gärtner, H., 2007a. Glacial landforms, tree rings: Dendrogeomorphology, in: Elias, S.A.  
388 (Ed.), Encyclopedia of Quaternary Sciences, Vol. 2, Elsevier, pp. 979-88.

389 Gärtner, H., 2007b. Tree roots - Methodological review and new development in dating  
390 and quantifying erosive processes. Geomorphology 86, 243-51.

391 Gärtner, H., Heinrich, I., 2010. Anatomie des cernes annuels chez les arbres et les  
392 arbustes, in: Payette, S., Filion, L. (Eds.), La Dendroécologie: Principes, méthodes et  
393 applications. Les Presses de l'Université Laval, Québec, Québec, pp. 33-60.

394 Gärtner, H., Schweingruber, F.H., Dikau, R., 2001. Determination of erosion rates by  
395 analyzing structural changes in the growth pattern of exposed roots.  
396 Dendrochronologia 19, 81-91.

397 Gärtner, H., Stoffel, M., Lièvre, I., Monbaron, M., 2003. Tree ring analyses and detailed  
398 geomorphological mapping on a forested debris flow cone in Switzerland, in:  
399 Rickenmann, D., Chen, C. (Eds.), Debris flow hazards mitigation: mechanics,  
400 prediction, and assessment, 1, pp. 207-17.

401 Garcia Gonzalez, I., Fonti, P., 2006. Selecting earlywood vessels to maximize their  
402 environmental signal. Tree Physiology 26, 1289-96.

403 George, S.S., Nielsen, E., Conciatori, F., Tardif, J., 2002. Trends in *Quercus*  
404 *macrocarpa* vessel areas and their implications for tree-ring paleoflood studies. Tree-  
405 Ring Research 58, 3-10.

406 Gers, E., Florin, N., Gärtner, H., Glade, T., Dikau, R., Schweingruber, F.H., 2001.  
407 Application of shrubs for dendrogeomorphological analysis to reconstruct spatial and  
408 temporal landslide movement patterns - a preliminary study. *Zeitschrift für*  
409 *Geomorphologie*, Supplement Band 125, 163-75.

410 Giardino, J.R., Shroder, J.F., Lawson, M.P., 1984. Tree-ring analysis of movement of a  
411 rock-glacier complex on Mount Mestas, Colorado, U.S.A. *Arctic and Alpine*  
412 *Research* 16, 299-309.

413 Haeberli, W., 1985. Creep of mountain permafrost: Internal structure and flow of alpine  
414 rock glaciers. *Mitteilungen der VAW/ETH Zürich*, 77 pp.

415 Hallinger, M., Manthey, M., Wilmking, M., 2010. Establishing a missing link: warm  
416 summers and winter snow cover promote shrub expansion into alpine tundra in  
417 Scandinavia. *New Phytologist* 186, 890-99.

418 Heinrich, I., Gärtner, H., 2008. Variations in tension wood of two broadleaved tree  
419 species in response to different mechanical treatments: Implications for  
420 dendrochronology and mass movement studies. *International Journal of Plant*  
421 *Sciences* 169, 928-36.

422 Heinrich, I., Gärtner, H., Monbaron, M., 2007. Tension wood formed in *Fagus sylvatica*  
423 and *Alnus glutinosa* after simulated mass movement events. *IAWA Journal* 28, 39-  
424 48.

425 Jakob, M., 1995. Dendrochronology to measure average movement rates of gelifluction  
426 lobes. *Dendrochronologia* 13, 141-46.

427 Kääh, A., Frauenfelder, R., Roer, I., 2007. On the response of rockglacier creep to  
428 surface temperature variations. *Global Planetary Change* 56, 172-87.

429 Kääh, A., Huggel, C., Fischer, L., Guex, S., Paul, F., Roer, I., Salzmann, N., Schlaefli,  
430 S., Schmutz, K., Schneider, D., Strozzi, T., Weidmann, Y., 2005. Remote sensing of

431 glacier- and permafrost-related hazards in high mountains: an overview. *Natural*  
432 *Hazards and Earth System Sciences* 5, 527-54.

433 Körner, C., Hoch, G., 2006. A test of treeline theory on a montane permafrost island.  
434 *Arctic, Antarctic and Alpine Research* 38, 113-19.

435 Kozlowski, T.T., 1979. *Tree growth and environmental stress*. University of  
436 Washington Press, Seattle, WA, 192 pp.

437 LaMarche, V.C., 1968. Rates of slope degradation as determined from botanical  
438 evidence, White Mountains, California. US Geological Survey Professional Paper  
439 352-I, 376 pp.

440 Owczarek, P., 2010. Talus cone activity recorded by tree-rings of Arctic dwarf shrubs: a  
441 study case from SW Spitsbergen, Norway. *Geologija* 52, 34-39.

442 Paolini, L., Villalba, R., Grau, H.R., 2005. Precipitation variability and landslide  
443 occurrence in a subtropical mountain ecosystem of NW Argentina.  
444 *Dendrochronologia* 22, 175-80.

