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1 **Subaqueous hot springs in Köyceğiz Lake, Dalyan Channel and Fethiye-**  
2 **Göcek Bay (SW Turkey): locations, chemistry and origins**

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27 **ABSTRACT**

28 In this study, horizontal temperature measurements along organized grids have been used to  
29 detect subaqueous hot springs. The study area, located in the southwest of Turkey and comprised  
30 of Köyceğiz Lake, Dalyan Channel and Fethiye-Göcek Bay, was scanned by measuring  
31 temperatures horizontally, 2-3 meters above the bottom of the lake or sea. After analyzing the  
32 temperature data along the grids, the locations with anomalous temperature values were detected,  
33 and divers headed here for further verification. Accordingly, among these anomalies, the divers  
34 confirmed seven of them as subaqueous hot springs. Three of these hot springs are located in the  
35 Köyceğiz Lake, three of them are located in the Dalyan Channel and one hot spring is located in  
36 the Fethiye-Göcek Bay. At the locations where temperature anomalies were detected, the divers  
37 collected samples directly from the subaqueous hot spring using a syringe-type sampler. We  
38 evaluated these water samples together with samples collected from hot and cold springs on land  
39 and from local rivers, lakes and the sea, with an aim to generate a conceptual hydrogeochemical  
40 model of the geothermal system in the study area. This model predicts that rainwater precipitating  
41 in the highlands percolates through fractures and faults into the deeper parts of the Earth's crust,  
42 here it is heated and ascends through the sea bottom via buried faults. Pervious carbonate nappes  
43 that are underlain and overlain by impervious rocks create a confined aquifer. The southern  
44 boundary of the Carbonate-Marmaris nappes is buried under alluvium and/or sea/lake water  
45 bodies and this phenomenon determines whether hot springs occur on land or subaqueous. The  
46 chemical and isotopic properties of the hot springs point to seawater mixing at deep levels. Thus,  
47 the mixing most probably occurs while the water is ascending through the faults and fractures.  
48 The gas geochemistry results reveal that the lowest mantle He contributions occur in the samples  
49 from Köyceğiz Lake, whereas the highest ones are found in samples from the Dalaman plain. For  
50 the first time, we made use of the micro-XRF sediment core scanning (ITRAX Scanner) for  
51 exploring the relation between subaqueous geothermal occurrence and chemical properties of the

52 surrounding sediments. The spatial elemental distribution of sea/lake bottom sediments suggests  
53 that depending on the surrounding rock units and the temperature of the hot spring, the sediments  
54 around the spring can be enriched with certain elements.

55 **Key words:** Water chemistry, Mixing, Gas geochemistry, ITRAX analysis, Saturation index,  
56 Conceptual model

## 57 **1. Introduction**

58 The world's energy demands are increasing (IEA, 2012) along with the global population and  
59 resulting rise in consumption, which in turn leads to further exploration of alternative and  
60 renewable energy sources by investors, researchers, and governments. Geothermal energy  
61 production has become popular since it is environmentally friendly when compared to stock-  
62 limited fossil fuels and the emissions of greenhouse gases associated with their burning. Along  
63 with the limited environmental concerns associated with geothermal energy, its sustainability  
64 will likely make it an indispensable energy source in the future.

65 Most of the on-land geothermal sites in Turkey are well studied. For example, there were a  
66 number of comprehensive studies carried out about the hydrogeochemistry of Turkish on-land  
67 geothermal sites (İlkışık, 1995; Balderer and Sýnal, 1997; Mutlu and Güleç, 1998; Vengosh et  
68 al., 2002; Tarcan, 2005; Süer et al., 2008; Baba and Sözbilir, 2012; Bundschuh et al., 2013;  
69 Karakuş and Şimşek, 2013; Özgür and Çalışkan, 2013). In addition, there exists an inventory  
70 of the geothermal sites issued by the General Directorate of Mineral Research and Exploration  
71 (MTA) in 2005. Tüfekçi et al. (2010) used this inventory together with the publicly available  
72 magnetic anomaly, Bouguer gravity anomaly, earthquake epicenter and lineament datasets and  
73 statistically showed the spatial distribution of potential geothermal occurrences in western  
74 Anatolia.

75 Geothermal energy started gaining popularity in Turkey after a geothermal law was passed in  
76 2007. Accordingly, geothermal investments in electrical power generation increased by 336%  
77 (Bertani, 2015), and the direct utilization of geothermal energy increased by 38% over the last  
78 five years (Mertoğlu et al., 2010; Mertoğlu et al., 2015). Moreover, the purchasing guarantee  
79 issued by the Turkish Government at a rate of 0.105 USD/kWh has provided the sector with a  
80 great incentive for investment. Mertoğlu et al. (2015) reported that there are over 2000 hot and  
81 mineral water sources, and 227 geothermal sites in Turkey, 70-80% of which are located in  
82 western Anatolia.

83 Not only in Turkey but also throughout the world, on-land geothermal sources have witnessed  
84 extensive exploration through geochemical, mineralogical, and hydrogeochemical  
85 investigations as well as expensive geophysical prospecting methods such as seismic  
86 stratigraphical methods. However, explorations for subaqueous geothermal resources have not  
87 yet had their deserved attention. There have been efforts to estimate the earth's heat loss and to  
88 assess shallow hydrothermal activity. For example, Williams (1976) stated that approximately  
89  $2 \times 10^{12}$  cal/s ( $= 8.37 \times 10^{12}$  W) of heat (20% of the earth's heat loss) is released through 1% of the  
90 Earth's surface area. This release is in the form of hydrothermal discharge from young rocks  
91 near active seafloor-spreading areas and submarine volcanoes. This amount of energy was  
92 roughly equal to the world's total energy consumption at that time (1976). Similarly, Prol-  
93 Ledesma et al. (2005) described "shallow hydrothermal activity" as an interesting but not-as-  
94 yet well-studied phenomenon, where shallow means a water depth less than 200 m, and they  
95 stated "Shallow-water vents provide accessible geological and chemical settings for studying  
96 the interaction of hydrothermal fluids with unconsolidated sediments, seawater, and basement  
97 rocks". More recently, Suarez-Arriaga et al. (2014) declared that submarine geothermal  
98 resources have not been sufficiently used and studied yet, but they could cover the world's

99 considerable energy demand of in an environmentally friendly way by integrating on-land  
100 geothermal prospecting experience and offshore hydrocarbon exploitation technology.

101 Although submarine and sub-lacustrine hydrothermal systems are not as well studied as on-  
102 land geothermal sources, recent technological developments have drawn many researchers to  
103 undertake studies on subaqueous hydrothermal systems. These studies include the ones carried  
104 out in lakes (Michel et al., 2001; Remsen et al., 2002; Ronde et al., 2002; De Batist et al., 2002;  
105 Duchkov et al., 2009), and in different parts of the world including Greece, Italy, Japan, the  
106 Antilles, Galapagos rift zone, Iceland, Mexico, the Pacific Ocean, Papua New Guinea, and  
107 Turkey (Williams et al., 1974; Corliss et al., 1979; Italiano and Nuccio, 1991; McMurtry et al.,  
108 1993; Michard et al., 1993; Fitzsimons et al., 1997; Hannington et al., 2001; Glasby and Notsu,  
109 2003; Prol-Ledesma, 2003; Prol-Ledesma et al., 2004; McCarthy et al., 2005; Pichler, 2005;  
110 Valsami-Jones et al., 2005; Esposito et al., 2006; Koschinsky et al., 2007; Pekçetinöz et al.,  
111 2009; Arango-Galván et al., 2011; Yamanaka et al., 2013). However, most of the hydrothermal  
112 systems in these studies were either discovered by chance or identified by local people.

113 This study presents the results of the exploratory research of subaqueous thermal systems that  
114 were expected to occur at the bottom of Fethiye-Göcek Bay as well as the Köyceğiz, Alagöl,  
115 Sülüngür and Kocagöl lakes (Fig. 1), according to the measurements of water temperature along  
116 horizontal lines taken to detect temperature anomalies. In addition, a hydrogeochemical  
117 conceptual model of the region was constructed by using water, gas and sediment geochemistry.  
118 The study area is located in southwestern Anatolia, which is a seismically active region.  
119 Additionally, the Fethiye-Göcek shoreline is located along one of the main fault zones of SW  
120 Anatolia, namely, the Fethiye-Burdur fault zone (on land continuation of Pliny-Strabo fault  
121 zone) (Fig. 1). There are lineations recorded by Yağmurlu et al. (2005), Ocakoğlu (2011) and  
122 faults recorded by Hall et al. (2014) in the south of continental shelf of Fethiye-Göcek Bay.  
123 Findings of these researches favor the idea that there are submarine buried faults at the bottom

124 of the Fethiye-Göcek Bay, these may enable submarine geothermal systems, making the  
125 Fethiye-Göcek Bay a potential site to investigate for submarine hot springs.

## 126 **2. Geological setting**

127 The Menderes Massif, the Beydağları Autochthon and the Lycian nappes are the main tectonic  
128 units of SW Anatolia (Fig. 1). The study area is located mainly on the Lycian nappes; the  
129 Beydağları Autochthon crops out in the region as tectonic windows. The Beydağları  
130 Autochthon is composed of Upper Cretaceous carbonate rocks and is the base rock in the region.  
131 The Lycian nappes are thrusting over the Beydağları Autochthon tectonically. Although there  
132 are many categorizations for the Lycian nappes, in this study we mainly use the one of Şenel  
133 (1997a; 1997b) where the nappes are divided into 5 main units: the Yeşilbarak, Tavas, Bodrum,  
134 Gülbahar, and Marmaris Nappes from the bottom to the top (Fig. 1).

135 The Late Lutetian-Early Burdigalian aged Yeşilbarak nappe is composed of shales and  
136 sandstones and located between the Beydağları Autochthon and the overlying Lycian nappe  
137 units. The overlying Tavas, Bodrum and Gülbahar nappes are composed of carbonate units and  
138 presented together (without differentiating) in the geological map as their hydrogeological  
139 properties are similar in the region. The Lower Cretaceous Marmaris nappe comprising  
140 peridotites overlies all rock units in the region and crops out as a NW-SE trending strip. The  
141 impervious to semi-pervious character of this unit greatly affects the hydrogeology, and  
142 consequently the geothermal system of the region. This subject matter will be discussed in the  
143 Hydrogeochemical Model section (Section 4.4.4) in detail.

## 144 **3. Materials and methods**

145 In order to detect the temperature anomalies, we measured the temperature of the water together  
146 with electrical conductivity (EC) and pH along horizontal linear routes. Where temperature

147 anomalies were detected, divers were sent to these localities to seek and confirm the presence  
148 of the hot springs. If there was a hot spring, water samples were taken, to be analyzed  
149 chemically. Together with the subaqueous samples, on-land hot and cold springs, stream water,  
150 seawater, lake water and channel water were sampled and chemically analyzed. The results of  
151 these chemical analyses were used to develop a hydrogeochemical model. Along with water  
152 chemistry, we also studied gas and sediment geochemistry. The following subsections present  
153 further details.

### 154 **3.1 Subaqueous hot spring exploration/prospecting**

155 In order to explore subaqueous hot springs in the study area, a total of about 900 km of cruising  
156 was carried out (Fig. 2), using a boat equipped with a sonar-mounted Global Positioning System  
157 (GPS) device (Garmin GPSmap 421s) and a multi-parameter conductivity-temperature-depth  
158 (CTD) sonde (YSI 6600), which allowed simultaneous real-time measurements of geographical  
159 coordinates, water depth and temperature. During the cruising, based on the water depth  
160 information from the sonar, the CTD was lowered down to a certain height above the  
161 sediment/water interface (2-3 m), and its depth was modified with respect to the bathymetrical  
162 changes. In this way, we ensured that we measured the water temperature as close as possible  
163 to the probable subaqueous hot springs while preventing the CTD from crashing to the ground.  
164 By setting the cruise speed to ca. 7 km/h and the devices to perform measurements every two  
165 seconds, we also ensured that the CTD and GPS measurements were done every 4 meters along  
166 the cruising routes. As a preliminary approach, the routes were designed in order to cover the  
167 complete target area of the bay and lakes. After analyzing the data, promising sites were  
168 identified, and more detailed measurements were done around these locations via fine grid  
169 routes.

170 In order to detect water temperature anomalies, which could be related to subaqueous hot  
171 springs, the datasets from the CTD and GPS were merged and synchronized to obtain  
172 temperature vs. time plots. Careful inspection of those plots allowed the determination of the  
173 approximate locations of the “potential” subaqueous hot springs. Afterwards, more detailed  
174 exploratory cruises with denser survey grids were conducted at the “potential” locations in order  
175 to confirm the anomalies and to provide the divers with more precise coordinates for water  
176 sampling. While measuring the temperature, EC and pH, sonar profiles were taken by the GPS.  
177 These sonar profiles were recorded by taking screenshots from the GPS device, and they were  
178 investigated and used as a tool to detect subaqueous hot springs. Moreover, some topographic  
179 anomalies, which may be related to the faults, were detected using the detailed bathymetric data  
180 generated by sonar GPS measurements. The CTD measurements were concentrated around  
181 these locations by designing a finer cruise grid.

### 182 **3.2 Hydrochemistry**

183 Whenever the divers confirmed the presence of a hot spring, they collected water samples with  
184 syringe-type samplers, measured the temperature of the spring water by putting the CTD at the  
185 outlet of the hot spring, and took photos as well as videos. The water samples were collected  
186 directly from the outlet of the spring to avoid or minimize mixing with sea/lake water. The  
187 water samples in the syringes were transferred into three different polyethylene bottles (100 ml  
188 each) for stable isotope, cation+trace element, and anion analyses, respectively.

189 Samples collected for cation and anion analyses passed through a 0.45  $\mu\text{m}$  filter, and the former  
190 ones were acidified to 1%  $\text{HNO}_3$  after filtering. Along with subaqueous hot spring samples,  
191 five on-land fresh water springs were also sampled (Fig. 1). In addition, three water samples  
192 were taken from the sea surface in order to delineate the chemical and isotopic properties of the  
193 seawater (Table 1). The water samples were kept refrigerated (4  $^{\circ}\text{C}$ ) and immediately sent to

194 dedicated laboratories. Major ion and trace element analyses were carried out at the Hacettepe  
195 University Water Chemistry and Environmental Tritium Laboratory (Ankara, Turkey), and  
196 stable isotope analyses were carried out at the Hacettepe University International Karst Water  
197 Resources Application and Research Center Stable Isotope Laboratory (UKAM, Ankara,  
198 Turkey). The titration technique was used to determine carbonate ( $\text{CO}_3^{2-}$ ) and bicarbonate  
199 ( $\text{HCO}_3^-$ ) ion concentrations. Fluoride ( $\text{F}^-$ ), chloride ( $\text{Cl}^-$ ), nitrite ( $\text{NO}_2^-$ ), bromide ( $\text{Br}^-$ ),  
200 nitrate ( $\text{NO}_3^-$ ), phosphate ( $\text{PO}_4^{3-}$ ), sulfate ( $\text{SO}_4^{2-}$ ), lithium ( $\text{Li}^+$ ), sodium ( $\text{Na}^+$ ), ammonium  
201 ( $\text{NH}_4^+$ ), potassium ( $\text{K}^+$ ), magnesium ( $\text{Mg}^{2+}$ ), and calcium ( $\text{Ca}^{2+}$ ) ion concentrations were  
202 determined by using a high performance ion chromatography device (DIONEX LC25, ICS-  
203 1000 Ion Chromatography system) and a conductivity detector. The spectrophotometric method  
204 (aT60U Spectrometer) was used to determine  $\text{SiO}_2$  and boron concentrations. Trace element  
205 (Ag, Al, As, Au, Ba, Be, Cd, Ce, Co, Cr, Cs, Cu, Dy, Er, Eu, Fe, Ga, Gd, Ge, Hf, Hg, Ho, Ir,  
206 La, Lu, Mn, Mo, Nb, Nd, Ni, Pb, Pd, Pr, Pt, Rb, Re, Rh, Ru, Sb, Se, Sm, Sn, Sr, Ta, Tb, Te, Th,  
207 Ti, Tl, Tm, U, V, W, Y, Yb, Zn, Zr) concentrations were measured in a Thermo Electron X7  
208 model ICP-MS (Inductively Coupled Plasma–Mass Spectrometer) and the detection limits  
209 range between 0.001 ppt to 0.3 ppb. Chemical analyses of the waters were conducted in  
210 compliance with Clesceri et al. (1989). Stable isotope analyses were conducted by using laser  
211 spectroscopic methods, where the reference standard was Vienna Standard Mean Ocean Water  
212 (VSMOW). The analytical precisions are 0.2‰ and 1‰ for  $\delta^{18}\text{O}$  and  $\delta\text{D}$ , respectively.

213 Helium analyses of eleven samples were conducted at the GFZ Potsdam noble gas laboratory  
214 in Germany. Four of those samples were collected from subaqueous springs, and seven of them  
215 from on-land springs. At seven locations, including the subaqueous spring SUBC-2, free gas  
216 bubbling through the water was collected in glass flasks by means of a funnel, while at four  
217 locations spring water was sampled in 1 cm outer diameter, 50 cm long copper tubes sealed by  
218 steel clamps. During sampling, special care was taken to avoid atmospheric contamination.



219 Except for the gas sample SUBC-2 (see above), the other subaqueous springs were sampled by  
220 installing a small water pump at the outlet of the spring. In the lab, total gas analyses were  
221 carried out for the free gas samples using a Pfeiffer Omnistar quadrupole mass spectrometer.  
222 Detection limits in these analyses are ~10 ppmv for N<sub>2</sub>, O<sub>2</sub>, CO<sub>2</sub>, and ~1 ppmv for H<sub>2</sub> and Ar  
223 (Wiersberg et al., 2011). The glass flasks were then attached to the noble gas line, and  
224 appropriate volume splits were analyzed for noble gas concentrations and isotopic  
225 compositions. Water samples from copper tubes were discharged into a glass bulb and degassed  
226 in an ultrasonic bath, and the gases were transferred to a cold trap by means of a laminar flow  
227 of gaseous H<sub>2</sub>O (evaporated from the liquid water sample) through a connecting capillary.  
228 Again, appropriate splits were transferred to the noble gas line. There, chemically active gases  
229 were removed using another cold trap, two Ti sponge getters, and two SAES (Zr-Al) getters,  
230 and the noble gases were adsorbed to a stainless steel frit or activated charcoal in two cryostatic  
231 cold heads in order to separate them from each other for the measurement in a VG5400 noble  
232 gas mass spectrometer. Further details about experimental procedures and data reduction  
233 methods can be found in Niedermann et al. (1997).

### 234 **3.3 Sediment geochemistry**

235 In order to spatially evaluate the geochemical composition of the bottom sediments, and  
236 therefore in order to detect the possible geochemical effects of subaqueous hydrothermal fluid  
237 emissions on sedimentation, gravity cores were collected and analyzed by an ITRAX  $\mu$ XRF  
238 core scanner. For an unbiased spatial comparison, it is crucial to do the geochemical analyses  
239 for the same stratigraphical horizon (i.e., on the sediments that were deposited in the same time  
240 window). For this reason, this study targeted the sediments at the water/sediment interface.  
241 Accordingly, rather than using an Ekman sampler that may significantly mix and disturb  
242 sediments, we used a gravity corer with a messenger-triggered vacuum system. Moreover, this

243 method helped to achieve minimum sediment loss and disturbance at the water/sediment  
244 interface.

245 A total of 233 gravity cores, 30-50 cm long, with a diameter of 7.5 cm, were collected (99 from  
246 Fethiye-Göcek Bay, 105 from Köyceğiz Lake, and 29 from Dalyan Channel). The cores were  
247 collected in two groups: 1) Cold Location Cores (CLC), which were almost uniformly  
248 distributed within the study areas to reveal the general characteristics of sediments, and 2) Hot  
249 Location Cores (HLC), which were collected in the vicinity of the detected subaqueous hot  
250 springs. The geochemical data from the CLCs and HLCs were then compared to investigate the  
251 possible effects of subaqueous hot springs on the chemical composition of sediments.

252 Since splitting and scanning 233 cores one by one would require a significant amount of labor  
253 and time, only the top 3-5 cm of each gravity core were sampled by using 2 cm diameter  
254 syringes, and put successively in PVC half pipes. In this way, ten artificial cores, being  
255 approximately 1 m long and containing sediments from 20-30 locations each, were prepared  
256 and sent to the GEOPOLAR Laboratory of the University of Bremen for ITRAX  $\mu$ XRF  
257 scanning to obtain geochemical data as well as optical and radiographic images.

258 Along the artificial cores, radiographic scanning was done at 200  $\mu$ m resolution. XRF scanning,  
259 on the other hand, was carried out with a Mo-tube at 1 mm increments for the Fethiye-Göcek  
260 cores, and at 2 mm increments for the Köyceğiz and Dalyan cores, with an exposure time of 12  
261 seconds. As a result of the XRF scanning, semi-quantitative distributions of 19 elements (Si, S,  
262 Cl, K, Ca, Ti, Cr, Mn, Fe, Ni, Zn, As, Se, Br, Rb, Sr, Y, Zr, and Pb) were obtained in units of  
263 kilo counts per second (kcps). For each element, average values for all sampling locations were  
264 calculated. In order to achieve comparable scales among the elements for spatial evaluation,  
265 average kcps values were standardized separately for the three study sites, which have different  
266 sedimentological environments.

267 Since ITRAX  $\mu$ XRF scanning is performed on bulk sediments, the scanning results are also  
268 influenced by the physical properties of the sediments, such as grain-size distribution, water,  
269 and organic matter contents. In order to minimize these effects and to better reveal the  
270 geochemical signal, inter-element ratios are used (e.g., Croudace et al., 2006; Kylander et al.,  
271 2011). As a first step, all possible inter-element ratios (i.e., 380 for 19 elements) were calculated  
272 for each location and their spatial distribution maps were prepared. However, visual inspection  
273 of spatial distribution maps does not always reveal the difference between the HLCs and CLCs.  
274 For this reason, the mean values of inter-element ratios from the two groups of cores (i.e., HLCs  
275 and CLCs) were determined and compared. Finally, those inter-element ratios of the HLCs that  
276 significantly deviated from the means of the CLCs were detected and were attributed to the  
277 effect of subaqueous hot springs on the chemical composition of the sediments.

## 278 **4. Results and discussions**

### 279 **4.1. Subaqueous hot spring exploration/prospecting**

280 The measurement routes and the bathymetric maps are shown in Fig. 2. Along with the  
281 temperature measurements, EC and pH values were also measured. As a result of these  
282 physicochemical measurements, on first inspection of the data, over one hundred temperature  
283 anomaly indications were detected (some of the selected anomalies can be seen in time vs.  
284 temperature plots provided in Supplementary Fig. 1). At several of these locations, detailed  
285 physicochemical measurements were performed by tracking along a finer mesh of routes (Fig.  
286 2a and b). As a result, divers were sent to 49 potential hot spring locations for confirmation of  
287 the presence of subaqueous hot springs. Among the 49 locations, they confirmed seven  
288 subaqueous hot springs: three in Köyceğiz Lake, three in the Dalyan Channel, and one in the  
289 Fethiye-Göcek Bay (Fig. 1). The temperature of the subaqueous hot springs in the Köyceğiz  
290 Lake ranges between 26.17 and 29.47 °C (Table 1). The discharge temperatures of the Dalyan

291 subaqueous springs are around 27-28 °C. The water depth and temperature of the subaqueous  
292 hot spring in the Fethiye-Göcek Bay was 19 m and 27.19 °C, respectively (VIDEO). No  
293 temperature anomalies were found at the bottom of the Alagöl, Sülüngür and Kocagöl lakes.

294 The other aid used in detecting hot springs were the sonar profiles. SUB-2 and SUB-3 springs  
295 in Köyceğiz Lake were detected by considering the lineations on the sonar profile as an  
296 indication of a subaqueous hot spring. These locations were also marked as anomalies, and  
297 divers were sent there in order to search for the hot springs under water. Bathymetry maps of  
298 Köyceğiz Lake and the Fethiye-Göcek Bay were prepared by using the sonar GPS data (Fig. 2c  
299 and d). Köyceğiz lake has two troughs in the north and the south and the deepest points are 25  
300 m and 32 m respectively. In the Fethiye-Göcek Bay the shelf is located around 125 m from the  
301 shoreline, and the water depth reaches 500 m in the outer shelf. Some lineations were detected  
302 in the bay through detailed investigation of the bathymetry map; they may be inferred as the  
303 faults at the bottom of the sea. Temperature anomalies that were detected at the nearshore  
304 section of these lineations made these areas potential subaqueous hot spring sites, so it was  
305 worth studying them in detail (Fig. 2b). Except for İnlice Bay, there were, however, no hot  
306 springs found at the remaining three locations.

#### 307 **4.2 Sediment geochemistry**

308 ITRAX  $\mu$ XRF scanning on basin-wide sediment samples allowed detection of the geochemical  
309 traces of subaqueous hot springs on the sediments. The locations of the cores collected in this  
310 study are presented in Fig. 3. The optical and radiographic images of the artificial cores, which  
311 are composed of the top 3-5 cm of sediment from 233 gravity cores, are shown in  
312 Supplementary Fig. 2. Along the radiographic images, especially, the boundaries between  
313 different cores can be clearly seen as horizontal white lines, which correspond to mm-scale low-  
314 density gaps. Before calculating the average “kcps” values for each element at coring locations,

315 several measurement points above and below these boundaries are excluded from the XRF  
316 dataset since they can be erroneous. After data purging, the average “kcps” values were  
317 standardized for each element in the three study sites separately (Köyceğiz Lake,  
318 Supplementary Table 1; Dalyan Channel, Supplementary Table 2; and Fethiye-Göcek Bay,  
319 Supplementary Table 3).

320 All possible elemental ratios derived from the standardized “kcps” values are plotted separately  
321 for the cold and hot locations of each site (Supplementary Fig. 3). In addition, these ratios are  
322 interpolated over the study areas to visualize their spatial distribution (Supplementary Fig. 4).  
323 Fig. 5 illustrates the Sr/Ti ratios at the coring locations in the Fethiye-Göcek Bay. The cores are  
324 sorted along the x-axis in the way that the 12 HLCs are plotted far-left in the graph and the  
325 CLCs in the order roughly from west to east (Fig. 4a). It is clear from the plot that the mean  
326 Sr/Ti ratio of the HLCs significantly deviates from the CLCs and is almost  $4\sigma$  higher than the  
327 mean of the CLCs. The spatial distribution map of the same ratio, on the other hand, enables  
328 visualization and can pinpoint the highest values of the HLCs (Fig. 4b). This observation  
329 implies that the sediments around the subaqueous hot spring in İnce Bay are enriched in Sr. It  
330 should also be noted that the high Sr/Ti ratio of the F-71 core, which was classified as CLC,  
331 may indicate another hot spring location in the vicinity of that core. Unfortunately, water  
332 sampling could not be done at this location since the depth is around 40 m there, which exceeds  
333 the safe diving limit.

334 Similar to the Sr enrichments illustrated in Fig. 4, other elemental enrichments were also  
335 detected. In order to find out which elements are more prominent, the deviations of the HLCs  
336 from the CLCs for all possible elemental ratios are presented in matrix form for the three study  
337 sites separately (Fig. 5). The colors of the cells represent the magnitude of deviation between  
338 the means of CLCs and HLCs. For instance, in Köyceğiz Lake, the means of the ratios of Cl to

339 almost all other elements for HLCs are higher by  $2\sigma$  than the mean of the same ratios for CLCs.  
340 It can be deduced, accordingly, that the sediments around the Köyceğiz subaqueous hot springs  
341 are enriched in S, Cl, and Br, while they are depleted in Mn (Fig. 5a). Similarly, in the Fethiye-  
342 Göcek Bay, the sediments around the subaqueous hot spring in İnce Bay are enriched in S, Cl,  
343 Ca, and Sr (Fig. 5c). No significant elemental enrichment related to subaqueous hot springs is  
344 detected in the Dalyan Channel (Fig. 5b). The plots and some of the representative spatial  
345 distribution maps for the above-mentioned elements are presented for Köyceğiz Lake and  
346 Fethiye-Göcek Bay in Fig. 6. In Köyceğiz Lake, not only the HLCs but also the CLCs in the  
347 southern basin show similar anomalies, which implies that the subaqueous hot springs affect  
348 the sedimentation throughout the southern basin. This can be attributed to the lack of significant  
349 tributaries in the southern basin, which may cause subaqueous hot springs to dominate the water  
350 chemistry in the basin.

351 Although some elemental ratios show significant anomalies in the Köyceğiz Lake and the  
352 Fethiye-Göcek Bay, no prominent elements could be determined for the Dalyan Channel. The  
353 absence of any elemental enrichment in the Dalyan Channel can be due to water currents in the  
354 channel, which may hinder regular sedimentation, thereby preventing any archive of  
355 sedimentary traces of subaqueous hot springs on the channel's bottom sediments.

