

AND ATMOSPHERIC SCIENCES

ASSOCIATION SYMPOSIA AND WORKSHOPS

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IUGG XXIV G	eneral Assembly	July 13, 2007	Perugia, Italy
Abbreviations			
	International Accessiati		t
	International Associati		ronomy
	International Associati		Dhorny
	International Associati		Sciences
	International Associati		Oceans
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	International Associati		stry of the Earth's Interior
	Climate and Cryospher		histry of the Earth's Interior
Ev-K2-CNR	Everest-K2 CNR Comm		~
GEWEX	Global Energy and Wa	te Exerciment	
HKH-FRIEND	Hindu Kush-Himalavar	hove eaimes from	Institutional Experimental
	and Network Data		
ABO	International Association	o for logic cean	raphy
ACS	International Associati	on of Cryospheric Science	es
CACGP	International Commiss	Den Alenemberie Sher	ristry and Global Pollution
CASVR	International Commiss	on n must e -s ill	e tation Relations
CCE	International Commiss	on Montinanta El s	
CCL	International Commiss	on a Clin	
CCLAS	International Commiss	ion on the Coupled Land	-Atmosphere System
ССР	International Commiss	contiouts and thech	ation
CDM	International Commiss	N on ynami. Met oro	gy 🗛
CGW	International Commiss	ion on Groundwa <u>ter</u>	
CIMOD	International Center fo	or Intrasted Mourain D	eveloptient
СМА	International Commiss	ion on the Middle Atmos	phere
CRS	International Celestial	Peterstein	
CSIH	International Commiss		lrology
CSW	International Commiss		
СТ	International Commiss		
CWQ	International Commiss		
CWRS	International Commiss		
GAC	International Global At	mospheric Chemistry	
GS	International Glaciolog	ical Society	
LP	International Lithosphe		
NQUA	International Union for		
ON	International Ocean No		



















not as well u nderstood or ob served as the ht number of temperature and salinity in situ p of Argo floats gave an opportunity for such Ocean region. We used the temperature an Ocean, spanning 0-25 N and 40-100 E, during 200

internal tides. Since different Argo floats pon up during different times of a day, a criterion has been used to take c are of these di urnal oscilla seasonal basis to study the t emporal and region has the highest temporal variability lue Sea has less variability compared to the Ba the changes in the wind magnitude and the ne changes in MLD in relation to the available altimeter derived



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yer due to the lac k of sufficient an b<u>e comp</u>uted. The deployment the o wise data sparse Indian tions in the north Indian terion has been used to

ve bee n averaged on Indian Ocean. Somali hitude. Similarly, Arabian locations mainly reflects se temporal and seasonal sea surface height anomalies.

nor







with strong modes of climate variation, including Ocean zonal dipole m ode and decadal v ariation. The challenge in designing IndOOS was to ge adequate observations to address outstanding scientific guestions, while containing the cost to a doable

level. The plan calls for an integrated deplo in the open ocean for long periods of time between different types of in strumentation mooring array that is now about 1/3 imple her instruments provide a frame work for large ьca integrates the streams of data for research and applica

ph ysi bons, i

dian Oc ean associated al variat ion, the Indian t

> hat can be maintained Intage of the synergies Ocean is a basin-scale ats, XBT lines and othe r Ocean-state estimation











New interest in infraso und monitoring spurred by the the installation of a n umber of infrasound poject include monitoring volcanic eruptions, extrem of these event s are different in terms of so than destruction) - the emission of long-range hearing, or infrasound. There is strong evid existing warnings systems to improve prep remote monitoring of atmospheric events.

national Mo sper evote her, large oce chanism but t und with frequ supp he and ide Some re ups ar

that

ring System (IMS) has led to ions. These applications appli wells, boildes and tsunamis. Each have a common feature (other es b<u>elow th</u>e perception of human asound can supplement n tists h an additional tool for g infraso und created by

infrasound h as I ed t o t he c reation o f an Int ernational In frasound Consortium. The c onsortium will provide members with data from a larger a interest. Treaty restrictions s till limit acces coupled with data from local networks prov time, data standardization is not as routine working towards that goal.

volcanic activity to do infraso und to mography of the atmosphere. The new emergen ce of interest in eir particular scientific a from those stations age of the earth. At this scientists in the field are

Keywords: infrasound, network, monitoring















Conventional remote sen sing products of studying the a erosol indirect effects (AIE): mismatch can lead to false relationship; 2) a cloud particle size represents a thin layer of clo LWP with large uncertainties. These limitati the A-train sensors by taking full advantage others, MODIS, CLOUDSAT and CALIPSO. T layer water clouds. From the multi-spectral channel

of clou d dr oplet effective radiu s (DER). Using the DER data, MODIS extinction data, we will test a hypothesis that the AIE is more significant to the DER at the cloud bottom than at the top. In combination with CLOU AIE are dictat ed by the stage of cloud d loading and DER is not sufficient to reveal to differentiate clouds into dif ferent catego ies objectives, new remote sensing algorithms vill

d an suffer fr) laç y in of syn ptical death is ops. 3) precir be ove rcom юf syner et strate dei cor DDIS,

on is often indirectly inferred from lessened with the constellation of meas our

Towards achieve these

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hents m ade by, among ly is limited to boundary ieve the vertical profiles AOD, and CALIPSO aeros ol that the DER and the at ion b etween aerosol e data, we shall be able

following limitations for

ime and location and

blumn quantity, whereas retrieved

Keywords: aerosol, cloud, precipitation





In this work aerosol and cl oud interact meteorological data, aircraft data, and satellite data. meteorological sounding data over eastern entry significantly reduced during the last 40 year an the h igh con centrations of aerosols. M ete stability in the troposphere has been incre aerosol layer in the lower troposphere aff atmospheric stability. The enhancem ent in and precipitation in this region. Using a large and

concentration is studied. The results indicate that the value of cloud spect ral relative dispersion varies between 0.2 to 0.8 when the cloud drop converges to wards a narro w range of 0. Because the distribution of t he cloud dro indirect radiative effect of aerosols on the forcing can be reduced by 10-40% under h

sis of observ vs th fina is reduction o tal sou nding uring the las rad psphe tab raft m

stern Chin a are studied using surface cinitation, MODIS data and pitation in this region is le pi cipitation is strongly correlated to indicates that t he atmospheri c ye ars. It is spec ulated that t he h lead t o changes i n epress upward motion ts of cl oud droplet siz e

distributions, the relationship between clou d spec tral relative dispersion and cloud droplet number

sses,

ends

Keywords: aerosol, cloud, china

about 50 cm- 3), and ncentration is higher . in estimating the first corresponding radiative









The global di mming and bri ghtening of s urface show several years. The trend of surface shortware in pyranometer data after the quality control p found also in China. The spatially averaged s until around 1 990, and the n increased slig and aerosols related to the irradiance trend aerosol data published by sci entific papers Project (ISCCP). The sensitivities of shortw 'ave i' vapor, and ae rosols, which we define as

combination of changes in these factors. Therefore contributions of these factors to the radiative forcing must be evaluated quantitatively by consid suggest that cloud optical thickness change amount and water vapor change are quite for cloudy sky conditions if aerosols are ab

n discussed for rece nt -2000 was studied with 19 g and brightening p roperties were ce in China decreased by 15W/m2 tors <u>such as</u> clouds, water vapor, adiativ ansfer calculation with ional llite Cloud Climatology nges in clouds, water

potential radiative forcings (PRF) are depe ndent on the

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Keywords: aeroso

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shortwave irra analyzed the

> variations. The results he effects due to cloud e is also important even










Clouds and their impact on transfer of sole aspect of climate and climate change. This is beca microscopic properties such as phase (i.e., qu precipitation particles. With the advent of cl models, representation of cloud microphysi needed for cloud microphysics that on one but on the other provide information not of their local mixing ratios), but also on particle sizes

and Ea rmal rac ons l inte versus sd d), olving and "s u mes the key lar re com htic e mas cl o 0 / import

the m ost challengin g ation and clouds involve leen i and concentration of cloud and parameterized" general circulation e. Arguably, new approaches are feasi such as bulk schemes), onder e and precipitation (i.e., radiative transfer. Thi s

paper will first present devel opment and v alidation of a two-m oment bulk microphysics scheme for warm clouds, with the emphasis on droplet_nucleatio_ n_drizzle/rain devel opment, and t ransformations

due to entrainment and mixing. The latter properties of a cloud field, such as the an three-variable bulk ice scheme will be discu diffusional as well as riming growth of the the overall goal to apply them in numeri atmospheric aerosols in "superparameterized" general circulation models

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itical impact on mean rt, a ne w two-moment approach to represent re being developed with fying indir ect effects of









The Pacific Dust Experiment (PACDEX) is transport of dust and pollution across the Pa of 2007. PACDEX will utilize the NCAR/NSF information from ground and satellite data. aircraft missions to sample portions of the over a one week period. In addition to char the plume, PACDEX will examine the interact using in situ sampling of ice and water size dist

forecasting and in post-mission studies to explore clouds. This is the first time t hat a dust ev entire transection across the Pacific Ocean transport of d ust and soot a nd their impa America. In this presentation we describe project, (b) An overview of t he major inst um that were accomplished, and (d) An overview of preliminary



follow the atmospheric nt during April and May dust sampling vehicle, with supporting nake s use o f long-range multiple ous travel times from the sourc e d radiative properties of e gas n cold warm cloud regions by leasurements of Clou d

Condensation Nuclei and Ice Nuclei near the clouds . Several numerical models will be used for both

the evolution of the plume and its interaction with

t event and across its sights in to long range acific O cean and North and thei r roles in the ummar y of the missions results and observations from the project.

















By considering cloud properties and meteorelogical properties (CERES) project estimates radiation at top-or atm its three broadband channels. From these fluxes under different cloud types and clear sky, ar types (deep convection, cirrostratus, cirrus, stratocumulus, and cumulus) base d on clou International Satellite Cl oud Climatol ogy P product with cloud and radiation information are u

150W/m^2. High and thick clouds usually ar e ice or ice over water clouds. They reflect more radiation back to the sp ace. Therefore, there are le more ATMO due to lim ited transmission between cloudy and clear conditions within heating and cooling effects for high a nd lo significant effects compared to clear skie studied.

, Cloud tei here A) an MOsphei rac ated in currer bostratus ess an CCP) RES tudv.

urface bstra<u>tus,</u> tical d le Sa

arth Radiant Energy System ing measurements from n (ATMC) fluxes, especially those dy. Clouds are classified into nine altocumulus, stratus, as defined by the te Fo otprint (SSF) data mean for ATMO is about

-115W/m^2. Under high and thick clouds, the A TMO fluxes are muc h higher and reach more than

Inface, but they ret ain differences calculated clearly the atmospheric evel clouds do not show nal differences are also





























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To analyze the effects of atmospheric aero SPRINTARS, has been used, which is couple mixed layer o cean. An incr ease in the microphysics by forming smaller and nume cloud lifetime. On the other hand, the insola effects, resulting in decreases in the surface form clouds. This study analyzes changes especially in the Asian region where economy has n

Institute for Environmental Studies (NIES). The model predicts mass mixing ratios of t and organic carbons), sulfate, soil dust, an dioxide and dimethylsulfide. The aerosol tra sulfur chemistry, wet deposition, dry depos CCSR/NIES/FRCGC AGCM is extended for t by aerosol particles considerin g refractive index in each wavelength, size distribution, and hygroscopic

growth of each aerosol. The cloud droplet number concentration is calculated for the aerosol indirect effect with the aerosol partic le number con properties of each aerosol s pecies, updraft change in the cloud droplet number concern leading to a change in the radiation budget. The increases and the precipitation decreases d ue to e in direct affect of anti-progenic aerosols all ove r the Asian region, especially along the coast, with ocean model is coupled, on the other hand, both the liquid water path and precipitation decrease by a reduction of t he solar radiat ion at the surf ace main ly due to the aer osol direct effect, leading to a decrease in evaporation of water vapor fro

large amount of black carbon aerosols eval However, the liquid water path still increa effect due to abundant anthropoge nic a simulated precipitation is in general agreen South Asia considering the aerosol effects.

aerosol climate model, ion model (AGCM) with al circi ration has the effect on cloud cing precipitation and increasing is reduced by the aerosol-induced ability of water vapor to and n d ue anthropogenic aerosols SPRINTA RS is driven by

eloped the AGCM developed by the Center for Climate Syst em Research (CCSR)/University of Tokyo, National

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and Frontier Research Center for G lobal Change (FRCGC). carbonaceous (black f sulfate, that is, sulfur on, advection, diffusion, he radiation scheme in

cattering and absorption

of the water vapor. A

and pre cipitation rate, ud ople i siz simulation indicate s that the liquid water pat h face te mp ature. If the mixed layer South Asia, this can be also because a to t he aerosol se mi-direct effect. trong aerosol second indirect e h ydrological cycle. T he le variations in E ast and

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The SKYIL method is a well-t ested in-situ constants, specifically created fo r the PREDE s instruments located at V alencia and El Arenceille AERONET network (Holben et al., 1998). The PREDE and CIMEL radiometers were take n characteristics of the CIMEL instrument. The previously been calibrated using a standard the determination of the solar calibration constants

were compared with the available sets of experi standard Langley plot method. Concerning the instrument located at Valencia, a marked influence of the external temperature on the retrieved ti

although the temperature dependence of through labo ratory tests and subsequently taken at the same wavelength were also co use of a Fourier analysis allowed the definition oscillations, as well as a d iurnal component. Subsequently, a generalized least square method was s et

up in order to remove the e two components of the calibration constants. The same analysis was performed for the El Arenosillo and Singar effects of the external temperature on the dependence is not a part of the SKY IL met such a nalysis, because the results obtained by a pplying the SKYIL met hod to the CI MEL data, wer e found to be of worse quality than those obtained the method was applied to a Prede instrument (Campanelli et al., 2004), which does not require correction was performed, the accuracy of the method was evaluated, finding that it is within 2.1% and

solar calibration constant values given by uncertainty produced by the SKYIL method constant values retrieved by the SKYIL me calibrations. T he agreement was c onsiste method was found to be suitable for being CIMEL instrum ents, provided that the radi radiance. A fur ther step of the present inv CIMEL sun- sky radio meters belo nging to

that appropriate corrections can be applie deviation occurred. Although I imited to thr results. For these reasons, it is worthwhil method to the AERONET database. Simulta recommended for CIMEL instruments invo

of the solar calibration aily dete mete plied t o three CIMEL ft wa apore, the last one being part of hd electronic characteristics of the (YIL <u>method</u> was adapted to the nly if ance. d to th

CIMEL radiometer has iterative procedure for dataset, and the results mental calibratio n constants deter mined with the

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0 nm was investigated adiance measurements (1998) procedure. The he season al temperature

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s necessary to perform

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une correcti •. Once the temperatur e 1.0%, depending on wavelength. This is an important result, since the nominal uncertainty affecting the ration m ethod (of 1 - 2%) and the The time-trend of the calibration calculated in the experimental s In conclusion, the SK YIL pration c onstants also for rated for the diffuse he method to several calibratio n metho d

independent of the AERONE T system would be very us eful to diagno se the condition of a sky radiometer, whose data analysis is sensitive to small errors in the measured data. Using an independent method, the variation of the calibration constant due to instrumental drift can be quickly identified, so e peri od in which the

y furnished interesting tter applying the SKYIL e present m ethodology is the recently c reated Spanish

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network RIMA, and can be compared to oth al., 2004). References M. Campanelli, T. Ne constant for a sun-sky radiom eter: Propos N. Holben, T. F. Eck, I. Slutsker, D. Tanr, J. Nakajima, F. Lavenu, I. Jankowiak and A. data archive f or aerosol cha racterization, Toledano, A. M. De Fruto s, M. So rribas, PHOTONS-AERONET sites in Spain. Calibr Apl., 37, 3401 3406 (2004).