445 Roer, I., 2007. Rockglacier kinematics in a high mountain geosystem. *Bonner*  
446 *Geographische Abhandlungen* 117, 217 pp.

447 Roer, I., Nyenhuis, M., 2007. Rockglacier activity studies on a regional scale:  
448 comparison of geomorphological mapping and photogrammetric monitoring. *Earth*  
449 *Surface Processes and Landforms* 32, 1747-58.

450 Roer, I., Gärtner, H., Heinrich, I., 2007. Dendrogeomorphological analysis of alpine  
451 trees and shrubs growing on active and inactive rockglaciers, in: Haneca, K.,  
452 Verheyden, A., Beekmann, H., Gärtner, H., Helle, G., Schleser, G. (Eds.), *TRACE –*  
453 *Tree Rings in Archaeology, Climatology and Ecology* 5, pp. 248-58.

454 Roer, I., Kääb, A., Dikau, R., 2005a. Rockglacier kinematics derived from small-scale  
455 aerial photography and digital airborne pushbroom imagery. *Zeitschrift für*  
456 *Geomorphologie* 49, 73-87.

457 Roer, I., Kääh, A., Dikau, R., 2005b. Rockglacier acceleration in the Turtmann valley  
458 (Swiss Alps): probable controls. Norsk Geografisk Tidsskrift - Norwegian Journal of  
459 Geography 59, 157-63.

460 Sahling, I., Schmidt, K.H., Gärtner, H., 2003. Dendrogeomorphological analysis of the  
461 enlargement of cracks at the Wellenkalk-scarp in the southern Thuringia Basin, in:  
462 Schleser, G., Winiger, M., Bräuning, A., Gärtner, H., Helle, G., Jansma, E.,  
463 Neuwirth, B., Treydte, K. (Eds.), TRACE - Tree Rings in Archaeology, Climatology  
464 and Ecology 1, pp. 125-30.

465 Sass, U., Eckstein, D., 1995. The variability of vessel size in beech (*Fagus sylvatica* L.)  
466 and its ecophysiological interpretation. Trees 9, 247-52.

467 Schweingruber, F.H., 1983. Der Jahrring. Standort, Methodik, Zeit und Klima in der  
468 Dendrochronologie. Haupt, Bern, 234 pp.

469 Schweingruber, F.H., 1996. Tree rings and environment. Dendroecology. Haupt, Bern,  
470 609 pp.

471 Schweingruber, F.H., Poschlod, P., 2005. Growth rings in herbs and shrubs: life span,  
472 age determination and stem anatomy. Forest, Snow and Landscape Research 79, 195-  
473 415.

474 Shroder, J.F., 1978. Dendrogeomorphological analysis of mass movement on Table  
475 Cliffs Plateau, Utah. Quaternary Research 9, 168-85.

476 Shroder, J.F., 1980. Dendrogeomorphology: Review and new techniques of tree-ring  
477 dating. Progress in Physical Geography 4, 161-88.

478 Sorg, A., Bugmann, H., Bollschweiler, M., Stoffel, M., 2010. Debris-flow activity along  
479 a torrent in the Swiss Alps: Minimum frequency of events and implications for forest  
480 dynamics. Dendrochronologia 28, 215-23.

481 Strunk, H., 1991. Frequency distribution of debris flows in the Alps since the “little ice  
482 age”. Zeitschrift für Geomorphologie, Supplement Band 83, 71-81.

483 Van Tatenhove, F., Dikau, R., 1990. Past and present permafrost distribution in the  
484 Turtmantal, Wallis, Swiss Alps. *Arctic and Alpine Research* 22, 302-16.

485 Zoltai, S.C., 1975. Tree ring record of soil movements on permafrost. *Arctic and Alpine*  
486 *Research* 7, 331-40.

487

488

489 **Figure captions**

490

491 Figure 1: Swiss willow shrub growing on an active rockglacier (left) and after sampling  
492 in the lab (right). The black line indicates the position of the root sample cut off for  
493 anatomical analysis.

494

495 Figure 2: Digital micro-photo of a lateral root of Swiss willow which grew on the active  
496 rockglacier, note the distinct annual rings and the thick surrounding bark.

497

498 Figure 3: Location of the sample areas on the rockglaciers and of the study area  
499 (Turtmann valley) within Switzerland.

500

501 Figure 4: Location of the sample sites on the active (white circle) and on the inactive  
502 rockglacier (black circle), note the extraordinary high horizontal velocities of the active  
503 lobe. Underlying orthophoto of 1993; based on aerial images © Swisstopo.

504

505 Figure 5: Micro-sections of Swiss willow roots from an active (left) and inactive  
506 rockglacier (right); black arrows indicate vessels (width of white bar in images = 0.5  
507 mm).

508

509 Figure 6: Comparison of vessel lumen area (VLA) values averaged for single roots from  
510 the active (N=17) and inactive rockglacier (N=17) (upper diagrams) as well as for the  
511 single shrubs (N= 10 per site) (based on the values of their single roots) (lower  
512 diagrams).