## 356 **4.3 Hydrogeochemistry**

### 357 ***4.3.1 Water chemistry***

358 For a better understanding of the origin of natural waters, it is crucial to characterize their  
359 chemical properties in detail. Major and minor ion concentrations of the waters are given in  
360 Table 2 and Table 3, respectively. According to the results, inland cold waters and stream waters  
361 are of the Mg or Ca – HCO<sub>3</sub> type (Fig. 7). The other samples (subaqueous, lake, channel etc.)  
362 are Na-Cl type waters. This is an expected result because a seawater contribution to the lake

363 waters and geothermal waters in the Köyceğiz-Dalaman region was reported by previous  
364 studies (Bayarı et al., 1995; Gökgöz and Tarcan, 2006; Avşar et. al., 2016). The total dissolved  
365 solids (TDS) of the waters range from 302 to 41528 ppm. When the Cl ion concentration and  
366 TDS are compared, it is obvious that the TDS of the waters are dominated by the Cl  
367 concentration (Fig. 8). The Cl content of the waters may be considered as seawater contribution.  
368 Accordingly, it is obvious that all the waters (including subaqueous springs) are mixtures of  
369 fresh and seawater, and the seawater contribution (Cl concentration) correlates with the distance  
370 to the sea. This phenomenon will be the key point of the hydrogeochemical model, which will  
371 be discussed in the following sections.

372 In order to find out the possible impacts of temperature on mineralization, Saturation Indices  
373 (SI) were calculated by changing the temperature from 20 to 165 °C with the help of the  
374 computer program PHREEQC (Parkhurst and Appelo, 1999). The results for SUB-1 and SUB-7  
375 are presented as temperature vs SI graphs in Fig. 9. In the Fethiye-Göcek Bay, the calculated  
376 SI of the SUB-7 hot spring in İnlice Bay reveals that aragonite, calcite, dolomite, and fluorite  
377 are highly saturated (positive SI) at the discharge temperature of the hot spring (i.e., 28 °C)  
378 (Fig. 9a). Hence, the Ca enrichments at this spot can be attributed to precipitation of the above-  
379 mentioned minerals from SUB-7 hot spring water. Sr has similar chemical properties as Ca, and  
380 thus it can replace Ca within the crystalline structure of the above-mentioned minerals (Hem,  
381 1985). Sr is a trace element and is not as abundant in nature as Ca. Yet, given their chemical  
382 similarities, it is not unusual to observe Sr anomalies in a system with Ca anomalies. The  
383 mineral anhydrite, the presence of which could be a possible explanation for high S values in  
384 İnlice Bay sediments, has positive SI values at temperatures higher than 80 °C. Given the  
385 discharge temperature of 28 °C, the SI of anhydrite for the İnlice Bay hot spring is negative,  
386 which eliminates the possibility of S-enrichment due to anhydrite precipitation. The presence  
387 of S and Sr anomalies may also be due to celestite precipitation; however, its SI is also negative

388 at all temperatures and hence, it is not likely that the high abundances of these elements are  
389 related to this mineral. Although the present chemical composition of the İnlıce Bay hot spring  
390 does not seem suitable for anhydrite and celestite precipitation, it might have been different in  
391 the recent past. Given that the elemental results are the average values of the top 3-5 cm of the  
392 sediments, which represents ca. 15-20 years of sedimentation, S and Sr enrichments may  
393 indicate that the chemical composition of the İnlıce Bay hot spring was possibly different and  
394 was suitable for anhydrite and celestite precipitation in the past.

395 It is well known that the Cl ion is highly soluble, thus it is almost impossible for Cl to make a  
396 bond with a cation and precipitate as a mineral once it enters the solution as an anion. Hence,  
397 highly negative saturation values of sylvite and celestite are not surprising in this study. To  
398 explain Cl enrichments in the sediments, a more extended and comprehensive study focusing  
399 on geochemistry is required.

400 Cl and Br enrichments in the sediments around the Köyceğiz subaqueous spring cannot be  
401 explained by halite or sylvite precipitation since their saturation index values are highly  
402 negative (Fig. 9b). In order to find out the reason for Cl and Br-enriched sediments, more  
403 detailed investigations and analyses such as pore water chemistry analyses are required in this  
404 area.

#### 405 ***4.3.2 O- and H-isotope compositions***

406 Stable isotope ratios are useful to understand the origin of the waters and delineate the  
407 subsurface processes that affect their chemical composition. In this study, two campaigns of  
408 sampling were conducted for stable isotope analyses; 1) in September 2013 (dry season) and 2)  
409 in May 2014 (wet season) (Table 4). Highly negative  $\delta D$  and  $\delta^{18}O$  values of on-land fresh  
410 waters are an indication of recharge from inland and high altitudes (Fig. 10a and b). The on-  
411 land fresh waters seem to be of meteoric origin. However, a seawater effect is obvious where



412 the  $\delta^{18}\text{O}$  and  $\delta\text{D}$  values of the affected waters plot on a line that connects fresh water and  
413 seawater instead of the meteoric water lines (Fig. 10 a and b). This line may be interpreted as a  
414 mixing line between the freshwater and the seawater in the study region.

415 Another point that should be mentioned here is the role of the geographic position of the  
416 sampling locations (waters) with regard to seawater contribution. The closer the sampling  
417 location is to the sea, the more positive the isotope ratio becomes. As mentioned above, the  
418 mixing phenomenon was inferred from the chemical composition (Cl concentration) of the  
419 waters, and the distance to the sea affects the amount of seawater contribution and thus the Cl  
420 concentrations. The waters from close to the sea are also enriched with heavy isotopes, while  
421 the distant ones mostly show properties similar to fresh waters and are enriched in light isotopes.

422 The seasonal variations in the stable isotope ratios are also an important point to be taken into  
423 account. The water samples are generally, slightly enriched in light isotopes of O and H during  
424 the wet season (Fig. 10a). The most probable reason for this is a dilution effect due to the  
425 contribution of rainwater, which is enriched in light isotopes. On the other hand, during the dry  
426 season the waters are enriched in heavy isotopes of O and H, most probably due to evaporation.  
427 There are two samples which do not follow the seasonal variation trend: 1) Sülüngür Lake water  
428 and 2) sample DAL-2 from the Dalyan Channel (Table 4, Fig. 10b). The Dalyan Channel and  
429 Sülüngür Lake are known to be connected to the sea. A heavy isotope enrichment (seawater  
430 effect) in these waters in the wet season, in spite of the diluting effect of rainwaters, may be due  
431 to a water current that occurs in the wet season from the sea to the lake and channel. Although  
432 Alagöl Lake is also connected to the sea in the same way, a light isotope enrichment in this  
433 lake's waters in the wet season is compatible with the other waters. This indicates that although  
434 Alagöl Lake is very close to the sea and connected to it by a channel like Sülüngür Lake, for  
435 some reason, it is physically more isolated.

### 436 4.3.3 Gas geochemistry

437 The geothermal potential is generally high in areas where there is recent magmatic or tectonic  
438 activity like volcanism, extension, or fluid transport through fault-hosted deep pathways. The  
439 most sensitive indicator of such activity is the isotopic composition of helium ( $^3\text{He}/^4\text{He}$ ) (e.g.  
440 Kennedy and van Soest, 2007). Variations in helium isotope ratios of the fluids give researchers  
441 clues about the origin of these fluids and provide evidence for the presence of mantle-derived  
442 volatiles in geothermal systems. Helium in geothermal waters is derived from two terrigenous  
443 components with distinct isotopic composition (crustal He and mantle He). The continental  
444 crust is dominated by isotopically heavy helium ( $^4\text{He}$ ) that has been produced in situ by nuclear  
445 reactions ( $^4\text{He}$  by radioactive decay of uranium and thorium,  $^3\text{He}$  by thermal and epithermal  
446 neutron capture of lithium) in crustal rocks and minerals. This crustal component has a  $^3\text{He}/^4\text{He}$   
447 ratio of  $\sim 2 \times 10^{-8}$  (Mamyrin and Tolstikhin, 1984; Andrews, 1985). In contrast, the Earth's  
448 mantle contains He enriched in the light isotope  $^3\text{He}$  that has been inherited from planet  
449 formation. The convective mantle as sampled by mid-ocean ridge basalts (MORB) has a  
450  $^3\text{He}/^4\text{He}$  ratio of  $\sim 1.1 \times 10^{-5}$  (Mamyrin and Tolstikhin, 1984; Ozima and Podosek, 2002), though  
451 other mantle domains may have higher or lower  $^3\text{He}/^4\text{He}$  ratios.  $^3\text{He}/^4\text{He}$  ratios are often  
452 expressed as a multiple of the present-day atmospheric  $^3\text{He}/^4\text{He}$  ratio,  $R_a$ , which is  $1.39 \times 10^{-6}$ .  
453 Thus, the  $^3\text{He}/^4\text{He}$  ratio of the atmosphere is 1  $R_a$ , that of the crust may vary between  $\sim 0.01$  and  
454  $0.1 R_a$ , and typical mantle values are  $\sim 8 R_a$ . In the vicinity of the study area, a mantle  $^3\text{He}/^4\text{He}$   
455 ratio of  $7.7 R_a$  has recently been measured in olivines from the Kula volcanic field (Heineke et  
456 al., 2016).

457 In addition to  $^3\text{He}/^4\text{He}$  ratios,  $\text{CO}_2/^3\text{He}$  ratios are also used to recognize mixing of mantle and  
458 crustal fluids (Barry et al., 2013; Karakuş and Şimşek, 2013). Mantle-derived  $\text{CO}_2$  gases usually  
459 have a  $\text{CO}_2/^3\text{He}$  ratio of approximately  $2 \times 10^9$  (Marty and Jambon, 1987), which will be

460 preserved unless mixing or fractionation occur during the journey of the mantle gases through  
461 the lithosphere (O’Nions and Oxburgh, 1988; Wiersberg et al., 2011; Karakuş and Şimşek,  
462 2013).

463 In this study, to contribute to the hydrogeochemical conceptual modeling of the geothermal  
464 system in the study area, gas geochemistry analyses and helium isotope data were evaluated  
465 accordingly. For springs where free gases could be sampled, the total gas compositions are  
466 presented in Table 5 together with the helium isotope,  $^4\text{He}/^{20}\text{Ne}$  and  $\text{CO}_2/{}^3\text{He}$  ratios. Water  
467 samples degassed in the lab were only analyzed for noble gases, therefore no total gas  
468 compositions and  $\text{CO}_2/{}^3\text{He}$  ratios are available for them. Air-corrected  ${}^3\text{He}/{}^4\text{He}$  ratios  
469 ( $({}^3\text{He}/{}^4\text{He})_c$ ) were calculated according to Craig et al. (1978) and are also given in the same  
470 table. According to Table 5, the majority of the gas samples are composed mainly of nitrogen  
471 ( $\text{N}_2$ ) with variable amounts of  $\text{CO}_2$ , Ar,  $\text{O}_2$ ,  $\text{CH}_4$ ,  $\text{H}_2$  and He, except for one sample collected  
472 from Köyceğiz Lake (SUB-2) in which the major contribution is from  $\text{CH}_4$  with considerable  
473 contributions of  $\text{H}_2$  and  $\text{O}_2$ .  $\text{N}_2/\text{Ar}$  ratios of the gas samples are up to 119 and are close to the  
474 atmospheric value (83). The gas compositions reported here are distinctly different from those  
475 reported by Wiersberg et al. (2011) for the Büyük Menderes Graben, where the majority of the  
476 samples have  $\text{CO}_2$  concentrations  $> 95\%$ .

477 The  ${}^4\text{He}/^{20}\text{Ne}$  ratios of the samples range from 0.35 to 130, indicating that contributions from  
478 atmospheric He are significant in some samples, but negligible in others when compared to the  
479  ${}^4\text{He}/^{20}\text{Ne}$  ratios of air ( $\sim 0.319$ ; Sano and Wakita, 1985) or air-saturated water (0.274 at 25 °C;  
480 e.g. Ozima and Podosek, 2002) (Table 5). The atmospheric contribution is highest in samples  
481 collected from Köyceğiz Lake, where  ${}^4\text{He}/^{20}\text{Ne}$  ratios are close to the ones reported for air,  
482 which might be due to admixing of lake water during sampling.

483 The ( $^3\text{He}/^4\text{He}$ )<sub>c</sub> of all samples are above the crustal value ( $\sim 0.05 R_a$ ), indicating a contribution  
484 of mantle origin. A mantle helium component has been reported in Western Anatolia by  
485 previous researchers (Güleç, 1988; Güleç and Hilton, 2006; Karakuş and Şimşek, 2013). In fact,  
486 Güleç (1988) stated that mantle helium could be observed in a large area in Western Anatolia  
487 depending on the distribution of fault systems related to the extensional regime. By using  
488 ( $^3\text{He}/^4\text{He}$ )<sub>c</sub>, which corrects for atmospheric helium contributions, relative fractions of mantle  
489 and crustal helium can be calculated using a simple mixing model between mantle and crustal  
490 endmembers (Equation 1):

$$491 \text{Mantle He (\%)} = 100 \times [ (^3\text{He}/^4\text{He})_c - (^3\text{He}/^4\text{He})_{\text{crust}} ] / [ (^3\text{He}/^4\text{He})_{\text{mantle}} - (^3\text{He}/^4\text{He})_{\text{crust}} ]$$

492 Equation 1

493 Assuming ( $^3\text{He}/^4\text{He}$ )<sub>crust</sub> =  $0.02 R_a$  and ( $^3\text{He}/^4\text{He}$ )<sub>mantle</sub> =  $8 R_a$ , mantle He percentages between  
494 2% and 17% are obtained (Table 5). The lowest mantle He contributions were found in the  
495 samples from Köyceğiz Lake, whereas the highest ones were observed in springs located on-  
496 land in the Dalaman plain near the Mediterranean Sea (THER, MUS, CUR, KAP-2). Small  
497 mantle He contributions also occur in on-land springs located near Köyceğiz Lake and the  
498 Dalyan Channel (SUL-2, KEL, DEL). The different helium contributions from air, crust and  
499 mantle in all samples are visualized in a  $^3\text{He}/^4\text{He}$  vs.  $^4\text{He}/^{20}\text{Ne}$  graph (Fig. 11). For instance, the  
500 Köyceğiz Lake samples with the highest atmospheric He contributions plot near air/ASW on  
501 the mixing line between ASW and crustal He endmembers (Fig. 11). Likewise, samples with a  
502 high crustal He contribution plot near the ASW-crustal He mixing line (Fig. 11). On the other  
503 hand, those samples with the highest mantle He contribution are located closer to the mixing  
504 line between ASW and mantle He. A relatively high contribution of mantle He in the Dalaman  
505 plain suggests that the faults of extensional tectonics in the area promote the escape of gases

506 originating from the deeper parts of the Earth through the brittle parts of the crust, and this  
507 affects the geothermal system.

508 Examining the differences in mantle helium contributions of subaqueous (SUB-2, 3, 4 and  
509 SUBC-2) and on-land (SUL-2, KEL, DEL) Köyceğiz Lake samples with the Dalaman plain  
510 samples (CUR, MUS, THER, KAP-2) indicates that these samples are products of two different  
511 geothermal systems. Köyceğiz Lake on-land samples have relatively high discharge  
512 temperatures (between 37 and 40 °C) when compared to Dalaman plain samples (26-30 °C).  
513 The CO<sub>2</sub>/<sup>3</sup>He ratios in free gas samples range from 0.05×10<sup>9</sup> to 1.097×10<sup>9</sup>. These ratios are  
514 lower than the average ratio reported for the upper mantle (2×10<sup>9</sup>) (Marty and Jambon, 1987).  
515 Moreover, these ratios are much lower than the ones reported by Mutlu et al. (2008), Wiesberg  
516 et al. (2011) and Karakuş and Şimşek (2013). This probably reflects the effects of gas  
517 fractionation processes (such as degassing and/or dissolution) in the hydrothermal system,  
518 which modified the He-CO<sub>2</sub> systematics. A sound discussion of He-CO<sub>2</sub> systematics, however,  
519 would require δ<sup>13</sup>C data in addition and is beyond the scope of the present study.

#### 520 ***4.3.4 Hydrogeochemical model***

521 Our findings from this study enable us to construct a conceptual model (Fig. 12), which shows  
522 that fresh water percolates from the fractures, joints, and faults and is heated at depth along the  
523 geothermal gradient of the earth. While ascending to the surface from either the sea or lake  
524 bottom or on-land via active faults, the waters appear to be mixed with seawater below the sea  
525 or lake bottom.

526 The study area is located on the Lycian nappes (Fig. 1), where rock packages are overlain by  
527 thrust faults. Generally, faults dip southward. The key point is that the carbonate rocks are  
528 pervious, and these rocks are underlain and overlain by the impervious Yeşilbarak nappe and  
529 Marmaris nappe, respectively. The southward-dipping nappes force groundwater to flow from

530 north to south (to the Mediterranean). The Yeşilbarak nappe is located at the bottom of all other  
531 nappes tectonically; and since it is buried, it crops out in only a very limited area in the study  
532 region (Fig. 1). This makes it an impervious mat lying at the bottom of the Lycian nappes. The  
533 Marmaris nappe is tectonically the uppermost layer of the Lycian nappes and may be  
534 understood as impervious with a peridotite composition. The Gülbahar, Bodrum and Tavas  
535 nappes, although differentiated as separate units by Şenel (1997a; 1997b), are made up of  
536 carbonate rocks and are sandwiched between these two impervious units, namely the Yeşilbarak  
537 nappe at the bottom, and the Marmaris nappe at the top. This arrangement of pervious and  
538 impervious nappes creates a somehow confined aquifer in the region (Fig. 12).

539 The Marmaris Nappe is situated on top of the carbonate nappes, lying as a NW-SE-oriented  
540 strip. The carbonate nappes underlie this impervious patch and crop out to the north and to the  
541 south of it (Fig. 1). The carbonate units are pervious and known to be karstic in character. The  
542 northern outcrop of the karstic nappes can be interpreted as a recharge area of the confined  
543 aquifer and the southern outcrop as the discharge area (Fig. 12). The groundwater flow is from  
544 the northern highlands towards the southern lowlands, namely the Mediterranean. The  
545 southward-dipping tectonic contacts between these units also enable this groundwater  
546 movement.

547 The rainwater precipitated in the northern carbonate highlands travels to the southern low  
548 elevation areas below the impervious Marmaris nappe (Fig. 12). The southern boundary of the  
549 Marmaris-carbonate nappe units is cut by young normal faults, leading to the formation of a  
550 NW-SE lying depression. This depression is filled either by water or by sediments. For  
551 example, in the Köyceğiz area it is filled with water forming Köyceğiz Lake. However, around  
552 Dalaman, it is filled with alluvium, forming the Dalaman plain. The groundwater traveling from  
553 north to south is not able to discharge to the surface due to the overlying Marmaris nappe.  
554 However, the heated groundwater immediately emerges at the southern boundary of the

555 Marmaris-carbonate nappe units. This is either due to the normal faults that enable the vertical  
556 permeability or to the lack of an impervious cap rock in the south.

557 The subaqueous hot spring of the Fethiye-Göcek Bay (SUB-7) most probably emerges at the  
558 southern Marmaris-carbonate nappe boundary. Since the boundary is covered by seawater and  
559 alluvium, its exact location is not visible. We do not know whether the spring emerges exactly  
560 at the boundary. The location may be on the carbonate side, since the boundary is covered by  
561 alluvium, and there may be a shift where the alluvium is thin enough to allow the groundwater  
562 to percolate. Alternatively, the water may emerge on the Marmaris nappe side. The Marmaris  
563 nappe may get thinner towards the boundary and become pervious to semi-pervious due to  
564 fractures and joints in the rock. In both cases, we can suppose that the hot water emerges not  
565 too far away from the boundary.

## 566 **5. Conclusion**

567 This study aimed to find subaqueous hot springs by making horizontal temperature  
568 measurements. Measurements revealed 49 temperature anomalies in the study area and among  
569 these anomalies seven of them were confirmed as subaqueous hot springs: three in Köyceğiz  
570 Lake, three in the Dalyan Channel, and one in the Fethiye-Göcek Bay. There are no temperature  
571 anomalies at the bottom of the Alagöl, Sülüngür and Kocagöl lakes. The temperature of the  
572 subaqueous hot springs in the study area ranges between 26 and 29 °C.

573 According to the chemical analysis results, inland cold waters and stream waters are of the Mg  
574 or Ca – HCO<sub>3</sub> type. The other samples (subaqueous, lake, channel etc.) are Na-Cl type waters.  
575 The on-land fresh waters are of meteoric origin and display highly negative  $\delta D$  and  $\delta^{18}O$  values.  
576 This is an indication of recharge from inland and high altitudes. However, there is a seawater  
577 effect, obvious in the  $\delta^{18}O$  and  $\delta D$  values of the remaining waters.

578 The geology of the region greatly affects the geothermal system in the study area. The Marmaris  
579 nappe acts as an impervious cap rock. Especially in the south of the study area, this unit is  
580 buried under alluvium and/or lake/seawater. Similarly, the Yeşilbarak nappe also underlies all  
581 units as an impervious base rock. Pervious and even karstic carbonate units are sandwiched  
582 between these two impervious rocks forming a somehow confined aquifer. The tectonic  
583 boundaries between these units are dipping mainly southward. This allows groundwater  
584 originating from the northern highlands to travel through the pervious carbonate rocks below  
585 the impervious Marmaris nappe. The water travels towards the lowlands in the south, is heated  
586 at depth, and mixes with seawater while traveling up through high angle normal faults, it  
587 emerges as hot springs from the southern boundary of the carbonate-Marmaris nappe units. The  
588 type of hot spring (either on-land or subaqueous) is determined by the local occurrence of  
589 alluvium or water. If the boundary is covered by alluvium, like in the case of the Dalaman plain,  
590 the hot springs (MUS, THER, KAP, CUR) emerge on land. If the boundary is covered by thin  
591 alluvium and a water body (lake/sea/channel), the hot springs emerge either on-land (SUL,  
592 DEL, KEL) or subaqueous (SUB-1-7).

593 According to the total gas analysis results, the majority of the gas samples are composed mainly  
594 of nitrogen (N<sub>2</sub>) with variable amounts of CO<sub>2</sub>, Ar, O<sub>2</sub>, CH<sub>4</sub>, H<sub>2</sub> and He. N<sub>2</sub>/Ar ratios of the gas  
595 samples are close to the atmospheric value (up to 119). The <sup>4</sup>He/<sup>20</sup>Ne ratios of the samples  
596 indicate that atmospheric contribution is generally negligible, but is significant in some samples  
597 such as those taken from Köyceğiz Lake, which might be due to admixing of lake water during  
598 sampling. The (<sup>3</sup>He/<sup>4</sup>He)<sub>c</sub> of all samples are above the crustal value (~0.05 R<sub>a</sub>) which indicate  
599 a mantle contribution. According to a mixing model between mantle and crustal endmembers,  
600 the lowest mantle He contributions were found in on-land springs located near Köyceğiz Lake  
601 (SUL-2, KEL, DEL), whereas the highest ones were observed in springs located around  
602 Dalaman plain (THER, MUS, CUR, KAP-2). The differences in mantle helium contributions



603 of subaqueous and on-land Köyceğiz Lake hot springs and Dalaman plain hot springs indicate  
604 that they are the products of two different geothermal systems. The  $\text{CO}_2/{}^3\text{He}$  ratios in free gas  
605 samples ( $0.05\times 10^9$  to  $1.097\times 10^9$ ) are lower than the average ratio reported for the upper mantle  
606 ( $2\times 10^9$ ) and much lower than the ones reported for western Anatolia, which may be attributed  
607 to the effects of gas fractionation processes in the hydrothermal system.

608 S, Cl, Ca, and Sr elements are enriched around Fethiye-Göcek subaqueous hot spring (SUB-7),  
609 whereas S, Cl, Br, elements are enriched, Mn element is depleted around Köyceğiz lake  
610 subaqueous hot springs. Fluid mineral equilibria calculations reveal that Ca enrichment around  
611 SUB-7 is probable from precipitation of aragonite, calcite, dolomite, and fluorite minerals from  
612 the hot spring. Sr has a similar atomic structure and chemical properties with Ca. It most  
613 probably replaces Ca and displays a parallel anomaly. Cl and S enrichment around both  
614 Köyceğiz and Fethiye-Göcek subaqueous hot springs cannot be explained by the fluid mineral  
615 equilibria study. S enrichment could be related to biochemical processes.

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## 827 **Figure Captions**

828 **Fig. 1.** Geological map of the study region (modified from Şenel, 1997a; Şenel, 1997b) with  
829 water sampling locations indicated. Sample labels are given in Table 1 along with sample  
830 information. Fethiye-Burdur and Pliny-Strabo fault zones are shown on the upper right inset  
831 map as FBFZ and PSFZ, respectively (Hall et al., 2014).

832 **Fig. 2.** Physicochemical measurement routes in (a) Dalyan Channel, Köyceğiz, Alagöl, and  
833 Sülüngür Lakes; (b) Fethiye-Göcek Bay and Kocagöl Lake (in the west). Bathymetry maps of  
834 (c) Köyceğiz Lake and (d) Fethiye-Göcek Bay.

835 **Fig. 3.** Locations of 233 gravity cores collected from a) Köyceğiz Lake, b) Dalyan Channel,  
836 and c) Fethiye-Göcek Bay. The white dots are “Cold Location Cores” (CLC), while the red dots  
837 are “Hot Location Cores” (HLC). The insets are close-up views around the HLCs.

838 **Fig. 4.** a) An example illustration of the comparison of the mean Sr/Ti ratios in the HLCs and  
839 CLCs from the Fethiye-Göcek Bay. The mean of the Sr/Ti ratio in HLCs (i.e., around the  
840 subaqueous hot spring) is significantly higher than the mean of the CLCs (by approximately  
841  $4\sigma$ ), which implies that sediments around the subaqueous hot spring are enriched in Sr. b)  
842 Spatial distribution of the Sr/Ti ratio over the Fethiye-Göcek Bay. The maximum around the F-  
843 71 core also implies the presence of another subaqueous hot spring around that location.

844 **Fig. 5.** Matrices illustrating the magnitude of deviations between absolute means of HLCs and  
845 CLCs for all possible elemental ratios; a) Köyceğiz Lake, b) Dalyan Channel, c) Fethiye-Göcek  
846 Bay. For example, in the Köyceğiz matrix, the cell at the intersection of the 4<sup>th</sup> column and the  
847 5<sup>th</sup> row shows that the mean Cl/K ratio of the HLCs is  $2.5\sigma$  higher than the mean Cl/K ratio of  
848 the CLCs.

849 **Fig. 6.** Plots and spatial distributions of the mean ratios of enriched elements in Köyceğiz Lake  
850 (left) and Fethiye-Göcek Bay (right). The gray lines in the plots stand for individual elemental  
851 ratios, while the black and red lines indicate the mean of these ratios for the CLCs and HLCs,  
852 respectively. The horizontal lines on the graphs indicate  $\pm 1\sigma$  boundaries for CLCs. Note the  
853 significant difference in elemental enrichments in the southern basin of Köyceğiz Lake. The  
854 basinwide difference is attributed to the lack of tributaries in the southern basin, which enhances  
855 the effect of subaqueous hot springs on the overall sediment chemistry.

856 **Fig. 7.** Piper diagram of waters.

857 **Fig. 8.** Cl vs. Total Dissolved Solids (TDS) graph.

858 **Fig. 9.** Temperature vs. Saturation Index (SI) diagrams for samples SUB-1 and SUB-7.

859 **Fig. 10.**  $\delta^{18}\text{O}$  vs.  $\delta\text{D}$  diagram of waters [GMWL: **G**lobal **M**eteoritic **W**ater **L**ine (Craig, 1961);  
860 MMWL: **M**editerranean **M**eteoritic **W**ater **L**ine (IAEA, 1981)] (a) wet, (b) dry season  
861 measurements.

862 **Fig. 11.**  $^3\text{He}/^4\text{He}$  ( $R_a$ ) versus  $^4\text{He}/^{20}\text{Ne}$  graph for hydrothermal fluids and gases. Binary mixing  
863 lines between air saturated water (ASW) ( $^3\text{He}/^4\text{He} = 1 R_a$ ), mantle ( $^3\text{He}/^4\text{He}_{\text{mantle}} = 8 R_a$ ) and  
864 crustal ( $^3\text{He}/^4\text{He}_{\text{crust}} = 0.02 R_a$ ) endmember compositions, 1% and 20% air-mantle mixing lines  
865 (mantle %) and 1%, 10%, 20% and 50% air proportions (air %) are also presented herein. The  
866 graph was prepared using the HNC-Plot application (Karakuş and Aydın, 2016).

867 **Fig. 12.** Sketch of a conceptual hydrogeochemical model for the study area. Not to scale.  
868 Modified from Gökgöz and Tarcan (2006) and Avşar et al. (2016).