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bcedures, like t he Kciclo method (Cachorr o et Liker, Determination of the solar calibration Appl. Opt., 43, 651 659 (2004). B. Liket, J. A. Reagan, Y. Kaufman, T. Liket instrume nt network and 16 (1998). Cachorro, V., C. De La Morena. T he in method. Opt. Pur. y

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elev. 3500 m) in March 2007 and will last app instrumentation will be deployed: a series to 1100 (cm-1), a m icrowave radiometer f profiling o f atmo spheric aero sol, tempera optical properties and coverage, temperatu (Michelson Interferometer for Passive Atmo (Environmental Satellite) platform (Europ MIUR, PRIN2005 project # 2005025202 mately 20 days. At this site the following Spectrometers, covering the range 100 one Raman Lidar systems for the Furt her info rmation about cloud will be obtained through MIPAS Spectro onboard ENVISAT/ 1

Keywords: radiation

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To estimate the effect of cloud condensati on nucle microphysical model composed of particle retho developed. Our model estimates the maximu of cloud droplets by using the parcel mode condensation, coalescence, b reakup, sedim using two-moment bin method on the grid hybrid microphysical cloud m odel, each gri nuclei. In the case that the r elative humid y of th

case that relative humidity of the grid point is larger than 100% and cloud water on the windward side of the point does not exist, air parcel including CCN_and vapor starts to rise from the windward side of the point. When droplets condensed on CC not become cloud droplets, the cloud dropl temperature in the parcel are given to the and coalescence on grid points are calculat moment bin method developed by Chen a droplet size distribution. There are 71 bins for radii between 1 mm and 3.25 mm. Coalescence efficiency

growth by condensation and coalescence step properly, two kinds of scheme are coalescence of large droplets, the other is scheme is used, very short time step (0.01 s) is a based on the model designed to test the warr ra

(domain: 9 km x 3 km, dx and dz: 50 m, experiments to estimate the effect of hygro

show that 10-minute seeding of NaCl partic small-updraft cloud is not effective in increase optimal seeding duration will be presented.

eccurately, hybrid cloud thod in grid model wa s d bin ion and the number concentration work. And our model estimates f cloud droplets and raindrops b y elian framework. In our ian or el to imate the activation of or the first time, or the

> m embryo, which can of vapor and potential growth by condensation work by using the two erical diffusion of cloud

developed by Seifert et al. (2005) is used to estimate coalescence and breakup. The time steps f or multicoales cence in one time stic d escenc e for rare lucky fre ue Fe fo coalescence of a large droplet and small droplets f ollowing to Chens doctor al thesis. If only general stochastic coales cence dynamical new work of this study was eeded. The microph model in Case 1 of the fifth WMO Cloud Modeling Workshop (Szumowski et al . 1998). Th is dynamical cloud model predicts an evolvin g flow for 150 minutes and performs a two-dimensional advection of the temperature and water variables in applica tion of this model, numerical

tation was carried out. The results dius under the cloud base for ptimal seeding particles and

Keywords:








humidity, or the water vapour super-satura profile can also strongly influence the aero thesize-distribution and the mixing state background ae rosols influence the impact measurements or relevant l aboratory studies uncertainties in the models. Hence we discuss the uncertainties. We findconsiderable differences in

lof fan

ne vertical distribution if we knew the mass, operties of the nat ural e. Unfort unately, in-situ relevantprocess-specifi c e implied range of aerosol impact s due to these

aerosol optical properties an d impacts on clou ds ey of d privective scavenging and

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between our test calculations with CAM-Ost vertical transport give higher aerosol burder moreabsorption on the top of the atmosphere

at the ground and ribution give significant

m brir ai changes in CCNs and in-direct forcing, with much smaller impact on total aerosol mass. We also find that although changes may be small on a globally average ed they n besignific regionally. For example, treating primary emitted sulphate as sulphuric and n mode particles reduces instea accumulat theanthropogenic optical depth over Europe up to 25%, compared to only 6% globally.









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Christos D. Papadimas 1,2

lopo Atmospheric aerosols, both natural and ant atmosphere system, since they perturb its ra hia radiation and indirectly by acting as clou microphysical and optical properties. However, forcing and he nce, the assessment of glob microphysical and optical pr operties and t optical properties are among the most important d

monitoring the se properties i s essential for better minimizing the associated uncertainties. The Medi suspended particulate matter. In this area due to the accumulation of aerosols, which the eastern Mediterranean, and because radiation. The Mediterranean Sea is one world. All these, together wit h the fact that Me

eastern Mediterannean basin an ideal region to study aerosol properties and radiative/c limatic effects. Lately, sophisticated satellite-based instruments have improved drastically the observation of aerosols, providing accurate data with comp lete s investigate the temporal and spatial distribution basin, using 6-year (2000-2006) daily satellit. accurate characterization of aero sol proper ties in the study area based available data enabling complete spatial coverage.

of possible changes in aerosol loads and pr effects on the regions climate . The study from MODIS-Terra Level-3 daily mean atm at 1x1 latitude-longitude resolution, for the given on the variability of Aer osol Optical as well as on mean ann ual and seasonal average value of AOT is found to be equal areas as well as over region s affected by

Moreover, the inter-annual variability and trends of AOT is investigated at both the geographical cell and mean regional scales. Overall, the regional mean aerosol optical thickness is found to have decreased over the period 2000-2006 by -18% in relative percentage terms. Similar analyses are also performed

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for the rest of the 18 aerosol parameter fraction), the Angstrm exponent, the mas factor. For m any p arameters, e.g. AOT, wavelengths, providing thus an insight on shows that the overall decreasing trend of

e clima te of the Earthant f ectly by cattering and absorbing nd in doing so modify cloud nty regarding the aerosol radiative eir rapidly changing s due ' r om nds t o weeks. Aer osol atic effects. Therefore,

hts of a understanding the climatic role of aerosols and

oulos 3, Ilias Vardavas 2,4



a unique area in terms of kes maximum values nditions prevailing over ged exposure to solar optical depths in the nsitive reg ion make the

ader ea of the Mediterrane an in the b the **IOL** rate resolution Imaging ata' fron Spectroradiometer (MODIS) instrument onboard the Terra and Aqua satellites. The aim is to provide an The most contemporary This stue ndates sin ar ones that have bee n performed based only on pr evious generation sate llite data, on surface-based meas urements, or on contemporary satellite data but of short temporal coverage. Our study period also allows the detection the study region, which can have severe liffer ent aerosol products derived over the Mediterranean region February 2006. Emphasis is

regin. In this study, we

over both land and ocean, el. The 6-year regional parent over large urban thern Africa deserts.

DT at 0.55 µm (fine , and the asymmetry estigate d at numerous region. O ur investigation ttributed to decreasing loads











The formation of droplets starting from cloud conden to the formation of ice crystal s is reflected in the microphysics are crucial to understand the measurements are limited in their temporal clouds. S atellite measu rements from passiv profiles of precipitation size drops are retriev leave a gap, especially in our understanding is sparse and remote meas urements difficult to

measuring the vertical profile of cloud microphysical properties. We present results that confirm the capabilities of this concept with respect to regions. 3D c loud microphysical in formati input into a 3D Monte Carlo radiative tran the visible, ne ar-infrared, and infrared spe ra set of simulated radiances and given cloud nic algorithm to derive a profile of cloud phase and

nuclei thro le of ical p t of precipit at ial extent a ors a<u>re con</u>c rada develd nt he CĽ

the liquid phase mechanisms ties. Insights into these ld pro or any aerosol effect. Insitu pplied mainly to s hallow types of ating on cloud tops while vertical ients b oudSat. These data still ction where in-situ data ep co ud aer osol interactio n

mission in 3-D) passive clou d side viewing sensor concept is planned to become a new means o f

nd mixed phase cloud ble m odel is used as ur notional channels in onstrated that this dat a or a Bay esian retrieval



















and changes in existing technologies, that is increased o pencast mining with highly m degradation in various sectors. Air environ emissions from mining activities. Water reso solids. The highly mechanised operat ions in changes in land use patterns are some of the other

large-scale coal mining, transportation and <u>small scale utilisation activities in the Manuguru Coal belt</u> area of Andhr a Pradesh has become a ma study area, Manuguru region consists of t Manuguru OC-II is the biggest mine yieldir mainly responsible for the environmental d environmental status of the study area ba noise levels in the stud y ar ea. Air around an o pencast coal mine is polluted due to various unit

teo logy transfer ed s ystems, deteriorated e po đ٢ ng ai ead mpact

is t echnology transfer has led to h has resulted in environmenta to dust pollution and gaseous id min ainage a nd suspended noise ution. Deforestation and st mining. This calls for

the study of e nvironmental impacts related to open cast mining. Environmental deterioration due to



al area in India. The derground mines. The ction of the area, and is necessary to assess the water characteristics and

operations like drilling, blasting, removal of over burd en or coal, tran sportation of over burden or coal, operations at coal handling plant etc. Air n the area covering important locations has evinced that pollution has bee ang impact will be certainly nts namely suspend ed adverse in the coming years. Seasonal and ur **PITCI** particulate matter, sulphur dioxide and oxides of nitrogen is presented in the paper. Overall air quality is also obtained from the Oak R idge Air Quality Ind . The prob ble environ. eptal impacts of mining on the air environment are predicted with the help of a ilable d

Keywords: oraqiindex, coaimining, airquiaityassessment





Keywords: radiogenic source, r

ation measurement



linear regression is used to estimate asymp wavelengths, based on AERONET retrievals is used to retrieve single scattering albedo (ssa). Numerical tests w ith the radiative transfer code SBDART show ed that the natural lo garithm of DDR inc reases linearly with single s cattering albedo. Thus, simulations studies were performed t retrieve single scattering al bedo, using M asymmetry factor, single scattering albedo and complex refractive indices. Comparisor ; b

estimated from the linear r elationships re approach was applied to real measurement s performed with a MFRSK collocated with radiometer at Sao Paulo city from 19 Aug ust to 28 September 2004. In this period , 47 AERONET retrievals were available. AERONET single

depth วับ tor fro ng The fr ffuse

as usual. A multi-variate t calculated at different expo DR) spectral irradiances

> ear relationships to as us ed to estimate

aerosol size distribution tering albedo and the 6 or a bout 0.02. This

an AERONET e d from measur ements of sky

radiances performed when solar zenith angle . MF performs global and are the results, MFRSR diffuse irradiance measureme nts every 1 te. rde tο, single scattering albedo estim ates were averaged, considering 15 minutes before and after AERONET measurement times. Good cor relation was observ d for the fo r waveleng. s although MFSRS results are systematically lower than the AERONET Other analy zed as well as the rei idds will be possible reasons for this underestimation of MFRSR single scattering albedo retrievals.











Asian tropospheric aerosols (dust, sulfate, optical properties (aerosol optical thickness during 2003-2005 (from 1 June, 2003 to 30 transport model (CFORS). CFORS model res and Skyradiometer optical ob servations. It many observed important features including

outflow and aerosol vertical profiles. In these comp were transported over Japan in the vertically wide range, especially during spring. To clarify the typical vertical characteristics of Asian aerosol layers during spring, aerosol scale heights for dust and spherical aerosols (including sulfate and carbonaced and CFORS model results. It was shown th with spherical aerosols. Further analysis o th and AOT fields shows this vertical characteri gradually increase along the ridge of its h rizo

irbo osols extinction coe 005) were sim e int<u>ensivelv</u> vn th mu hig se T le ve fo u

t) transport and their ent, and single scattering albedo) ed using a regional-scale chemical nined with the Mie scattering Lidar aer o felds acc urately capture asso ed with the continental my dense aerosol layers

> the Lidar observation s over Japan compared ed aerosol scale height that dust scale height s sp herical scale height

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maintains the altitude level of 2-3k m. Therefore, we found that Asian dust is mainly transported while increasing its transport altitude and that spherica I aerosol are mainly transported within the boundary layer.

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Keywords: particle size distribution, <u>multiwavelenght lidar</u>, convective atmosphere





Aerosol particles have a major impact on solar radiation, changing cloud properties and so on. industrial activity especially over Asia. Then his r of density of aerosol concentrations. It is in the climate. In order to investigate aerosol in model, called SPRINTARS [Takemura et al. Atmospheric G eneral Circulation model [N mineral dust and sea salt aerosols from var particles impact on the radiation budget by directly

Then investigation of radiation budget variations of radiation budge t are analyze Because aero sol particles change the amo absorption of solar radiation, we calculate analysis of trends in surface radiation. Nex effects on the radiation budget are investig through radiati on budget change. This st industrial activity is growing rapidly.

thro ugh causing air pollution, reducing il fuel usage in fu incre to investigat 0 on the climate 200<u>2, 200</u>51 et al Ŋ٦ sion and a

en increasing with growing resulted in an increase sade w these erosols in the air affect rosol transport model is used. This cou _pled with CCSR/NIES/FRCGC fbonaceous, sulfate, l treat said

the incr ease of aerosol e solar ra diation and by indirectly mod ifying the optical properties [Twome y, 1974] and lifetimes of clouds [Albrecht, 1989].

variation is_important for estimating aerosol effects. Initially climate change in

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mode I simulations. und by scattering an d e atmosphere from the and these cloud change ct on the dynamical field East Asia, where







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factors in the estimation is a sensor calibration. procedures. In this study, we have develo ped a vi carious calibration technique based on the original sensor information of GMS-5/VISSR, using the radiative transfer code RSTAR6 (system for transfer of atmospheric radiation). The atmospheric pa

aerosol and cloud using REAP (Retrieval Analysis Program for Cloud Optical Measure JRA-25 for temperature, pre ssure, water v apd surface was decided from the wind velocity of s standard product of the MO DIS data (16day mean product). Aerosol optical characteristics in the land

between four detectors were e enough small Than analysis of radiation budget on Refined-VISS wa has been performed using R-VISSR data set Studies) field experiment period. This study was

ssues in the calibration is, by MODIS data for PCOM (Comprehensive y objective analysis data The albedo in the se a

lbedo in Australia used a region used the gro und observation data (Sunp hotometer). The stripe n oise by sensitivity difference

> tion algorithm f or a long-ter m hple, adiation budget analysis 4 ic P ruc Environmental Change performed based on a collaboration of Center for

Climate System Research, University of Tokyo and Meteorological Satellite Conter, Japan Meteorological Agency. Aerosol optical characteristics data by ground observation in A ustralia provided by GAW PF R network, at the Australian Bureau of M eteorology (BoM). The authors would like to acknowledge Mr. Christoph Wehrli (PMOD/WRC, Davos, Switzeriand) and Dr. Bruce W Forgan (BOM).

AP


















The collision-coalescence of sedimenting particles and computational problem. Most previous numerical, on the motion of individual point particles ir given in terms of the geometric col lision keilel interactions. More recently, we extended th local aerodynamic interaction s of dr oplets particles (J. A tmos. Sci. 62: 2433-2450, undisturbed turb ulent flow and an analytical repl

flows induced by the particles. While this method represents a consistent approach for a turbulen t suspension of aerodynamically-interacting particles, it does not resolve fully the boundary condition and flow at the scale of droplet siz e. Here we particles suspended in a v iscous fluid. Prosperetti and co-workers (J. C omp. Ph interactions. In this appr oach, the fluid flo Stokes flow representation is used to simul te t its neighborhood. Particle Reynolds numbers up to a rew

lets in a turb rticle les d se to the carr bring particle pa r<u>ticle rep</u>r t simu δn e app h c bf loca

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t gas presents a challenging in turbulent flows focus gulat uid turbuence. Collision rates are icle hydrodynamic or aerodynamic tation to include some aspects of urbule ollisi on of sedimenting irect sim ulation of the e disturbance (Stokes)

approach for colliding bach developed by enting t he particle-fluid h a nd a local analytical ticle with the fluid flow in nundreds could be handled by this appro ach and the particle size may overlap with the flow length scale. This paper will focus on the development of e gravitational collision of two

a ran dom su spension.

undary conditions and

the numerical method when applied to t sedimenting droplets and the second is Specific imple mentation issues such as incorporation of sub-grid lubrication force will be discussed, along with results designed to establish the validity of the approach.