513

514 Figure 7: Box-plot diagrams indicating the variation of vessel lumen area (VLA) values  
515 of all single roots on the active and inactive rockglacier (upper diagrams), the lower  
516 diagram shows the difference between the average VLA-values of all single roots  
517 growing on the active and inactive rockglacier.

518

519 Figure 8: Box-plot diagrams based on the data from within the range of the upper and  
520 lower quartile around the respective median values in Figure 7. Upper diagrams indicate  
521 the variation of VLA-values of single roots on the active and inactive rockglacier, the  
522 lower diagram shows the averaged VLA-values of all roots growing on the active and  
523 inactive rockglacier.

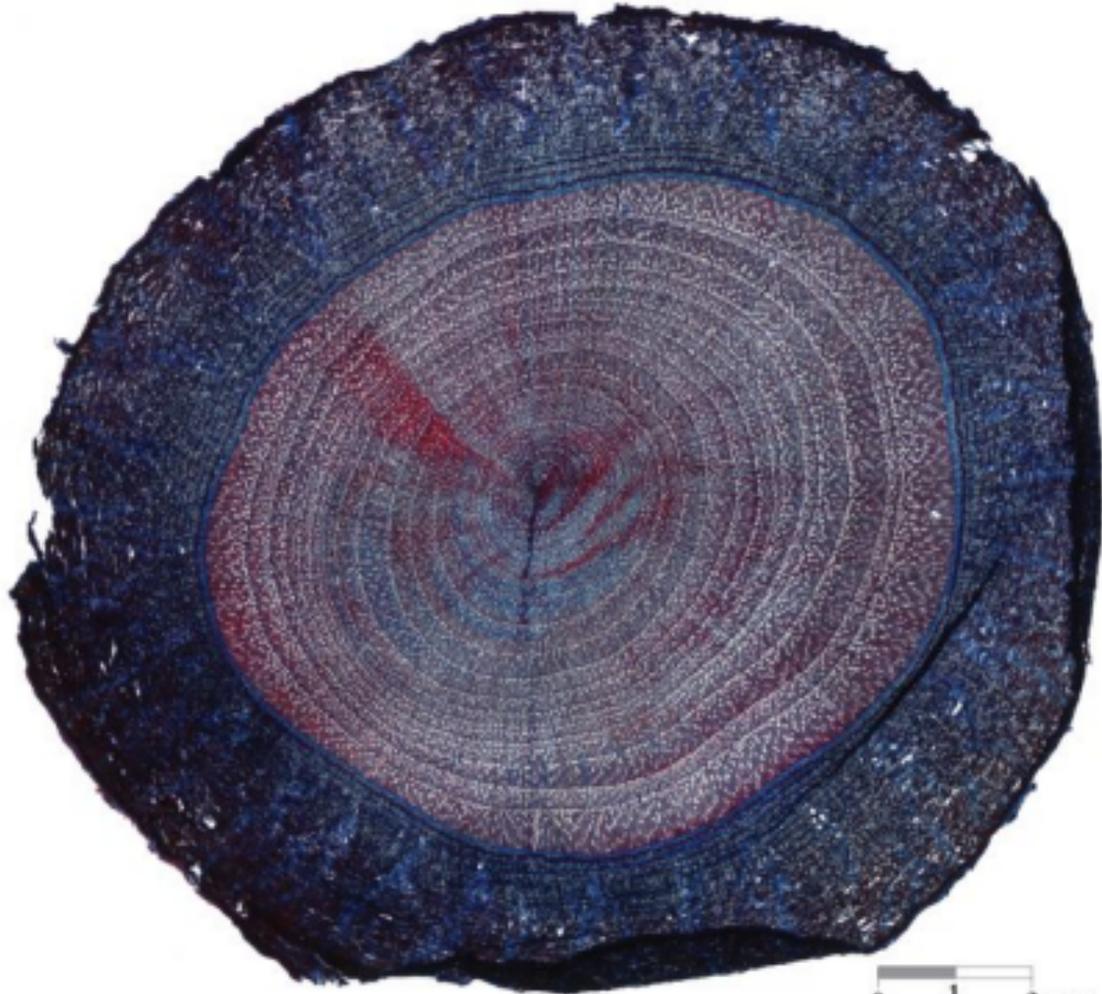
524

525 Figure 9: Box-plot diagrams showing the size variations of the 10% biggest (left) and  
526 35% biggest vessels (right) in roots for both rockglaciers. The upper two diagrams  
527 indicate the variation of VLA-values of single roots on the active and inactive  
528 rockglacier, the lower diagram shows the averaged VLA-values of all roots growing on  
529 the active and inactive rockglacier.