869 **Supplementary Fig. 1.** (a) Temperature and probe depth measurements taken around the  
870 submarine hot springs. The anomalies of SUB-1, SUB-4, SUB-5 and SUB-7 hot springs are  
871 marked by circles. Temperature anomaly at 18:02:14 on 26.08.2013 indicates SUB-7 hot spring.  
872 A second measurement on 31.08.2014 verified the existence of SUB-7 hot spring. See the  
873 anomaly at 15:21:00 marked by a circle. (b) Sonar profiles used to detect SUB-2 and SUB-3.

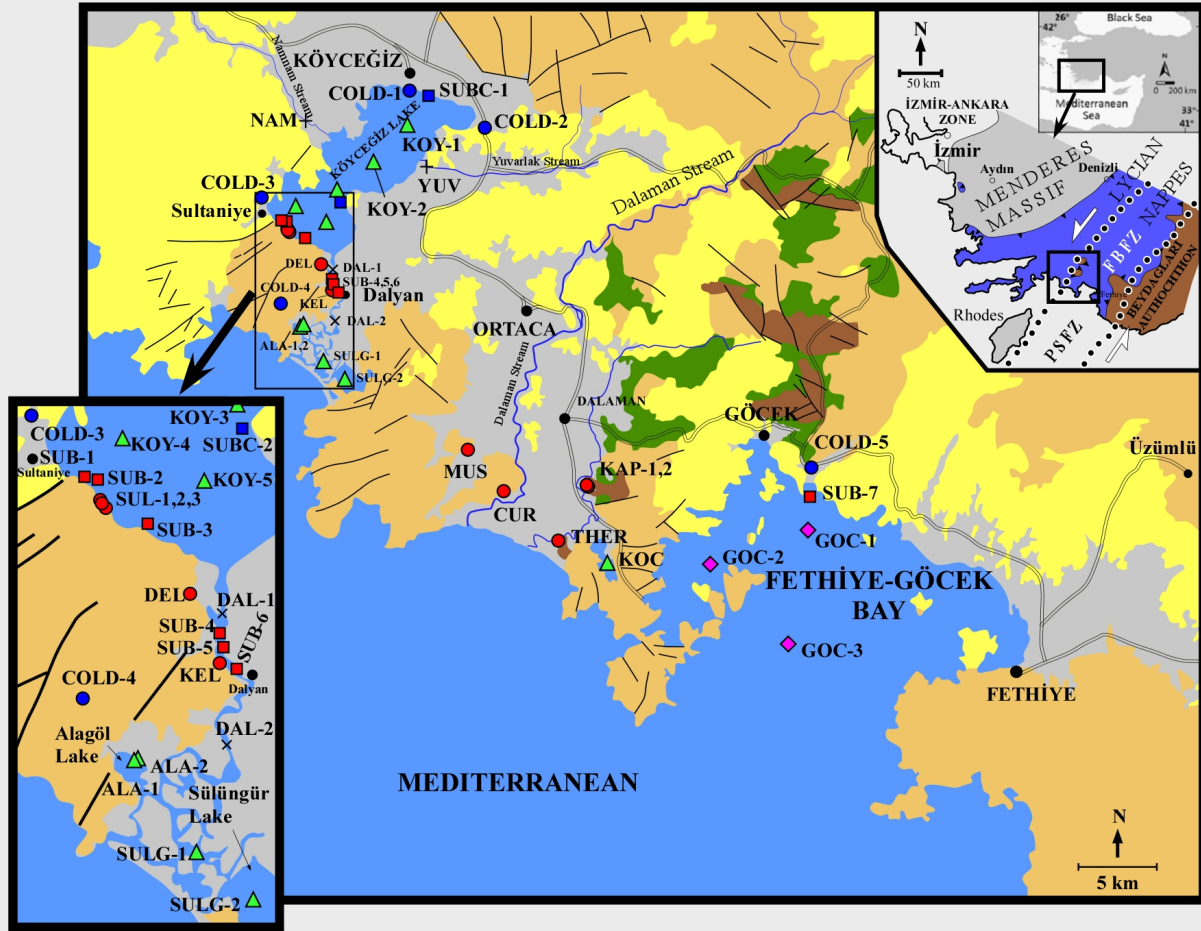
874 **Supplementary Fig. 2.** Optical and radiographic images of the artificial cores composed of the  
875 top 3-5 cm sediments from 233 gravity cores. Labels and coordinates of each gravity core are  
876 presented as well. “K”: Köyceğiz, “D”: Dalyan, “F”: Fethiye-Göcek, and “I”: İnlice Bay in  
877 Fethiye-Göcek Bay.

878 **Supplementary Fig. 3.** Plots of elemental ratios based on the standardized “kcp<sub>s</sub>” values for  
879 (a) Köyceğiz Lake, (b) Dalyan Channel, (c) Fethiye-Göcek Bay.

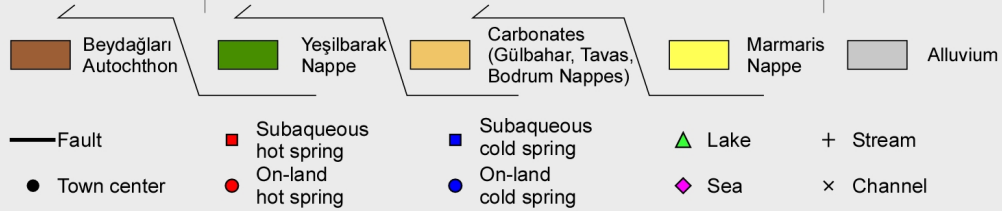
880 **Supplementary Fig. 4.** Spatial distribution maps of elemental ratios over the study area. (a)

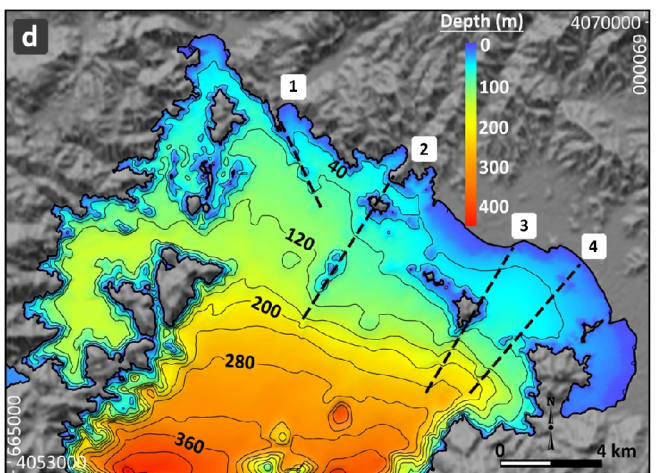
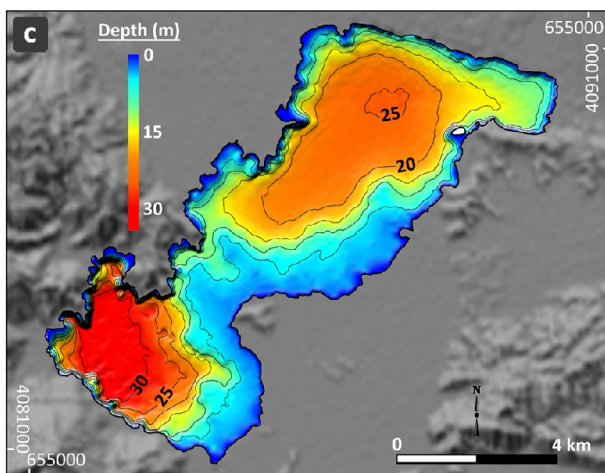
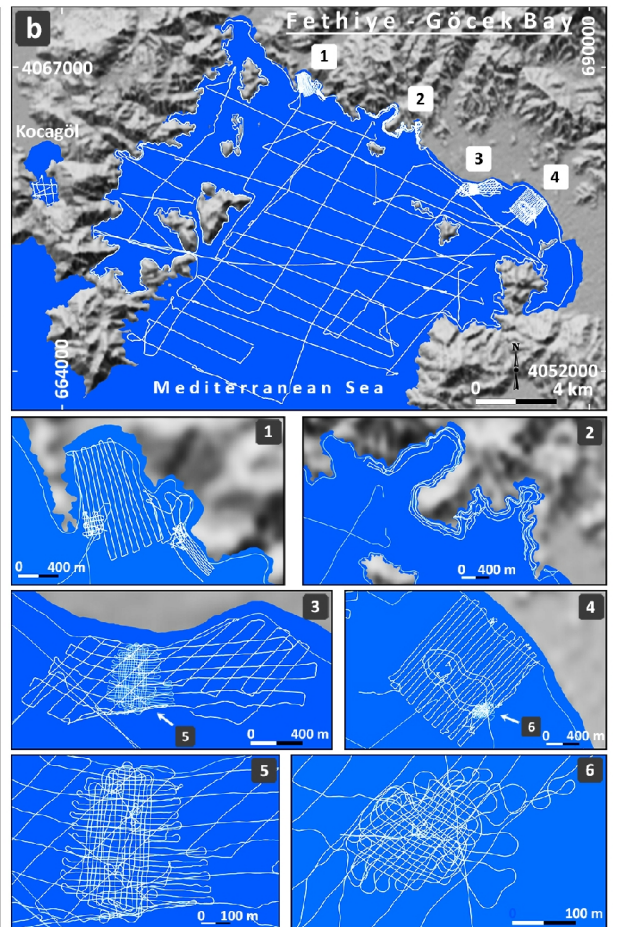
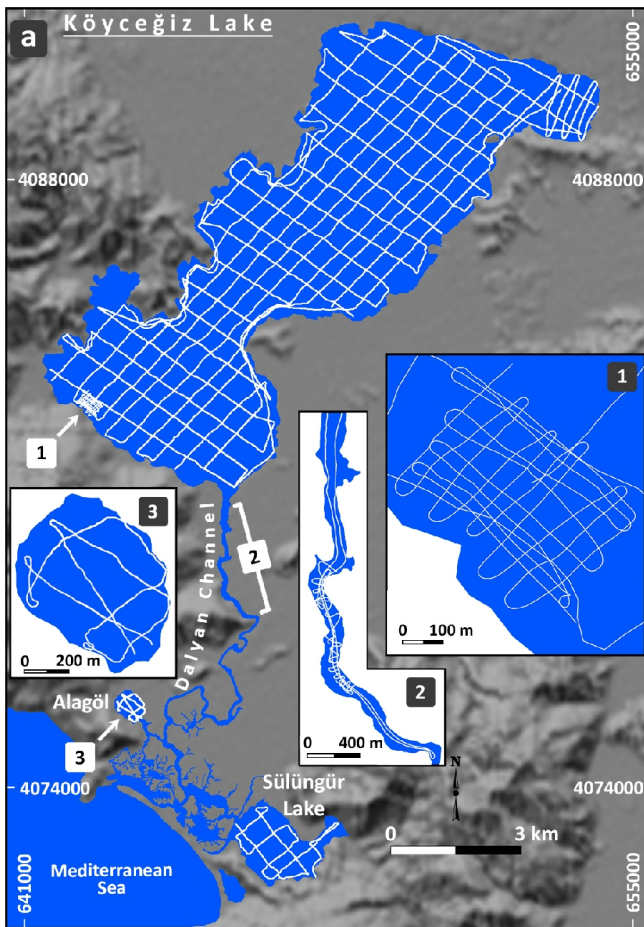
881 Köyceğiz Lake, (b) Dalyan Channel, (c) Fethiye-Göcek Bay.

882



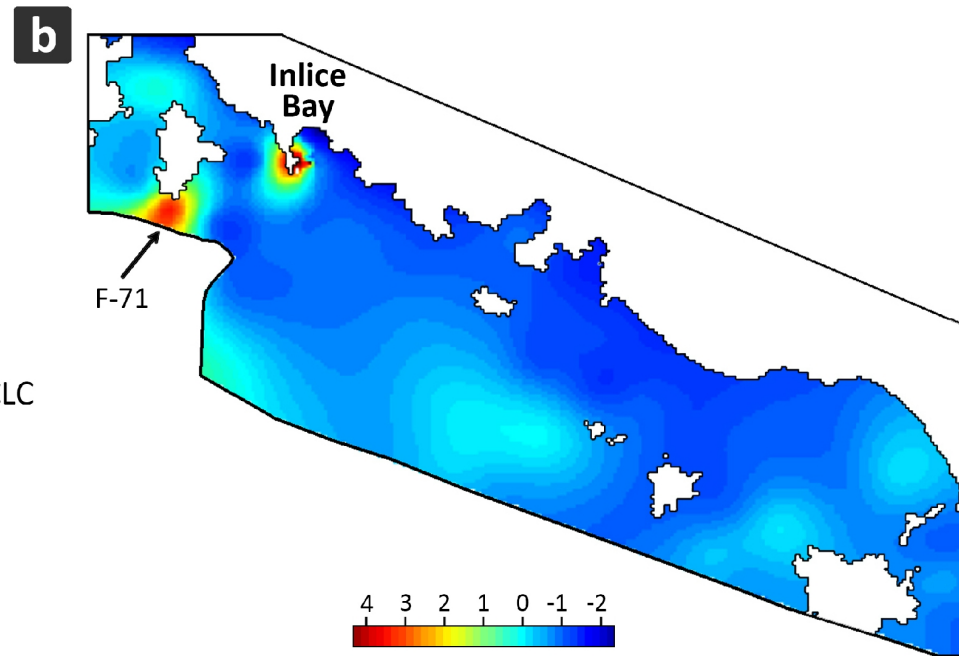
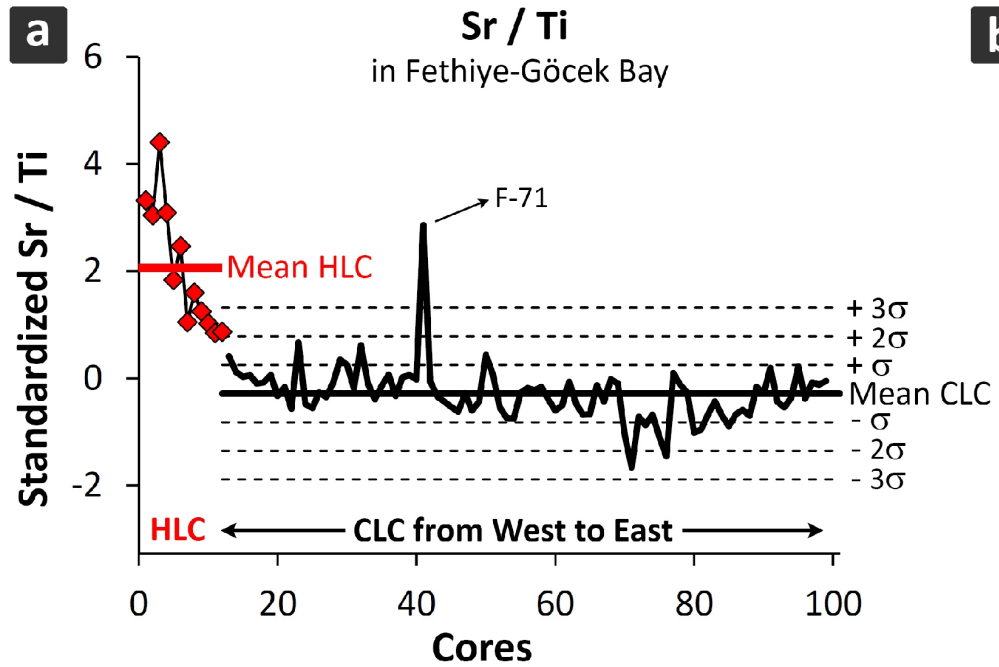
### Lycian Nappes

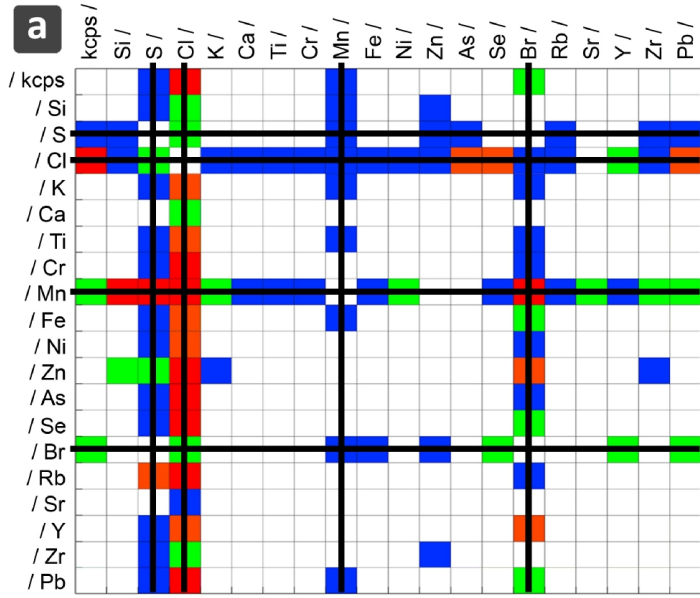




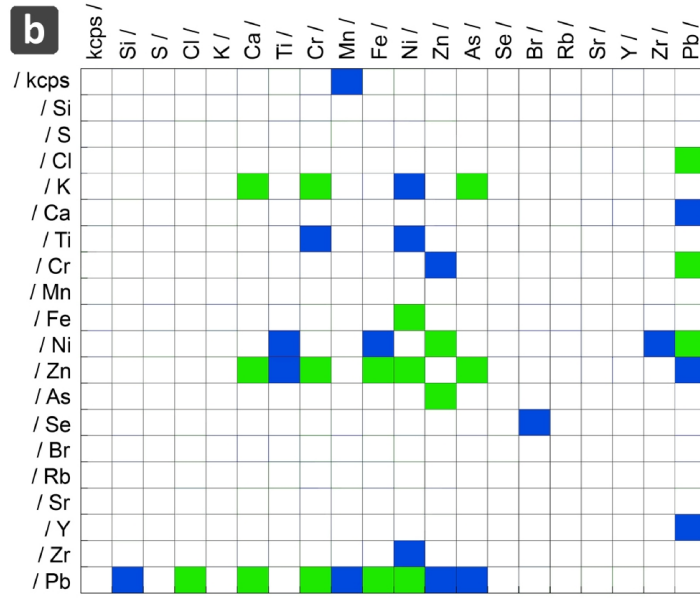




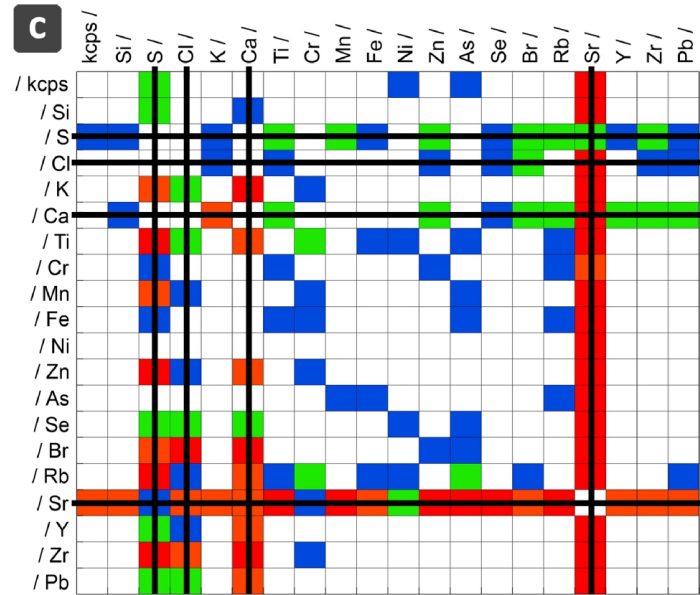




**KÖYCEĞİZ LAKE**



**DALYAN CHANNEL**



**FETHİYE-GÖCEK BAY**

**Köyceğiz Lake**

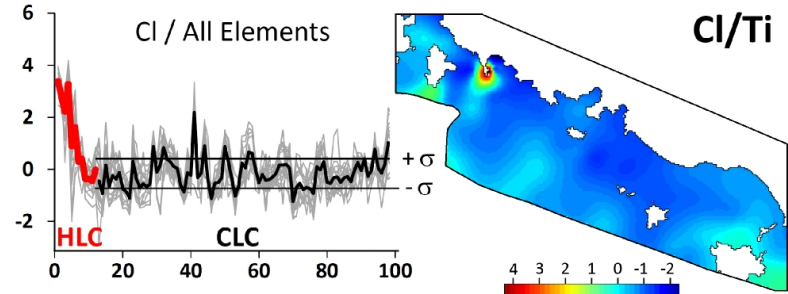
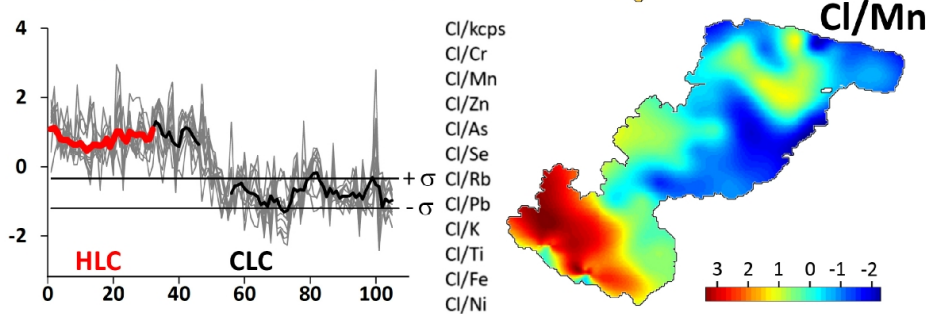
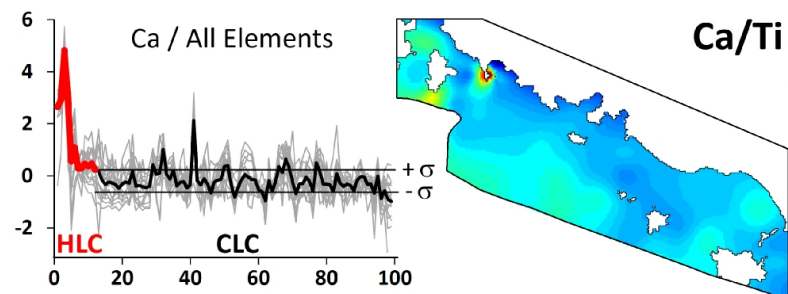
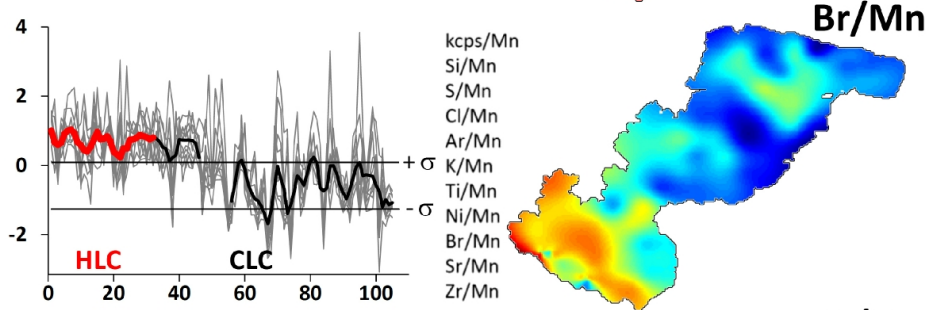
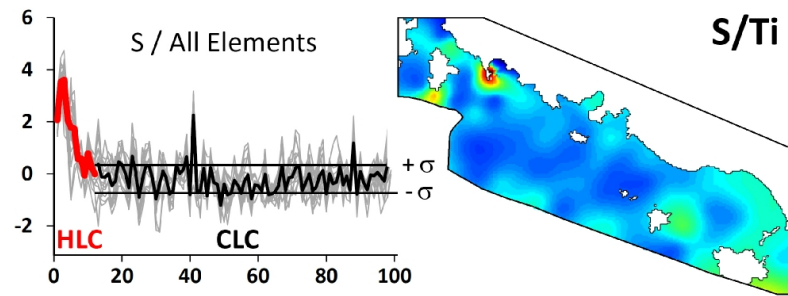
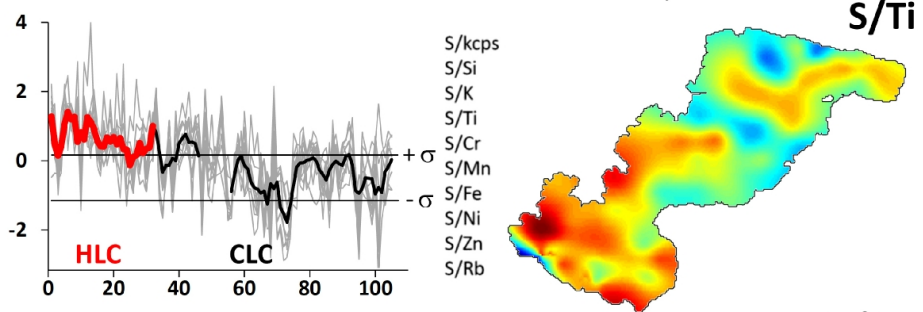
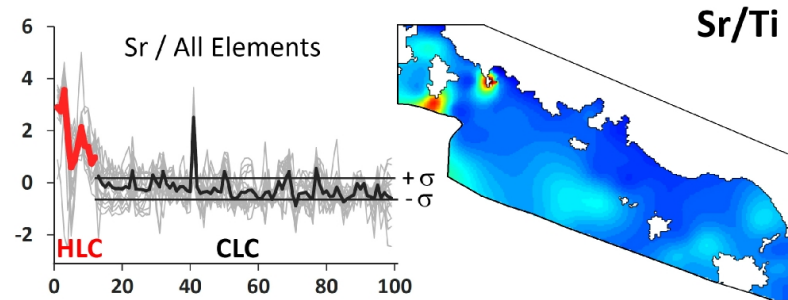
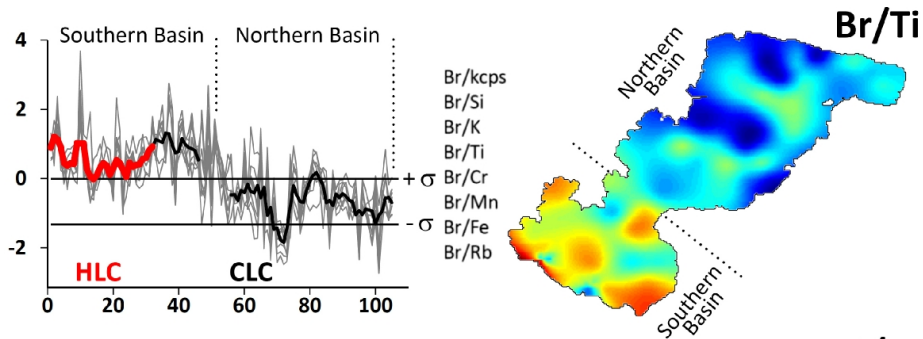
- $|Ave_{CLC} \pm 2.5\sigma_{CLC}| \leq |Ave_{HLC}|$
- $|Ave_{CLC} \pm 2.0\sigma_{CLC}| \leq |Ave_{HLC}| < |Ave_{CLC} \pm 2.5\sigma_{CLC}|$
- $|Ave_{CLC} \pm 1.5\sigma_{CLC}| \leq |Ave_{HLC}| < |Ave_{CLC} \pm 2.0\sigma_{CLC}|$
- $|Ave_{CLC} \pm 1.0\sigma_{CLC}| \leq |Ave_{HLC}| < |Ave_{CLC} \pm 1.5\sigma_{CLC}|$
- $|Ave_{HLC}| < |Ave_{CLC} \pm 1.0\sigma_{CLC}|$