Keywords: collision coalescence, moving boundary, simulation







Extinction of radiation in the marine bo undersalayer atmospheric aerosol . This is important to o environmental monitoring, backscatter of lig In unpolluted regions the greatest effects or from breaking waves and v ariations in rela modulated by wind, tide, swell, wave spect superimposed upon aerosol generated by o focus of our study is the extinction and optical effe

predictive algorithms that de scribe 3-D aer model of the marine and coastal atmosph basis of the long-term experimental data re (dN/dr) in the band particles sizes in 0.01 of heights is 0 - 25 m. Bands of wind spee dN/dr of the model is characterized by the r1 = 0,03; r2 = 0,24; r3 = 2; r4 = 10

amplitude and width of the modes as f unctions of fetch and wind s peed. In the paper the dN/dr behavior depending at change meteorolog speed (U) and RH is show. T he received r and ANAM. On the basis of the developed m step $\Delta \lambda = 0.0001$ m is pres ented. Also $a(\lambda)$ profile experimental data of $a(\lambda)$ received by a tra

[1]. G. Kaloshin, J. Piaz zola. The Coastal Laser Radar Conf., Nara, Japan, pp. 423-42



ng and absorption due to remot e sensing at satell ate forcing), cloud properties etc. tinction will be a result of sea-s alt f b reaking waves appears to be ese influences will be tion s. that v s with wind speed. The ific coastal region. This

involves linking coastal physical properties to oceanic and meteorological parameters in order to develop osol structure and variability. The aerosol microphysical



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nodel is made on the es distribution function resent time for the band 20 km, RH = 40 - 98 %. ith modal radiuses, equal

b ve s ea level, fetch (X), wind lable crophysical models NAN ed to av or scheres the description of e th el v sad bry_ the last version of developed code MaexPro 5.0 (Marin e Aerosol Extinction Profiles) for spectral profiles of aerosol extinction coefficients $a(\lambda)$ calculation the wave right band, and $\lambda = 0.2 - 12$ m, with for vari 🖬 n d mode (combi nations X an d U) calculated by MaexPro 5.0 code are given. Results of $a(\lambda)$ profiles calculations are presented at chang e RH = 40 - 98 % and heights H = 0 - 25 m. The calculated spectrums of $a(\lambda)$ profiles are compared with arious geographical areas. References I, Proc. of the 23rd International





In the report description of the last version calculation spectral and vertical profiles of a roso area distribution, volumes dis tribution, mod a computer pro gram under constantly dev location place in which a fetch is key entran as functions of atmospheric effects using s structure, a spectral band an d a height of submitted as graphically, and as tables. Commands

Windows program with a user-friendly inte

depends on the necessary wave length re

height, and do es not exceed tens seco nd characteristics, such as a erosol sizes distrit

extinction spectra, are performed in a few

of Ma le (Marin oeffig inctid ol extinc bn nt to e stima mete<u>r [1]. T</u>ł l mete bgi sor lo n p for su

ogl extinction Profile) for erosol sizes distribution, ΄ α(λ) tra is submitted. Code MaexPro is E O systems signal power at a ogra<u>m carries</u> out calculation $a(\lambda)$, aeroso I microph ysical arama S pe al behavior $a(\lambda)$ can be n or change of figu res;

profiles extrapolation; a lens; all kinds of p ossible copyings; the data pr esentation, convenient for a n

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input in code MODTRAN, and etc. are stipulated. The code MaexPro is a completely mouse-driven PC litid eco

vertical profiles of $a(\lambda)$ nd the location place ition. Other calculatio ns tribution, modes aerosol shin, S. A. Shishkin, S.A.

Serov. Development of the co de MaexPro to r calculation atmospheric aerosol extinction in the m arine and coastal surface layer. Atmospheric and Oceanic Optics, 2007 (in press)















components of the studied parameters i n the peasured values. For example, value of R between G and Ta, S, da a correspondingly is equal:-8 ar 0.30, -0.21, 0.05, -0.06, 0.32, 0.17 and 0.26. The same for the random components (time series of D and Bg are random) is equal: -0.24, -0.62, -0.01 42, 0.29, .02, 0.75. given results will be used een the a bove-indicated for the multidimensional regression analysi tions bet atmospheric parameters.













has, to our k nowledge, not appeared in transportation and sinks of five major aero soot and sulfate) and is well suited to simu where biomass burning, ind ustrial pollution Model calculations where per formed for March and

season in which large amounts of dust are swept into the atmosphere over the Chines e deserts. The Asian Dust network consists of several stations (12_where used in this s tudy), mainly in China, South Korea and Japan. Observations are made a hours a da y, 7 days a week. The networ calculations, in particular with respect to th and dust events. Additional AOD observation are also considered in this study. A first an predicting several of these events. Due t

tends to be overestimated by the model. Fi

simulated backscatter, it appears there is a

ature. The S c ies <u>(carbon</u>a s of a comp 'ni bute d stor nd)4 a n ไ

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NTA RS model sim ulates sources , s ae rosol, mineral dust, sea s alt, ol over S outh-East Asia arge atmospheric loads. before and during the

> every 15 minutes, 24 omparison with model ibution of both pollution

tworks of skyradiometers RS model is capable o f vertical resolut ion (20 layers), the height

distribution in the simulation is more diffuse than in the observations. The height of the boundary layer observed LIDAR backscatter to sc at ing properties, especially Mie calculations) m ay eu y t

at 1064 nm. In particular the backscatter pr fent bag need to be r evised. Our ultimate goal is to de velop an assimilation system for SPRINTARS tha incorporates o bservations fr om both ground st stions (skyra Consequences for such an assimilation system dervel from the iometers and DARs) and satellites. will also be discussed. present study

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Aerosol number size distributions have been mea aerosol optical spectrometer. The data pres hts >0.3 to 20.0 mm during the period from N seasonal patte rns o f the ae rosol particle variation of the particle number concentration highest total number concentrations have b 1700-2200 hours). Seasonally, the highest total nur

dominating and is contributing more than <u>98%</u> for the total number concentration. Particles in the submicron size regime con tribute domin Interesting features are observed in the be foggy, hazy, a nd clear-sky c onditions prev le aerosol concentration exhibited a larg e di rib higher on ha zy and foggy days. Backwa transport process of the air mass in Delhi.

Nove 5d sir distribution 2004 to De r co <u>ncentra</u>t wed rved ۱q centra

t NPL, New Delhi using 207 h the particle diameter size range per 2006. The daily, monthly a nd have been i nvestigated. Diurnal lation h traffic activities, the urs (i.e., 07001000 a nd pea observe d during winter

than in summer and monsoon. At NPL, submicron mode (DP < 1.0 mm) particles concentration is

diati ve interactions. ution under intermittent udy. In all size regim es bilities bein g particularly ployed to examine the

















The inference of particle microphysical d effective radius, refractive in dex etc., from problem. The optical parameter set is conne the microphysical parameter set. Various methods and constraints have already been the first time, we propose to tackle the pro statistics, one can use probability to quantify unce

In order to do so, one needs to summari experiment in a probability distribution calle da priori distribution. Once that is accomplished, one can integrate that information with the informa distribution is called posterior distribution. one can expect analytical results only in a obtained appr oximately by simulating dra (MCMC) technique. This is basically a distribution, which involves solving repeatedly the Fredholm integral equation for different combination

of the physical parameters (ie. we solve the inverse problem by stoch astic solutions of the so called forward problem). Bayesian theory offers t information outside of the experiment with context, and pro vides a natural measure demonstrate the application of this method to the inversion of extinction and optical t hickness data in order to derive the accum ulation mode of atmospheric particles. Starting logarithmic normal distribution for the particles fize geometric sta ndard deviation and particle concentr atmospheric measurements using Mie-theory. For simplification the complex refractive index is assumed to be known. However it will be shown that method and an un known re fractive index Furthermore an outlook for the application be given.

Keywords

vo lume-concentration, well-hown ill-posed inversion bral equation of the first kind with uares fitting, based on different this i sion problem. Here, for stics. ne conte xt of Bayesian f a physi cal parameter.

ze all information about each parameter o utside of the

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theorem. The resulting y very c hallenging and posterior distribution is kov Ch ain Mo nte Car lo e unknown posterior

se problem s: it integrates ng the ment coherent mathematical parameters.We w ut t ph' sica The will be a monomodal ke posterie distributions mode radius, ations will be o btained for several different

> e lifted with no loss of generality of our eral l uncertainty of the results. (or additional) optical data will











Automated all sky imagery system would observation. A recent World Meteorological Organization offers the potential to understand and quant y clo (All Sky Imager) has been developed by present 2004. In our plast works, we have mainly made from the all-sky images. We found that the decision of cloud pixels, espe cially for the ci

than other places mostly. In other words, A

we investigate a new approach to get the First, the relat ion between AOD and the di stribution of b lue/red radiative ratio of ASI images was analyzed by numerical model simulations. We find that the radiative ratio value distribution is sensitive to the AOD under different aerosol types a we set up an inversion algorithm for AOD we compare the AOD retrieve d from all-s instrument at the same site. Results show provided by the AERONET/CIMEL. This diff atmospheric aerosol profiles. On the o ther hand,

or in ground-based macro-scale cloud t mentio por effe more hors and has earches on h opt<u>ical dep</u>t east Ch he aei ŵ٢ s con info aeroso bth in

the use of all-sky imagery In China, a type of ASI urate cted series of all-sky images since b get the parameters about cloud DD) made a critical impact on the op tical depth is bigger aerosols. In this paper, ion a om the all-sky images.



between them. Then, on. Using this formula, m the AE RONET/CIMEL tic bigger than the AOD liative transfer model and

obtain the AOD from the ra diative ratio of the all-s ky image. In order t o improve the retrieved result, the paper established another inversion for ata and the radiative ratio value of the images. This is a formula with regiona show that this regional combar ve may establish mor e formula improved the inversion results of A Ra tiga ЮП continuous AOD series for non overcast sky which is valuable to investigate aerosols direct and indirect effect to climate/radiative interaction.

Keywords: all sky, images, aod










and 57.135.1 ug/m3 in sum mer, fall, and wir 4.32.6, 4.02.7, and 5.23.2 ug/m3, respectively. From the regression analysis between CO and BC, two distinct regres sion trends were observed. Regr

me ession s lope (BC/CO) was determined to be 6.8 Vember 2006 to February 2007. orio

was 33.428.1, 28.130.9, h BC con centration was

of CO. Differences in the

ug/m3/ppm during May to October 2006 a This result can be used as an indicator for characteristics of aerosol light absorption h condition.

mass types and haze

aethalometer Keywords: absorption coefficient, bla















Ozone Mapping Spectrometer (TOMS) AOT ol properties and other related variables, clim atological (for the 12-year page) lv me aerosol DREs wer nhoni generated at 1x1 latitude-longitude resolution Tł ed at the top of the e cç atmosphere (TOA, Δ FTOA), within the atmosphere (Δ Fatmab) and at t he Earths sur face (Δ Fsurfnet). Subsequently, the magnitude of the aerosol effect was quantified by the on therma lynamics (A difference AETD = Δ Fatmab- Δ Fsurfnet. We four gr Is has a significant effect the pre nge of aero on the thermal dynamics of the Earth-atmosphere system. More specifically, by cooling the surface and warming the atmosphere, aerosols act to produce <u>more stable atmospheric conditions</u> by decreasing convective activity. They also reduce evapo ce, and so can have a significant effect on the hydrological cycle by suppressing cl y the convective) and associate d

datasets (NCEP/NCAR). Global distributions of aerosol optical thickness (AOT) were taken from the Total

precipitation. This aerosol red istribution of the atmosphere highlights the role of incre largest effect of aerosols on thermal dynamic threatened areas, such as the Sahara and i

tween the Earths surface and ulate matter to climate. The and nearby desertification-

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Keywords: aer











radiation, and precipitation. These effects a cloud top and base height, thickness, variab improve the accuracy of parameteri zations and clouds. In this work, two cloud models th brightness temperatures from radiosonde da a ai water density, models are necessary to estimate cl

transfer algorithms that simulate electroma gnetic parameters such as brightness temperatures. Clouds are identified fr exceeds a suitable threshold function (whi used). We first evaluate and optimize the clouds, by comparing the cloud base heigh a ceilometer. Then, the performances of simulated brightness temperatures in the a microwave radiometer. Two cloud models that are currently in use in pro pagation and remote sensing

simulations in the presence of non-precipitating clouds are evaluated: the model proposed by Decker et al. and the model proposed by Salonen a suggest an improvement of the Salonens h new cloud density function for computing clo of the three models are assessed at the Atmospheric Radiation Measurement (ARM) Programs Southern Great Plaints (SGP) site in O klahoma, USA, by using data from Vaisala RSS and a dual-channel microwave radiometer at 22-8 a will be used to distinguish different types of clouds and to infer inter-annual variations and trends from

(KNMI).Statistics of cloud boundaries (base presented for each cloud model and m investigation will be also carried out by using Mare, in Italy, by the Italian Air Force Serv

ly dep el den cloud wate scrib<u>e_atmo</u>s an be' ide analyz Sind id and



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cloud venical structure, such as tent. Therefore, it is important to c radiative transfer, water vapor, the ra ive transfer to simulat e dios o s do not measure cloud ty profiles for radiative

atmospheric attenuation and

when the re lative humidity humidity sensor being rrectly the presence of with those provided by zed by comparing the asured by a dual-channel

madel is proposed: we both ng th resence of clouds, and a vithi n a oud. The performance s adiosondes, a ceilomete r

, the three cloud models

a long record of radioson de observations operated by the Royal Netherl ands Meteorological Institute thickness as a function of height will be g es will be performed. A similar cted at the station of Pratica di fferent European area.

Keywords:









temperature errors. We have investigated th Centre for Medium range W eather Forecas Centre Atmospheric (RCA) model radiation resolution downscaling of climate scenarios and for

scheme in RCA was originally developed for forecast purposes. The scheme is computationally fast but highly simplified with only one wavelength band for the longwave region and one for the shortwav e region and empirical coefficients for most a RRTM longwave spectrum is divided into bands. The ECMWF radiation scheme is ab mode. In the ECMWF model the full radiat bn (step of 20 mi nutes) at every 4th latitudina po the fluxes at each grid point and time step. We have run both radiation schemes in single column mode using observed data of humidity and clouds from

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emporal sampling in the European ne and i n the Rossby ation s mo del is used for high l clim t climate. The radiation

> impacts. The ECMWFspectrum has up to 6 heme in single column very 3 hours (for a tim e ime and space to obtain

the Cloudnet proje ct and studied the impact of reduced temporal sampling on the calculate adi



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Negusini Monia, Sarti Pierguido, **Tem**asi.

The Italian Mario Zucchelli Station (-74 41 5 5.6 Northern Victo ria Land, is e quipped with continuously observing since 1998. Mario Zu of scientific observations are carried out, eit er ' devoted to atmospheric physics is located a Icar where several atmospheric measuremen Vts are

analyse the GPS data set acquired over a set acquired over acq station TNB1 with the aim of retrieving the vapour radiative effects on the thermal rad the energy bal ance of the surface-atmos assumes in general appreciably lower values the measurements of the mean longwave radiation a function of precipitable wat er to show the relationship existing between this radiative term and the

derived IPWV values. The radiosounding dat

7 E), situated at Terra Nova Bay, ti oning System re ceiver (TNB1), ntific facility where a large number nally. Ì articular, an observatory the b : it is a Clean Air Facility long-wave radiation

Lanconelli Christian

measurements have been p erformed and recorded using a Kipp&Zone n CNR-1 net r adiometer. We period spanning 20.00-2005 by the permanent GPS

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IPWV) content. Water of basic importance for here pre cipitable water ccurate calculations and ova Bay are presented as

temperature data for the

total atmospheric content of water vapour. GPS-derived IPWV values are compared with radiosondenhg th

errors due to radiation and heat exchange r relative humidity data ecelly de veloped.We are for the errors and various dry bias follo ng a ire een applied to GPS and radiosonde presenting the analysis strategies that have b data sets for computing IPWV, the relation between the measured radian flux density of the GPS-derived IPWV and the comparison with the predicted irradiance derived a mod el with different profiles of temperature and humidity computed from radio soundings performed at Terra Nova Bay during the austral summer season.. GPS and radiosonde derived water vapour contents at Terra Nova Bay show a good agreement over the whole sample pe













Speculations on the impact o f variations concerns that so lar dimming has largely marshed the that the recent reversal fro m solar dimmin responsible for the observed warming. Todis nta warming, trends in diur naltemperature rand

overglobal land surfaces show, after decade 1980s. They suggest that so lar dimming, p masking green house warming, but only u brightening. The reversal from dimmingto brighten

moretransparent atmospheres. With this transition, the uncovered greenhouse effect started to reveal its full dimension, as manifested in a rapid te Recent surface so lar brightening canno t warming, since land temperatures increase did not fully o utweigh solar dimmingwithir 2007: Impact of global dimming andbrig h Wild, M., and Co-authors 2005: From dim ming Earths surface. Science, 308, 847-850.

full magnitude of to irathe ighter surface sola nalyzed. The c line, a disti aused 1980 nen

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aration on global warmi ng range fro m nhouse warming, to claim s reenhouse effect was á n fr d greenhouse influences on global nal temperature r anges averaged

endency to level o ffsince the mid lution, was effective in lually t ransformed into e effective air pollution measures and the breakdownof the economy in the former communist countries, leading to cleaner and