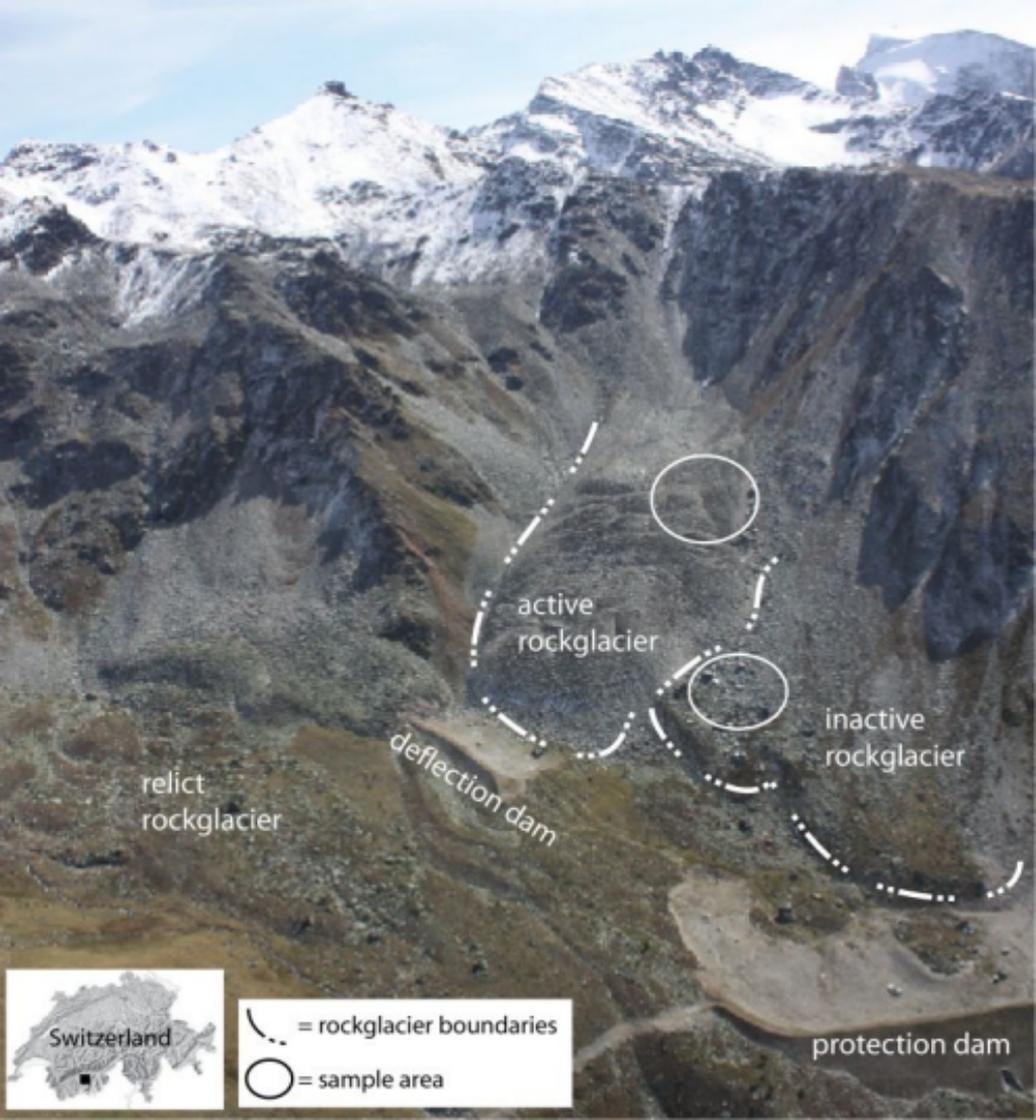
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531