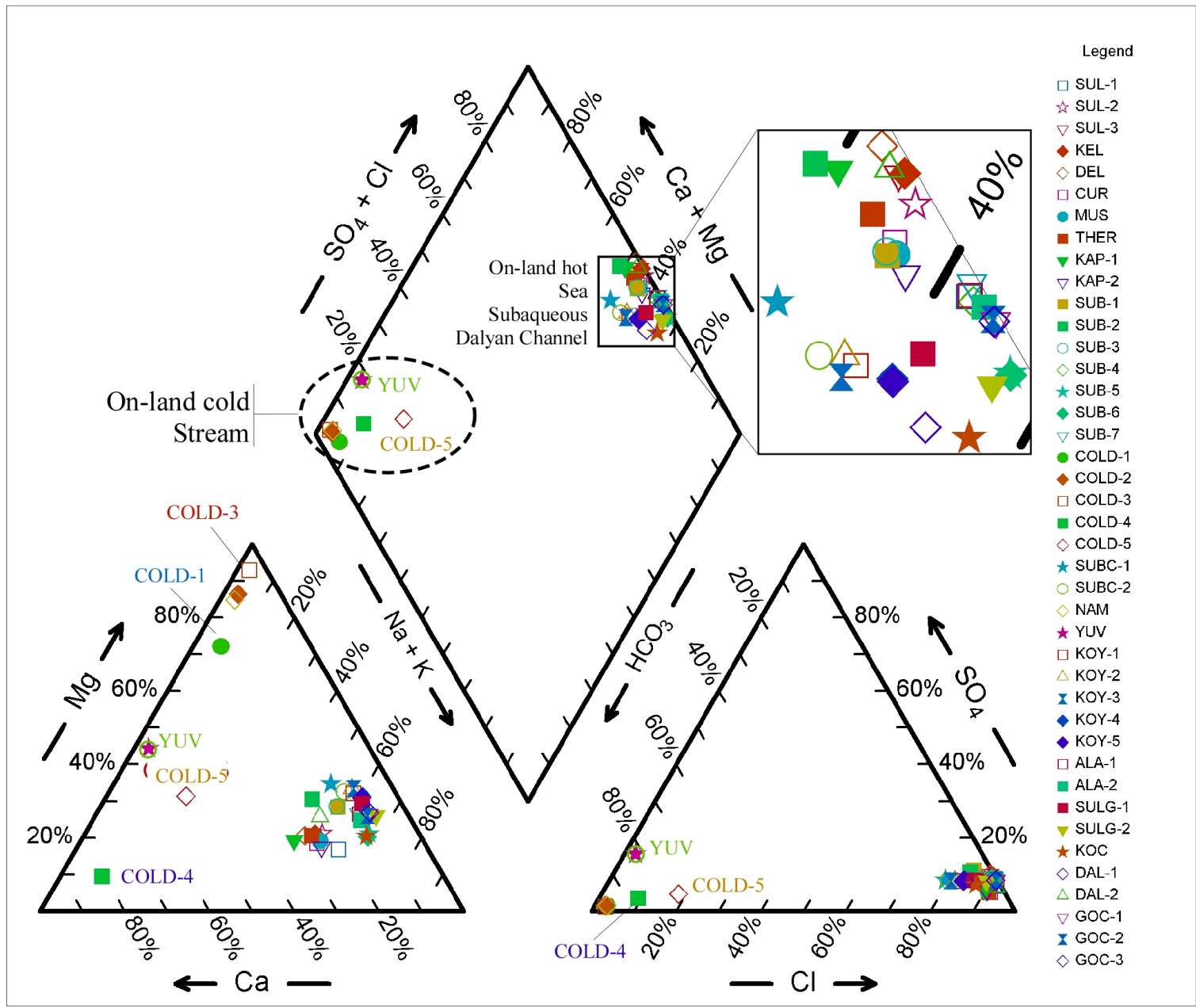
**Dalyan Channel**

- $|Ave_{CLC} \pm 2.5\sigma_{CLC}| \leq |Ave_{HLC}|$
- $|Ave_{CLC} \pm 2.0\sigma_{CLC}| \leq |Ave_{HLC}| < |Ave_{CLC} \pm 2.5\sigma_{CLC}|$
- $|Ave_{CLC} \pm 1.5\sigma_{CLC}| \leq |Ave_{HLC}| < |Ave_{CLC} \pm 2.0\sigma_{CLC}|$
- $|Ave_{CLC} \pm 1.0\sigma_{CLC}| \leq |Ave_{HLC}| < |Ave_{CLC} \pm 1.5\sigma_{CLC}|$
- $|Ave_{HLC}| < |Ave_{CLC} \pm 1.0\sigma_{CLC}|$

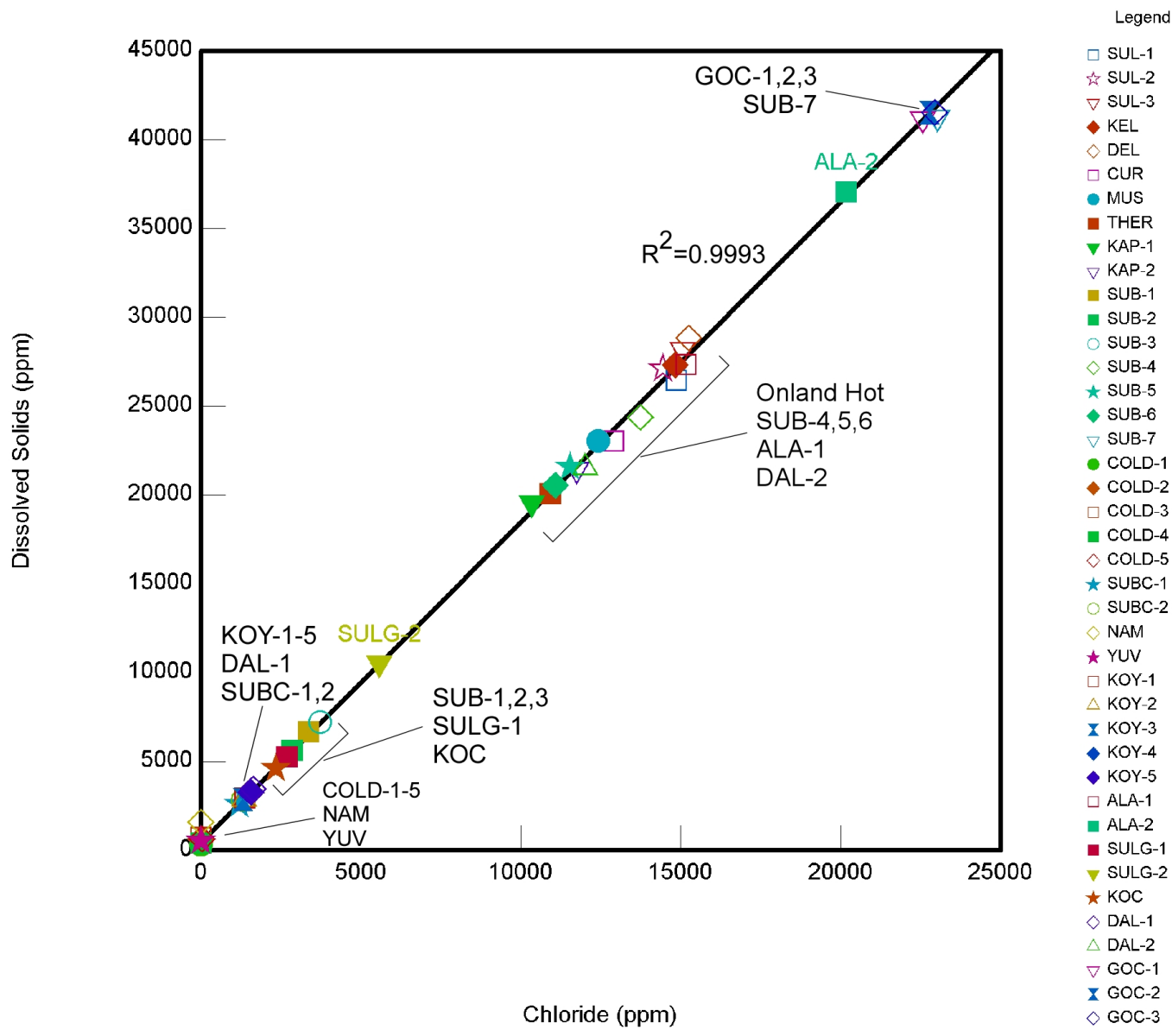
**Fethiye-Göcek Bay**

- $|Ave_{CLC} \pm 5.0\sigma_{CLC}| \leq |Ave_{HLC}|$
- $|Ave_{CLC} \pm 4.0\sigma_{CLC}| \leq |Ave_{HLC}| < |Ave_{CLC} \pm 5.0\sigma_{CLC}|$
- $|Ave_{CLC} \pm 3.0\sigma_{CLC}| \leq |Ave_{HLC}| < |Ave_{CLC} \pm 4.0\sigma_{CLC}|$
- $|Ave_{CLC} \pm 2.0\sigma_{CLC}| \leq |Ave_{HLC}| < |Ave_{CLC} \pm 3.0\sigma_{CLC}|$
- $|Ave_{HLC}| < |Ave_{CLC} \pm 2.0\sigma_{CLC}|$

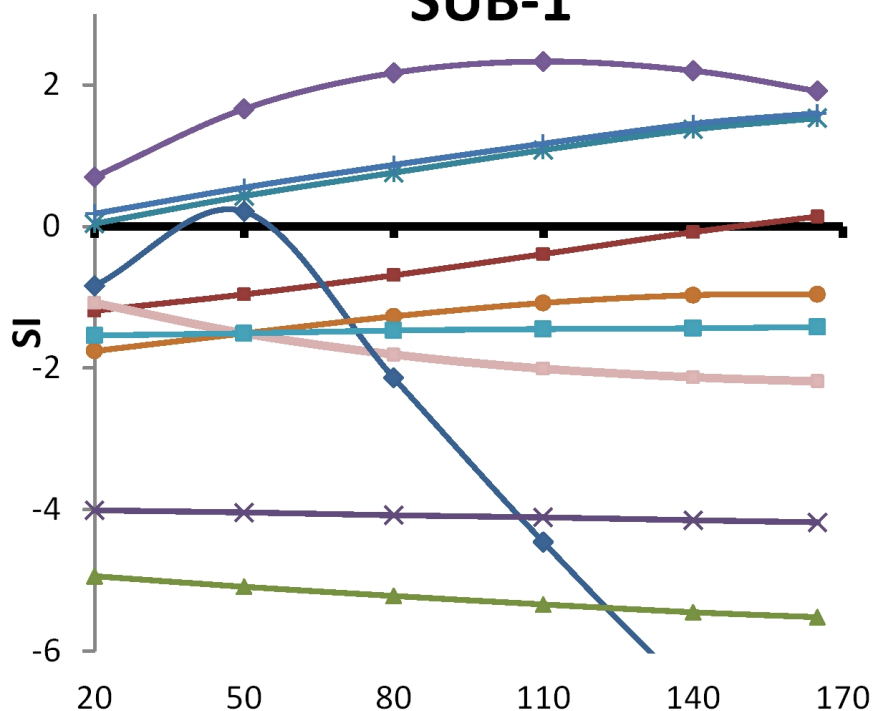




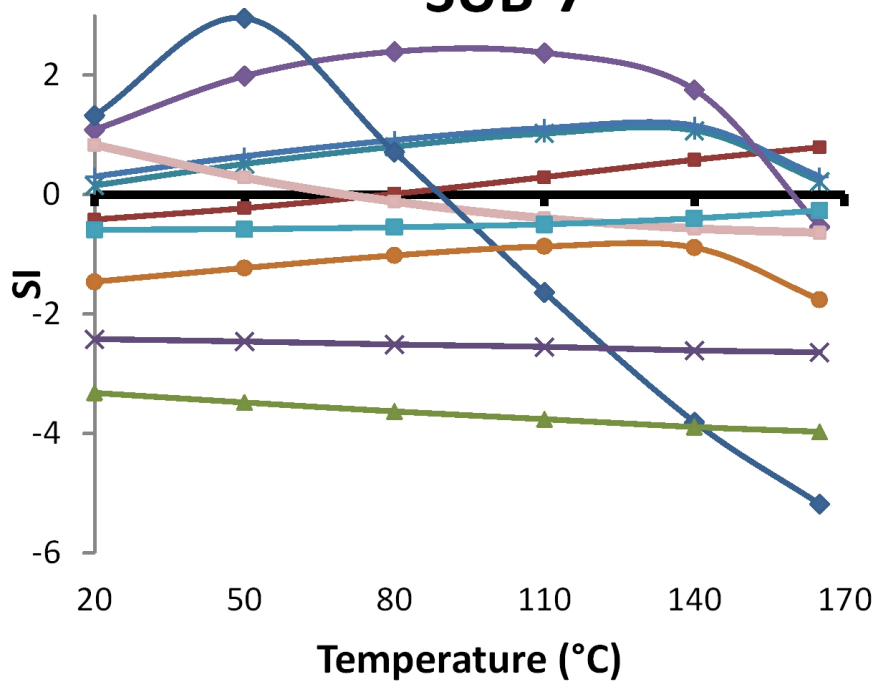


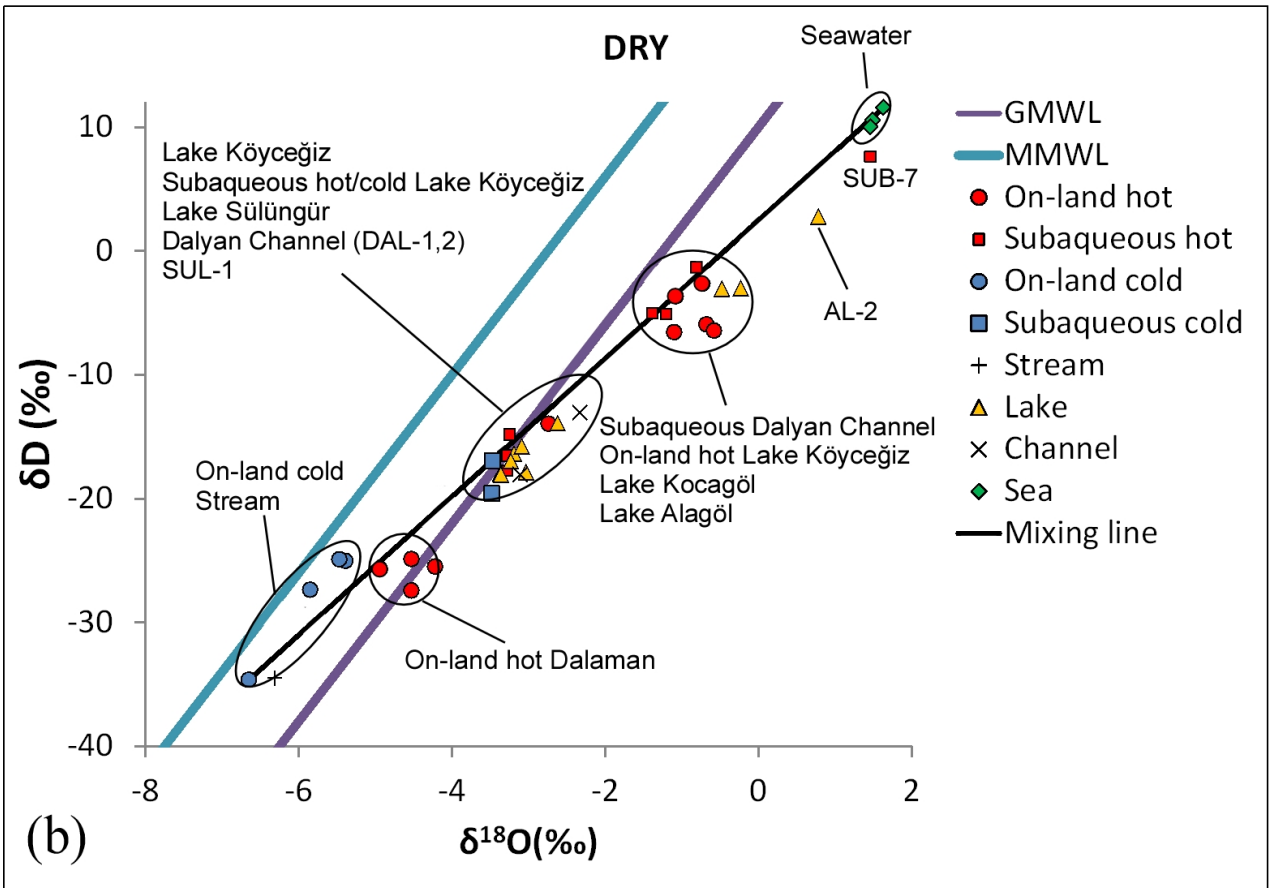
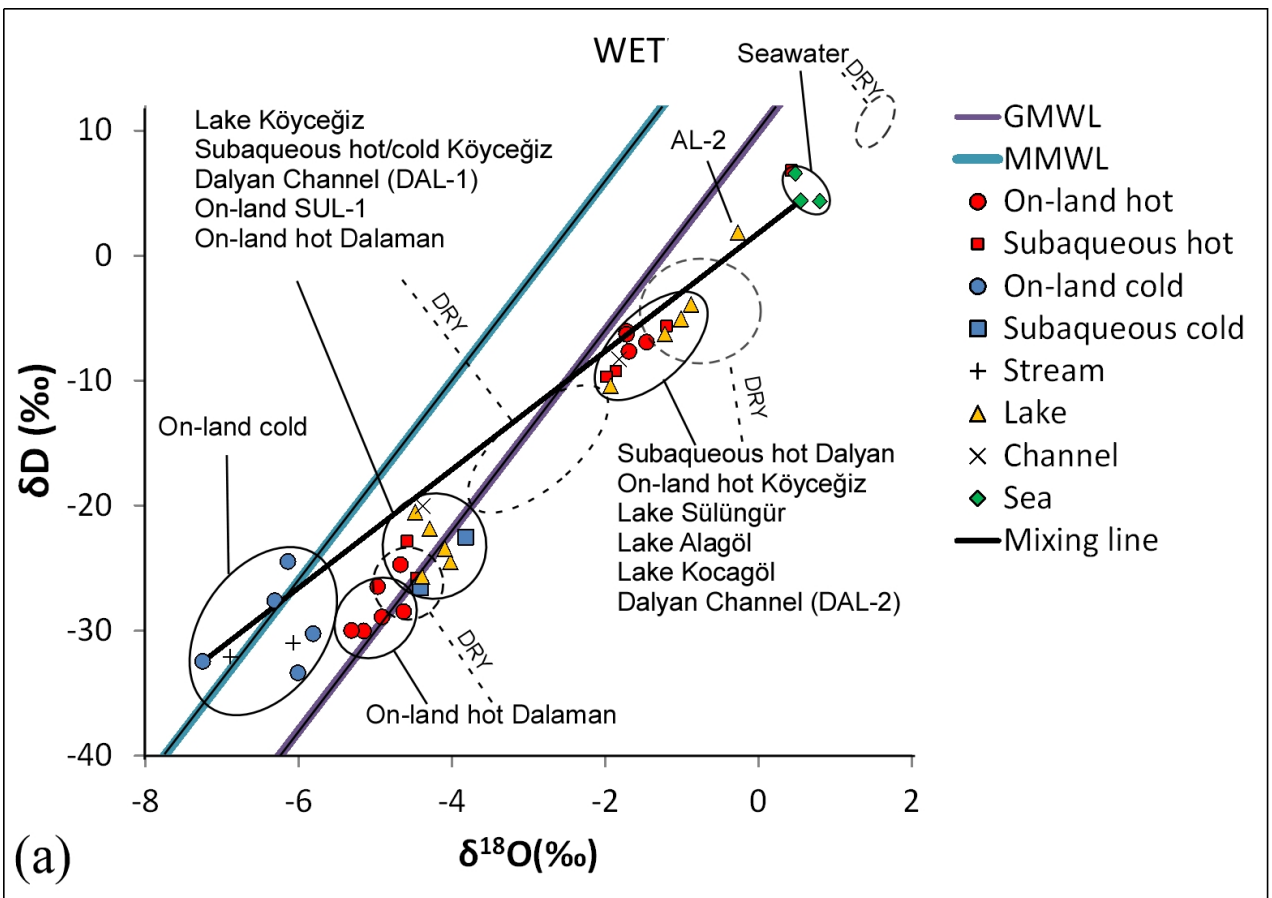


## SUB-1

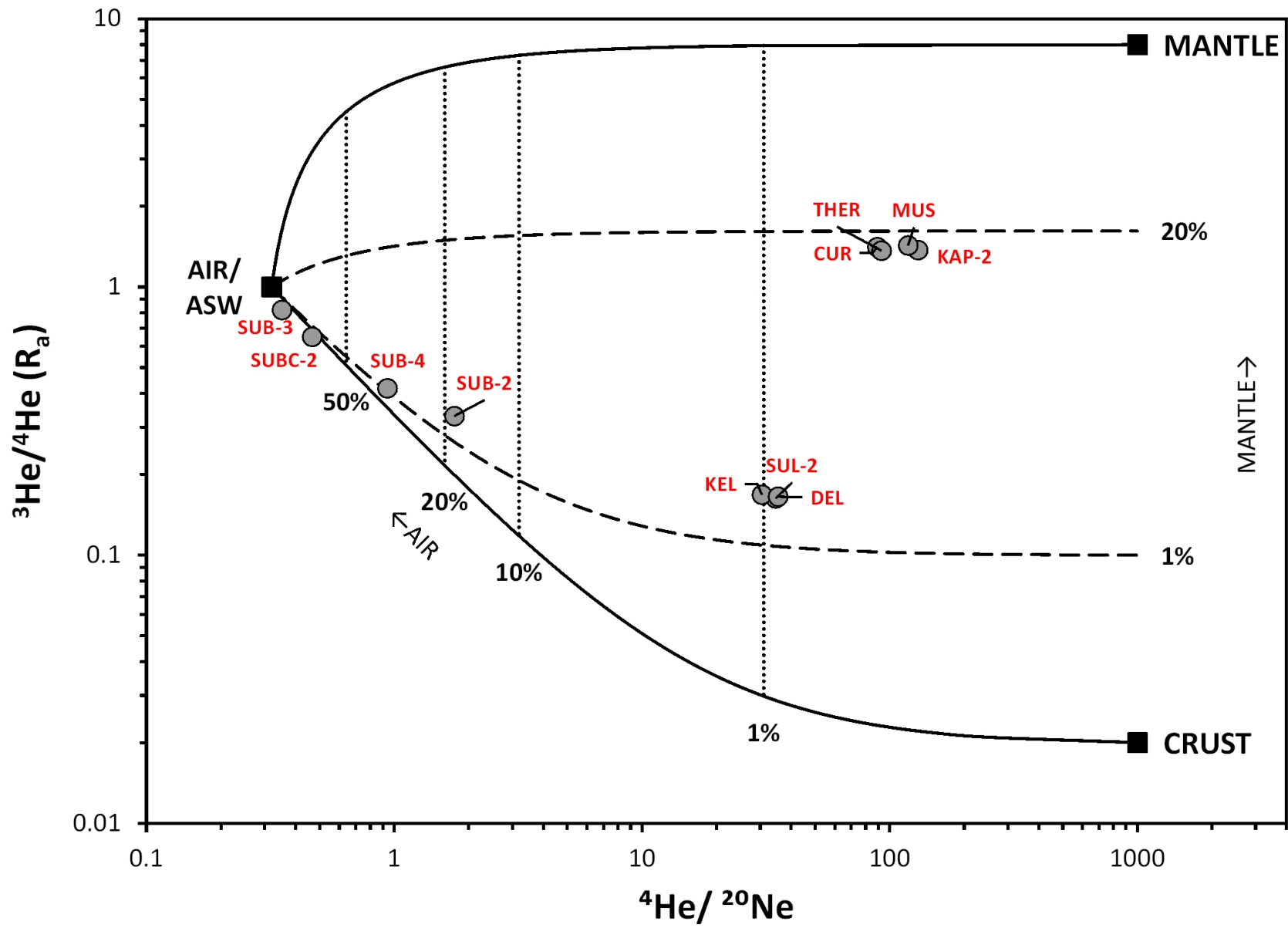


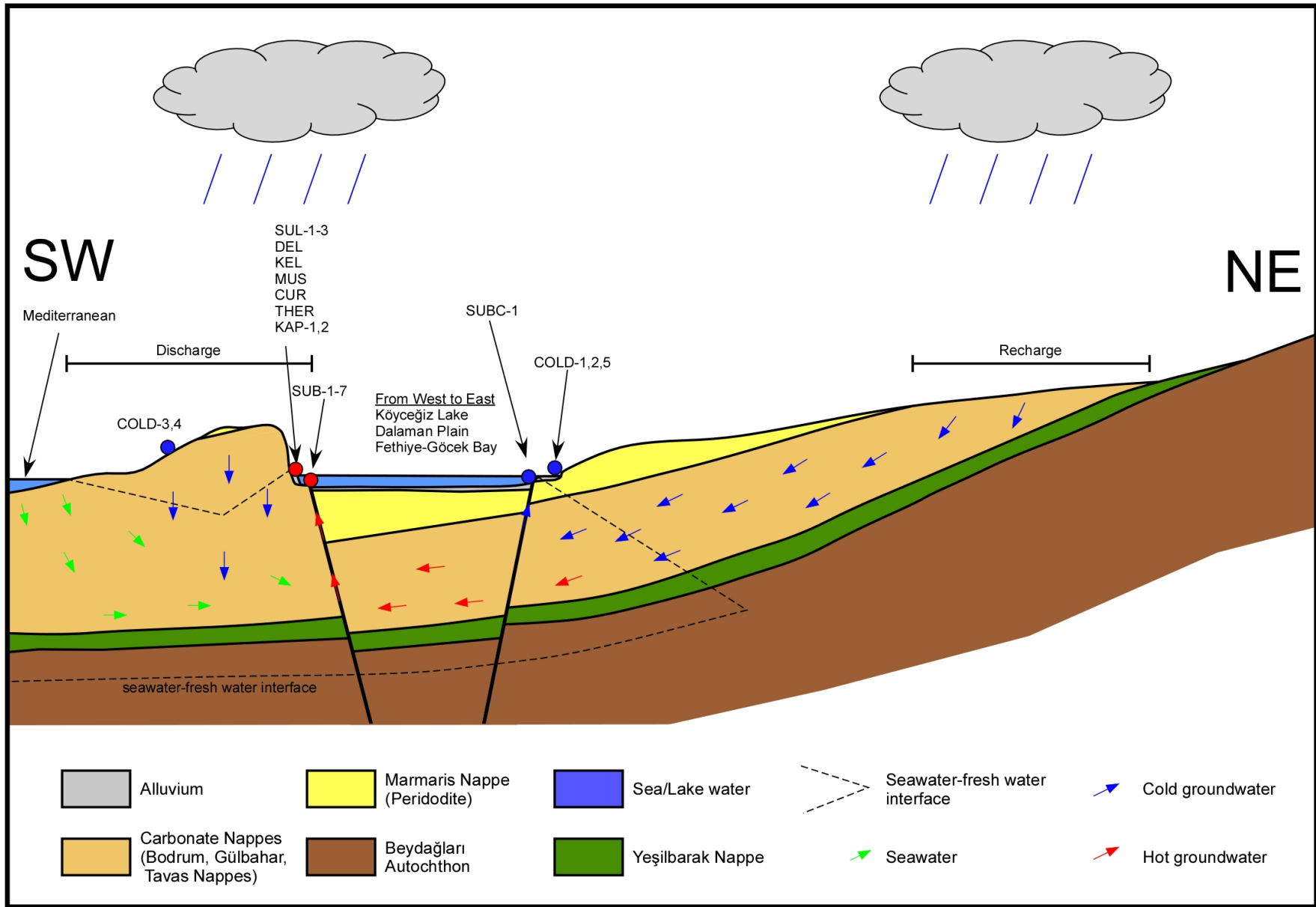
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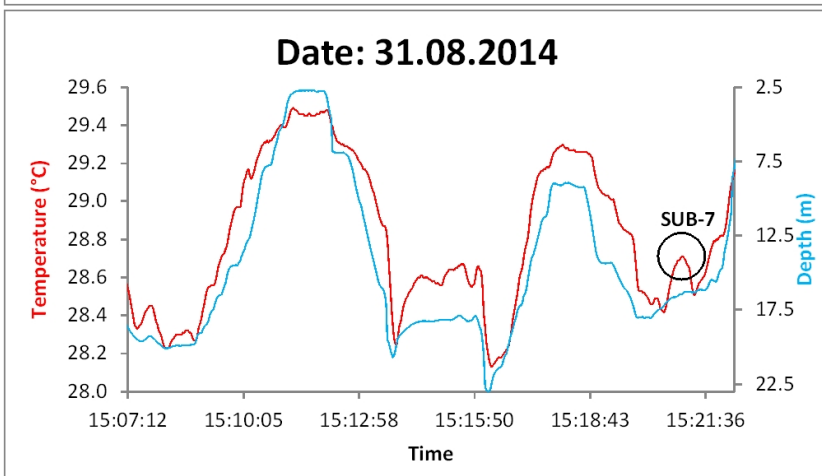
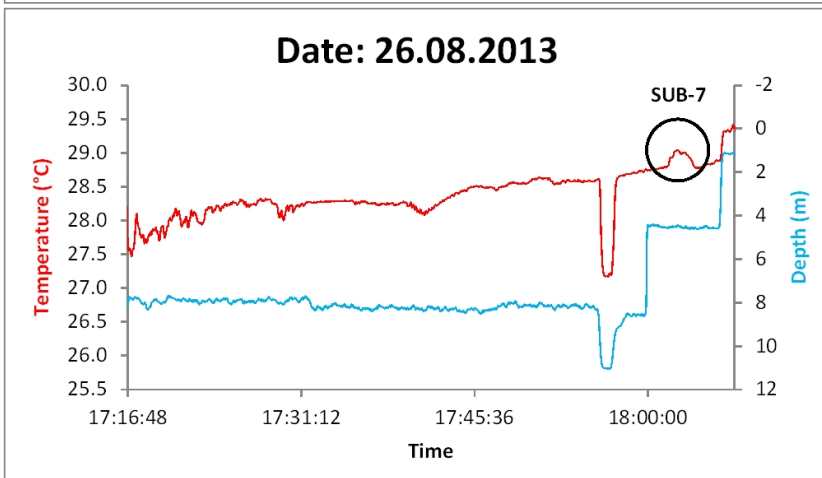
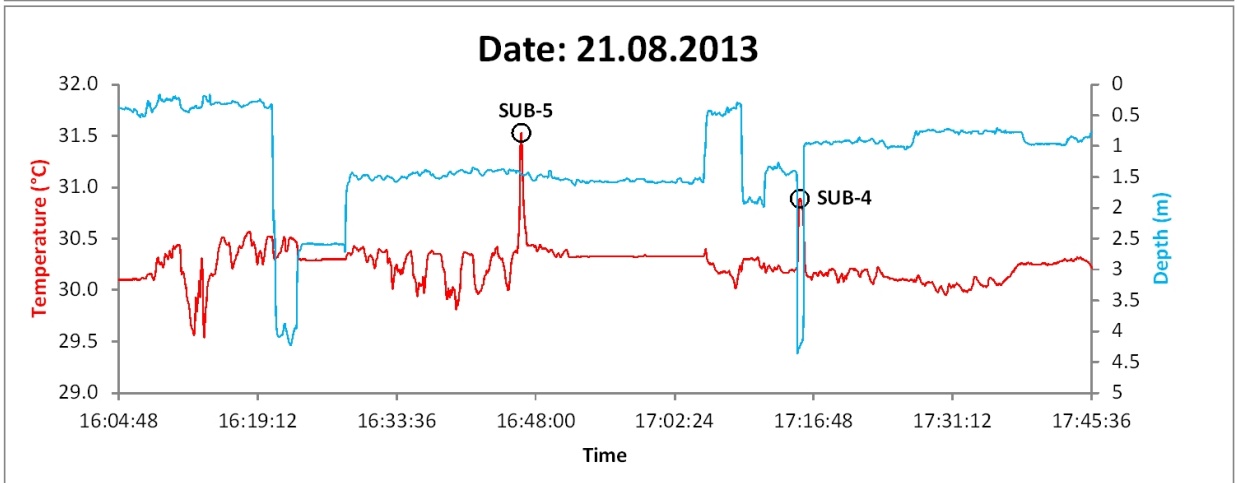
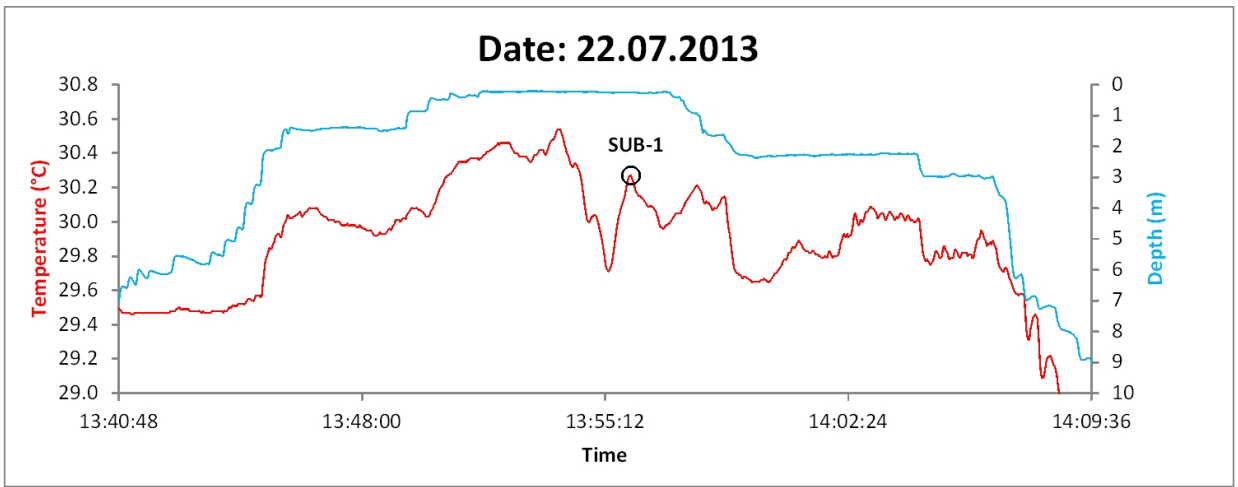