> and since mid-1980s). main c ause of global hough sol ar brightening hmura A., Makowski, K ., . Res. Lett., 34,L04702. changes insolar radiation at the



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Substantial uncertainty still exists regarding system, and it s representation in General Circulation

surface observations, the maj ority of the GC 1s p (AR4) overestimate the surface insolation, b This is in line with an analysis of 20 ear Intercomparison Project AMIP II and sugges solar radiation. Based on observational clear from the Base line Surface Radiation Netwo Program (ARM) it is shown, that the surface insola

global mean). This identifies an overly transparent cloud-free atmosphere as a key error source for the long known problem of excessive su rface participating in IPCC-AR4 with higher atmo averaged) and more realistic aerosol treatn clear-sky climatologies. This underlines the | pr the observation and diagnosis of solar rad model in the longwave radia tion budget is the downwa

clear sky conditions. A comparison with ave IPCC AR4 GC Ms tend to underestimate the Ohmura, A., Gilgen, H., Roeckner, E., Gio tta. radiative energy in the global climate system: GCM 14, 853-869. Wild, M., Ohmura, A., Gilgen, H., longwave radiation in Gener al Circulation Modes radiation budgets in atmospheric model in tercomparisons from a surfac e perspective. Geophys. Res. Lett., 32, L07704, doi:10.1029/ 2005GL02 2421. 850. Wild, M., Long, C.N., and Ohmura, A., 2006: Evaluation of clear-sky solar flu xes in G perspective. J. Geophys. Res., 111, D01104 2006: Radiative fluxes in ECHAM5. J. Clima

Keywords

adiative energy within the global climate dels. Comp the ipatin St IP 2 on average GCMs partici the GCM atmo imato or Idw ēς N) ar sp he ie A

forth assessment report lile the bas is smaller at the TOA. g in the Atmospheric Model eres are still overly transparent for distribute d anchor sites Radiation Measurement ler cloud-free conditions

comprehensive set of

in many GCMs with comparatively low atmospheric clear-sky solar absorption (around 60 Wm-2 in the

tio longwave flux at the surface. Accordingly, ra

now several model s m- 2 and up, globally It with the observational modeling as well as in A difficult component to

large discrepancies exist in the global mea ns of this component in the GCMs, both under all sky and BA and BSRN suggests that the d References: Wild, M.,] 98: The disposition of <u>- 1</u> versus observational estimates. Climate Dynamic s, rcrette, J., and Sin to A., 2001: Downwar 9. Wild, M., 2005: Solar Climate 22 7-32 AMIP an d IPCC- AR4 from a surface

6118. Wild, M., a nd Roeckner, E.,



















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Karl Froyd, Daniel Cz

Mineral dusts are known to be good ice nuclei and c studies have investigated the ice nucleation montmorrillonite, kaolinite, alumina and iron sample and evaluate the composition of t performed during a measure ment campaign 2005. The AID A chamber consists of a large simulated. An adiabatic expansion in the ch with respect to ice, and therefore induces the form

instruments sensitive to the particle size, o test aerosols consisted here of Arizona Te hematite. Sulphuric acid, ammonium sulph time the Particle Aerosol by Laser Mass chamber to investigate the role of the dust particles sampled from the AI DA chamber were

with a Pumped Counterflow Virtual Impactor (PCVI) and melted. The resulting ice residue was then sent to the PALMS instrument. Particles from 0. ablated by a p owerful desorption/ionization analyzed on- line with a time- of-flight mas hematite show ed that hematite nucleates i ce b showed that particles with a large silicate for a better k nowledge of the relevant min properties.



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Another ppr

ate indirectly. Many recent ure dusts such a baical is to consider a real mineral dust the lot. Such an approach was Karlsruhe, Germany in November of an ice cloud can be vapour supersaturation I experiment consists of

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injecting a known test aeroso I and inducing ice crys tals during the expansion. The properties of the aerosols and resulting ice cr ystals are monitored over the course of the experiment by a wealth of



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ice crystal shape. The mixtu re of ill ite and times used. For the first was connected to the

ring the experiment, ice aerodynamically separated from the int erstitial aerosol namically sized and then

06 V resulting i ons were then print of the particle analyzed on- line with a time- of-flight mass spect or etailwhich gives a manupprint of the particle composition. As a result the composition of single particles was accessible. External mixture of illite and s a, Experiments with Ariz ona Test Dust er than illi or etric signal nucleat ed ice less than nective nass spe other particles . No other co mpositional tr ends for the major elements of Arizona Test Dust wer e evidenced. The use of a single particle aerosol mass_spectrometer for AIDA chamber experiments allows istics with r egard to their ice nucleation

> Keywords: aerosol ma e nucleation


There have been numerous recent publication solar radiation in addition to its capability (IFN), and could lead to reduced cloud cov investigated by using a cloud model with de The model is i nitialized using measured dis results show t hat when the dust layer with, below 3 km, where the temperature is war mer the

is st ng th hiner condensation brecip itation i ni cro<u>physics</u> ns and ce ncen on a apora

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be a good absorber for clei (CCN and ice fo rming nuclei reg ion it presents. This effect is th warm and ice p hase processes. eral dust particles. Our cloud- base height and by higher temperature

inhibits the development of cloud particles and precipitation, and together with early activation of larger cloud droplets on giant cloud condensation nuclei, which accelerates drizzle formation thr ough collision

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coalescence process, reduces the cloud opt dust layer locates at altitudes with temper ice nuclei and intensify the ice-forming pro either increase or decrease cloud opt ica chemical composition of the absorbing con atmosphere.



d that only when the ols can act as effective stence of dust layer can he con centration and rosols s uspended in the





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SERGIO RODRGUEZ 1-2, Emilio Cuevas 2, Ir T Huelva, Department of Geology, Huelva, Sp Meteorologa, La Marina, 20, 6 planta, 3 Laboratoire d'Optique Atmosphrique, Obs AERONET", Universit des Sciences et techno baie atmospheric aerosols have e xperienced a significa

number of parameters determined by these tec performing this mo nitoring. These a erose knowledge of the aerosol pr operties whic reduce the uncertainties in the climate ford may provide highly reliable aerosol data, in errequired (among other reasons) for assess column-integrated and the in-situ aerosol mea these may be affected by different processes, for example the presence of aerosols at low altitude (local

or regional) mostly affecting the in-situ measurements or aerosol layers at high altitude mostly affecting the column-integrated but not in-situ meas techniques can be properly compared only phenomenons are not frequent. One of these above the mixing layer (i.e . in the free tropospher e) at 2 367 meter above the sea level in Tenerife

(Canary Island s) and is not affected by local or rangional sources of aeros compared the aerosol volum e tempor al variations and ae ros equipment). Time series of the aerosol

(um3/cm2 and um3/cm3) exhibit highly co concatenation of long range transport of a latitudes, which results in high and low aer shows that the aerosol volume si ze disti um3/cm3) exhibits a high correlation in all concentrations). Under low volume concentrations distribution exhibits a predomin ant 2.5-4

distributions, such as the contribution (%) the inherent differences in the measurer suitability of the Izaa Global Atmospherid measurements.

Philippe Goloub3 1. University of Observatory, Instituto Nacional de e rife Canary Islands Spain. 3. rosols "PHOTONS sur ed re te sensing monitoring of Plast years both in the hniques and in the worldwide number of sites

contribute to obtain detailed bsol mo delling and to ren te sensing technique s bsol me asurements ar e а algorisms. However, the dir tly comparable becaus e

tained with these t wo data g here the above cited atior eric Observatory which is located

At Izaa site we have distributions obtained by olume siz column-integrated measurem ents performed into the AE RONET framework (CIMEL sun photometer instrument) with those obtai ned by in-situ aerosol charact erisation techniques (optical particle counter ns obtained by the t wo techniques

> tions mainly prompted by the d from the North Atlantic mid ctively. A statistical analysis chniques (um3/cm2 and tration (from low to high ne aeros ol volume size tion be tween volume

concentrations and coarse mode volume increase, time series of the aerosol volume size distributio n mode obtained with the two techniques (c olumn-integrated and in-situ) exhibits correlated temporal variations with frequent incr eases from 0.3um up to 13um. So me minor differences in the size

load, are attributed to s study evidence the erosol r emote sensing









Surface crusts play an important role in d ust emiss surface and d ecreasing the number of grain characterized by various measurement technique strength which tend to vary considerably d penetrometers, and beam balances). Recent strength meas urements at v ertical and hor regard to the effect of abrasion of sal tatin present study, a newly-designed portable field

conjunction with a portable field wind tun nel (0.75 m 1. 0 12.0 m open floored working section) to measure dust emissions in r elation to crust strength from three crusted desert surfaces. Tests wer e undertaken on a playa site in Nevada with with a complex biological -chemical crust. A the playa with a second set near the mar interior location. At all three lo cations, 14 feed that introduced separately, sand of tw Prior to each wind tunnel test, 50 crust stren gth measurements were made down the edge of the tests

plot. Maximum strength values varied cons iderably at all three sites s panning almost two orders of magnitude. The distributions of the maximum skewed, with a mean value of 18.2 mPa for margin site. In contrast, the sandier New Mee to cast trassmark terized by a stongly positively skewed distribution of strength with a mean value of 0.98 mPa. These mean strength values are much higher than comparable values pre sented in the literat emissions were lowest at the sandy New Mexico su clay-rich interior playa site (mean mg m-2 s-1). The highest emissions (mean 7.65 mg m-2 s-1) were

emissions. Similar to other published liter rates and crus t strength me asured with t provides a measure of near surface strer bombardment of a sur face by a saltating suggest that tensile strength, no matter evaluate the role of crusting on dust emiss

ncreasing ava ble fo ansp which typically g on the inst ch su<u>aqests </u>t to ch cales ire illime on th ara

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entrainment threshold of a crust strength can be Surfa vide a masure of the bulk tensile nt used (⁻i.e. shear box, torvane, nost_available instruments provide terize the surface with of the surface. In the liameter) was used in

> ich site in New Mexico ken near the centre of xture than the clay-rich carried out using a sand range of shear velocities.

site , 16.6 mPa for the siltier ntr d soils. As concred, measured dust 1 n a m-2 sand the heavily crusted found at the silt-r ich playa margin site indicating the importance of textural characteristics for dust elations were found between emission pit e the fact that this technique to that assoc iated with the unction with recent studie s appropriate parameter to

wo playa sites were negatively

er

Keywords: dust



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This paper presents results on the mineralegical and samples from Africa, China, and the Middle ast. modeling of visibility impairment and radiati Globally, the most abundant g roup of ambie of silicate and oxide minerals. Wind-blo wn comprise one of the hig hest atmospheric m aerosol). Large emissions of mineral dust f input to the earth. Approxima tely one half of all of

ion concentrations, opt ical properties of pr ime importance in radiative transfer processes and global climate change issues depend on the com mineral abundances are needed. To model use of satellit e imaging techniques) of ent needed. These properties are a function of distribution. The refractive index is not just the dust, but strongly depends on its miner alogy. The De

roperties al ı erties ical p er as indut fo sol particles t rom <u>de serts</u> a ings d tic ca and ina origina

suspende d fine desert dust dust are needed for t he mine obal and regional climate models. ass (approximately 50%) consists nth<u>ropogen</u>ically disturbed areas n a pl ary scale (after sea salt stantia modify the solar energy e Sahara, making it the

most important source region to be investigated. While common filter analysis provides elemental and



components, therefore e effects (including the orption properties are particle size and sha pe elemental composition of sert Research Institute (DRI) developed the

necessary instrumental capabilities for aeros ol sampling as well as mineralogical and real- time optical passed through PM2.5 size characterization of dust aerosols. Dust is selective impactors, and opt ical absorption are fi ed the DRI photoacousti c ðuar detacu absorption instrument and the DRI cavity extinction instrument. dd ĥce en Filter samples are simultaneously collected for mineralogical analysis by X-ray diffraction. The resulting XRD patterns are quantitatively analyzed using Rietveld p e perimental procedure ocedure. Th allows for rel ating aerosol absorption, extinction and since to mineralogical seattering a edo composition.

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Various airborne and ship based studies ow dust and pollution aerosol from near its so urce to la entrained into the marine boundary layer (M species such as sulfate and nitrate during pa into the free troposphere near Japan. At tim subsidence in high pressure r egions that bro the MBL. S hipboard measurements and lida activity, associated with the passage of log inversion. High temperature volatilizati on d

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sulfates, nitrates, carbon an d sea-salt to leave only dust sized by an optical particle shipboard data and concurre nt chemical measurements revealed the relation between entrainment of pollution and dust into the MBL associated of low pressure systems also revealed scav ocean surface. This process appears to be mechanism for supplying the ocean surfa Measurements in the free troposphere and Airborne missio ns flo wn no rth o f Hawaii dun characterized the vertical structure of subsidin

dust/pollution layers and structure using the air-mass entraining into the marine bo unda distribution, chemistry, optical properties a humidity [f(RH)] were able to characterize and discriminate between M BL air, FT Asian dust/pollution and an external mixture of the two airtransported from Asia to the Pacific Ocean north on

alt.). In situ measurements confirmed that regional and global chemical transport models (CTM) successfully predicted several Asian dust/p is put into br oader context using models Observatory. We include a discussion of potentially important source of iron to the

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ons up to 10

ookm downwind where it was up to half of the soluble umula egions in Asia bef ore being lofted oposphere encountered regions of and pollution down to the top of ainment and convective tra nsport through the used to remove most counter. These

years have allowed us t o measure Asian

. Subsequent passage gh pre cipitation to the t over the Pacific and a h to the ocean surfac e. ects of these processes.

NASA PEM-Tr opics and IMPEX missions g dust and pollution. In -flight mapping of the

show a sloping, subsiding Asian lasun ents of the aero sol size atter s a function of relative пg Justy air-masses were entrainme here (FT) and were even e free tropo evident as a pronounced event in data at the NOAA Mauna Loa Observatory (19.54N, 155.58W, 3,397m

d subsidence into the MBL. One episode and data from the Mau na Loa ce due to wet-deposition, a Cacific Subtropical Gyre.

Keywords



















An intense cyclone in the northern Sahar incorporated into a slopin g rising cloud laye that 2007. The cloud band was observed to glacite -11C isotherm level, as dep icted by two ir e METEOSAT Second Generation (MSG) geos observations. The glaciation front kept prod sloping ascent. This pro vided a continuous (every 15 minutes) during several ho urs. Higher

glaciated at temperatures ranging between -31C and -35C. Animation of the progress of the glaciation fronts within the clouds will be shown in the pres measurements of the glaciation temperatur way to meas ure the glaci ation tempera measurements of the CALI PSO overpass s being at level s that should have reached havereached the higher clo uds that remain glaciate also at -11C, but the glaciation progressed over a range of several degrees and not as abruptly

reated se dust vards iraled bst instar lane hdent microp / sat<u>ellite.</u> upwin on of eď ob İmilar

that asce nded and visibly cyclo center on 22 February y as the loud top rose above the cal multispectral schemes of the rob orated by AVHRR and MODIS nts of cloud band made their ation every new MSG image s above the dust layer

entation. This time sequen ce of geostation ary based vides us with a new

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sol conditions. Lidar al infere ncesof the dus t at -11C, but could no t vective clouds started to

as in the sloping layer clouds. The cloud drops effective radii just above cloud base were larger and increased faster with height for the conve ar by loping layer clouds. This suggests that larger dust particles rise with fee d convective clouds. Yet, the onset of glaciation occu rred for both ture, revealing the ic e d t s a' me ipe nucleation temperature o f - 11C for that desert dust . This is in agreem ent with previous laboratory measurements for ice nucleation activity of deservedust. TRMN Precipitation Redar measurements over with additional insights to these clouds (not yet analyzed at the time of the st mission) 🛚 🖉 rovide u the role of the dust in the formation of precipitation.