0 1 2 mm



relict  
rockglacier

active  
rockglacier

inactive  
rockglacier

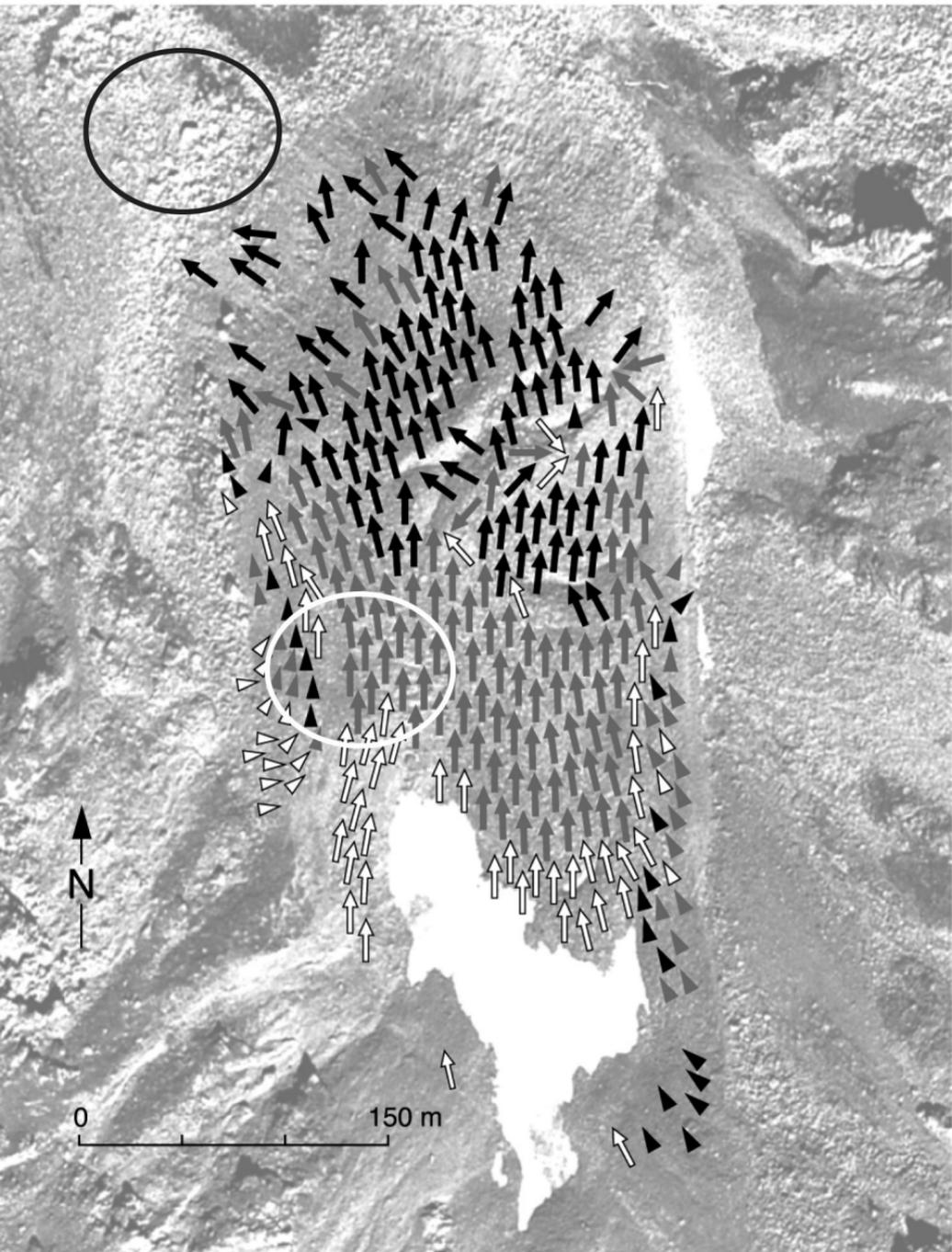
deflection dam

protection dam



Switzerland

-  = rockglacier boundaries
-  = sample area



**Annual horizontal velocity (m/a) 1987 - 1993**

△ 0 - 0.15

▲ 0.16 - 0.3

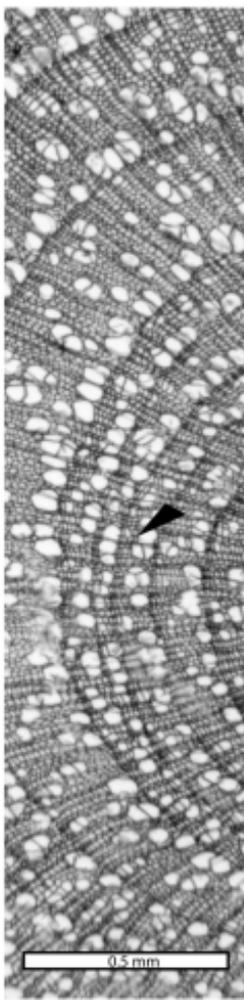
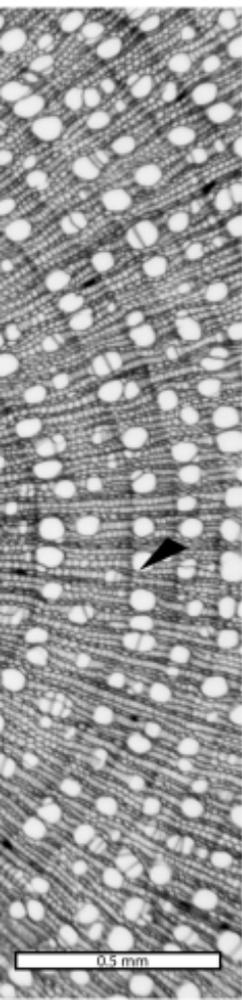
▲ 0.31 - 0.5