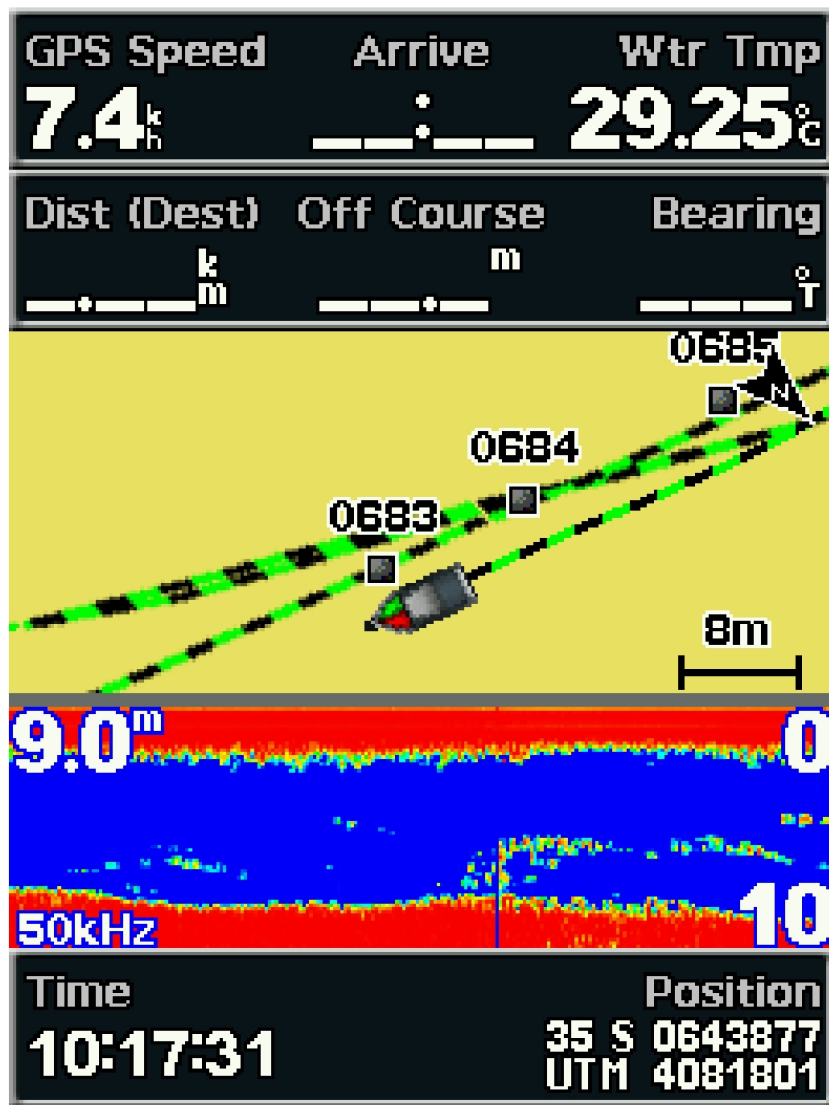




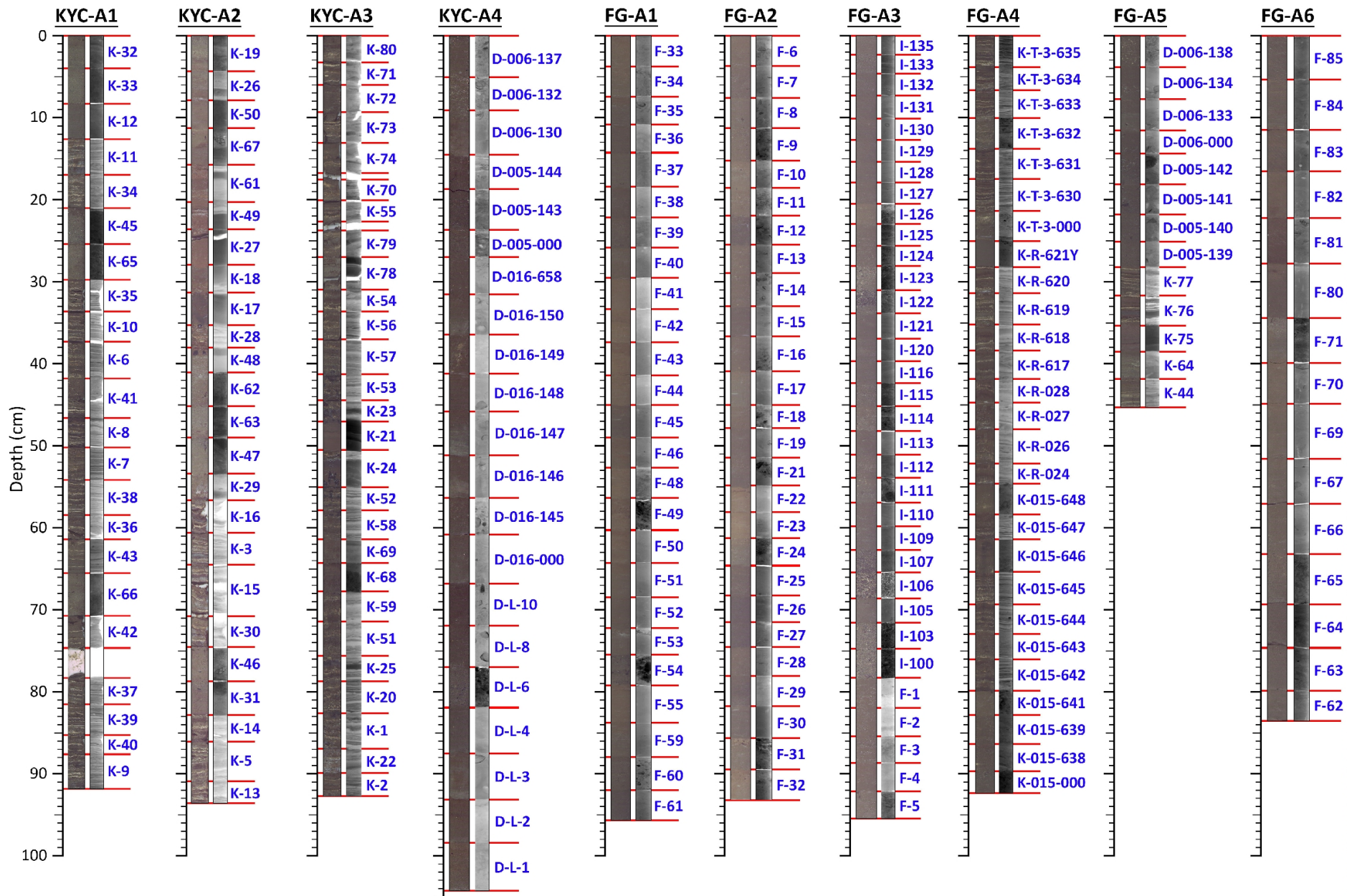




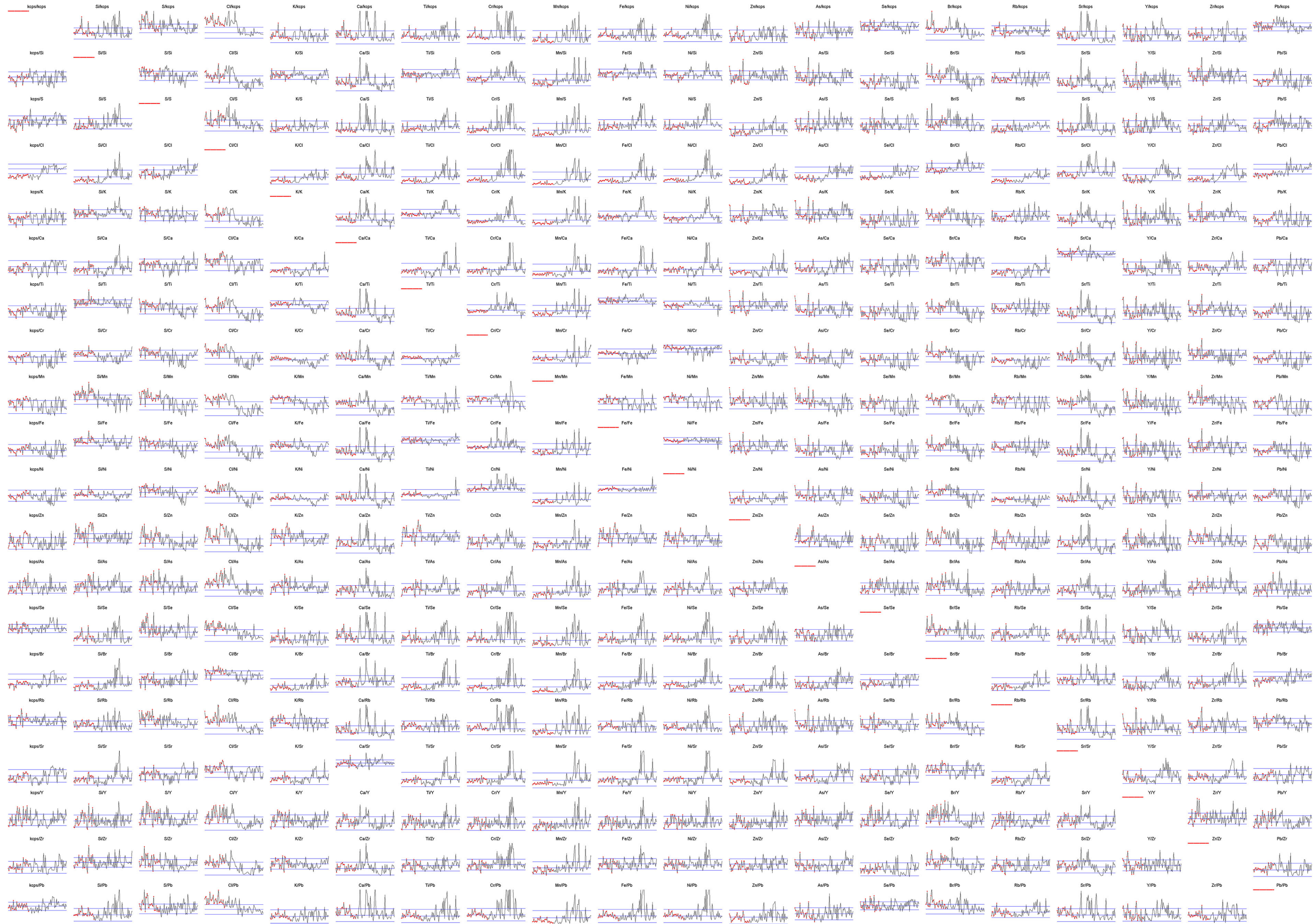
**SUB-2**



**SUB-3**



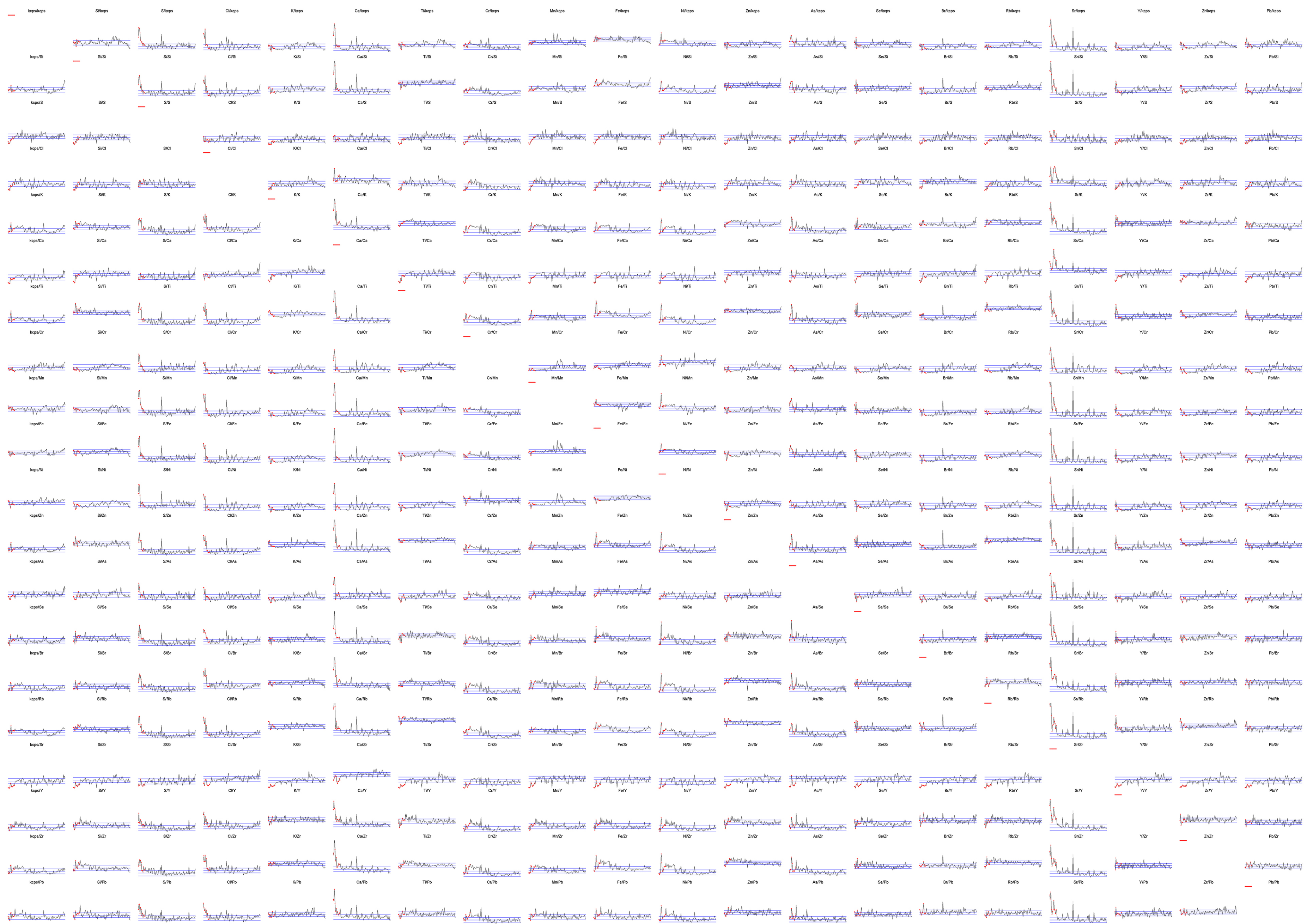




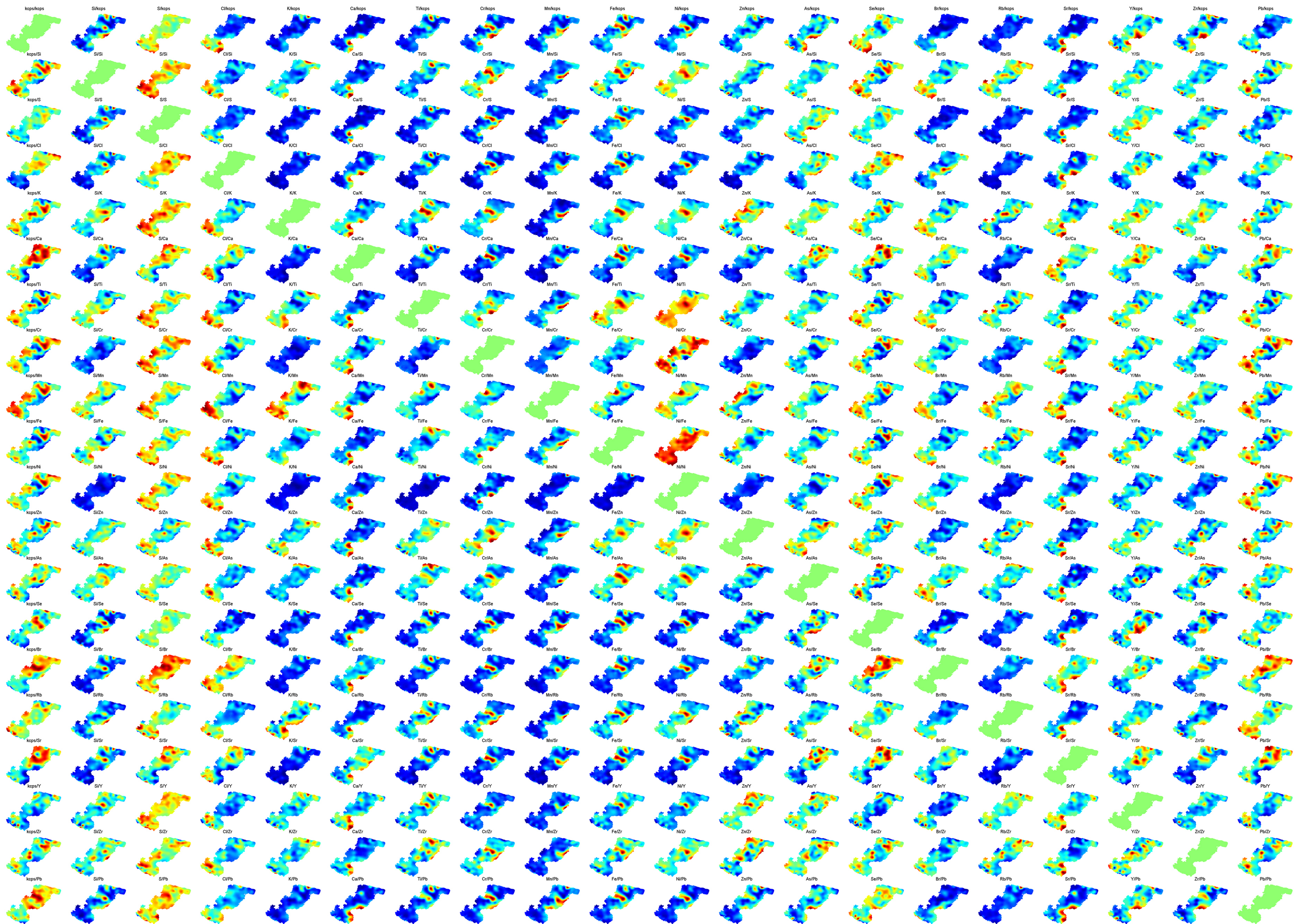




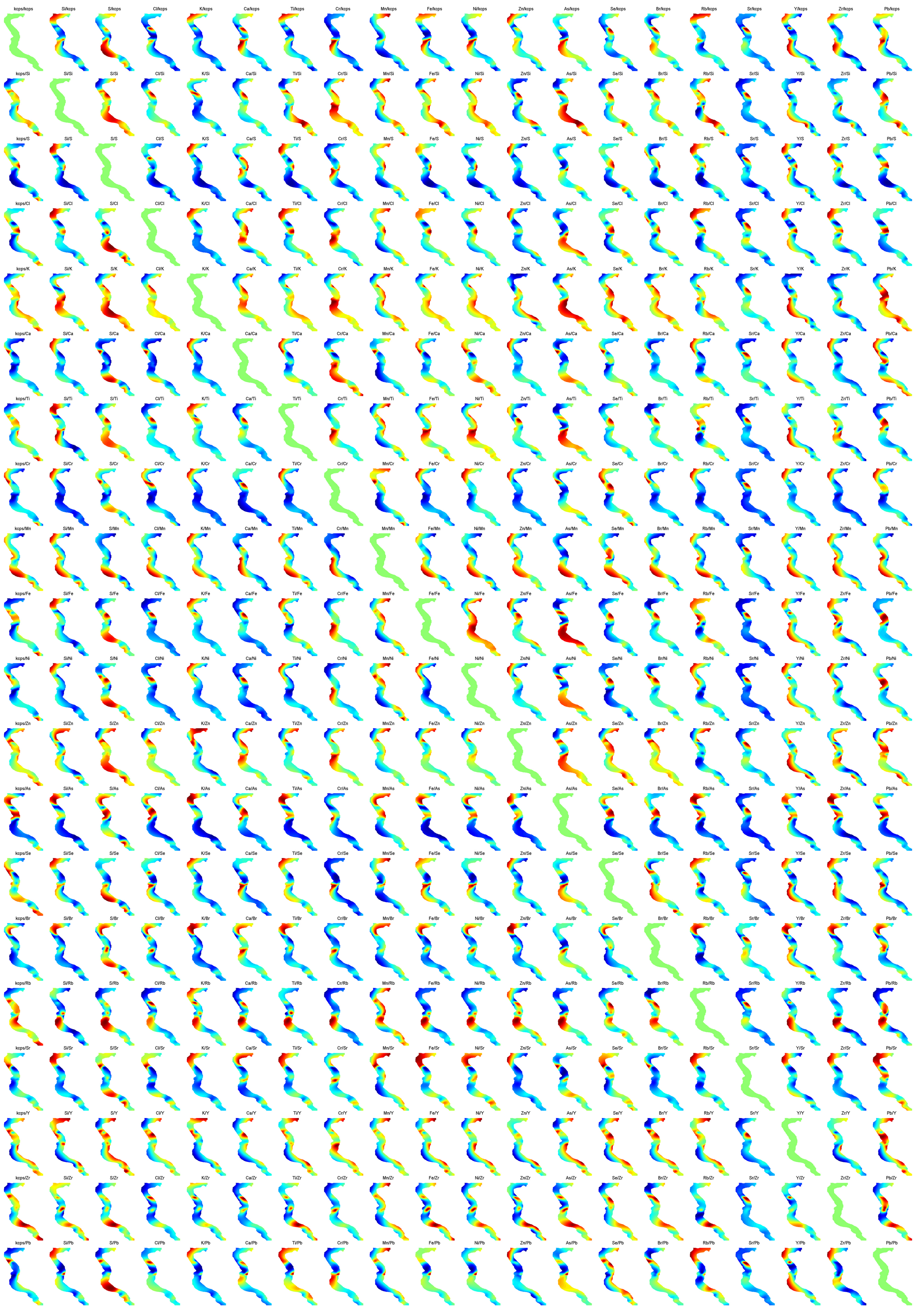














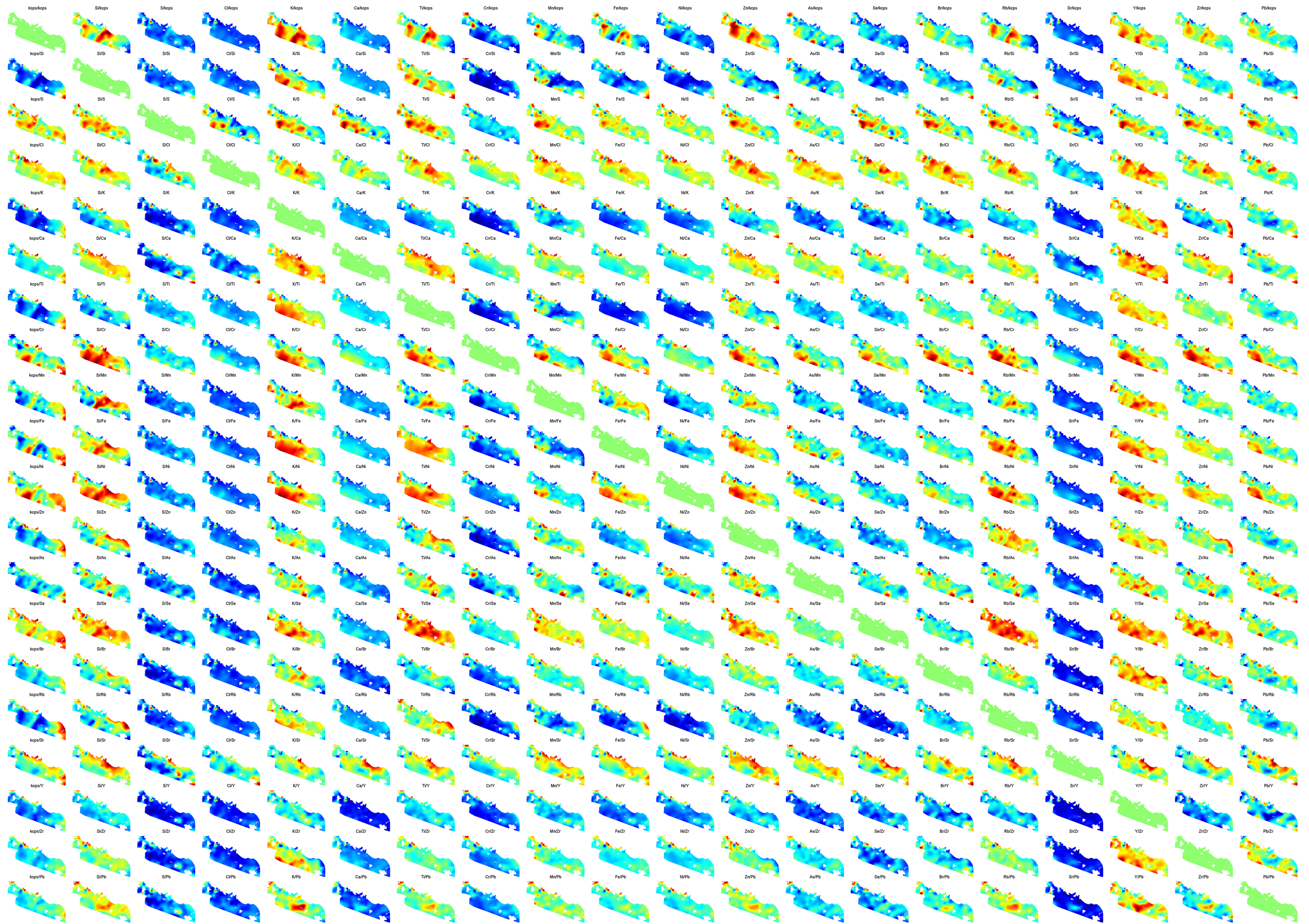




Table 1. Sampling locations and water parameters.