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Dust particles ability to scatter and absorb radiation a lead to an important radiative impact on th clin various regio ns can have di fferent so lubil interact with cloud development in a variety of interaction of several types of dust aerosol, with water vapor at a wide range of tempera determined using a humidified tandem diffe 95% and cloud condensation nuclei (CCN) activity

temperatures was then compared to ice nucleatio diffusion cham ber (CFDC). The CFDC was measure the i ce nucleation ability of the d above water s aturation. In all experiment quasi-monodisperse particle stream, gen samples showed no hygroscopic gr owth at

supersaturations than are predicted for insoluble, wetta bie particles. Due to the larger size distribution for dust aerosol than the background aerosol, the potential for the dust particles to serve as giant CCN is important in determining the role of this the dust showed little dependence on tem particles at a similar RH for any temperature solution particles; however none of the sam ples solution particles; however none of the sam ples 30C. The particles which initiated the ice phase at t

freezing for use in cloud-resolving models.



teract with water vapor may tly. Further, dusts from finð cal and surface properties, and v understood. We investigated the west_U.S. and the Saharan region, copic th of the particles was relat humidities (RH) from 5 nuous flow CCN counter

at supersaturations up to 2 %, both at a temperat ure of ~30C. The interaction wi th water at warm



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n behavio r, examined us ing a continuous flow of -6 0<T<-20oC to ion to several perce nt to sampling, yielding a nm. While most

00 erve as CCN at lowe ation_Ice formation behavior of -400 nucleating ice on 1% of

article size. Initiation of ut sti pen enceron the ice phase was at RH values significantly lower than required for homogeneous freezing of aqueous operatures warmer than ~nation at te bear to be those with the lowest conditions a smallest content of soluble material. This was verified by coating the most IN active dust sample with secondary organic compounds, which required substantial increase in RH for ice phase initiation. The collected ice n ucleation data was then us y parameterization of heterogeneous

> Keywords: mineral dus microphysics





In the open o cean, dust de position undoubtedly pr surface ocean waters. Associated with dust partie such as iron. Via dissolution processes, som of During transp ort away from the dust source photochemistry and possibly organic complete particles modify the dust c hemical and mine solubility increases with distance from the s the atmosphere is critical to understanding the fra

n import ral el are s these e le hen atmo spher coup<u>led to s</u>e i l con tio reo it a that i

y of micronutrients to bgeoch emical interest nts o come bicavailable for the biota. bud processes o facid cycles, ve losses of coarser mode aeroso 1 l load is, it ap pears that iron dox speciation of Fe in rs th in surface waters after

deposition and consequently has implications for the bioavailability of this atmospherically derived Fe. The impact of tran sport process on the <u>chemical speciation</u> of iron from du st has been study from laboratory experiments on dust collected in during atmospheric transport also induces content. An extraction proced ure was perf aerosol sample collection and measuremer preconcentration on C18 phase colu mn. E greater solubility of iron resulting from clays (~4



. The size segregation osition in favor to clav present in dust and on ine method coupled with (h ydr)oxide iron show a h of the others minerals (





hundred times the concentration on clean d recorded over the 49 year period 1958/9-20 visibility of les s than 5000 m eter and the p comprise 18000 days of obse rvation. From than 10 micron (PM10) were measured by Ministry for Protection of the Environment. The red

duration and concentration. This dat a may help explain the meteorological and ph ysical processes involved in the Dust phenomena. Results : 1) The number of days with DE in Tel-Aviv per year has increased 1958-2006 with a s lope of 2.3 d Automatic continuous measur ements are a Duration of DE was 3 to 72 hours, most begin in Tel-Aviv between 5:00 and 8:00 in between 8:00 and 15:00 hours. 4) 13%-4 DE in Tel-Aviv over a year was 172 to 495 nours,

servation and 6 in Tel-Aviv e of <u>vellow-o</u>r 2006 ier Telby analv

asurements of mineral dust were Jerusalem. Dust was identified by desert du st in rain-water. These ons of osols with diameter les s ctric Company and the stics su ch as beginning,

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to 27 days in 2006. od our results are : 2) st DE were observed to the inversion, and ended . 5) The total duration of

caused PM10 values to exceed the Israeli standard of 60 mgr/m3/year in 6 of the 12 years. Using our DE identificati on algorithm we computed without DE. However, even without the DE contribution PM10 exceeded nended W (20 mgr/m3/year) ever y ar) even without DE. 7) year. PM2.5 values were over the Israeli red me Ē15 57 The contribution of DE to the an nual PM10 average was 9%-27%. These results may indicate the effect of global or local change, such as desertification, on dus he dust could in turn affect intrusions. weather and climate through direct (redistribution) heat in e atmo sphe and indir ect (cloud and precipitation) effects. DE statistics are relev ant for transportation eg. th rough visibility, and for publi c health concerns.

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Pedro Campos 1, SERGIO R ODRGUEZ 2-1, Em Observatory, Instituto Nacional de Meteorolo a, Canary Islands Spain. 2. University o f Huel d'Optique Atmosphrique, Observatoire de Re he des Sciences et technologies de Lille In this tudy column-integrated and in- situ aerosol measurement

Global GAW (Global tmospheric Watch) station mixing layer and representative of the Subtropical North Atlantic free troposphere. The objectives are :

1) to characterize aerosol phy sical propert properties ove r Izaa - free troposphere measurements. The study w as performed lin 1 concentrations reaches high er concentrations column-integrated aerosols measurements (Aer

concentrations of particles le ss than 10um were performed by an OPC - Optical Partic concentrations by experimentally determined first an intercomparison between the two c olumn-integrated aerosol datasets was performed. In the second part a detailed time s eries study of the intersitu aeroso concentration

column-integrated aerosol load (Beta) and in-site (FP1/PM10meteorological charts and maps of the Aer the AOD measurements performed with b good agreement (r2 between 0.93 and 0.9 was lower because a signal overestimation

aerosol mass and volume concentrations e the study allowed us to understand the ori summer at Izaa. Day- to-day variations in t the Beta vertically integrated aerosol load ratio) and column integrated (Alfa) fine-to-total aerosols fraction exhibited highly correlated day-to-day

variations which, in turn, showed a high anti-correlation with the PMx and Beta aerosol concentrations. The analysis of these time series, together the back trajectories, the meteorological charts and the AI-TOMS, shows that the large variations in th meteorological scenarios prompting Southe mineral dust with strong NW subsiding a predominant fine size. The concatenation of ug/m3 in PM10 concentrations and from ≤

e Golo up3 1. Iz aa Atmospheric ta, 38071, Santa Cruz de Tenerife, gy, Huelva, Spain. 3. Laboratoire PHOT / AERONET", Universit ts of tercomparison between nospheric Observatory

(28.47N, 16.24W; 236 7m a.s. l.), located above t he



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of the aerosol physical inte rated with the in-situ), because the aerosol y in this period. The

hd e Alfa and Beta ngstrom parameters) were performed with two equipments: an AERONET/PHOTONS Cimel sunphotometer and a Precision Filter Radiometer-PFR (World Ra diation Centre). The in-situ aerosol measurements (mass

> PM_5 and PM1, respectively) volume to aerosol mass aero udy is chided in two main parts:

PM10, PM2.5 and P M1), and colu n-integrated fine aerosols fraction (Alfa), was performed. Owing to the lack of local aerosol sources, day-to-day variations of these aerosol parameters were int erpreted in terms of lo_ng-range transport by using air-ba ck trajectories, In the first part, the study showed that sunphotometers exhibits a ver y for the 870nm channel, which peters. Moreover, the in-situ

well. The second part of erosol c oncentrations in sol concentrations and the in-situ (PM1/PM10

he c oncatenation of unts of coarse Sahara n I co ncentrations with a

creases from <10 to >125 ases from 0.45 down to 0.15









Photochemical processes are affected by the atmosphere in various ways. On e is the alt absorb and sc atter the incoming solar rad modification of the j-values (photo lysis rate particles surfa ce that can lead to reduce photochemical reaction, the result is a sh ift the main focus will be on the quantitative a processes, especially on ozone, sulfur dioxide, part

For this purpose, advanced atmospheric and photochemical models are implemented with the aid of air pollutant measurements from stations in the region. The models used are the RAMS atmospheric model, the SKIRON/Eta atmospheric modeling syst CAMx photochemical model. New model de formation in the air quality model. Results the available measurements and h ave als photochemical processes in the M editer ane particulate sulfate, depending on the atmospher

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le conc entrations in the luxes b ecause particles ctions and the n affected by the uptake of gas molecules into the wh <u>en that</u> gas t akes part in a products. In this work t has on photochemical e nitrate concentrations.

> ust module and t he o new aerosol ina t sonable agreement with

t dust particles to the levels an d in creasing













Matthias Tesche, Albert Ansmann mthaler, Dietrich Althausen, Birgit Heese, Matthias Wiegner, Michael E Petzold, Markus born, Andr Garhammer, Ina Mattis, hluca isan bovik leg

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The Saharan Mineral Dust Experiment is a si of characterizing optical, physical, chemical region and in the outflow re gime. The data ĸе aerosol retrieval schemes and atmospheric t ansi the CALIPSO (Cloud-Aerosol Lidar and Intrared Pa

contribution we discuss results on optical and microphysical dust properties obtained with three groundbased Raman lidar systems. The systems_measured during the first intensive field phase of SAMUM which was carried out in May/June 2006 in Ouarzazate (30.93 N, 6.9 W). A high-spec the German Aerospace Agency (DLR), sour to the field site at Ouar zazate to a second place of intensive ground-based in-situ ob properties wer e investigated with aerosol coefficients and particle extinc tion-to-back particle backscatter coefficients at several systems provide linear depolarization ratios

1064 nm. In addition, sun photometer ob instruments belongs to the A erosol Robotic Netwo

measurement days. The plumes reached t op heigh 0.4 in the wavelength range from 380-1640 nm. Preliminary analysis shows particle depolarization ratios around 30% at visible wavelengths, which

dust transport events to Ce ntral Europe. transport event carried dust from Northwest Raman lidar stations of the European Aero transport event. This work is supported by

al German institutes with the goal s of <u>Sahara</u> n dust at the s ource ta to validate satellite nd trut on on st properties will support Satell vation) mission. In th is

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ems were stationed at ch aircraft, operated by ern coastline of Morocco W), the latter being the t vertically resolved dust

lidar sy stems stationed in one of the source regions of Saharan mineral dust. Altogether, the lidar systems provide information on particle volume extinction

> a nd 32 nm wavelength, and 64 nm. The four lidar 3 a n ngths between 355 and em ht wave servations were carried out at Ouarzaz ate. One of the

eved on 25 out of the 28 k. Dust plu hes were ob of up to mab ove a level. Optical depth at 500 nm wavelength was as high as 0.9. Daily co lumn-mean Angstrm exponents avera ged between 0-

> is usually observed during long-range SAMUM field phas e a long-rang e urope, and Italy and Greece. erved the plume during that haft DFG.

Keyword





























performed in order to investigate particular events, like dust intrusions, volcanic eruptions and forest fires. Particular attention is devoted within EARLINET to Saharan dust intrusions in the European region:

an alerting system was established in order and monitoring the distribution of dust particles and their optical properties ı, be se of the short distance between our s ite and the Sahara desert, a lus intesion every 10 da ys is ut arar observed. Dust particles are observed between 1.8 and 9 km above the sea level and the optical depth of the dust layer is on average 0.13 reaching a maximum value of 0.68. The large number of observed dust cases allows a climatological analysis of Salar eions in ter n dust in o f optical properties as well as the id entification of some peculiar cases for whi ch a more det ailed analysis is carried out . ACKNOWLEDGMENTS The financial support of this work by the Euro pean Commission under grant RICA-025991 is gratefully acknowledged.










































Even though studies of Arctic ice forming particle from open lea ds could be im portant for cloud f polar marine psychrophiles or viruses has no b in the atmosphere. Here, we examined the i isolates that were representatives for most of and a polar C olwellia lysoge n phage virus solutions containing bacterial cells or virus part

determine the fraction of f rozen droplets at a depolarized light scattering intensity from the drople_ts in free-fall. Our tests revealed t hat all sea- ice isolates and t he virus nucle ated ice at te temperature for the nucleation medium (droplets were frozen) for artificial seawater might prove important for ice nucleation pr loc psychro-active bacteria and viruses are not ice or polar clouds.

at ion in the examined un ation acti own d ns Hig oluti ce n a free

I or viral source derived bact c, the ice nucleation potential for onditions closely resembling those INA) of several sea- ice bacteria ntarctic s ea-ice bacteria ctic ar ctra were o btained for atio n g tube technique to

particular droplet temperature by measuring the

nogeneous nucleatio n e at which 50% of the fived from other sources far indic ate that mari ne cleation processes in sea

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This study explores the applicability of molecular biogenic aerosol (PBA) particles in the atmos here suspended par ticulates (TSP) have been col cte sampling locations. From filte r aliquots load could be extracted and genet ic sequences d Molecular techniques (e.g., DNA sequ encing organisms, and to estimate diversities and blank and background samples showed that filter i

genetic analyses. Mass fractions of DNA in <u>PM2.5 we re found to be aro und ~0.05 %</u> in all sampled locations. The average concentration of DN indicating that human adults may inhale a haploid human genomes). Most of the bac lte and some from Actinobacteria and Firmicut and Basidiomycotas, which are known to a sequences could be attributed to green plants an one unicellular eukaryote (protist). Over 80% of

about 40% of the 19 T-RF peaks (58 to 49 The results de monstrate that the T- RFLP sequence an alysis. Shan non-Weaver indices that the bacterial diversity differs among sampling locations

characterizat ion of primary natter (PM2.5) and total cicula es of filte materia Is at German am of air particulate matter, DNA acteria, fungi, plants and animals. etermi he identi ty of biological hisms. Investigatio ns o f nicro d hinated prior to use, and

that the sampling and handli ng procedures have to be carefully controlled to avoid a rtifacts in the



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e order of ~7 ng m-3, corresponding to ~105 re from Proteobacteria acteristic for Ascomycota atmosphe re. The pla nt d moss spores, while animar DNA was found only for

the 53 bacterial sequences could be matched with the investigated PM2.5 samples.

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erial diversity than the d T-RFLP data indicate

Keywords: da, flp, micro reanisms

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Key to improving the predicta bility of earth system be is an improved understanding of the couplin multi-scale modeling framework. Credible prediq observations and process studies that explicit cycles at multiple temporal and spatial scal landscapes where biogeoche mical cycles ar drying associated with climate chan ge. Wa and include some of the f astest growing population

project, called BEACHON (Bio-hydro-atm osphere in teractions of Energy, Aerosols, Carbon, H2O, Organics & Nitrogen) has recently been ini study will be the linkage bet ween biogenic and field studies are proposed ed to address reported, including plans for a long term field

Keywd

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ale of months to a decade peochemical cycles in a 'a nɗ les require coordinated modeling, vater, energy and biogeochemical ticul<u>arly important</u> for semi-arid ut rie and are threatened by the earths land surfac e r hal aerosols emitted by a

water-limited biosphere are key points of interaction between the atmosphere and biosphere. A new

A key element of this cycles. Both laboratory to date will b e gress







In the mid-1960s Gabor Vali refined the dree for guick and uneguivocal detection of act temperature range. His early work suggested to -6 C it contained. But it took a year of coming from the organic component of the leaves were the probable source of the ice In an attempt to follow what was thought w leaves as they decayed, to our great surprise ice no

slurry. In an as yet un explained action, the author put t hese samples into frozen storage and never thought to test them again for some living entity. It was one year later that a repeat of the earlier leaf decay experiment led to the identification These nuclei were named bacteria derived year project, took four. That was 30+ yea (assuming they do) is still unclear. Also atmospheric ice nucleation pro cesses eve the address some of these unknowns? It would seem th nucleus pro tein pro duced b y B DN from its for rmation in the co at of the bacteria thro ugh to its

of dete freezin e ic huclei nd th th le more fortile h beyond that d an<u>other 6 </u>r [hese] j slow uld emio 'at -1.

ctivity ectrum, over a wide bil, the more ice nuclei active at -5 eal ize the active i ce nuclei were hs to determine that well decay ed name f derived nuclei (LDN). lease he ice nuclei from green

oduced in the moist leaf

freezing nuclei that allowed

cing active ice nuclei. buld have been a tw o w the BDN become LDN ole of BDN and LDN in hportant. How then to at one could trace a genetic marker for the ice

phere and in ice nucleation sites

appearance in LDN, and by extension look in young ice crystals collected within clouds



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plants, soils and ocean surfaces, these ice-n cleat even in the polar-regions. These ice-nuclea ng and have been observed in clouds and hails on any other ice nuclei in ice formation of clouds crystals exceeding background ice nuclei we co were caused mainly through collisions of grappels graupels is responsible for ice multiplication process

highly depends on the initiation process of ice cumulus clouds). Thus, we hypothesize the ice and the subse quent ice multiplication proces resolved microphysics was d eveloped to in reice nucleation process by bacteria was simulat formation was determined. On one important improve estimation accuracy of radiation bidg ent ice anclei at temp attach bacte, were iteria are eadi bacteria thus s emperatures obser in t is th close roph cess order watable c

and in most all climate regions, disseminated into t he atmosphere d play a more important role than re -10C. High con centration of ice varm-band cumulus clouds, which triming occss). The formation of dition of graupels

highly depends on the initiation process of ice crystals at relative warm temperatures (growing stage of cumulus clouds). Thus, we hypothesize the <u>ice</u>-nuc<u>lea</u>ting bacteria plays a key role in gra upel formation

A 10-French-horestaan gale the illes of the et a d the Placht illes a fit of the global et so leer that the version of the global of the so role in gra upel formation cloud model with binrosols and clouds. The this process and gra upel ect effect, this finding will limate model.