↑ 0.51 - 1

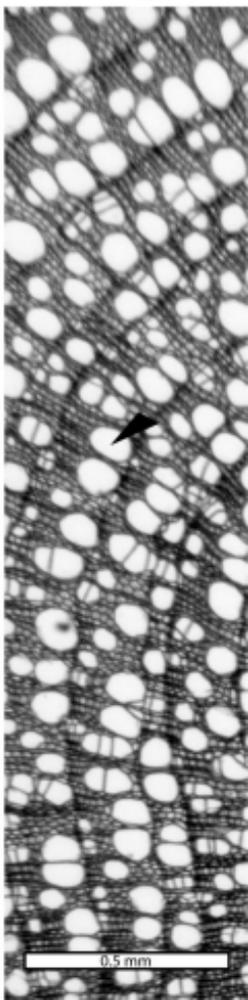
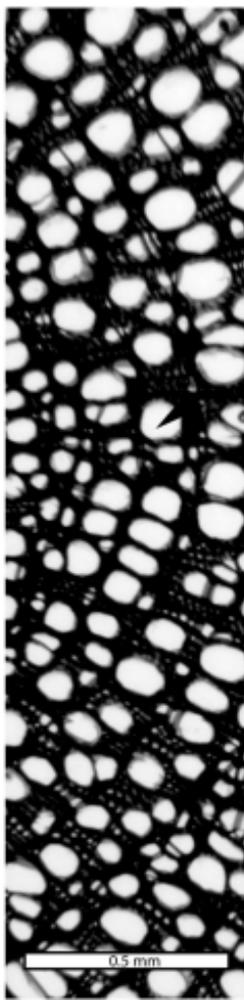
↑ 1.01 - 2