	Type	Location	Sample label	Sampling date	Coordinates (UTM / ED50)		Temp. (°C)	TDS <sup>a</sup> (ppm)	EC <sup>b</sup> (mS/cm)	pH
					E	N				
1	On-land hot	Sultaniye-1	SUL-1	01.06.2014	642788	4082305	24.30	26438	18.700	6.92
2	On-land hot	Sultaniye-2	SUL-2	21.06.2014	642853	4082265	39.93	27127	44.250	6.74
3	On-land hot	Sultaniye-3	SUL-3	21.06.2014	642805	4082293	41.04	28163	44.100	6.70
4	On-land hot	Kelgirme	KEL	02.07.2014	645593	4078489	37.71	27302	27.410	6.76
5	On-land hot	Delibey	DEL	21.06.2014	644921	4080156	38.98	28841	44.390	6.75
6	On-land hot	Çürükardı	CUR	22.06.2014	656671	4065767	29.27	23029	24.400	6.43
7	On-land hot	Musalar	MUS	22.06.2014	654393	4068427	26.58	23031	25.630	6.5
8	On-land hot	Thermemaris	THER	26.06.2014	660168	4062670	26.60	20042	22.740	6.55
9	On-land hot	Kapukargın-1	KAP-1	22.06.2014	661828	4066361	30.02	19505	22.590	6.44
10	On-land hot	Kapukargın-2	KAP-2	22.06.2014	661869	4066291	30.16	21434	22.410	6.43
11	Subaqueous hot	Köyceğiz Lake	SUB-1	05.09.2013	642409	4082916	29.47	6671	11648	-
12	Subaqueous hot	Köyceğiz Lake	SUB-2	28.06.2014	642725	4082889	26.29	5629	5.102	8.67
13	Subaqueous hot	Köyceğiz Lake	SUB-3	28.06.2014	643870	4081804	26.17	7200	4.951	8.69
14	Subaqueous hot	Dalyan Channel	SUB-4	28.06.2014	645593	4079204	27.37	24354	34.406	8.23
15	Subaqueous hot	Dalyan Channel	SUB-5	28.06.2014	645679	4078895	27.47	21572	33.951	8.26
16	Subaqueous hot	Dalyan Channel	SUB-6	28.06.2014	646005	4078338	27.61	20547	30.134	8.25
17	Subaqueous hot	Fethiye-Göcek Bay	SUB-7	28.06.2014	675787	4066071	27.19	6671	57.734	8.35
18	On-land cold	Köyceğiz	COLD-1	21.06.2014	650135	4091604	15.03	302	0.305	8.65
19	On-land cold	Köyceğiz	COLD-2	21.06.2014	654921	4089336	18.72	611	0.627	7.92
20	On-land cold	Sultaniye Village	COLD-3	26.06.2014	641110	4084397	23.77	718	0.709	8.57
21	On-land cold	Çandır Village	COLD-4	21.06.2014	642438	4077514	17.94	409	0.470	7.02
22	On-land cold	İnlince Village	COLD-5	26.06.2014	675818	4067938	18.47	613	0.800	7.32
23	Subaqueous cold	Köyceğiz Lake	SUBC-1	28.06.2014	651338	4091283	25.07	2581	2.780	8.84
24	Subaqueous cold	Köyceğiz Lake	SUBC-2	28.06.2014	646048	4084218	27.70	2802	3.698	8.71
25	Stream	Namnam Stream	NAM	22.06.2014	644396	4087801	22.72	473	0.539	8.12
26	Stream	Yuvarlakçay Stream	YUV	22.06.2014	650966	4086369	20.07	540	0.594	7.83
27	Lake	Köyceğiz Lake	KOY-1	28.06.2014	650033	4089319	27.60	2917	3.635	8.88
28	Lake	Köyceğiz Lake	KOY-2	28.06.2014	648020	4086829	27.52	2865	3.702	8.91
29	Lake	Köyceğiz Lake	KOY-3	28.06.2014	645751	4084993	27.84	2849	3.848	8.89
30	Lake	Köyceğiz Lake	KOY-4	28.06.2014	643216	4083815	28.60	3247	3.940	8.82
31	Lake	Köyceğiz Lake	KOY-5	28.06.2014	645147	4082863	27.73	3258	4.056	8.82
32	Lake	Alagöl Lake	ALA-1	28.06.2014	643431	643431	32.15	27337	50.565	8.18
33	Lake	Alagöl Lake	ALA-2	10.09.2013	643727	4076043	32.90	37065	53.285	8.22
34	Lake	Sülüngür Lake	SULG-1	01.07.2014	645231	4073821	28.92	5254	35.698	8.26
35	Lake	Sülüngür Lake	SULG-2	01.07.2014	646606	4072741	29.75	10498	35.138	8.44
36	Lake	Kocagöl Lake	KOC	22.06.2014	663646	4061694	28.51	4621	7.261	8.31
37	Channel	Dalyan Channel	DAL-1	28.06.2014	645606	4079815	28.09	3433	5.316	8.75
38	Channel	Dalyan Channel	DAL-2	28.06.2014	645861	4076524	27.83	21549	37.126	8.21
39	Sea	Fethiye-Göcek Bay	GOC-1	02.06.2014	675724	4063884	21.13	41133	57.562	8.38
40	Sea	Fethiye-Göcek Bay	GOC-2	05.06.2014	669676	4061469	22.10	41528	56.654	8.38
41	Sea	Fethiye-Göcek Bay	GOC-3	03.06.2014	674596	4056408	22.71	41526	57.482	8.38

<sup>a</sup> Total dissolved solids

<sup>b</sup> Electrical conductivity

Table 2. Major ion concentrations of the waters. CBE: Charge Balance Error.

Sample No	Ca (ppm)	Mg (ppm)	Na (ppm)	K (ppm)	HCO <sub>3</sub> (ppm)	SO <sub>4</sub> (ppm)	Cl (ppm)	CBE (%)	Type
SUL-1	1994	946	6510	284	302	1453	14874	1.30	Na - Cl
SUL-2	2294	1280	6216	355	302	2136	14453	4.27	Na - Cl
SUL-3	2588	1289	6123	389	302	2233	15067	3.26	Na - Cl
KEL	2478	1262	6028	348	178	2068	14843	3.38	Na - Cl
DEL	2928	1321	6161	377	308	2285	15262	4.55	Na - Cl
CUR	2055	895	5009	269	847	934	12894	0.24	Na - Cl
MUS	2059	975	5353	212	900	991	12431	4.16	Na - Cl
THER	1881	907	4403	196	794	804	10932	3.53	Na - Cl
KAP-1	2183	833	4012	232	865	895	10348	4.50	Na - Cl
KAP-2	1857	785	4815	218	882	948	11749	0.50	Na - Cl
SUB-1	378	414	1514	64	302	596	3369	3.36	Na - Cl
SUB-2	414	369	1076	79	320	477	2863	1.93	Na - Cl
SUB-3	414	451	1636	82	317	498	3740	4.60	Na - Cl
SUB-4	916	1378	5929	318	255	1743	13752	-0.47	Na - Cl
SUB-5	944	1020	5962	283	267	1481	11551	4.72	Na - Cl
SUB-6	946	904	5534	304	267	1440	11081	3.12	Na - Cl
SUB-7	1739	2400	10293	532	184	2948	23037	2.05	Na - Cl
COLD-1	17.14	35.09	5.79	0.62	219	2.57	4.04	2.67	Mg - HCO <sub>3</sub>
COLD-2	16.59	84.70	6.25	0.70	444	5.01	6.91	3.18	Mg - HCO <sub>3</sub>
COLD-3	8.74	114.64	6.58	0.38	503	4.15	8.60	4.11	Mg - HCO <sub>3</sub>
COLD-4	86.03	6.03	11.58	0.99	267	8.17	16.39	2.76	Ca - HCO <sub>3</sub>
COLD-5	84.13	31.94	35.51	1.60	361	16.35	49.76	4.21	Ca - HCO <sub>3</sub>
SUBC-1	127	190	513	29	320	172	1204	2.34	Na - Cl
SUBC-2	121	196	617	31	326	184	1305	3.92	Na - Cl
NAM	16.04	70.49	5.18	0.77	354	4.22	6.31	2.09	Mg - HCO <sub>3</sub>
YUV	77.87	40.11	5.57	0.81	302	45.43	5.79	3.38	Ca - HCO <sub>3</sub>
KOY-1	108	204	681	34	284	185	1375	4.55	Na - Cl
KOY-2	100	207	653	30	290	182	1358	3.66	Na - Cl
KOY-3	96	207	661	30	314	180	1335	4.80	Na - Cl
KOY-4	100	224	804	34	264	207	1566	4.98	Na - Cl
KOY-5	104	221	806	36	267	207	1568	4.92	Na - Cl
ALA-1	1164	1626	7038	347	267	1604	15173	4.01	Na - Cl
ALA-2	1669	2056	9747	494	207	2616	20174	4.61	Na - Cl
SULG-1	179	332	1286	67	302	352	2702	2.81	Na - Cl
SULG-2	298	603	2847	141	290	690	5585	4.01	Na - Cl
KOC	210	204	1232	49	267	261	2355	3.21	Na - Cl
DAL-1	113	204	893	36	267	218	1647	4.84	Na - Cl
DAL-2	1715	1262	4843	215	302	1112	12023	4.84	Na - Cl
GOC-1	1426	2454	11037	629	130	2752	22574	4.83	Na - Cl
GOC-2	1529	2450	11166	555	178	2768	22782	4.94	Na - Cl
GOC-3	1378	2486	11056	554	184	2812	22940	3.94	Na - Cl

Table 3. Trace element concentrations of the waters. Piccogram/gram

Sample No	Al (ppb)	As (ppb)	Ba (ppb)	Br (ppm)	Cr (ppb)	Cu (ppb)	F (ppm)	Fe (ppb)	Mn (ppb)	Ni (ppb)	Se (ppb)
Detection Limit	0.3	0.025	0.002	0.01	0.003	0.02	0.01	0.05	0.03	0.03	0.01
SUL-1	775	258	165	31	96	<0.02	6	3279	<0.03	125	202
SUL-2	<0.3	365	109	42	99	<0.02	11	2963	<0.03	165	334
SUL-3	<0.3	234	116	44	110	<0.02	15	3052	<0.03	49	473
KEL	<0.3	405	98	36	118	440	8	3323	<0.03	23	357
DEL	<0.3	290	104	47	130	13070	7	3574	<0.03	87	309
CUR	<0.3	319	95	22	119	<0.02	10	3439	<0.03	45	235
MUS	<0.3	121	80	25	136	<0.02	8	4602	102	35	311
THER	<0.3	216	119	18	141	<0.02	10	3897	<0.03	24	320
KAP-1	<0.3	373	97	20	154	<0.02	19	4924	115	54	210
KAP-2	<0.3	375	103	29	165	<0.02	15	5895	<0.03	47	324
SUB-1	<0.3	103	34	9	60	<0.02	1	1451	<0.03	<0.1	79
SUB-2	<0.3	98	20	5	70	<0.02	<0.01	1851	<0.03	<0.1	68
SUB-3	<0.3	113	33	8	72	<0.02	1	2025	<0.03	<0.1	72
SUB-4	<0.3	386	47	41	236	<0.02	2	11100	<0.03	50	249
SUB-5	<0.3	186	36	34	131	<0.02	4	3558	<0.03	<0.1	223
SUB-6	<0.3	253	36	36	183	<0.02	1	4925	<0.03	<0.1	267
SUB-7	<0.3	229	17	67	175	<0.02	10	4350	<0.03	<0.1	350
COLD-1	<0.3	9	24	<0.01	11	<0.02	<0.01	134	<0.03	<0.1	2
COLD-2	<0.3	6	1	<0.01	9	<0.02	<0.01	125	<0.03	<0.1	2
COLD-3	<0.3	33	2	<0.01	26	<0.02	<0.01	614	<0.03	20	11
COLD-4	<0.3	8	97	<0.01	3	<0.02	<0.01	45	<0.03	<0.1	3
COLD-5	<0.3	6	84	<0.01	4	<0.02	<0.01	43	<0.03	<0.1	<0.1
SUBC-1	<0.3	28	38	4	33	<0.02	1	1015	<0.03	<0.1	28
SUBC-2	<0.3	44	33	4	31	<0.02	<0.01	768	<0.03	2	31
NAM	26	1	11	<0.01	4	<0.02	<0.01	<0.05	<0.03	2	<0.1
YUV	<0.3	4	105	<0.01	6	<0.02	<0.01	66	<0.03	1	3
KOY-1	<0.3	50	39	4	26	<0.02	2	655	<0.03	3	61
KOY-2	<0.3	48	95	4	27	31	<0.01	735	<0.03	2	49
KOY-3	<0.3	47	34	4	30	<0.02	<0.01	726	<0.03	5	47
KOY-4	<0.3	41	36	4	39	<0.02	<0.01	1271	<0.03	9	50
KOY-5	<0.3	55	37	5	30	<0.02	1	776	<0.03	7	42
ALA-1	<0.3	215	54	41	291	<0.02	11	10280	<0.03	101	326
ALA-2	<0.3	290	26	62	173	<0.02	3	4713	<0.03	<0.1	392
SULG-1	<0.3	47	37	8	25	<0.02	3	638	<0.03	7	64
SULG-2	<0.3	140	44	16	56	<0.02	4	1394	<0.03	12	122
KOC	<0.3	63	76	6	37	<0.02	1	1470	<0.03	7	45
DAL-1	<0.3	66	37	5	28	<0.02	<0.01	771	<0.03	1	45
DAL-2	<0.3	282	69	38	210	<0.02	3	8355	<0.03	32	164
GOC-1	<0.3	266	30	68	173	549	3	5710	<0.03	17	424
GOC-2	<0.3	251	12	70	175	<0.02	5	5171	<0.03	98	344
GOC-3	<0.3	432	25	67	162	<0.02	2	4111	<0.03	18	430

Table 4. Results from the stable isotope analysis of the waters.

Sample no	DRY		WET		DIFFERENCE	
	$\delta^{18}\text{O}$ (‰)	$\delta\text{D}$ (‰)	$\delta^{18}\text{O}$ (‰)	$\delta\text{D}$ (‰)	$\delta^{18}\text{O}$ (‰)	$\delta\text{D}$ (‰)
SUL-1	-2.74	-14.0	-4.67	-24.7	1.93	10.8
SUL-2	-0.68	-5.9	-1.72	-6.1	1.04	0.1
SUL-3	-0.58	-6.4	-1.46	-6.9	0.88	0.5
KEL	-1.10	-6.6	-1.69	-7.7	0.59	1.1
DEL	-0.73	-2.7	-1.72	-6.3	0.99	3.7
CUR	-4.53	-24.9	-4.91	-28.9	0.38	4.0
MUS	-4.22	-25.5	-4.63	-28.5	0.41	3.0
THER	-4.53	-27.4	-4.97	-26.5	0.44	-1.0
KAP-1	-4.94	-25.7	-5.15	-30.1	0.21	4.4
KAP-2	-1.08	-3.7	-5.31	-30.0	4.23	26.3
SUB-1	-3.25	-14.8				
SUB-2	-3.28	-17.7	-4.59	-22.9	1.31	5.1
SUB-3	-3.28	-16.5	-4.46	-25.8	1.18	9.3
SUB-4	-0.80	-1.3	-1.86	-9.2	1.06	7.9
SUB-5	-1.38	-5.1	-1.98	-9.7	0.60	4.6
SUB-6	-1.20	-5.1	-1.20	-5.7	0.00	0.5
SUB-7	1.46	7.6	0.43	6.8	1.03	0.8
COLD-1	-6.65	-34.6	-7.25	-32.5	0.60	-2.2
COLD-2	5.86	-28.4	-6.01	-33.4	11.87	5.0
COLD-3	-5.39	-25.0	-6.14	-24.5	0.75	-0.6
COLD-4	-5.47	-24.9	-5.81	-30.3	0.34	5.4
COLD-5	-5.85	-27.4	-6.31	-27.6	0.46	0.3
SUBC-1	-3.47	-16.9	-4.41	-26.6	0.94	9.6
SUBC-2	-3.48	-19.6	-3.82	-22.5	0.34	2.9
NAM	-	-	-6.07	-31.0	-	-
YUV	-6.31	-34.5	-6.89	-32.1	0.58	-2.4
KOY-1	-3.37	-18.0	-4.02	-24.5	0.65	6.5
KOY-2	-3.36	-18.1	-4.39	-25.7	1.03	7.6
KOY-3	-3.03	-17.9	-4.09	-23.5	1.06	5.6
KOY-4	-3.19	-16.4	-4.29	-21.9	1.10	5.5
KOY-5	-3.09	-15.8	-4.48	-20.5	1.39	4.7
ALA-1	-0.48	-3.07	-1.93	-10.4	1.45	7.3
ALA-2	0.78	2.76	-0.27	1.8	1.05	0.9
SULG-1	-3.23	-17.0	-0.88	-3.9	-2.35	-13.1
SULG-2	-2.62	-13.9	-1.22	-6.3	-1.40	-7.7
KOC	-0.23	-3.0	-1.01	-5.1	0.78	2.1
DAL-1	-3.11	-18.0	-4.38	-20.1	1.27	2.0
DAL-2	-2.33	-13.1	-1.82	-8.3	-0.51	-4.8
GOC-1	1.63	11.6	0.80	4.3	0.83	7.2
GOC-2	1.49	10.5	0.55	4.4	0.94	6.2
GOC-3	1.46	10.0	0.48	6.6	0.98	3.5

Table 5. Total gas composition,  $^3\text{He}/^4\text{He}$ ,  $^4\text{He}/^{20}\text{Ne}$  and  $\text{CO}_2/{}^3\text{He}$  ratios for free gas samples and  $^3\text{He}/^4\text{He}$  and  $^4\text{He}/^{20}\text{Ne}$  ratios for dissolved gases from hot springs. Error limits are  $2\sigma$ .

Sample ID	Sample type	Sampling date	Ar (%)	CH <sub>4</sub> (%)	CO <sub>2</sub> (%)	N <sub>2</sub> (%)	O <sub>2</sub> (%)	He (ppm)	H <sub>2</sub> (ppm)	N <sub>2</sub> /CH <sub>4</sub>	N <sub>2</sub> /Ar	CO <sub>2</sub> / <sup>3</sup> He (*10 <sup>9</sup> )	<sup>4</sup> He/ <sup>20</sup> Ne error	<sup>3</sup> He/ <sup>4</sup> He (R <sub>a</sub> )	<sup>3</sup> He/ <sup>4</sup> He (R <sub>a</sub> ) error	( <sup>3</sup> He/ <sup>4</sup> He) <sub>c</sub> <sup>a</sup> (R <sub>a</sub> )	( <sup>3</sup> He/ <sup>4</sup> He) <sub>c</sub> (R <sub>a</sub> ) error	Mantle Helium <sup>b</sup> (%)	
SUL-2	Free Gas	7/8/2014	1.17	0.1	5.1	93	0.2	220	165	650	79	1.026	34.6	±1.7	0.163	±0.012	0.156	±0.012	2
KEL	Free Gas	7/17/2014	1.17	0	4.6	92	2.5	180	165		78	1.097	30.5	±1.6	0.168	±0.012	0.160	±0.012	2
DEL	Free Gas	7/8/2014	1	0.1	4.5	94	0	210	300	870	94	0.932	35.4	±1.8	0.165	±0.011	0.158	±0.011	2
MUS	Free Gas	7/9/2014	0.76	0	4.2	91	4.3	420	50	9100	119	0.050	118.7	±6.0	1.426	±0.027	1.427	±0.027	18
THER	Free Gas	7/9/2014	0.94	0	10.3	88	0.3	460	170	8800	94	0.118	92.6	±4.7	1.365	±0.029	1.366	±0.029	17
KAP-2	Free Gas	7/9/2014	0.72	0.1	21.4	78	0.3	440	135	1270	108	0.255	130.1	±6.5	1.372	±0.027	1.373	±0.027	17
SUB-2	Free Gas	7/17/2014	0.36	62	0.19	34	3.1	28	1460	0.55	96	0.150	1.75	±0.09	0.330	±0.027	0.194	±0.036	2
SUB-3	Dissolved Gas	7/17/2014											0.352	±0.023	0.822	±0.058	<0.61		<7.4
CUR	Dissolved Gas	7/9/2014											89.0	±4.5	1.408	±0.041	1.409	±0.041	17
SUB-4	Dissolved Gas	7/17/2014											0.937	±0.051	0.419	±0.033	0.152	±0.062	2
SUBC-2	Dissolved Gas	7/17/2014											0.466	±0.028	0.651	±0.032	0.05	±0.19	1
Atmosphere/ ASW (25 °C)											83/38	0.319/0.274							

<sup>a</sup> assuming  $(^4\text{He}/^{20}\text{Ne})_{\text{air}} = 0.295 \pm 0.025$ , i.e. a range including air (0.319; Sano and Wakita, 1985) and ASW (25°C) (0.274; Ozima and Podosek, 2002)

<sup>b</sup> assuming  $(^3\text{He}/^4\text{He})_{\text{crust}} = 0.02 R_a$  and  $(^3\text{He}/^4\text{He})_{\text{mantle}} = 8 R_a$



Supplementary Table 1. Standardized ITRAX analysis results for Köyceğiz cores.

	ID	Type	X	Y	kcps	Si	S	Cl	K	Ca	Ti	Cr	Mn	Fe	Ni	Zn	As	Se	Br	Rb	Sr	Y	Zr	Pb
1	K-9	HLC	642097	4083621	-0.74	-0.63	0.43	0.96	-0.65	-0.12	-0.74	-0.62	-0.80	-0.77	-0.55	0.41	1.52	0.34	0.17	-1.05	0.35	1.09	-0.26	-0.53
2	K-40	HLC	642452	4083142	-0.75	-0.60	-0.24	1.16	-0.94	-0.58	-0.72	-0.71	-0.73	-0.67	-0.53	-0.06	0.64	-0.83	0.38	-0.36	-0.16	1.52	0.42	-0.62
3	K-015-648	HLC	642433	4082872	-0.21	-0.63	-0.84	1.24	0.10	0.59	-0.12	-0.36	-0.31	-0.22	-0.03	-0.35	0.59	-0.98	2.77	-0.44	0.93	0.42	-0.38	-0.64
4	K-015-647	HLC	642325	4082985	0.32	-0.06	-0.17	0.79	-0.40	-0.32	-0.35	-0.36	-0.62	-0.35	-0.30	-1.06	-0.39	0.47	0.82	-0.65	-0.35	-0.41	-0.12	0.54
5	K-015-646	HLC	642370	4082945	-0.43	-0.06	2.17	1.29	0.31	0.30	0.14	-0.18	-0.48	0.00	0.11	-0.64	0.53	-0.18	0.73	-0.40	0.43	-0.43	-0.09	0.68
6	K-015-645	HLC	642428	4083010	-0.46	-0.52	1.10	0.85	-0.32	0.09	-0.40	-0.52	-0.71	-0.52	-0.33	-1.06	0.44	0.08	0.21	0.22	0.36	-0.03	-0.34	-0.71
7	K-015-644	HLC	642424	4082965	-0.29	-0.34	1.30	0.69	-0.22	-0.19	-0.37	-0.55	-0.67	-0.39	-0.22	-0.37	-0.94	-1.08	0.35	-0.37	0.04	-1.13	0.17	0.32
8	K-015-643	HLC	642514	4082948	-0.71	-0.22	0.50	0.38	-0.47	-0.05	-0.38	-0.51	-0.72	-0.51	-0.35	-0.55	-0.81	0.02	0.32	-0.04	0.10	1.11	-0.56	-0.26
9	K-015-642	HLC	642461	4082929	-0.36	-0.20	1.97	0.80	-0.24	-0.24	-0.16	-0.33	-0.59	-0.15	-0.06	-0.98	1.31	-0.29	0.57	0.25	-0.41	0.90	-1.27	-1.72
10	K-015-641	HLC	642446	4082822	-1.20	-0.36	-1.37	0.87	0.15	0.69	-0.20	-0.42	-0.29	-0.35	-0.24	-0.18	0.11	-0.87	3.40	-0.34	0.94	0.13	-0.39	-0.46
11	K-015-639	HLC	642450	4082891	-1.20	0.05	1.97	0.60	-0.07	0.68	-0.27	-0.32	-0.50	-0.35	-0.24	-0.29	-0.48	-1.16	0.60	-0.25	0.79	0.21	-0.56	-0.59
12	K-015-638	HLC	642491	4082865	-1.12	-0.22	1.37	0.47	0.06	0.20	-0.07	-0.20	-0.36	-0.17	-0.11	-0.03	-0.15	-0.74	0.54	-0.24	0.15	-1.47	-0.65	-0.96
13	K-015-000	HLC	642409	4082916	-1.95	0.52	1.93	0.50	0.74	0.95	0.13	-0.02	-0.09	-0.10	-0.08	0.25	0.18	-1.53	0.40	-0.16	0.58	0.02	0.06	-1.22
14	K-T-3-635	HLC	642694	4082932	0.99	0.70	2.70	1.71	0.31	0.60	0.12	-0.06	-0.51	0.41	0.54	-0.04	0.20	-0.57	0.42	-0.12	0.61	-0.98	-0.27	-0.33
15	K-T-3-634	HLC	642663	4082973	1.24	-0.10	0.93	1.28	0.04	0.05	-0.05	-0.17	-0.54	0.08	0.07	-0.15	-0.37	0.18	0.56	0.04	-0.01	-1.37	1.40	1.02
16	K-T-3-633	HLC	642804	4082807	1.21	0.24	0.43	1.06	-0.09	-0.14	-0.18	-0.09	-0.53	0.19	0.22	-0.32	-0.18	2.84	1.12	-0.08	-0.20	-0.82	0.12	0.57
17	K-T-3-632	HLC	642764	4082855	0.18	-0.33	1.10	1.44	0.22	1.01	-0.20	-0.22	-0.36	-0.27	-0.02	0.38	0.24	0.58	0.82	-0.72	0.78	0.09	0.73	0.27
18	K-T-3-631	HLC	642795	4082968	0.59	-0.25	0.37	1.16	-0.22	0.04	-0.23	-0.21	-0.60	-0.23	-0.12	-0.26	0.06	1.30	0.69	-0.54	0.09	-1.47	1.03	-0.30
19	K-T-3-630	HLC	642763	4082920	0.58	-0.09	1.10	1.17	-0.04	-0.07	-0.20	-0.21	-0.63	0.03	0.04	-0.83	-1.25	0.41	0.44	-0.34	0.00	-0.94	-0.23	0.22
20	K-T-3-000	HLC	642725	4082889	0.11	0.05	1.37	0.74	0.55	0.24	0.24	0.02	-0.14	0.14	0.20	-0.04	0.68	-0.29	0.23	0.37	-0.07	-0.60	-0.66	0.31
21	K-39	HLC	643071	4082508	-0.17	-0.71	0.00	1.33	-0.56	-0.02	-0.66	-0.61	-0.37	-0.52	-0.51	-0.11	0.85	-0.46	0.60	-1.30	0.60	1.41	0.49	-0.63
22	K-37	HLC	643718	4082070	-0.57	-1.03	-0.90	1.43	-0.64	0.49	-0.70	-0.67	-0.70	-0.81	-0.73	0.44	2.20	0.67	0.02	-1.16	0.63	1.70	1.47	-1.39
23	K-R-621Y	HLC	643940	4081750	-0.45	1.03	3.04	0.73	1.18	0.59	0.57	0.39	0.25	0.38	0.18	0.51	-0.66	-1.33	0.34	0.61	-0.22	0.22	0.08	0.11
24	K-R-620	HLC	643987	4082011	1.32	0.98	-0.07	1.34	-0.05	-0.51	-0.33	-0.29	-0.57	-0.21	-0.24	-0.39	-0.78	0.46	1.10	0.16	-0.41	-1.35	-1.10	0.73
25	K-R-619	HLC	643861	4081923	0.77	-0.17	-0.10	1.48	-0.11	0.19	-0.36	-0.44	-0.59	-0.35	-0.39	-1.00	-1.03	-0.55	0.42	-0.07	0.59	0.03	-0.56	0.84
26	K-R-618	HLC	644120	4081843	0.88	0.18	-0.57	1.26	-0.01	-0.61	-0.19	-0.17	-0.52	-0.06	-0.14	-0.99	-1.47	0.26	1.58	0.12	-0.46	-0.11	-0.55	0.07
27	K-R-617	HLC	644065	4081747	0.95	0.05	0.97	1.14	0.14	-0.35	0.07	-0.21	-0.54	0.07	0.10	-1.06	-0.05	-0.28	0.64	0.20	-0.43	-0.60	0.19	0.50
28	K-R-028	HLC	643751	4081904	1.45	0.22	0.23	1.68	-0.11	-0.56	-0.37	-0.36	-0.64	-0.36	-0.38	-0.90	-0.87	1.48	0.83	0.14	-0.48	-1.01	-0.77	1.54
29	K-R-027	HLC	643975	4081886	1.31	-0.22	0.33	1.39	-0.58	-0.59	-0.31	-0.44	-0.66	-0.44	-0.41	-0.80	1.72	1.29	0.89	-0.45	-0.47	1.33	-0.13	0.29
30	K-R-026	HLC	644100	4081976	1.16	0.34	-0.30	1.17	0.04	-0.42	-0.01	-0.23	-0.55	-0.16	-0.09	-0.88	0.59	0.26	1.39	0.04	-0.24	-1.47	-0.77	1.05
31	K-R-024	HLC	643871	4082002	0.89	0.07	0.77	0.85	-0.24	-0.51	-0.30	-0.30	-0.56	-0.21	-0.16	-0.80	-1.93	-0.02	0.94	-0.17	-0.31	0.01	-0.22	1.76
32	K-42	HLC	644494	4081634	-0.39	-0.86	-0.20	0.89	-0.73	0.17	-0.68	-0.74	-0.71	-0.87	-0.73	-0.33	-0.77	0.31	0.91	-0.42	0.87	-0.49	-0.57	1.22
33	K-7	CLC	642862	4083653	-0.11	-0.33	1.43	1.75	-0.44	0.11	-0.55	-0.50	-0.69	-0.26	-0.25	-0.64	-0.73	-0.16	0.90	-0.34	0.60	-0.90	-0.98	0.97

34	K-38	CLC	643387	4082903	0.26	-1.10	-1.24	1.01	-0.85	-0.50	-0.83	-0.76	-0.82	-0.91	-0.75	-1.06	-1.14	0.59	0.35	-0.42	0.25	0.07	-0.48	1.14
35	K-36	CLC	644081	4082513	0.65	-0.61	-1.77	0.87	-0.80	-0.68	-0.65	-0.42	-0.68	-0.70	-0.61	-0.60	-0.85	1.96	1.27	-0.05	-0.35	-1.47	-0.56	1.39

Supplementary Table 1 contuniued.