Keywords: bacteria, ice nucleation, modelling





Cimbal Julia, Klose Melanie, C

Aerosol particles are of central importance for atmo health. A significant fraction of the atmosph and fungal spores, pollen, plant and animal sources, prop erties and effects of biological characterized. The use of molecular genetic methods for the analysis of biolog ical aeros with a High Vo lume Sampler separating fin over a period of one year 2006/2007 in Marnz, Ge of fungal, plant, bacterial a sequenced. The sequences were blasted in the National Center for Biotechnology Information databank to find the closest match and determine t originated. Fungal DNA was d etected on d taxonomic diversity for the 161 fungal sequ were characteristic for different groups of sco discharge spores into the atmosphere. Ter ger of them belong to the class Homobasidiom ycetes and some of them include species which can act as plant pathoge ns or huma n allergens (e.g. Stereum) . Few sequences belong to the genus Itersonilia within in the class of the Het erobasidiomy allergenic gen us S porobolomyces was detec sequences within the Ascomycota belong to Epicoccum (a soil and litter fungus and allergen) as well a s to some other genera which include plant pathogen species (e.g. Stemphylium). The plant quences co the families Betul aceae (e.g. birch), Po aceae e.g. (e.g. nettle), Taxaceae (e.g. yew), Cupressaceae (e.g. fir) and moss spores with in the Bryophyta moss superclasses. Bacterial seq uences could be attr <u>ibuted to Proteob</u> acteria, Actin obacteria and Bacteroidetes, which are known to be wid detected on the coarse part icle filters and Archaea and C renarchaeota e nvironmenta represent the third domain of life and are p

are bacteria-like microorganisms which are have the capability of living in almost all k water and sediments.

einrat O., Pschl Ulrich

hemistry ic (fic p bf bio les a hts, etc. o fa rticles in the resolves m air fil cles. I arse p odyna cles same

physics, clim ate and public e.g. bacteria, bacterial cal or bwever, the abun dance, diversity, bsphere have not been well limitations of traditional detection amples were collecte d cut-off diameter 2.5m) alyzed for the presence

nd archaeal D NA. Al I P CR products were clon ed an d several clon es

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h which the DNA had nary results show high samples. The sequences ch are known to actively nycota were found. Eight

eal plant pathog ens. Further the ota s Urediniomycetes. The im / ani portant allergen) and Id be attributed to green plants, within ntagiaceae. g. plantain), Urt icaceae

sphere. Archaeal DNA could also be vsis points to high similarity with Pr o- and Eukary otes, Archaea anism group on Earth. They ronmental conditions and rent kinds of soil, ocean



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Highlights of recent publications (Baker an and cirrus clouds are presented. Wave cloud results in the potential simplicity of wave cl ouds and the co large data set (17 wave cloud missions) is

condensation and glaciation within the sam cycle and may be considered natural laborat evidenced by the replication of earlier result and only if data from well-behaved clouds a width of columnar ice are contradicted. Ne

the onset of side plane growth are derived from the same data set. Examples of copious ice production coincident with the evaporation of the super cooled_ studies (Cooper 1995, Field et al. 2001, He is well do cumented and ne arly ubiquito u observationally well established. However clouds a pri ority of cl oud ph ysics researc understanding the general problem of ice i of cirrus cloud microphysics to wave cloud microphysics but on larger temporal and spatial scales; the

existence of high concentrations of small ice crystals; the application of a large (15000 Km) in situ cirrus data set to global climate model parametering situ observations of the microphysical prope J. Atmos. Sci. 63, 3160318 5. Cooper, W. enhancement during evaporation. Preprints: AMS Conf. On Cloud Physics, January 1995, 147-152. Field, P. R., R. J. Cotton, K. Noone, P. Glantz, P. H. K Reiner, M. An dreae, C. P. R. Saunders, A. Archer, Bandy, D. Johnson, and A. H eymsfield, 2001: Ice nucle ation in orographic wave clouds: Measureme nts made during INTACC. Q. J. R. Meteorol. Soc., 129, 19031927. Heymsfield, A. J., and L. M. Miloshevich, 1993: Homogeneous ice nucleation and su Sci., 50, 2335-2353. Lawson, R. P., B. A. E

microphysical properties of wave, cirrus a 3186-3203. Ono, A., 1969: The shape and Sci, 26, 138-147.

son et al. 2006) on microphysics of wave



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e, E. Hirst,

s natural laboratories. A cloud he observed number of cycles of ave clou ds exhibit on ly on e su ch ud m<u>icrophy</u>sical processes. This is r crystals (Ono 1969), if (1969) on the maximum rosette crystals and for

of both the complexity and

liquid have been r eported in earlier wave cloud hysterious observation s, the phenomenon is aking the study of wave go a lon g way toward s Its include: The similarity

> A., and R. P. Lawson, 2006: In hvil c ds. Part 1: Wave Clouds. s in clouds: Observed

tor, M. Smill B. Brooks, C. Hoell, B. T. Ch ou er in orog raphic wave clouds. J. Atmos. 2006: In situ obs ervations of the Clouds. J. of Atmos. Sci., 63, in na tural clouds. J. Atmos.

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Keywords: wa



This study pr ovides a theoretical evidence for th nucleation processes, which is one of the michan 1997). Recent observations in clouds and la crystals are a dominant habit especially in t are known to have large impacts on the ene polycrystals with spatial structure can be i developed a Spectral Habit Ice Prediction Sy particles (habits and types) based on t her histo

side planes and irregular polycrystals) based on the habit frequency dat a constructed from laboratory experiments conducted by Bailey and Halle sedimentation and mixing can change the Two-dimensional simulations with different nucleation rates were co nducted for a wi ter stratiform cloud from FIRE A CE. The habit and discussed with available observations of had

ation of poly ed in pro indicate that le and upper lget <u>of the e</u>a nt in t I PS)

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tals through the freezing (Pruppacher and Klett, litera vcrystals ncluding rosette-shaped Is $(T < -20\overline{C})$. Clouds in this region hrough radiative for cing. Also, the ation p ess. The authors have ates the evolution of ic e h a n Eul erian dy namic

framework. SHIPS nucleates either monocry stals (hex agonal crystals) or polycrystals (bullet rosettes,

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ophysical processe s, at a particular location. ributions , and freezing ring I MPROVE-2 and a lations will be compared







During last ten years there has been an ongoing deb ice particles (5-50um) in ice clouds. Small ic the evolution of ice clouds, f ormation of mi transfer in clouds. In-situ observations indic few per cubic centimeter, independent of t arisen due to the fact that small ice particles precipitation and ice sublimation, where small Recent studies suggest that small ice particles may

t the pre 00 part ireat s hā bd e, and lav the average ture <u>in the r</u>a ally o ll ice d rtic

and concentration of small n the understanding of ortañ a si gnii cant impact on radiative entration of small ice particles is a ro m - 5C to -40C. Confusion has s, including zones of ice ions are not expected. oncer of larg er ice partic les

impacting the inlets of aircraft instruments, the fragments of which may bounce into the sample volume of the probes. These studies are balanced by work sug gesting that instruments like FSSP are not

affected by shattered ice if ice particles are problem Envir onment Canad a has develop coefficient of clouds. The pro be consists o tł and therefore any shattering at the edges other instruments with much smaller samp Convair 580 and flown during the CloudSat



o attempt to solve this hents of the extinction ted a few meters apart, b be negligible relative to as deploy ed on the NRC

s of these measurements

The extinction coefficient measured by the transmissiometer was compared to that calculated from size distribution measured by PMS FSSP, OAP-2 have permitted some conclusions about the be presented at the Ice Microphysics Session



























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Philip J. Rasch, Dorothea C.

The concentrations of small crystals in ice clouds have (1) difficulties in measuring these crystal si ls an high concentrations (typically 500-5000 L-1) (SD) of these small crystals include the FSS not clear what percentage of the small cryst the instrument inlet. We have aske d the q impact the performance of Global Climate M a realistic treatment of ice particle shapes, the SD

here was not designed with this qu estion in mind, but the results do provide some insight on what the answer might be. In situ FSSP/2DC measur in mid-latitude cirrus differs appreciably fr these measurements have been incorporation part of the Community Climate Systems scheme are bimodal, with crystals having

small mode in the mid-latitude SD scheme study was how does the differing temperate impact GCM simulations? Another guestion e for mid-latitude cirrus or tropical anvil cirrus mid-latitude cirrus predicted from cloud temperature treatment of Mitchell and Heymsfield (2005) for acc

modes were based on CPI d ata. Using rea rates and the cloud life cycle were better r CAM simu lations were performed u sing d scheme. In the tropical SD si mulation, the anvil cirrus was strongly gov erned by the anvil cirrus, the small crystals have a stron water path (IWP), cl oud lifetime and cloud direct differences in SD bimodality, dramat

the tropical SD scheme due to greater temperatures in the upper tropical troposp latitude SD scheme. These results suggest region may be primarily due to an inadequ their fall velocities) in tropic al anvil cloud particles affect cirrus radiative properties. A far, Timo Nousiainen

ngth D < 60 emains controversial due to abilita theory to explain thei r existi d to measure the size distribution the $2D\overline{S}$, and others. It is often shat<u>tering of larger ice particles at</u> are real, how will they crys this q ion, the GCM must have ice particle fall velocities

(i.e. SD mass sedimentation rates) and ice cloud radiative properties. The GCM experiment described



dependence of the SD s. Parameterizations of Model (CAM) at NCAR, d the mid-latitude SD de. The amplitude of the

mode of the tropical SD scheme increases with decr easing temperature. The question evaluated in this mode in these two SD schemes s it mat r wh er one uses a SD scheme lula f cli ate treatment of ice clouds 11 in the CAM was modified by implementing the following schemes: (1) SD schemes for tropical anvil and tent; (2) the fall velocity e a nd ī water c rate ic edimentatio rates; (3) the Modified Anomalous Diffraction Approximation (MADA; parameterized for GCM use) for accurate treatment of ice cloud radiative properties. I ce crystal shape recipes_ representative of the small and large particle SD le shape information, ice sedimentation

aud radiative properties. One-year and only the mid-latitude SD per regions (T < -50 oC) of r higher concentrations in rates and hence the ice factors, along with the d long wave (LW) TOA

cloud forcing in the tropics (up to -26 and +20 W/m2, respectively, for annual zonal mean) relative to simulations using the mid-lati tude SD scheme. Moreover, SW and LW heating rates were greater using IWP and SD projected area, respectively. This resulted in

edicted using the midby some GCMs in this e. small ice crystals and ide clues to how aerosol irst affect the small mode of





In mixed-phase clouds it is possible to des ice water cont ent and liquid water content The st precipitation formation, climate, and the radiative remote sensing, satellite ret rievals and cloud e microphysics is considered an important top liquid water droplets evaporate to keep the ice crystals grow by vapour deposition. Thu Wegener-Bergeron Fin deisen mech anism. water droplets are determined by the vapour dens

regions, which can also mak e a significant contribution to the diffusional vapour field. We solved the diffusion equation to calculate the vapo u cloud. The real problem was simplified by particles and cloud droplets are represente positions of the particles are fixed in space

equation is reduced to Laplaces equation

takes the ice s aturation value at the surface of surface of t he droplets; (2) vapour de nsity takes t he water saturation value at infinity. The results obtained allow analysis of the influence on

as a function of the proportion between

phase af fects the rate of in numerical modelling, nport ese reasons mixe d-phase cloud s. In some atmospheric conditions pect to liquid water and the sm all e of water by the so-called tion of ic e crystals an d ev ap es, which in turn may

depend on both the configuration of the neighbouring particles and the particles placed at distant



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on of a mixed- phase y state, (b) both ice ize a nd shape; (b) the this case the diffusio n are: (1) vapo ur density

and accrystal in the presence of

ensit neighbouring crystals. Also, this model allow mong the cloud droplets and ice particles and to dete rmine the resu ng the p articles for the de ity/ q атп configuration of a particular mixed- phase cloud. The ambient vapour density is para meterized as a function of the variables of the system, namely number Dncentratio **T**cloud droplets and ice crystals.









shock wave of electrical disc harge in the cloud. The electrical discharge is assumed to occur in the odd frozen drops, as a result of the dynamic stree of discharge. The second mechanism of the clack in between ice particles and water drops. Several different range of temperatures, liquid water count, lightning channel, values of drop charge, and the fast - almost immediate, freezing of one payortion of spectra distribution.

The checiation process of innert with: The case of of the sinic rules g e in is contact for ing, fr erical similar ons ha it, aveined downadius, e of the nargo area. T

cess of su percooled drops du e to at process is the initiation of i ce, es g enerated from the lightning ing, freezing c aused by collision s have been car ried out for a radius, and al energy density of the area. The model results show very the test op and c hanges in drop







The present experiment aims at investigating the ice of various insoluble aerosol species. Droplets with dry aerosol p articles at warm temperatures subsequent fre ezing of the c loud dro plets simulates the atmospheric pathway in a mix the immersion mode. Activati on of aerosol particle counter which activates any kind of tube. The experiments on ice nucleating ability in

g abilitie ierse roso tivation of for pre cise k se cl<u>oud fro</u>rr is doi ĺth due bart le v sion m

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nmersion freezing mode ticles e produced by activating nodispense aerosol particles and edge of the droplet content and aerosol particles to ice crystals in odified er-based condensation igh s saturati on in its growth refore not restricted to

section which cools them to a desired experimental temperature down to approximately -40 C. Droplet freezing is initiated in a par allel-plate con Chamber). The conditions of this chamber with respect to water. This prevents evapo depolarisation detector distinguishes betwee at the end of the chamber. The setup for t



Zurich Ice Nucleation e exposed to saturation ro plet temperature. A for a given temperature

Keywords: ice nucleation, immersion freezing, droplet activation

particles which act as CCN. The droplets with a size of roughly 5 micr ometers are led into a cooling













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In anal yzing weather observational data definitions, and interpretation of freezing precip definitions of weather phenomena by differe Enat to be some what inconsistent, or these different char acteristics. Su ch problems phenomena code groups (WMO Doc. No. 30 snow group, fog group, driz zle group, usin negative temperatures in cold seasons, these glaze and rime or other types of solid hydrometeor

observational r esults obtained in different countries differences between weather phenomena d_efinitions and interpretation: which of the phenomena is actually recorded using a certain code a phenomenon in a given geographic al regio how practically important some of its chara in Russia, the term freezing rain is primari supercooled rain or drizzle, and only secon

networks and longest railways. This interpret and the way of handling climato logical data daily reports on freezing rain represented a conventional in Russia.

some inconsistency in terminolo gy, phono mena lifferent countries. Simila r ore closely, either prove wh iewed d differently or descried, using ut hor analyzed WMO weather ed in aviation, such as rain group, n etwo of the former US SR. At azardous events such as recor ice pellets. For analyzed ns, sr

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to be compatible , it is important to c larify the ly as signed to the he phenomenon is and mer USSR and presently formed by precipitating The characteristics of ice

crust such as thickness (diameter), weight per unit le ngth, ice persistenc e on an object (e.g. electr ic wire) are considered the most important ones for a cold country such as Russia wit hits vast wire bservations at weather station s nthl eorol al Tables, which include

ion 211 res. Als o discussed in Epo per phenomena, their definitions and interpretation this paper are other WMO code groups of weather

Keywords: glaze, rime, terminology





Cirrus clouds are important modulators of Earth's r clouds can contribute to a net warming of the Ea cirrus clouds a re treated as homogeneous internal structure or "patchiness" of cirrus cl additional radiative effects. I n recent measu neutral mixing layers within the ice supersat November 2004). These layers were probab first simulations with a cloud resolving morel it wa

to 2 m/s). This leads to ice cr ystal formation by homogeneous nucleation and to discontinuities in the ice crystal number densities within the su represents a sort of patchiness. This effect 2D/3D anelast ic, no n-hydrostatic mo del microphysics, including nucleation (homo and sedimentation, is used for idealized 2D km, which is lifted by a syno ptic scale vertical verocity



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soumed that (thin) cirrus ly, in large scale models h. Uš en from surface o bservation the ese inhomogeneities could lead to bresence of slightly stable or eve n duri ng CIRRUS II, r the jet regions. Fr om elease from growing ice

> of magnitudes, which hsitivity studies. The two moment bulk ice sion grow th/evaporation er is prescribed at 10-1 2

(1-10 cm/s). Within the supersat uration layer a "mixing" layer with a thermal stabilty different to the stable environmental profile is prescribed. The eakly ansta ble. The simulatio ns stability of this layer ranges between weat 50 g were carried out using high spatial (dx = and tem poral (dt ~1s) nts within these cells is resolutions. The formation and evolution of atio Inu nve C ar те investigated. Additionally, the effect of aerosols is studied: Using inho mogeneous distributions of tion of ice crystals on the heterogeneous ice nuclei we investigate the effector f hetero neous nucle following homogeneous nucleation event. This yack mracts on t dynamics which triggers o vari ou the patchiness of cirrus clouds.