↑ 2.01 - 4.82

**Root vessels, active rockglacier**

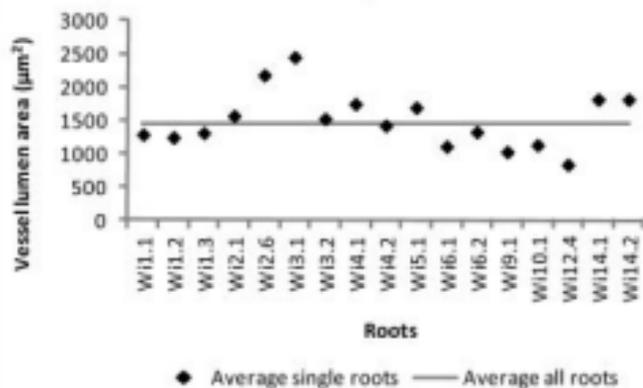


**Root vessels, inactive rockglacier**



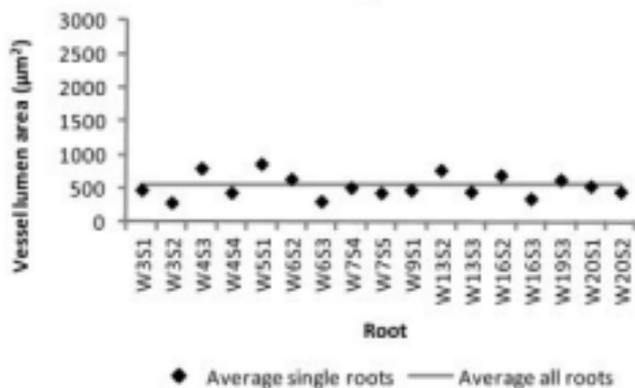
## Inactive rockglacier

Average VLA-values of single roots

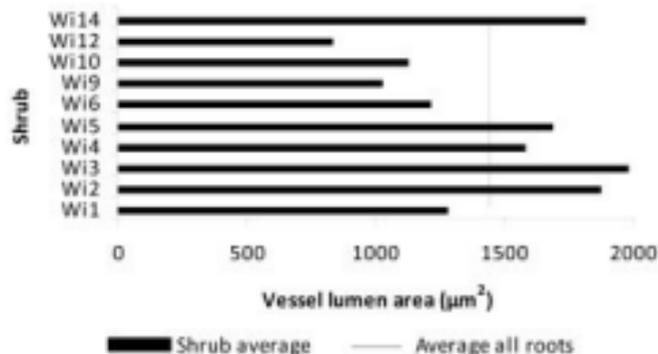


## Active rockglacier

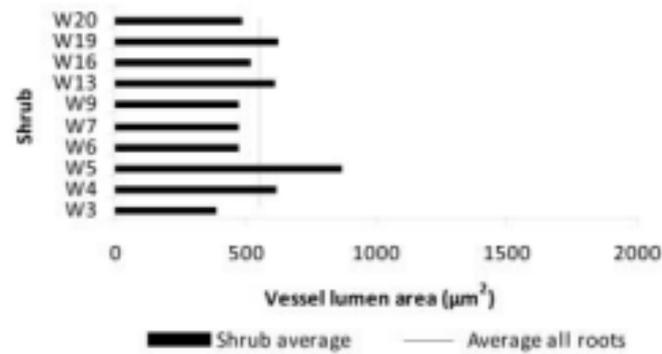
Average VLA-values of single roots



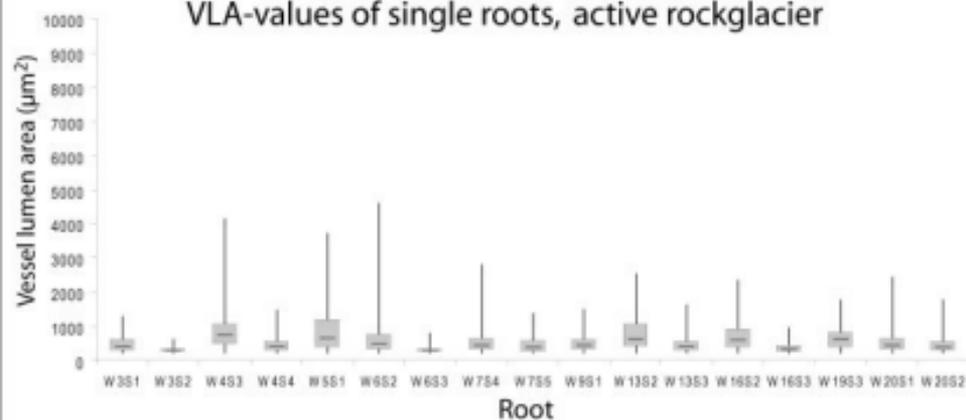
Average VLA-values of single shrubs



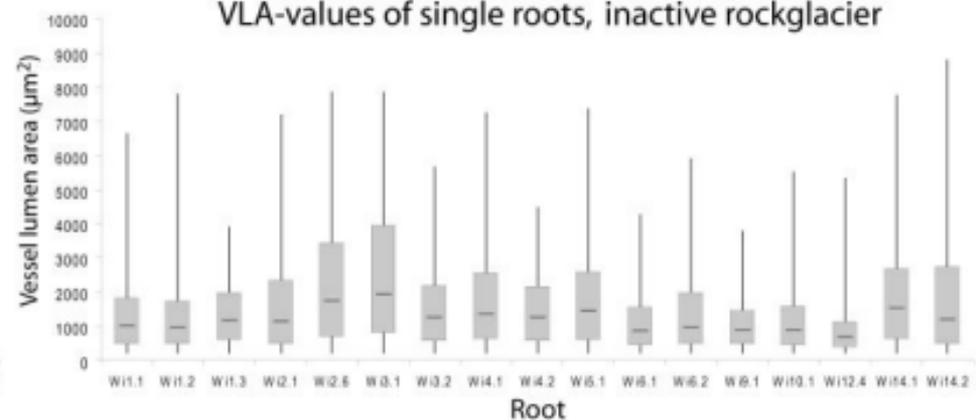
Average VLA-values of single shrubs



VLA-values of single roots, active rockglacier



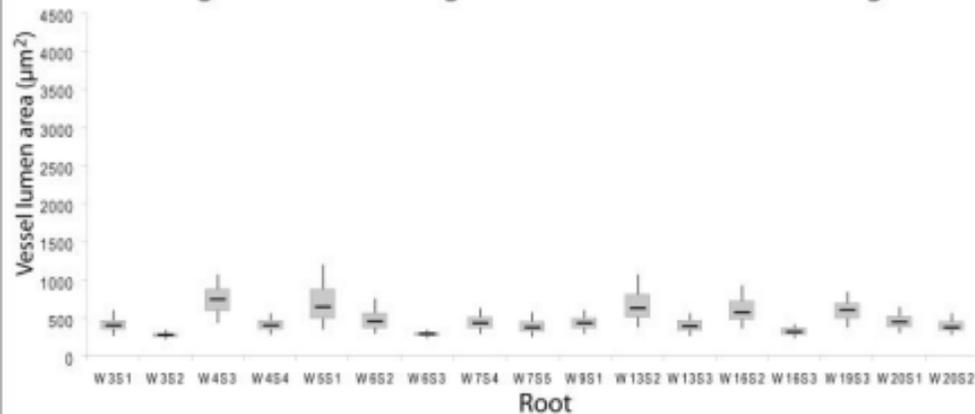
VLA-values of single roots, inactive rockglacier



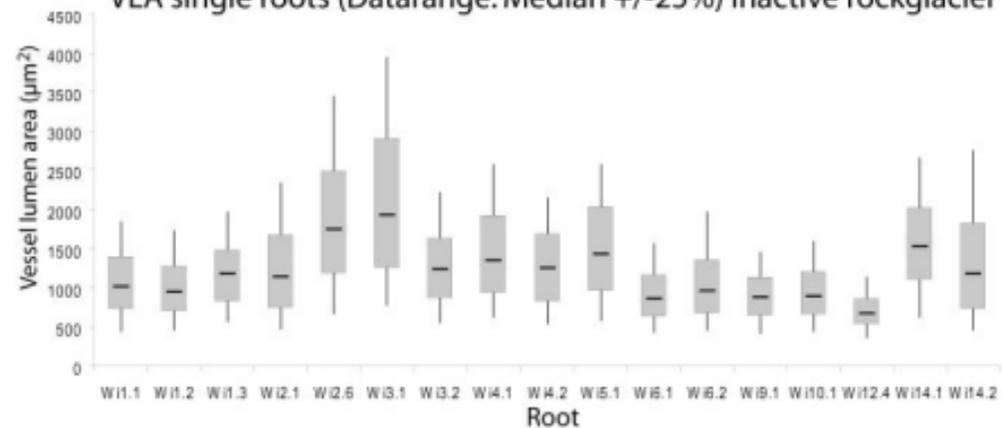
Root analysis: Average VLA-values on rockglaciers