	ID	Type	X	Y	kcps	Si	S	Cl	K	Ca	Ti	Cr	Mn	Fe	Ni	Zn	As	Se	Br	Rb	Sr	Y	Zr	Pb
36	K-43	CLC	644872	4082142	-0.50	-0.57	-0.77	0.75	-0.31	0.45	-0.44	-0.33	-0.56	-0.36	-0.24	-0.29	-0.22	0.37	0.76	-0.21	0.64	1.21	-0.35	-0.39
37	K-66	CLC	645460	4081551	-1.83	-0.95	-0.54	0.69	-0.40	1.69	-0.75	-0.47	-0.60	-0.90	-0.65	-0.86	-0.85	-0.32	1.39	-1.06	2.22	-0.27	0.27	0.07
38	K-65	CLC	645875	4082013	-1.95	-1.17	-2.17	0.83	-0.39	3.28	-0.82	-0.29	0.03	-1.04	-0.92	-0.69	0.63	-0.77	0.54	-0.91	3.45	-0.51	-0.51	-1.21
39	K-44	CLC	645379	4082750	-0.40	0.64	1.10	-0.04	0.62	0.32	0.00	0.06	-0.38	0.30	0.28	0.37	-0.92	-0.54	0.11	-0.47	0.08	-0.70	0.92	0.02
40	K-35	CLC	644430	4082925	-0.08	-0.87	-0.37	0.74	-0.74	-0.03	-0.63	-0.52	-0.75	-0.62	-0.53	-0.38	-1.03	0.47	0.90	-0.57	0.46	-0.17	-0.07	1.18
41	K-10	CLC	643798	4083413	-0.18	-0.95	-0.03	1.11	-0.69	-0.29	-0.66	-0.60	-0.76	-0.44	-0.32	-0.42	-1.93	0.36	1.15	-0.27	-0.15	-0.70	-1.63	2.25
42	K-6	CLC	643321	4084181	0.01	-0.64	0.37	1.39	-0.47	-0.03	-0.46	-0.45	-0.68	-0.39	-0.30	-0.81	-0.80	0.29	0.86	-0.18	0.53	-0.55	-1.05	0.99
43	K-41	CLC	642996	4084726	0.34	-0.84	0.57	1.40	-0.50	-0.35	-0.52	-0.23	-0.68	-0.04	0.09	-0.48	-0.10	-0.01	0.67	-0.61	-0.02	-1.47	-0.45	0.25
44	K-8	CLC	643029	4085344	0.01	-1.03	-1.00	0.91	-0.82	0.16	-0.77	-0.32	-0.80	-0.67	-0.48	-0.93	-1.35	-0.01	0.80	0.19	0.56	-0.78	-0.88	1.72
45	K-11	CLC	644188	4084101	-0.06	-0.68	0.23	0.68	-0.65	-0.15	-0.51	-0.47	-0.69	-0.38	-0.26	-0.83	-0.91	0.61	0.48	-0.48	0.25	-1.43	-0.80	0.95
46	K-34	CLC	644873	4083484	-0.35	-0.79	0.10	0.47	-0.32	0.58	-0.51	-0.37	-0.64	-0.35	-0.22	-0.83	-0.28	-0.15	0.29	-0.75	0.87	-0.43	-0.73	-0.09
47	K-45	CLC	645775	4083244	-2.42	-0.22	-2.04	0.27	-0.02	4.40	-0.61	1.08	0.05	-0.84	-0.66	-0.21	-0.31	-1.65	-0.15	-1.32	2.59	0.36	-0.52	-1.01
48	K-64	CLC	646220	4082702	-1.19	0.02	1.17	0.11	0.12	1.90	-0.35	0.09	-0.26	-0.42	-0.20	-0.09	0.55	-0.18	-0.31	-0.66	1.18	0.11	-0.06	-1.72
49	K-33	CLC	645333	4084028	-1.55	-0.66	-1.30	-0.02	-0.18	2.96	-0.64	-0.08	-0.34	-0.78	-0.48	-0.93	-0.28	-1.32	1.49	-1.13	3.40	0.50	-0.35	-0.64
50	K-12	CLC	644828	4084650	-1.15	0.07	0.43	-0.09	0.23	0.85	0.02	1.08	0.31	0.24	0.41	0.15	0.28	-1.00	-0.32	-0.14	0.40	-0.62	0.06	-0.62
51	K-32	CLC	645786	4084749	-0.86	-0.52	-0.84	-0.37	-0.17	1.87	-0.49	-0.33	-0.34	-0.54	-0.30	-0.86	-0.99	-1.24	1.54	-0.34	2.89	-0.76	0.82	-0.05
52	K-13	CLC	645320	4085261	-3.77	-1.33	-0.34	-2.13	-1.74	-1.38	-1.26	-1.17	-0.98	-1.67	-1.48	-0.77	-1.93	-0.10	-0.90	-0.06	-0.68	1.63	0.23	1.24
53	K-5	CLC	644934	4085676	-1.01	-0.90	-0.37	-1.33	-1.04	-0.83	-0.90	-0.88	-0.76	-1.00	-0.96	-0.03	-1.30	0.28	-1.03	-0.13	-0.13	0.11	-1.05	0.18
54	K-14	CLC	645876	4085956	-3.40	-1.25	-1.04	-1.86	-1.61	-1.35	-1.22	-1.07	-0.90	-1.52	-1.35	-0.41	-1.93	0.30	-1.07	0.34	-1.16	1.84	-0.85	2.41
55	K-31	CLC	646267	4085354	-0.12	1.15	-0.64	-0.96	0.27	0.67	0.44	1.25	1.91	0.46	0.60	1.51	0.39	-0.91	0.35	-0.61	0.11	0.95	0.03	-0.89
56	K-46	CLC	647507	4085404	-1.02	0.21	-1.44	-0.69	0.17	0.68	0.03	0.16	1.18	0.18	0.17	0.91	2.17	-1.13	0.13	-0.53	0.74	0.70	1.22	-2.54
57	K-30	CLC	646970	4086044	1.11	-0.28	-0.50	-0.86	-0.49	-0.69	-0.33	-0.35	-0.34	-0.39	-0.39	-0.53	-0.09	1.38	-1.12	-0.25	-0.85	-0.74	-1.63	0.22
58	K-15	CLC	646450	4086636	0.75	-0.25	0.43	-0.27	-0.15	-0.69	-0.20	-0.46	-0.35	-0.15	-0.32	0.07	1.34	0.72	-0.58	0.12	-0.45	0.20	-0.70	-0.06
59	K-3	CLC	646262	4087710	0.76	0.07	0.03	-0.48	-0.08	-0.03	-0.16	-0.31	-0.49	-0.07	-0.09	0.72	0.20	0.32	-1.17	-0.66	0.05	-0.02	0.77	-0.17
60	K-16	CLC	646992	4087300	0.54	-0.87	-0.90	-0.98	-1.11	-1.02	-0.82	-0.54	-0.63	-0.98	-0.73	-0.53	-0.78	1.76	-0.64	0.14	-0.95	0.45	-1.63	2.35
61	K-29	CLC	647464	4086657	0.54	0.38	-0.10	-0.80	-0.22	-0.29	-0.09	0.21	0.45	0.26	0.26	0.00	-0.62	-0.42	-0.34	0.49	-0.42	1.57	1.17	0.60
62	K-47	CLC	648001	4086000	-1.43	0.53	0.10	-0.95	0.18	1.63	-0.02	1.06	0.62	-0.06	0.26	0.50	-0.17	-1.59	0.59	-0.24	1.50	1.54	1.42	-1.19
63	K-63	CLC	648601	4085443	-1.75	0.33	-1.34	-1.06	-0.18	3.00	-0.32	1.75	0.43	-0.39	-0.22	-0.12	0.70	-1.54	-0.59	-0.83	2.15	1.61	1.58	-1.92
64	K-62	CLC	649048	4085973	-1.64	0.91	-0.64	-1.08	0.20	0.86	0.34	1.85	1.54	0.54	0.79	0.47	-0.37	-2.00	-0.42	-0.39	0.68	0.34	1.16	-1.07
65	K-48	CLC	648447	4086580	0.47	-0.36	-0.40	-1.12	-0.45	-0.58	-0.10	0.03	0.55	0.12	0.35	-0.14	-0.26	0.16	1.47	-0.23	-0.52	2.51	0.91	-0.46
66	K-67	CLC	649761	4086044	0.05	2.53	0.17	-1.20	0.56	-0.02	1.33	2.23	1.72	1.93	2.77	1.23	1.08	-0.70	-1.10	-0.11	-0.27	1.19	0.56	-0.94
67	K-61	CLC	649544	4086587	0.85	-0.11	-0.30	-0.67	-0.10	-0.44	0.13	0.13	3.57	0.64	0.62	-0.03	-1.93	-1.31	1.61	0.70	-0.51	1.72	0.97	0.08
68	K-49	CLC	648992	4087228	0.38	1.26	-0.24	-1.03	0.23	-0.91	1.50	3.38	3.45	3.00	4.26	0.78	1.54	-1.11	-1.70	0.67	-1.30	-0.31	0.32	-1.33

69	K-28	CLC	648103	4087422	1.38	-0.33	0.43	-0.66	-0.86	-0.90	-0.42	-0.26	-0.34	-0.22	0.11	-0.97	-1.53	-0.06	-1.24	0.86	-1.11	2.23	0.89	0.98
70	K-17	CLC	647583	4088032	0.57	1.24	0.23	-0.88	0.29	-0.88	1.63	3.77	0.17	2.77	2.99	1.99	0.81	-0.87	-1.78	0.51	-1.26	-0.31	1.30	-0.89

Supplementary Table 1 contuniued.

	ID	Type	X	Y	kcps	Si	S	Cl	K	Ca	Ti	Cr	Mn	Fe	Ni	Zn	As	Se	Br	Rb	Sr	Y	Zr	Pb
71	K-18	CLC	648079	4088632	1.18	1.21	0.83	-0.71	0.37	-0.94	1.57	2.73	0.28	2.82	2.79	1.83	0.42	-0.61	-1.51	0.85	-1.27	1.07	1.68	-1.26
72	K-27	CLC	648616	4088021	0.49	1.48	-0.57	-1.07	0.11	-0.88	1.39	4.16	1.62	2.98	4.15	1.67	0.06	-1.50	-1.85	-0.02	-1.29	-1.25	1.31	-1.55
73	K-50	CLC	649472	4087826	0.44	3.53	-0.37	-1.34	0.72	-0.72	2.07	3.16	1.49	2.99	4.05	1.92	1.49	-0.68	-1.96	0.87	-1.26	1.16	1.32	-1.64
74	K-68	CLC	650723	4087255	-0.54	0.01	-1.94	-0.90	-0.09	0.29	0.34	1.53	3.61	0.56	0.61	0.89	0.53	-0.44	0.13	-0.73	-0.59	-1.07	-0.35	-0.29
75	K-59	CLC	650504	4087831	1.28	-0.44	-0.84	-0.78	-0.88	-0.89	-0.52	-0.60	-0.28	-0.59	-0.45	-0.65	-0.18	1.92	-0.27	0.25	-0.99	0.72	-0.66	0.89
76	K-51	CLC	650001	4088492	1.07	-0.75	-1.14	-0.61	-0.89	-0.91	-0.71	-0.70	-0.70	-0.91	-0.78	-0.41	0.72	0.93	-1.04	0.00	-1.03	0.42	-0.59	1.05
77	K-26	CLC	649064	4088587	0.84	1.18	2.00	-1.05	0.50	0.01	0.86	1.22	0.85	1.29	1.37	0.56	0.74	0.18	-0.78	0.94	0.05	0.06	2.29	-1.08
78	K-19	CLC	648575	4089249	-0.69	-0.22	0.57	-0.49	0.04	2.14	-0.12	0.72	0.48	0.00	0.40	-0.01	0.07	-2.15	-0.35	-0.39	2.12	0.00	1.24	-1.23
79	K-2	CLC	648150	4089979	0.49	-0.44	-0.37	-0.44	-0.47	-0.95	-0.17	-0.40	-0.25	-0.37	-0.36	0.41	1.40	0.58	-0.60	-0.53	-0.89	-1.47	-0.19	-0.66
80	K-20	CLC	649088	4089900	0.70	-0.76	-0.57	-0.58	-0.44	-0.39	-0.37	-0.68	-0.59	-0.55	-0.65	0.01	0.83	0.38	-0.80	-0.42	-0.37	0.34	0.02	-0.72
81	K-25	CLC	649609	4089254	0.95	-0.71	-0.84	-0.48	-0.70	-0.67	-0.66	-0.72	-0.73	-0.92	-0.81	-0.74	0.83	1.99	-0.36	-0.14	-0.90	0.18	-1.63	0.43
82	K-52	CLC	650482	4089108	1.15	-0.38	-0.90	-0.39	-0.67	-0.91	-0.66	-0.78	-0.61	-0.87	-0.70	-0.91	-0.64	2.03	-0.38	-0.69	-1.11	-0.76	-1.44	0.91
83	K-58	CLC	651018	4088464	0.71	-0.49	-0.90	-0.34	-0.63	-0.85	-0.52	-0.59	-0.42	-0.66	-0.53	0.09	1.12	1.82	-0.43	-0.32	-0.91	-0.96	-0.86	0.27
84	K-69	CLC	651220	4087886	0.53	-0.33	-0.03	-0.63	-0.17	-0.60	0.21	0.03	2.38	0.56	0.36	0.90	1.05	1.14	0.42	0.38	-0.81	-0.41	-0.44	0.31
85	K-57	CLC	651356	4089125	0.85	-0.75	-0.77	-0.59	-0.46	-0.64	-0.27	-0.50	-0.41	-0.25	-0.21	-0.55	0.48	0.09	-0.62	-0.25	-0.57	-0.87	-1.08	0.43
86	K-24	CLC	650173	4089890	0.23	0.67	0.97	-0.71	1.56	-1.16	2.51	0.29	0.16	1.85	0.45	1.95	1.78	0.89	-1.21	2.00	-1.30	-0.37	1.21	-0.53
87	K-53	CLC	650932	4089676	0.96	-0.76	-0.44	-0.39	-0.69	-0.90	-0.52	-0.60	-0.57	-0.67	-0.33	-1.06	0.07	1.15	-0.63	0.14	-1.11	-0.52	-0.89	0.77
88	K-56	CLC	651919	4089598	0.38	0.21	1.03	-0.91	0.58	-0.08	0.31	-0.28	-0.03	0.21	0.06	0.43	0.72	0.10	-0.53	0.18	0.03	1.19	0.15	-0.53
89	K-70	CLC	652739	4089574	-0.15	-0.52	-0.74	-1.18	0.30	0.23	-0.01	-0.52	0.08	-0.26	-0.49	-0.24	1.25	-0.42	-0.71	-0.21	0.51	-0.67	-0.62	-0.17
90	K-74	CLC	653614	4089248	0.16	-0.47	-0.03	-1.01	0.08	-0.05	-0.30	-0.44	0.41	-0.42	-0.45	-0.35	-0.77	0.13	-0.52	0.31	-0.12	-0.41	-0.74	0.73
91	K-73	CLC	653410	4089703	0.34	-0.36	-0.14	-0.97	-0.07	0.60	-0.42	-0.57	0.23	-0.46	-0.53	-0.17	-0.40	0.40	-0.57	-0.10	0.69	-1.47	-0.10	-0.48
92	K-55	CLC	652306	4090116	0.53	0.59	0.93	-0.98	0.15	-0.64	-0.13	-0.45	0.74	-0.12	-0.25	-0.10	1.25	1.04	-0.71	0.25	-0.66	-0.67	-0.94	0.40
93	K-54	CLC	651381	4090220	0.63	0.02	0.30	-0.90	0.49	-0.53	0.00	-0.47	-0.21	-0.20	-0.32	-0.86	1.10	-0.07	-0.90	0.05	-0.87	-0.41	-1.15	0.52
94	K-23	CLC	650589	4090434	0.72	-0.40	-1.17	-0.40	0.05	-0.49	-0.05	-0.49	-0.54	-0.16	-0.23	-0.35	-0.56	1.66	-0.64	0.38	-1.03	0.55	-1.35	1.17
95	K-21	CLC	649632	4090562	-0.44	2.22	-0.84	-1.11	2.73	-1.06	4.09	1.02	0.70	2.84	0.71	3.08	1.16	-0.69	-1.71	1.66	-1.27	0.32	2.05	-0.96
96	K-1	CLC	648955	4090607	0.70	-0.17	-0.37	-0.51	-0.37	-1.02	0.01	-0.56	-0.51	-0.41	-0.73	0.52	-0.29	1.80	-0.74	-0.51	-1.16	-1.33	-0.79	1.29
97	K-22	CLC	648280	4090888	0.35	0.16	0.10	-0.90	0.24	-1.03	0.38	-0.34	0.49	0.18	-0.31	1.14	-1.29	0.93	-0.72	0.45	-1.17	-0.19	-0.28	1.05
98	K-75	CLC	649456	4091218	-0.28	5.16	0.60	-1.15	3.85	-0.91	5.62	1.35	1.11	3.39	1.25	3.35	0.54	-0.99	-1.95	2.60	-1.25	0.93	3.24	-1.82
99	K-76	CLC	650035	4091048	-0.33	1.72	0.43	-0.15	2.11	-0.82	1.81	0.28	0.40	1.47	0.32	1.46	-0.53	-0.73	-0.95	1.42	-1.12	0.71	1.60	-0.52
100	K-77	CLC	650813	4090888	-0.49	-0.52	0.10	0.87	0.85	1.14	0.30	-0.53	-0.18	-0.03	-0.38	0.72	-1.40	-1.93	-0.56	0.62	0.28	0.01	0.29	-0.80
101	K-78	CLC	651312	4090737	0.22	3.21	-0.70	-1.51	6.84	-1.12	3.17	-0.44	4.08	2.04	-1.14	4.96	1.82	-0.76	-1.57	7.86	-1.29	1.39	3.11	-0.46
102	K-79	CLC	651919	4090534	-0.22	-0.55	-1.10	-1.20	0.60	0.28	-0.36	-0.74	0.14	-0.77	-0.95	0.23	2.26	0.13	-1.32	-0.79	0.62	2.53	1.38	-1.08
103	K-80	CLC	652488	4090455	0.89	-0.41	-1.04	-1.43	0.07	-0.04	-0.38	-0.78	0.61	-0.51	-0.71	-0.56	-0.20	1.00	-0.29	0.22	-0.08	-1.35	-0.71	0.31

<b>104</b>	<b>K-71</b>	<b>CLC</b>	653093	4090108	0.86	-0.28	-0.03	-1.06	0.11	-0.18	-0.37	-0.63	0.44	-0.39	-0.48	-0.64	-0.45	0.26	-1.15	0.38	0.31	-1.12	-1.01	0.65
<b>105</b>	<b>K-72</b>	<b>CLC</b>	653719	4089989	0.66	-0.20	0.17	-1.30	-0.07	-0.29	-0.22	-0.50	0.57	-0.30	-0.46	-0.29	-0.81	0.19	-0.58	0.69	-0.26	-0.52	-0.14	1.09

Supplementary Table 2. Standardized ITRAX analysis results for Dalyan channel cores.

	<b>ID</b>	<b>Type</b>	<b>X</b>	<b>Y</b>	<b>kcps</b>	<b>Si</b>	<b>S</b>	<b>Cl</b>	<b>K</b>	<b>Ca</b>	<b>Ti</b>	<b>Cr</b>	<b>Mn</b>	<b>Fe</b>	<b>Ni</b>	<b>Zn</b>	<b>As</b>	<b>Se</b>	<b>Br</b>	<b>Rb</b>	<b>Sr</b>	<b>Y</b>	<b>Zr</b>	<b>Pb</b>
<b>1</b>	<b>D-L-1</b>	CLC	645642	4080700	0.43	-1.82	-0.38	-0.70	-2.69	3.43	-3.14	-1.12	-1.41	-3.20	-2.36	-2.60	-1.93	1.19	2.18	-1.68	3.61	0.77	-1.57	1.14
<b>2</b>	<b>D-L-2</b>	CLC	645589	4079949	1.02	-1.62	-0.33	-0.66	-1.17	-0.52	-1.32	-0.95	-1.85	-1.45	-1.25	-1.24	-1.95	0.50	1.57	-1.25	0.65	1.80	0.88	2.05
<b>3</b>	<b>D-L-3</b>	CLC	645622	4079284	-0.76	0.90	-0.13	-0.73	1.35	0.03	0.70	-0.65	1.43	0.50	0.00	-0.43	-0.56	-0.46	-0.93	1.38	-0.53	-0.84	0.51	-0.37
<b>4</b>	<b>D-006-130</b>	CLC	645505	4079077	0.61	1.93	0.07	-0.08	1.49	0.00	0.89	-0.38	1.42	1.15	0.74	-0.12	-1.19	0.58	-1.31	1.29	-0.62	0.58	1.28	0.11
<b>5</b>	<b>D-006-133</b>	CLC	645460	4079062	1.22	1.53	-0.01	-0.18	1.90	-1.11	1.64	-1.04	0.22	1.21	-0.13	1.99	-0.11	0.54	1.05	1.79	-0.59	1.02	1.81	-0.22
<b>6</b>	<b>D-006-000</b>	HLC	645551	4079019	0.12	-1.18	-0.12	0.65	-1.29	-0.07	-0.23	0.22	0.40	0.00	0.60	-1.39	0.24	-0.31	0.32	0.01	-0.22	-0.34	-0.30	-1.29
<b>7</b>	<b>D-006-134</b>	CLC	645496	4079018	1.57	-0.15	1.61	0.59	1.00	-1.37	0.52	-0.98	1.62	0.49	-0.73	-0.48	1.24	0.74	1.45	0.34	-0.64	-0.98	-0.09	-1.16
<b>8</b>	<b>D-006-132</b>	CLC	645542	4078991	-1.07	-0.76	-0.88	2.08	0.78	0.93	-0.39	-0.75	-0.67	-1.08	-1.44	-0.06	0.73	-0.90	-0.04	-0.98	1.70	1.93	-1.21	-1.45
<b>9</b>	<b>D-006-138</b>	CLC	645578	4078952	-0.92	-1.53	0.05	1.41	-0.42	1.61	-1.70	-0.94	-0.60	-2.23	-2.02	0.06	-0.88	1.09	0.40	-1.28	2.41	-1.25	-1.07	-0.48
<b>10</b>	<b>D-006-137</b>	CLC	645522	4078943	0.81	1.13	0.27	1.20	1.52	0.13	0.84	-0.47	-0.08	0.11	-0.69	0.04	-0.22	0.37	0.00	1.44	0.06	-0.51	1.47	1.43
<b>11</b>	<b>D-005-139</b>	CLC	645650	4078782	0.30	0.34	-1.85	-1.81	-0.42	-0.27	0.20	-0.28	0.99	0.09	0.54	-0.45	-1.10	0.58	-0.22	-0.18	-0.42	0.76	0.49	0.90
<b>12</b>	<b>D-005-140</b>	CLC	645613	4078736	-0.54	0.88	0.83	-0.92	0.04	-0.27	0.67	0.28	-0.91	0.44	0.79	0.32	0.27	0.22	-1.11	-1.08	-0.30	-0.58	-0.72	0.68
<b>13</b>	<b>D-005-141</b>	CLC	645561	4078729	-0.48	0.04	-0.51	-0.01	-0.32	0.79	-0.42	1.15	-1.37	-0.22	-0.08	0.51	-1.11	-1.00	-0.49	0.82	0.08	1.54	0.87	1.56
<b>14</b>	<b>D-005-000</b>	HLC	645629	4078718	-1.28	0.73	-1.43	-0.70	-0.34	0.67	-0.45	0.62	-0.05	0.35	0.97	-0.70	1.87	-1.76	-1.00	-1.28	-0.32	-0.95	-0.88	-2.67
<b>15</b>	<b>D-005-142</b>	CLC	645594	4078690	-1.26	0.08	-0.01	0.27	-0.50	0.70	-0.66	1.59	-0.34	0.22	1.04	-0.56	1.45	-1.87	-1.41	-0.41	-0.32	-0.18	-0.98	0.00
<b>16</b>	<b>D-005-144</b>	CLC	645635	4078680	0.59	-1.22	-0.23	-0.68	-1.28	0.04	-0.85	0.06	0.66	-0.41	0.46	-0.59	0.45	0.33	-0.02	0.01	-0.20	-1.52	1.21	-0.62
<b>17</b>	<b>D-005-143</b>	CLC	645591	4078648	-0.97	-0.04	0.22	-0.44	-0.64	1.02	-0.66	1.24	-0.35	-0.06	0.83	-0.48	1.44	-0.78	-0.67	-1.62	0.26	-0.09	0.50	-0.58
<b>18</b>	<b>D-L-4</b>	CLC	645736	4078344	0.17	-0.36	2.18	0.05	-0.25	-0.91	0.46	-0.40	-1.30	0.06	-0.11	0.37	1.22	-0.30	-0.19	1.20	-0.38	0.55	-1.08	-0.40
<b>19</b>	<b>D-016-146</b>	CLC	645901	4078252	0.07	0.33	0.27	-0.70	0.20	-0.15	0.80	-0.31	0.61	0.32	0.19	1.36	0.99	0.58	-0.10	0.63	-0.25	-1.11	-0.21	-0.16
<b>20</b>	<b>D-016-145</b>	CLC	645825	4078250	-0.21	-1.36	-1.13	-0.55	-0.78	-0.26	-0.52	-0.43	-0.23	-0.51	-0.50	-0.51	0.18	1.50	0.06	-0.43	0.20	0.74	0.60	0.63
<b>21</b>	<b>D-016-147</b>	CLC	645882	4078189	0.21	1.13	2.18	-0.28	0.72	-0.65	1.04	-0.09	-1.22	0.61	0.52	-0.32	-0.38	0.04	-0.64	0.80	-0.51	0.88	0.80	0.49
<b>22</b>	<b>D-016-000</b>	HLC	645955	4078161	-0.34	0.39	-0.33	0.14	0.22	-0.03	0.47	0.14	0.86	0.42	0.50	0.35	-0.24	0.41	-0.65	0.03	-0.37	0.88	0.22	-0.67
<b>23</b>	<b>D-016-658</b>	CLC	645981	4078103	1.08	0.10	0.37	-0.62	0.50	-0.24	0.50	-0.12	0.81	0.41	0.04	0.69	-0.15	0.67	0.17	0.55	-0.14	0.36	-0.25	0.22
<b>24</b>	<b>D-016-148</b>	CLC	645926	4078102	0.06	-0.36	1.12	-0.37	-0.10	-0.79	-0.26	-0.45	-0.67	0.02	-0.23	0.37	0.18	-1.07	0.21	0.06	-0.32	-1.34	0.26	-0.32
<b>25</b>	<b>D-016-149</b>	CLC	645942	4078052	1.67	-0.99	-0.53	-0.17	-0.44	-0.84	-0.49	-0.97	1.37	-0.41	-1.01	-0.35	0.02	1.10	-0.02	0.37	-0.39	0.05	0.55	0.41
<b>26</b>	<b>D-016-150</b>	CLC	646012	4078027	0.68	0.21	0.37	-0.39	-0.11	-0.66	0.08	-0.10	-0.28	0.23	-0.04	0.14	0.24	-0.09	-0.17	-0.31	-0.28	-0.06	-0.63	0.10
<b>27</b>	<b>D-L-6</b>	CLC	646067	4077417	-2.88	0.67	-1.78	-0.55	0.22	0.73	0.95	3.33	-0.40	1.48	2.34	0.30	0.53	-1.26	-1.63	-0.95	-1.05	-0.75	-2.32	-0.41
<b>28</b>	<b>D-L-10</b>	CLC	645015	4076334	-0.21	0.44	0.67	2.80	0.70	-1.09	0.90	0.60	1.39	0.92	0.61	1.92	-0.13	1.33	1.78	-0.08	-0.62	-1.35	-0.45	0.44
<b>29</b>	<b>D-L-8</b>	CLC	645930	4076095	0.32	0.56	-0.63	1.38	0.10	-0.84	0.44	1.21	-0.05	0.54	0.43	1.85	-1.12	-2.01	1.40	0.80	-0.51	-0.02	0.32	0.63

Supplementary Table 3. Standardized ITRAX analysis results for Fethiye-Göcek bay cores.