IUGG XXIV General Assembly 13, 2007 July

Perugia, Italy

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Symposium Theoretical advances in atmos

Convener : Dr. Eyal Heifetz Co-Convener : Dr. Nili Harnik

This session will focus on the applic ation of the atmospheric circulation, with an emph in the troposphere and stratos phere. Conti and cyclone growth and evolution, eddy-m from it, the dynamics of storm tracks and it streams the dynamics of the stratospheric polar vorte x, and its interaction with the troposphere

teorology and Atmospheric Sciences

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to the understanding of

mospheric circulations, both

a range of topics including wave

and possible dynamical r egimes arising

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particular weather prediction system is example conservation of the PV differ ence field sign es flows PV-gradient; misrepresentation of dial ati field at the verification time. Like wise inverio attribution) account for error of the primary ilo of the char acteristic structure and evolution of inadequate development of the Rossby wale amplitrajectories calculated from the dominant regions of

and contemporaneous a and corpretention a V-p jentropic diversion of the irror frictional process, and /or et a particular course of the error riable bottone in-silonid he Polifere field esh

no craneous anal ysis fields of a from a V-perspective. Noni of the rror across an ambien t ses, and /or error in the analysis are of the error PV-field can (vi a e in-site and far-field. Illustrations field as shown and point to opical oppopause, and backward

trajectories calculated from t he dominant regions of PV-error shed further insight on the nature of the "error".












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A three-box cascade type of energy conver atmospheric general circulation based on wa circulation. Mass-weighted is entropic zonal diagram by Lo renz (1955). The new sc he potential energy PZ with the eddy availab kinetic energy KZ with PE. Mass weights conservation properties and lower boundary KE in order to gain physical insights into energetic

to a three-box cascade type of energy conversion diagram composed of zonal mean available potential energy, PZ, zonal mean kinetic en ergy, KZ and produces PZ and it is converted into KZ th mainly converted to W through wave-mean friction. W is generated not only by the heating and it is rapidly dissipated through the only two d ynamic conversion terms, C climate in bore al winter and s ummer using the NCEP/NCAR reanalysis. It is interesting that C(PZ, KZ) equals to C(KZ,W) under the quasigestr

relationship almost holds in the extratropic subtropics, indicating the importance in no dynamically converted to W and the rest is a is substantially different between the two solstice s. Both C(PZ, KZ) and C(KZ,W) are about 30 percent larger in the boreal winter than those in the bore summer. active in the boreal winter. Stationary w average

do mainly in the so uthern hemispheric winter. It is confirmed that the life time of stati onary waves are quantitatively estimated to be considerably application of the scheme is the diagnosis will be briefly presented. R eferences Iwa mean-flow interaction and Lagrangian m Iwasaki, 2006: A cascade type global energy interactions. J. Atmos. Sci., 63, 32773295.

ha ke differen s no t directh tial e<u>n erav P</u> ole to ຸາຣ. W lond the no

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prose of diagnosing the and D angian-mean meridional energy ycle from the four-box hange the zonal mean available does exchange the zonal mean ostrophic wave e ffects, s no V is d

ed as the sum of PE and tion theorem. This leads

ulation C(PZ,KZ). KZ is st is dissipated through he dia batic differential gram is characterized by ry study i s made of the

er than C(KZ,W) in the

bout 60 percent of KZ is the brme lote that the energetics e s tro ics. iis is beca us ationary waves are very ninic i nst bility waves considerably contribute to energy conversions in the no rthern hemispheric winter, while baroclinic instability waves aroclinic instability waves. An important eliminary result of GCM diagnosis

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gy cycle viewed from wave

58, 3036-3052. Uno an d d on wave-mean-flow

teraction

Keywords: energetics, ise



















Baroclinic interaction is usually considered synoptic eddies while barotropic processes a synoptic eddies give indeed generally their e shows that in specific regions the situation diagnose thes e particular r egions where s reanalysis data. The processes involved will, new measure of the constraint exerted by t introduced foll owing recent results on perturb

quasigeostrophic flows . This new diagnostic is deformation tensor of the large-scale flow. It allows transient eddies can be barotropically rege to apply this diagnostic to reanalysis data, frequency part by applying a temporal filte corresponds to the synoptic s ignal and the two apparently similar zo nal-like weather r place and can act in a very different way on upper-level high-frequency eddles. The second part of our

the anism to sible the d the meth flo the opposite eddies baro arotro simu scale leforn izor in

in the grow th of transie nt heir amplitude. Although Sase č ia barotropic processes, our study r results will first present how t o ically gain energ y using ERA4 0 hode I on the sphere. A n on synoptic eddies is d temporally comple x

called e ffective deformation and is based on the

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us to indicate regions of the atm osphere where tical regions). In order into a high- and a low-The high-frequency part regime. It is sho wn that tical regions at the same

the same zonal-like jets as in the reanaly them synoptic-scale perturbations, the same perturbations. As they evolve close to the ba regenerated. Linear and nonlinear simulations will be compared to each other.

stationary and by adding to ata is recovered for the alvs mporarily barotropically reı

Keywords: barotresic,













Unstable periodic orbits (UP Os) are an im having positive Lyapunov exponents and contracting the

Anosov or Axiom A systems UPOs are de ine or system can be approximated by UPOs with a of such a system can be calculated by set of UR likely do not have Axiom A property. In this atmospheric system approximate its attracto Galerkin approximation for ba rotropic vortig the help of da mped Newton and Gauss-Newton n

characteristics (variance and leading EOFs) can be calculated by UPOs with very high accuracy. Other system properties like Kaplan-Yorke attract can also be reconstructed using UPOs. The attractor and may be important in understa ndi

maotic dissipative syste ms (i.e. systems hase space) ı attra SV accuraty. A Atmospheric : we tr<u>y to un</u>d yst er ter sphere on o ota were

me of chaotic systems like any t rajectory of the result all statistical characteristics ms are dissipative and chaotic but nd to what extent UPOs of simple der consideration is a th T12 truncation. With more than 500 of the

system UPOs. It was shown that average s tate of the system as well as its second order statistical

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Lyapu nov exponents skeleton of the system

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This study presents results from probabili ty density zonally averaged zonal wind from a series examination with the PDFs finds that bimoda if (1) its time scale is very long, (2) it is m the background potential vor ticity gradient i which over most of the para meter space are is not found. These results link the occurren flow and the type of z onal mean flow variability.

bimodality was observed. The power spectral analyses of the PC1 time series found that when the zonal index (poleward zonal mean flow anomaly more closely resembles that for an AR(1) can be better described by a first order or second order ODE. These model data are responsible for the characteristics of the po mean zonal wind.



er pectral analyses of the uation model runs. The nitive cipal component (PC1) time series nal mean/eddy feedback, and (3) e ot her hand, if all three criteria, are n atisfied, then bimodality he pr ties of the background sence of bimodality was

found to depend on the height and the number of mountains, whereas for the double jet state no

nding power spectrum t that the zonal index eward propagation by a the physical processe s thern He misphere zo nal





Understanding the atmo spheric lo w-frequency variability climate studies, climate change detection, a Hemisphere climate features the planetary v Several observations and the oretical argume indicator obey a non- Gaussian statistics a thus characterizing the lo w-frequency po tropospheric jet strength is a critical parame exhibits a uni- or bim odal behavior, and ve dete

results are ob tained by considering the d overlapping period. Our results agree with the nonlinear orographic theory, which explains the statistical non-normality of the low-freq uency variability since the intensity of the jet is related to the tropical and the mid-latitude climate. Data very high-resolution global climate mo dels the overall properties of the Northern He only partially this picture, thus providing some h features.

is of crucial ande tend a releva t ind m to support pre<u>sent a m</u> the c etern

portance in fields suc h as brecas t. The Northern eathe ent of the atmospheric variability. idea that winter p lanetary waves od al probability density function, m. Wè w th at the uppe r anetary waves indicator er t h value of t he jet. These

ata of the N CEP/NCAR and ECMWF reanalyses for t he



bimodality. Moreover, onnection between the tions performed by two h to represent very well pheric variability, match representing large scale













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There has been a long-standing interest to by linearized dynamical equations of atmost herig (1966) showed that the growing and dec ay two-layer and Eady model can be unders propagating Rossby waves (CRWs). Evers generalize Bretherton's CRW-perspective to theoretical CRW-framework exists even for (2006)]. Requirements for the CRW-perspective to

state and the set o f GNMs. A pair o f CRWs can then be constructed unambiguo usly. The CRWperspective comprises an effective, concise_<u>description of</u> the linear evolution of initial conditions that can be constructed by arbitrary superpositi a subspace of all possib le perturbations, project on this subspace. At least since Fa non-CRW part) is an essential part of the s the system are investigated. Since the CR the evolution of initial c ondition described by the GNM and DNM, the aim then is to include th continuous spectrum also in a physically i ntuitive way. One method of doing that is to view any

perturbation as a superposition of P V-build PV-structure at one single level in the vertige boundary potential temperature) everywhere the advantage of the PVB-approach compared to the address, among others, the following que versa? How should we partition a given init

properties of t he system (for instance, grd modified significantly once we include (or spectrum?

Keywords: baroclinic instabilit

inic inst perturbati ons, described complexity. Bretherton gree d DNM) in the quasi-geostrophic s and ns on that or

a

locked and inter acting countern undertaken to systematically dels. At present a solid sphere [Methven et. al. identif y a backgroun d M and DNM cover only

alistic initial conditions ntinuous spectrum (the mal growth properties of hysical understanding of

VBs have a Dirac delta-function

he h ontal) and zero PV (and e ir e f w b oclinic netability of perturbations lse. is understood as the interaction of all PVBs (rather than just the two CRWs) with each other. Considering these two views on baroclinic instability, w e work like to investigate how by are linked. Obviously, ective is the CRW-pe the continuous spectrum is automatically included. The disadvantage is that we have to take into account all interactions (rather than just the interaction of the CRWs). To more cl_early assess the relevance of both views, we will are PVBs in exciting C RWs and vice part and a non-CRW-part? What and structural changes) are sentations of the continuous

ing rossbywave






























As part of the activities of the THORPEX We Interest Group has been set up on Theoretic have participated in the discussions of the I the goals that can be achieved by ensemble degree to which those goals are actually ad is of a fundamentally different nature from (viz. a probabi lity distributi on) has no objective

kina 5up d edic ts of Ensemb Group. The di re ge<u>nerallv</u> j an be °C istic p ctio lete and cai

ynamical Processes, an itv an ediction. Several tens of scientists sions have concentrated mostly on bilistic prediction, and on how the evalu Probabilistic prediction object of the prediction that erved. Two statistical

properties that make the quality of a probabilistic prediction system are reliability and resolution. The

degree of reliability and resolution of a probabilistic. evaluated to any desired accuracy, provide shows that the scores that are routinely prediction systems saturate at ensemble sit clearly elucidated. And such sizes are in an Ca probability dist ributions. Prob abilistic predic tid

prediction_system_can_in_principle be objectively

available. Experience teor ological en semble br that fact has not been or instance, variances of tions, am ong which the

unavoidably (relatively) small size of the verification on sample. P recise evaluation of those limitations would be extremely useful for the design of ensemble prediction systems.



























An approach to the general evaluation of p concept of proper scores. The definition and the implications of propri shown that proper scores allo w for decomposing,

similar to but more general than the well kn sharpness and reliability rela ted terms eme score to be proper, thereby motivating the u encourage 'honesty' among forecasters. Fur system (which can arise for example throug taken into account, which in general lead, to

statement of f orecast perfor mance, forecasters have developed various tools t o investigate forecast quality conditioned on different synoptic situations. One example of such a tool is the reliability diagram. The more d etailed information a reliability variability, diminishing the power of the to be properly acknowledged. To this end, it i diagram, giving the user a hint as to how actually reliable.



represented in detail. It is positive definite terms, corel Brier Score. A decomposition into hese decompositions require the ntependently of any attempt to ariability of the forecast of the ecast system) has to be res give only a general

is presented, which rests entirely on the

the p rice of a higher d and therefore has to added to the reliability even if t he forecast was

Keywords: scores, reliability, sharpness







No matter what computing resources we throw at the will always be probabilistic , constituting an consistent with the information we have av climate forcing s in the future. This is a famil forecasts are routine, but wit h an unfamilia probabilistic weather forecasting system by situations. In a forecast of 50 -year climate our estimate of initial condition uncertaint, by invo

variability, but uncertainty arising from model error, or uncertainty in the simulated response to external forcing, appears to be intractable. Using examples from recent perturbed physics ensemble experiments with simple and complex m odels, we will a is, for a partic ular forecast variable of inte across models and across model-gene ratio those aspects of the models that are not probabilistic climate forecasting that will ex will take time. Patience is important, since an illusion or convergence could be (and, arguably, is being) created artificiallyby incorporating a strong

convergence is uninfo rmative, "papering or current climate that was achi eved by flux responses of a particular clim ate variable, change with successive mod el gene rations, there appears to becomplete for that particular variable have converge d in the tails, corresponding to iowsensitive to details of model formulation.

has converged requires a succession of per models in order to compare the resulting reassessment of the allocation of resour resolution and resolving m ore processes, advent of distributed ("grid") computing m computing resources, leaving traditional reactivities, such as the pursuit of higher spa

n, a fore ble distrik stim e of 1 lał oday and som ituation in we We c<u>an eval</u>u ng iti ce ov fo alida there ۱ly he sp

f long-term climate change sible future climates on of sum ptions about the evolution of r forecasting, where probabilistic and improve the performance of a su ccession of forecast point. We can evaluate odel -simulated internal

> sticclimate forecasting of forec ast likelihoods bservations and varyi ng for a methodology for gnisign that this process

element of prior opin ion in to our forecasts, but such convergence of simulations of that distribution of forecast ate . vations and scenario of ¢lim future forcing, is insensitive to our choice of current clim ate model and (more importantly) ceases to linate m odelling problem ve can cla m that the Natura there are s matte rs of dearee : convergence may be achieved across the bulk of the forecast distribution, but forecasts may still not

probability ou teomes that may be particularly essing whether a forec ast distribution xperiments with a range of climate elihoods. T his may require a focussing on higher sp atial mbles. Fortunately, the ay be achieved with new nal cli mate modelling

Keywords: climate, ensembles, bayesian

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boundary conditions from global EPS system with the presentation of the enclosed le initial conditions from the limited area ensemble system itself, has been allow investigated. The results will be presented at the conference.















simulation length is increased, the ensemb the relative entropy. This c ontrasts with r varies with the ensemble spread. Some of the definition of the climatological distribut

er to the dispersion in the dispersion often ticular choices made for

Keywords: ensemble predictability, relative entropy, signal dispersion

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Gaffney et al. (2005) was also presented. that differences in trajectories between rea was here interpreted to be r elated to the previous two studies depend on the surface eddies. The pr esent analysis on the GFDL-GCM is to tally eulerian and based on the up per level eddy

activities (300 mb). However, a si milar conclusion has been drawn from the analysis of the band pass frequency of energy and momentum for the the momentum and energy of the very hig the reanalysis. The variance of meridional is in the latitude area where the reanalysis shows it to be position ed in the storm track: the model displaces it south of that. There is also a s uggen on that to chieve the c frequency baroclinic eddies, models should have en depends on the lower level circulation of the frontal circulation system. The meso scale circulation

1/40 resolution. It is clear that to adequate the resolution but also to im prove the bou resolution of climate models, this improver sea ice over the sub-polar regions are impo because no sophistication in the moist con to the unresolved dynamics.

this study also showed ch cluster of trajectories ddies of the GCM. The ies of the high frequency

uns, where it is quite clear that p del simulation than in deficiency of the high frequency

Tct intensity of the high

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associated with cyclones could be adequately represented in models with resolution equal or superior to ale, it is necessary to not only improve uxe s. Clearly, at t he present low lowever, if the cloudiness and should be an attainable goal ect those deficiencies due

Keywords: c





The quality of numerical weather prediction has impr remarkable achievement has to be considered predictability and model errors are hig hly conditions may not imply the same improve numerical weather prediction is perceived to events. There is therefore a growing need events, something that cannot be easily info of these events. For these re asons, a study has be

Centre for Medium-Range Weather Forecast (ECMWF) gl obal forecasting system in predicting a specific flow configuration that is believed to be ass region. Despite quantitative predictions of the large-scale flow conducive to major rai This is perhap s surprising since it is a com predictability.

orsiderably l tru con avel bendent l ere ٥v specific con ner i<u>n some s</u> the ca ugh a ed age 'out to

its beginning, however this ns. It known that atmospheric an increase in ski II for averag e s. Moreover the potential value of fic cases. like high-impact weather accur of significant weather ast because of the rarity es, no he skill of the European

ociated with e xtreme precipitation events over the Alpine ing, it was found that an average conditions. vere weather with low

skill

Keywords: predictability, planetary waves

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This study provides a dynamical analysis of the tropical Guineo-Sudanian zone of West A impacts on the local hydrology and human a conditions. The event was a ccompanied by Libya, and a major dust storm in the eastern Medium-Range Weather Fore casts (ECMWF reveal that the event is preceded by a several-day-

of tropical West Africa north of about 10N, with a rather pronounced pressure drop on 19 January 2004. The press ure fall causes a northward shift and intensification of the weak wintertime heat lo w, and allows low-level moist south erlies from the Africa. On 20 J anuary daytime heating trid supported by near-surface convective out evolution a sp ecial form of the pressure integrals of (1) horizontal advection of vir factor, and (3) a diabatic term. While the dynamica calculated from ECMWF analysis data, the third term is e stimated from ECMWF operational forecasts.