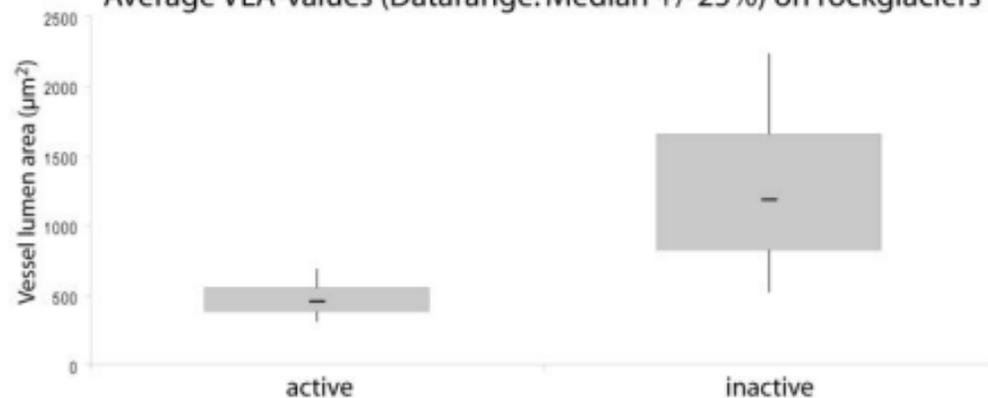
VLA single roots (Datarange: Median +/-25%) active rockglacier



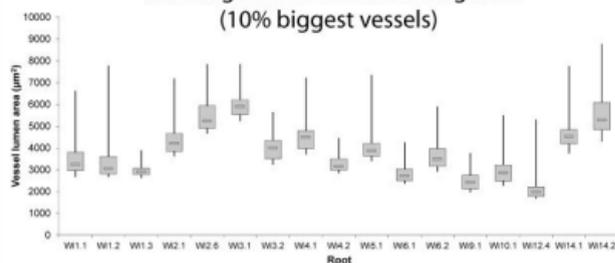
VLA single roots (Datarange: Median +/-25%) inactive rockglacier



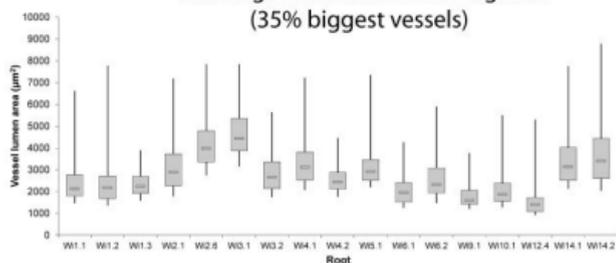
Average VLA-values (Datarange: Median +/-25%) on rockglaciers



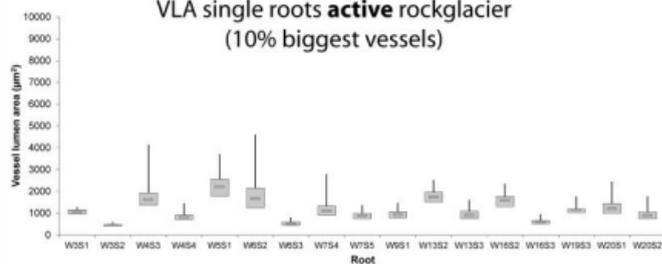
VLA single roots **inactive** rockglacier  
(10% biggest vessels)



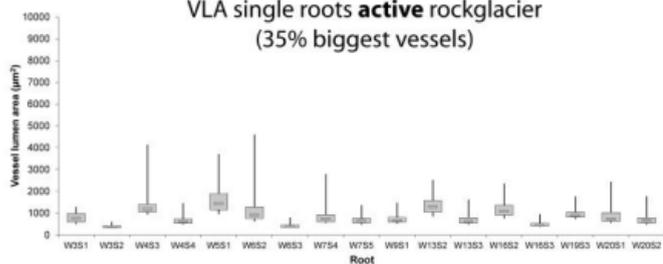
VLA single roots **inactive** rockglacier  
(35% biggest vessels)



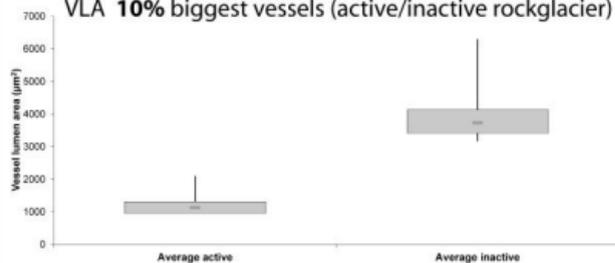
VLA single roots **active** rockglacier  
(10% biggest vessels)



VLA single roots **active** rockglacier  
(35% biggest vessels)



VLA **10%** biggest vessels (active/inactive rockglacier)



VLA **35%** biggest vessels (active/inactive rockglacier)