	ID	Type	X	Y	kcps	Si	S	Cl	K	Ca	Ti	Cr	Mn	Fe	Ni	Zn	As	Se	Br	Rb	Sr	Y	Zr	Pb
1	I-100	HLC	675787	4066071	-1.40	-2.01	1.63	3.45	-0.98	3.10	-2.17	-0.58	-2.57	-2.56	-1.59	-2.61	-0.52	-0.69	-0.91	-2.58	3.65	-0.84	-1.13	-1.60
2	I-103	HLC	675803	4066045	-1.90	-2.26	3.70	2.47	-1.03	3.33	-2.09	-1.21	-2.39	-2.65	-1.72	-1.91	-1.17	-0.74	-0.58	-2.63	3.28	-1.97	-0.52	-1.25
3	I-123	HLC	675819	4066069	-2.03	-3.08	3.21	1.08	-1.51	5.50	-3.16	-1.15	-2.79	-3.29	-1.85	-3.51	-0.46	-2.96	-1.98	-3.06	4.11	0.65	-1.72	-1.67
4	I-106	HLC	675722	4066166	-2.38	-3.33	1.39	3.03	-1.78	3.43	-3.24	-0.88	-2.77	-2.45	-1.35	-2.80	-0.10	-0.06	-1.68	-2.68	1.86	-0.70	-1.56	-0.44
5	I-122	HLC	675744	4066234	-0.53	-0.65	1.63	0.02	-2.17	-0.21	-3.15	0.35	-1.08	1.10	2.95	-3.05	3.20	-3.41	-2.65	-1.68	-0.07	-1.76	-2.72	-2.10
6	I-107	HLC	675770	4066297	0.46	-1.84	1.75	1.36	-2.43	0.94	-3.36	0.11	-0.47	0.03	1.22	-2.92	1.73	0.19	-1.67	-2.32	0.70	-1.19	-2.16	-1.50
7	I-111	HLC	675811	4066096	-1.10	-0.12	0.54	0.07	-0.32	0.16	-0.26	-0.31	-0.76	-0.51	-0.16	-0.24	-0.21	-1.19	0.35	-0.58	1.83	-0.52	-0.36	-0.44
8	I-126	HLC	675848	4066058	-0.11	-0.59	0.66	0.36	-0.16	0.27	-0.17	-0.38	-0.73	-0.60	-0.34	0.70	0.30	-1.19	1.52	-0.84	2.99	-1.10	-0.14	-1.33
9	I-114	HLC	675878	4066055	-0.42	0.87	-0.31	-0.86	-0.81	0.44	-0.55	1.01	-0.22	0.07	0.47	-0.53	0.05	-0.74	-0.79	-1.36	1.89	-0.73	-0.61	-0.29
10	I-125	HLC	675868	4066089	-0.08	0.56	0.90	-0.99	-1.01	0.13	-0.37	1.19	0.16	0.30	0.67	-0.54	-1.23	-1.17	-1.02	-1.15	1.66	-0.71	-1.58	-0.23
11	I-112	HLC	675863	4066127	-0.24	0.56	0.17	-1.05	-1.10	0.36	-0.82	1.45	0.43	0.54	0.89	-0.78	0.18	-1.19	-1.27	-0.88	0.83	-1.10	-1.21	-0.54
12	I-124	HLC	675826	4066143	-0.26	0.05	-0.19	-0.30	-1.04	0.13	-0.47	2.07	0.45	0.46	0.82	-0.56	0.84	-0.58	-1.30	-1.27	1.25	-0.64	-1.38	-1.19
13	I-115	CLC	675794	4066195	0.34	1.04	0.36	-1.09	-1.25	-0.14	-0.81	1.57	0.43	1.16	1.40	-0.92	3.62	-0.40	-1.72	-1.40	0.04	-2.41	-1.98	-2.60
14	I-120	CLC	675849	4066189	0.47	1.57	-0.07	-1.46	-1.23	-0.40	-0.48	1.74	0.84	1.20	1.37	-0.09	0.32	-0.33	-1.06	-0.53	-0.17	-0.87	-0.93	-0.90
15	I-109	CLC	675925	4066101	0.46	-0.18	-0.25	0.09	-1.26	-0.64	-0.54	1.43	0.76	1.18	1.28	-0.59	0.25	-0.47	-0.74	-0.51	-0.41	-2.18	-0.93	0.27
16	I-127	CLC	675923	4066051	0.37	0.17	0.05	-1.08	-0.99	-0.54	-0.37	1.36	0.55	1.17	1.33	-0.62	0.59	-0.38	-0.75	-0.19	-0.18	-0.94	-1.33	-0.48
17	I-135	CLC	675938	4066008	1.25	-0.86	-0.56	-0.35	-1.12	-0.41	-0.15	1.77	1.23	1.65	1.68	0.01	1.31	-1.78	-1.14	0.09	-0.25	-0.73	-1.50	0.22
18	I-110	CLC	675968	4066030	0.61	0.05	-0.43	-0.64	-1.18	-0.66	-0.44	1.30	0.59	1.20	1.28	0.01	-0.48	-0.47	-0.92	-0.56	-0.50	-1.51	-0.15	0.22
19	I-105	CLC	675952	4065951	0.63	0.01	1.02	-0.01	-0.92	-0.60	-0.45	1.69	0.36	1.00	1.21	0.17	0.55	0.94	-0.88	-0.40	-0.25	0.19	-0.88	-0.29
20	I-121	CLC	675987	4066125	0.84	0.42	0.66	-1.20	-0.98	-0.65	0.03	1.32	0.78	1.46	1.51	-0.34	0.66	-0.76	-1.31	0.01	-0.53	-1.33	-1.20	-0.79
21	I-113	CLC	675938	4066182	0.33	0.58	-0.07	-1.53	-1.24	-0.69	-0.31	0.95	0.60	0.99	1.09	-0.21	0.11	-1.37	-0.55	-0.70	-0.53	-0.63	-0.92	-1.06
22	I-116	CLC	675869	4066262	0.57	1.57	-0.74	-1.71	-0.80	-0.60	0.59	1.28	0.73	1.28	1.15	0.25	-0.17	-1.67	-1.09	-0.35	-0.46	0.01	-0.74	-1.04
23	I-133	CLC	675774	4066397	0.73	0.19	0.90	0.16	-1.30	0.05	-0.79	1.13	0.33	1.10	1.35	-1.17	0.81	0.21	-1.26	-1.15	0.53	-2.01	-1.07	-1.06
24	I-132	CLC	675906	4066352	1.05	0.38	0.42	-0.41	-0.95	-0.52	0.59	1.76	1.36	1.60	1.34	0.26	-0.64	0.60	-1.55	-0.16	-0.29	-0.31	-1.04	-0.44
25	I-131	CLC	675971	4066261	1.11	-0.10	-1.35	-1.13	-1.16	-0.71	0.47	1.19	0.87	1.63	1.53	-0.01	0.94	-1.39	-0.96	-0.40	-0.54	-1.46	-1.05	-0.40
26	I-130	CLC	676045	4066186	1.20	-0.16	-0.31	-0.96	-1.28	-0.85	-0.21	1.09	0.62	1.48	1.66	-0.70	0.49	0.05	-0.84	-0.26	-0.63	-0.59	-1.36	-1.83
27	I-129	CLC	676008	4066072	0.55	0.13	0.66	-1.00	-0.91	-0.77	0.02	1.35	0.46	1.43	1.44	-0.27	-0.06	-0.10	-0.80	-0.18	-0.58	-0.48	-1.22	-0.01
28	I-128	CLC	676052	4065986	0.30	0.01	0.54	-0.95	-1.18	-0.78	-0.51	1.32	0.67	1.31	1.52	-0.43	0.40	0.39	-0.88	-0.16	-0.59	-1.40	-1.32	-0.44
29	F-70	CLC	672579	4065446	-0.37	-1.20	-0.61	1.14	0.51	0.61	-0.31	-0.54	-0.74	-0.84	-0.66	0.03	-0.74	0.31	0.46	-0.08	0.45	-0.87	-0.07	0.49
30	F-63	CLC	673395	4068021	-0.98	0.47	-1.70	-0.41	-0.91	-0.10	-0.65	1.57	0.66	0.33	0.37	0.51	0.88	-1.34	-0.93	-0.76	-0.10	-0.77	-1.05	-0.67
31	F-64	CLC	672902	4068383	-0.17	-0.28	-0.80	-0.02	-0.97	-0.03	-0.11	2.68	0.33	0.67	-0.18	-0.28	2.67	-0.28	-1.23	-0.92	-0.38	-2.04	-0.98	-2.14
32	F-65	CLC	672794	4067762	-1.35	-0.62	-0.21	0.88	-0.44	1.40	-0.61	0.50	-0.67	-0.55	-0.80	-0.88	0.98	1.49	-0.79	-1.25	0.64	0.05	-0.92	-1.18
33	F-66	CLC	672606	4067052	-0.83	-0.57	-0.31	0.62	0.46	0.20	0.13	-0.13	-0.65	-0.63	-0.78	0.72	-0.36	0.23	0.19	0.27	0.03	0.30	0.86	-0.12

34	F-67	CLC	672154	4066275	-0.11	-0.11	-1.10	0.69	0.95	0.08	0.63	-0.23	-0.96	-0.56	-0.80	0.08	-0.91	1.33	0.68	0.80	-0.07	1.29	0.19	0.40
35	F-69	CLC	671900	4065635	-0.29	-0.03	-1.00	0.29	0.95	0.50	0.36	-0.46	-0.70	-0.73	-0.81	0.84	-0.65	-1.02	0.62	0.41	0.16	1.22	0.28	1.10

Supplementary Table 3 continued.

	ID	Type	X	Y	kcps	Si	S	Cl	K	Ca	Ti	Cr	Mn	Fe	Ni	Zn	As	Se	Br	Rb	Sr	Y	Zr	Pb
36	F-85	CLC	673753	4067631	0.67	0.39	0.78	0.24	-0.78	-0.42	-0.56	0.62	1.01	0.50	1.02	-0.09	0.39	0.04	-0.34	-0.41	-0.35	0.02	-0.16	0.21
37	F-62	CLC	674501	4066666	0.03	0.55	-0.01	-1.03	-0.03	-0.57	0.13	0.11	0.23	0.18	0.54	0.60	-1.13	-1.02	-0.40	0.02	-0.43	-0.46	-0.85	0.34
38	F-84	CLC	674888	4066139	0.72	1.48	0.38	-1.29	-0.55	-0.43	-0.45	-0.29	0.86	0.58	0.96	-0.11	-0.25	-0.24	-0.63	-0.35	-0.33	-0.57	-0.17	0.19
39	F-61	CLC	676075	4065876	0.64	-0.06	1.08	-0.29	-1.42	-0.83	-0.80	1.23	0.51	0.88	1.31	-0.37	0.55	0.64	-0.51	-0.52	-0.60	-0.82	-1.13	-1.08
40	F-60	CLC	674636	4065442	-0.65	-0.65	-1.10	-0.14	-0.13	-0.37	-0.44	1.32	0.01	-0.30	-0.10	1.08	-0.65	0.17	0.44	-0.13	-0.41	0.76	0.44	1.32
41	F-71	CLC	673064	4065097	-1.58	-2.58	2.26	2.15	-0.22	2.56	-1.84	-0.94	-2.40	-2.22	-1.58	-3.00	-0.14	-1.81	2.99	-1.70	3.27	-0.18	-0.84	-0.84
42	F-59	CLC	673793	4064755	0.04	-0.08	-0.92	-0.30	0.47	-0.18	-0.12	0.07	0.03	-0.36	-0.30	1.08	-0.69	0.30	1.09	0.73	-0.17	-0.09	0.09	-0.05
43	F-55	CLC	674162	4063171	-0.28	-0.94	-0.80	0.12	0.81	-0.27	0.16	-0.99	-0.31	-0.70	-0.92	0.86	-0.44	1.80	1.44	0.80	-0.43	0.26	0.72	2.11
44	F-54	CLC	674871	4064254	-1.60	0.05	0.17	1.19	0.05	-0.20	0.20	1.57	1.75	-0.41	0.16	0.26	-1.19	-1.15	-0.53	-0.28	-0.56	-0.52	-0.49	-0.48
45	F-53	CLC	675750	4064933	-0.66	0.13	-1.29	-0.52	0.65	-0.44	0.47	0.02	0.01	-0.12	-0.20	0.89	-2.14	-0.24	0.53	0.90	-0.50	0.08	1.08	0.60
46	F-83	CLC	676498	4065157	1.21	0.64	0.78	-1.42	-0.51	-0.93	0.53	-0.64	1.07	0.84	0.82	-0.21	-0.18	0.74	-0.25	0.02	-0.62	0.99	-0.20	0.57
47	F-52	CLC	676183	4064459	-1.17	-0.36	-0.68	-0.93	-0.02	-0.58	-0.09	-0.10	-0.82	-0.22	-0.02	0.96	-1.03	0.57	0.50	0.36	-0.55	-0.17	-0.34	-0.36
48	F-51	CLC	675922	4063775	-1.05	-0.06	-0.56	-0.63	1.11	-0.19	0.67	-0.87	-0.81	-0.39	-0.63	1.03	-0.36	-0.26	0.45	0.96	-0.45	0.44	1.01	0.66
49	F-50	CLC	674894	4062595	-1.06	-0.67	-1.71	-0.17	1.13	0.08	0.51	-1.10	0.26	-0.55	-0.94	1.27	0.49	0.12	0.74	1.10	-0.28	0.95	1.01	1.43
50	F-49	CLC	674129	4061555	-0.91	-0.73	-0.74	1.29	0.82	0.69	-0.13	-0.93	2.77	-1.02	-1.11	-0.13	0.65	-0.81	-0.08	-0.55	0.80	0.88	-0.17	-0.25
51	F-48	CLC	675028	4060955	-1.25	-0.36	-1.41	0.46	1.33	0.69	-0.03	-1.45	-0.11	-1.10	-1.35	-0.72	-0.31	0.89	0.45	0.26	0.20	2.03	-0.24	0.53
52	F-46	CLC	677029	4063255	-0.11	0.48	-0.74	-0.01	0.85	-0.12	0.56	-0.54	1.41	-0.15	-0.44	0.97	-0.12	0.39	0.91	0.96	-0.42	0.95	1.11	0.66
53	F-82	CLC	677457	4064661	0.84	0.89	0.28	-0.96	0.42	-0.86	0.76	-0.73	1.02	0.59	0.57	0.46	0.07	1.10	-0.27	1.39	-0.64	0.50	1.53	1.66
54	F-45	CLC	677798	4063913	0.12	0.64	-0.25	-0.57	0.77	-0.35	0.88	0.04	0.47	0.26	0.18	1.16	0.14	-0.31	0.45	1.04	-0.55	0.70	0.40	0.06
55	F-44	CLC	676980	4061670	-0.37	-0.82	-1.04	0.67	0.98	0.06	0.14	-1.35	0.72	-0.89	-1.16	0.45	-0.73	0.44	1.27	0.74	-0.30	1.29	1.04	1.57
56	F-43	CLC	676003	4060805	-0.23	-0.45	-0.07	0.56	1.39	0.54	0.15	-1.48	1.85	-1.01	-1.22	1.40	0.61	1.12	1.22	0.61	-0.12	1.18	0.70	0.56
57	F-42	CLC	676705	4060313	-0.01	-1.41	-0.01	1.22	1.31	0.18	0.13	-1.59	0.16	-1.01	-1.37	0.17	0.25	0.30	1.33	0.64	-0.22	-0.17	0.72	1.78
58	F-41	CLC	677700	4061345	0.34	-1.43	-1.35	1.28	0.86	0.15	0.13	-1.34	1.19	-1.07	-1.18	0.44	0.69	0.53	1.31	1.48	-0.09	0.69	0.61	0.31
59	F-40	CLC	677976	4062819	0.00	0.83	-0.31	0.38	0.84	0.16	0.42	-0.37	0.20	-0.28	-0.38	0.39	0.55	1.52	0.19	0.66	-0.28	0.37	0.55	0.14
60	F-39	CLC	678685	4063745	0.33	0.29	-0.19	-0.25	0.95	-0.38	0.55	0.30	0.40	0.29	0.32	1.11	0.17	1.07	0.21	0.96	-0.56	0.17	0.95	0.45
61	F-38	CLC	679251	4064155	0.79	0.25	-0.68	0.09	0.05	-0.65	0.23	0.35	0.40	0.43	0.58	0.36	-0.06	-0.15	1.03	0.72	-0.68	-0.91	0.61	0.99
62	F-81	CLC	680053	4064611	2.15	-0.03	0.18	-0.56	-1.40	-1.52	-0.70	-0.56	0.91	2.13	2.12	-0.07	0.55	-0.36	-0.33	0.34	-0.74	-1.02	-0.13	-0.99
63	F-37	CLC	678793	4062439	0.26	0.87	-0.92	-0.20	1.03	0.05	0.69	-0.23	-0.56	-0.06	-0.06	1.08	-0.92	0.39	0.39	0.82	-0.19	0.97	0.85	0.72
64	F-36	CLC	679773	4063445	0.90	0.52	-0.68	0.14	0.61	-0.44	0.73	-0.04	0.25	0.31	0.22	0.92	0.67	1.09	0.49	1.08	-0.56	0.28	1.10	-0.19
65	F-35	CLC	679600	4062065	1.03	1.04	-0.80	0.11	1.02	-0.03	1.03	-0.25	-0.45	0.05	-0.22	0.55	-0.19	-0.29	0.16	0.71	-0.24	-0.06	0.89	0.22
66	F-34	CLC	678659	4060908	0.28	1.18	-0.56	0.50	1.26	0.87	0.60	-0.79	-0.84	-0.70	-0.75	0.80	-0.81	-0.33	-0.02	0.24	0.42	1.75	1.49	-0.17
67	F-33	CLC	677723	4059825	0.84	0.62	0.90	0.70	1.52	0.68	0.91	-1.01	-0.62	-0.63	-0.88	1.21	-0.73	-1.33	0.46	0.55	0.12	1.75	1.14	2.76
68	F-32	CLC	678667	4059442	-0.96	0.69	-1.22	0.15	1.27	0.94	0.23	-1.15	-1.31	-0.96	-1.16	-0.83	0.60	-0.33	0.33	0.18	0.28	0.14	-0.03	-0.19

69	F-31	CLC	679543	4060525	-1.19	0.75	-0.43	-0.06	1.47	0.40	0.69	-1.05	-1.46	-0.86	-0.96	0.32	-1.17	-1.42	0.16	0.78	0.58	1.78	0.68	-0.83
70	F-30	CLC	680315	4061729	-1.21	1.51	0.05	-1.66	1.19	-0.29	1.46	-0.65	-0.52	-0.18	-0.64	0.64	-0.75	0.30	-0.10	1.61	-0.63	0.01	0.97	1.10

Supplementary Table 3 continued.

	ID	Type	X	Y	kcps	Si	S	Cl	K	Ca	Ti	Cr	Mn	Fe	Ni	Zn	As	Se	Br	Rb	Sr	Y	Zr	Pb
71	F-80	CLC	681072	4063573	0.08	1.01	1.08	-0.34	2.53	-0.48	2.70	-0.95	0.01	-0.27	-0.78	1.41	-1.18	0.82	0.22	2.22	-0.88	1.55	2.52	0.99
72	F-29	CLC	680862	4062964	-0.67	0.29	0.42	-0.86	0.38	-0.63	0.52	-0.74	-1.05	-0.67	-0.76	0.10	-1.55	0.35	0.96	0.71	-0.83	0.65	0.50	-0.21
73	F-28	CLC	680903	4061463	-1.46	0.29	-0.92	-1.04	0.72	-0.44	1.08	-0.55	-0.60	-0.47	-0.80	0.44	-0.37	0.69	0.34	1.13	-0.62	0.61	0.87	0.18
74	F-27	CLC	681579	4062136	-1.28	1.43	-0.74	-1.28	0.08	-0.38	0.41	-0.55	-0.87	-0.60	-0.60	-0.43	-0.71	-1.35	-0.18	0.50	-0.85	-0.55	0.05	1.26
75	F-26	CLC	682323	4061719	-0.99	1.73	0.05	-1.08	0.46	-0.39	1.28	-0.24	-0.61	-0.45	-0.73	0.01	-1.47	1.39	-0.91	0.73	-0.88	0.69	0.78	-0.13
76	F-25	CLC	681554	4061164	-0.40	0.87	-0.68	-1.00	2.00	-0.68	2.29	-0.76	-1.13	0.07	-0.89	1.43	-1.28	1.03	0.07	2.44	-0.76	1.29	1.13	-0.01
77	F-24	CLC	680482	4060097	-1.14	0.34	-1.22	-0.23	0.80	0.46	0.29	-0.82	-1.28	-1.02	-0.78	-0.27	-0.82	-0.62	-0.02	-0.11	0.57	0.31	0.12	-1.62
78	F-23	CLC	679671	4059085	-1.70	-1.06	-0.31	-0.12	0.67	0.17	-0.02	-1.47	0.18	-1.17	-1.29	-0.53	1.08	0.48	0.48	0.10	-0.17	1.78	0.14	-0.52
79	F-22	CLC	680736	4058572	0.24	-1.02	0.42	0.77	0.71	-0.19	0.03	-1.37	0.54	-0.87	-1.00	0.51	0.40	0.53	1.34	1.01	-0.40	1.15	0.19	0.83
80	F-21	CLC	682368	4060784	-0.99	0.87	-1.10	-0.62	1.54	-0.41	1.61	-0.35	-0.39	0.15	-0.53	0.26	0.43	-0.22	-0.24	0.75	-0.43	0.28	0.78	-1.31
81	F-19	CLC	683306	4060356	-0.32	0.42	0.17	-0.64	0.61	-0.43	1.12	-0.50	-0.86	-0.15	-0.66	0.45	-0.56	-0.42	0.29	1.06	-0.75	1.18	0.95	0.18
82	F-18	CLC	681776	4058139	-0.18	-0.28	-0.19	-0.01	0.79	-0.58	0.75	-0.55	1.83	0.10	0.09	0.77	-1.68	1.57	0.72	0.44	-0.52	0.72	1.35	0.91
83	F-17	CLC	682873	4057768	-0.83	0.13	-0.68	-0.57	0.80	-0.04	0.35	-1.09	-0.41	-0.54	-0.46	-0.27	-2.50	0.39	0.31	0.46	-0.41	0.15	0.50	0.68
84	F-16	CLC	684254	4059913	0.12	0.95	0.66	-0.56	0.45	-0.33	0.70	-0.53	-0.85	-0.13	-0.55	0.53	-0.43	-0.17	0.45	0.62	-0.60	0.22	0.67	0.91
85	F-15	CLC	684122	4060834	0.50	1.41	-0.19	0.00	0.29	-0.18	1.01	0.12	-0.04	-0.07	-0.50	0.14	0.42	0.53	-0.17	0.23	-0.73	2.00	1.50	0.41
86	F-14	CLC	685241	4060507	0.78	0.83	0.29	0.02	0.15	-0.11	0.65	-0.06	-0.67	0.07	-0.39	-0.24	0.00	1.98	0.40	0.27	-0.61	0.19	0.87	1.57
87	F-13	CLC	685292	4059431	0.66	0.73	-0.56	-0.26	0.17	-0.19	0.61	-0.23	-0.08	0.11	-0.31	0.96	1.02	1.23	0.29	0.27	-0.48	0.56	0.52	0.51
88	F-12	CLC	684102	4058485	0.92	0.50	2.36	0.28	0.14	-0.59	0.84	0.61	1.66	1.05	0.69	0.07	0.19	0.71	-0.37	0.09	-0.48	-1.12	-0.15	-0.01
89	F-11	CLC	683750	4057461	-0.26	-0.03	-1.16	0.55	0.28	-0.08	0.33	-0.54	0.02	-0.45	-0.49	0.07	0.70	1.71	0.35	-0.38	0.09	-0.06	-0.34	0.18
90	F-10	CLC	684523	4057009	-0.43	-0.34	0.29	0.14	0.47	0.15	0.28	-0.90	0.09	-0.53	-0.42	0.40	-0.68	0.14	1.10	0.35	-0.18	1.36	0.32	0.95
91	F-9	CLC	685095	4058061	0.39	-0.06	-0.37	0.84	0.13	0.21	-0.24	0.31	-0.28	-0.08	-0.05	-0.09	0.37	0.89	0.12	-0.30	0.22	0.15	0.24	0.29
92	F-8	CLC	686399	4058994	1.18	0.83	-0.56	0.00	0.26	-0.13	0.36	-0.12	0.07	0.32	-0.10	0.55	1.94	0.35	0.25	0.29	-0.42	0.37	0.15	-0.36
93	F-7	CLC	686293	4060059	0.77	0.75	0.29	0.31	0.16	-0.14	0.49	-0.08	-0.01	0.27	-0.23	-0.09	1.18	0.26	1.05	-0.19	-0.50	0.37	0.28	0.10
94	F-6	CLC	686644	4058081	1.96	0.17	0.54	1.91	0.26	-0.34	0.30	-0.31	1.37	0.69	0.15	0.45	0.63	0.12	2.13	0.23	-0.33	-0.09	0.25	1.37
95	F-5	CLC	687413	4059172	0.90	0.64	0.29	0.26	-0.52	0.20	-0.41	0.98	-0.26	0.33	-0.28	-0.64	1.28	0.14	-0.68	-1.18	0.09	0.79	0.87	0.10
96	F-4	CLC	688268	4057693	2.73	-0.32	0.29	0.19	-0.39	-1.08	-0.07	0.14	0.62	1.46	1.20	-0.42	-0.49	0.24	1.02	0.37	-0.72	-0.59	0.92	-0.05
97	F-3	CLC	688386	4056817	1.67	-1.02	-0.07	0.72	-0.30	-0.53	-0.24	-0.49	0.26	0.92	0.09	0.04	0.55	1.84	0.42	-0.01	-0.30	0.33	1.89	-0.59
98	F-2	CLC	688363	4056177	2.14	-2.01	0.78	1.85	-0.64	-1.12	-0.52	-0.71	-0.54	0.43	0.30	0.58	-1.26	1.03	2.34	0.24	-0.67	-0.31	0.65	1.26
99	F-1	CLC	687939	4055601	1.97	-1.84	2.24	2.00	-0.55	-1.06	-0.57	-0.76	0.27	0.67	0.40	1.27	-0.78	0.76	2.24	0.37	-0.58	0.19	0.89	1.93