Climatological values show a balance subsidence, and radiative cooling. Betwee upper-level low that moves slowly eastward associated with this feature are rather small, bu falling pressure over the re gion. On 19 Janua

Algeria and m erges with the prior system. This tendencies across West Africa that is mainly caused <u>by</u> s ubsidence in its western part and by warm advection in the east. Together with the er and finally enables the unusual precipitatio the ECMWF indicates some skill in predic comparably large influence of the usually w higher predictability than for ordinary sum

Keywords: tropical extratr in

on precipitation event in abunda t rainfall had substantial tting harvests to improved grazing in the arid parts of Algeria and of Europ ean Centre for An al and s tic station observations pressure over large parts



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tha n usual into West northward, most likely amical r easons for this hat consi sts of vertical on times a static stability T contributions (terms T and 2) can be direct ly

tendencies, mainly due t o

al ai is disturbed by a weak this cal pressure tendency anomalies t the formation of clouds and the higher moistu re content on the eastern side of the disturbance weaken the radiative energy is and t hereby support more in upp er-disturbance penetrates into creates a band of negative dy namical pressure

> results in the observed pressure drop erational precipitation forecasts by in advance. Most likely th е tion on the Tropics leads to

> > ency equation







K. De Sanctis, L.Molini, A. Pare

Numerical weather forecast of severe weather h meteorological community. It is well known the most important mechanisms to trigger d these mechanisms a numerical meteorologic spatial resolution can be achieved. In this w brk I have been used to simulate a convective structure

simultaneously measured a rainfall event on May 20, 2 003 from two different loca tions, S.Pietro Capofiume and Gattatico, revealing strongly localized run in the same co nfiguration, with a sp parameterizations. Sensitivity tests have be aim of investigating the different hydromet or and hydrometeor classification from dual-p performed with pluviometric network and t and discussed.

receiv d a e Planetary B Ther ve cells lha s be I on H bsta o Valley

rea sing attentio n in the hydro ary Layer (PBL) fluxes are one of , in order to be able to reproduce nydros s o that relatively high odels SMO-MODEL and MM5, two C-band radars have

convective cells. The mesoscale models have been ons and micro physical hysical schemes to the nicr the available radar data comparison have been study will be presented









swept through several outer suburbs of the justra 2003, a series of large pyro-cumulonimbi de elog pyro-cumulonimbi lasted for about 3 hours formed near the leading edge of the fire bef 2005). The local weather radar showed that sooty, black hail was reported to have faller of Canberra. One of the pyro-cumulonimbi spawne

is to use the Canberra wildfires to explore the dynamics and predictability of pyro-cumu lonimbi and the tornadoes they occasionally spawn. This is <u>achieved through a series of large eddy</u> simulations using the Weather Research and Forecasting (WRF) vertical resolution of 150 m. The model i January 2003 (0900 EST 18 January 2003) method described by Cunningham et al. (2 observed pyrocumulonimbi, including the f work addresses: the role pla yed by the heat of

Canb capit est to the individual cel subsequently eached ro-cu nir e one roł the ornada

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afternoon of 18 January anberra irport. The complex of a bout 10 km in diameter, and agating eastward (Fromm e t al. ights of 14-15 km, and imbi about 300 km east of the work presented

blution of 200 m and a ing on 2300 UTC 17 to the m odel using the ain characteristics of the re one was reported. The combustion in the development of the convection; the

importance of the water produced during combustion in deepening the convection; the back-reaction of the convection on the fire; and the dynamic tornadogenesis.Cunningham, P., S. L. Goodrick , M. Y. Huss aini, and R. R brtica ructures in numerical ohelent 1-75. Fromm, M., A. simulations of bu oyant plu mes from wildla Tan Fire fir 14 Tupper, D. Rosenfeld, R. Ser vranckx and R. McRae, 2006. Violent py ro- convective storm devastates 1029/2005GL0251611. Australia's capital and pollutes the stratosphere. ophys. Res. ett., 33, do

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Keywords: pyro cumulonimbus, tornado, wildfire























and rad ar simulation modules feeded with he models. This work focuses primarily on the eval the simulated co-polar and differential reflectivity (RSM). Since t he latter is able to provide C

second import ant task is co nstituted by observed reflectivity fields so as to assess weather conditions with a particular attent Particularly, a severe event occurred over I brtł above mentioned numerical models and results conce presented and discussed.

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h desired characteristics nentioned atmospheric he af fferen t microph ysical

parameterizations embedded into two atmospheric limited area model (COSMO-MODEL and MM5) on

datasets computed by a radar simulation software erent hy drometeors, a d and the av ailable ducing deep convective tion processes . cipita en simulated through the polarimetric RSM measurements will be



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The sensitivity of the MJO in the NCAR CAM3 w evaporation feedback and convective rain r both wind-evaporation feedback and convertive with amplitude comparable t o observations evaporation are minimized or removed, the moist static energy budget is examined to h and wind-evap oration feedback. These bud and convective rain re-evaporation may not be unit

used to develop observational evidence for the climate model. These satellite and buo induced latent heat flux and MJO precipita Partitioning of QuikSCAT-derived wind spec Ыi contribute sig nificantly to t he west Pacif intraseasonal latent heat flux variability.



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bert conve ction to windrol simulation in whic h 50 A ' re active produces a robust MJO bn fe edback a nd c onvective rain iability in the model collapses. The to rain re-evaporation evaporation-feedback ng MJO variability, and

suggest a more generalized approach to understanding why some climate models produce an MJO while others do not. TRMM precipitation, QuikSCAT ocean vector winds, and TAO buoy latent heat fluxes are

anism that is active in elati on between windas in the climate model. of 10 days and shorter alies that dominate







motion, and hence reduced (enhanced) rain variations in the Yangtze rainf all are mainly subtropical high is originated over the central North Pacific no to the SCS-Philippine Sea, and finally modulated the

. Dyr ically, the intraseasonal Bas l by g be twe the low-level relative eter upl vorticity and the upper- level divergence. In the mi ddle troposphere, the 15-35-day oscillation of the ten propagates westward h of Hawah the Yangtze rainfall. ariations












Keywords: global nonhydrostatic model, convectively coupled waves, aquaplanet









The Madden-Julian Oscillaiton (MJO) has convention from coupling between the large-scal e circulation theories and hypotheses on t he MJO have central mechanism. Glo bal climate mo del incoherence between the circulation and con raise a question on the MJO mechanism. extratropics has been propos ed and discus mainstream research effort. To revisit this

into a tropical channel model with its lateral (latitudinal) boundaries at 22N and S. Several case studies were conducted to simulated observed MJQ events for three to four months. The NCEP/NCAR reanalysis was used as the initial and lateral boundar were tested. It is found that the gross feat conditions nor sea surface temperature. T MJO events appears to the lateral boundar a time independent lateral boundary cond and when moisture effects on diabatic hea

erceived as moi with s op ch simulatio ns o , in <u>contract</u> sib ilit eratur d i ut l Jossib ili ale r

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omenon r esulted primarily the tropi cs. Numerous éc tið atio n-convection coupling as the e MJ O, ho wever, sho w stunning O theories and observations. This by in fluences from the nuch attention from the el (MM5) was converted

> ions to various factors sitive to neither initial o del can reproduce the ility was simulated when of the m odel was used model but a r ealistic lateral

ting pronibited in the were boundary condition was used, the model produced intraseaso nal variability that did not propagate initiated by influences from the eastward. These numerical simulations sug extratropics but its eastward propagation mi tion he tropics.

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Keywords: madden julian oscillation, mm5, tropical







The Madden-Julian oscillation (MJO), Kely intraseasonal oscillations (ISOs), are investigated (OLR) measurements as we II as the associa applying band-pass filters to the OLR data i tracking algorithm is applied to each ISO mo are traced in space and time in an objective to the identified convection centers of each, resembling the shallow-water solution, while a slig

suggests the impact of frictional boundary composite analysis reveals that the MJO is a ccompanied with a Kelvin wave a pproaching from the west preceding the MJO convective maximum in ER waves con structively interfere to enha leaves a Kelvin wave emanating to the ea wave prior to the development of MJO con vec ER wave, loosely concentrated around the MJĊ are also analy zed using Tro pical Rainfall Measurin g Mission

possible relationship among different ISO modes.

vave atorial 🛛 g a 2 vnamical requency-way that <u>convect</u>i The d mi i Oi de. T low-le

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R) wave, or collectively d of o oing longwave radiation he ISO modes a re detected by mber space. An auto mated waveenter<u>s accompanied</u> with the ISOs composited with respect uctur /e e> ts a low-level wind field gence over convection

-layer convergence on Kelvin wave dyna mics. A lagge d

aks as the Kelvin and o nvergence. The MJ O The appr oaching Kelvin and fall. The composite ut. Individual MJO events (TRMM) data to furthe r investigate the











A linearized, two-dimensional, non-rotating model of incorporates the saturation fraction of the toposition of convective inhibition into its convective logication unstable modes develop in this model. A solution saturation fraction anomalies and is a candidatione easterly waves and monsoon depressions. A court anomalies in convective inhibition caused by buck or boundary layer. The model is vertically resolved an

consists of two sinusoidal components with to the impose d heating profile while the o speed of the convectively coupled gravity r the earth's atmosphere and the predicted phase speeds of convectively coupled equ growth rate peaks at zonal wavenumbers Furthermore, the computed vertical structur observed Kelvin waves.



nearbore is presented here which entrop fluxes, and the strength hion. Two types of large-scale ture mode" is drive n primarily by any tro pical disturbances such as "gravitur ode" is governed by t above top of the planetary

except that the heating profile has the form of the first baroclinic mode. The calculated vertical velocity consists of two sinusoidal components with different vertical wavelengths. One component corresponds to the impose d heating profile while the one of the convertical velocity wavelengths are a statement of the convertical velocity results are a statement of the convertical



elengths. One compon ent corresponds where name ad gov erns the phase ruppento rupatorial Kelvin waves in the rules re close to the obser ved or man on alle parameter values the example the greatest s pectral energy. emperature structure, matches that of



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The Madd en-Julian Oscillation (MJO) is the dominar atmosphere, with a typical timescale of 30-9 floats, we show that changes in at mospheri wind stress, s urface heat fl ux, and precipi structure. Although changes in sea surface t that a significant effect of the MJO has been as temperature. Temperature and salinity fr from 2003 to 2005 inclusive were mapped to a red

to 50 m. Below that, an MJO-related signal out of phase with the surface signal, becau waves in response to surface wind stress. to the existence of the surface fres h pool i the central Pacific. Salinity anoma lies at t haloclines are raised or lowered in response to s

de of intrase day lv av ing associated v fo evaporation, ture (SST) are ent ed le floa lv-me

I veriability in the tropical vation provided by Argo ole ob the MJO, e. anomalies in surface ce changes in the upper ocean I do<u>cumente</u>d, this is the first time , and in salinity as well Western Pacific Oceans one-de gree horizontal

and five-metre vertical resolution. Composi tes of SST anomalies during MJO events show the wellknown pattern in which SST variability lags surface heat flux by a quarter of a cycle. This signal extends

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quatorial Pacific that is were d by propagating w complex patterns due ace salinity maximum in f eac h oth er wh en the







equatorial Indian Ocean, which is favorable for the organization of convections. On the other hand, these values in September were not so hig h. During each convectively a ctive period, the Rossby wave

disturbances developed over the western a gated westward, while new convections generated to the east of them , hq er, convections over the eastern Indian Ocean did not become organ t is interesting to note that MJO se OD he in September had the second largest amplitude, although the environmental state was not so favorable. In September, near-equatorial SST, especially solution of the equator, was relatively higher than that in October. This meridionally-extended warm SST 45 st ggested facil itate t convectiv e organization centered on the equator and lead to equatorial convective heating, which resulted in the MJO excitation.







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Shunsuke Shigeto, Takum

The eastward propagating 30- 60 days va dominant in the Tropics. It is known that the dipole mode event, etc. An understanding MJO is important because it relates not only also to the cloud organization, the vertical h precipitation m echanisms of the clouds acc Indian Ocean where in situ cloud m icrophysica

participated in the R/ V Mirai observation cruise MR06-05 (MISMO project), and the vertical distributions of the precipitation particles in the clouds were observed with videosondes. Videosonde is a balloonborne radiosonde which images of precipitation p consists of a CCD camera, a video amplific also has a stroboscopic illumination, which of particles are transmitted by the 1680 navigation deck of the R/V Mirai, and the videosondes were launched into the stratif

bands. Particle images transmitted from v (Snowflakes) near and above the freezing aggregations of nearly round graupel obse (Takahashi et al., 1995) and the R/V Mirai M

maritime stratiform clouds over the western Pacifi formation process in the upper level of stratiforms cl that. The number concentrations of ice cry larger than that observed i n the first hal suggested that a large-scale circulation mig

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an Oscillation (MJO) is lden soon, El **N**io, tropical cyclones, the anisms in clouds a ccording to the tation of water vapor directly, but this s aiming to clarify the banc at developed over the do ne up to now, we

> . Videosonde system and a control circuit. It size and shape. Images stem e quipped at the e MISMO project, seven

orm clou ds that the Doppler radar detected clear bright ideosondes were ice crystals, graupel, and aggregate S

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ds during TOGA-COAR E brm tra ter Pa fic Ocean (Suzuki et al., e w 2006). The n umber concentrations of ice crystal and graupel were gre ater than that observed in the region. It was found The stronger ice crystal ds was in ant over eIndian Ocean.Moreover, in the latter half of the obse rvation period, the large-scale upper air circulation had been dramatically changed, and the vertical pr ecipitation particle<u>distributions were greatly different</u> before and be hind ad been observed in the latter half was

her analysis is needed, it wa s physical features in clouds.

aggregates were different from

Keywords:



























We present results concerning the determination of in an idealized moist general circulation model. 7 sphere, a z onally symmetric aquapla ret le parameterizations including gray radiative trans simple model framework allows us to stude to convection scheme in a clean manner. A con coupled Kelvin waves which are remarkably ersis. These waves propagate with an equivalent depth

ion of the open and str lel. The mode, consist the open boundary, and issue of a simplifie set of the le pendence down of the le pendence down erst, at an opmin the depth open oximately

reference of equatorial Kelvin waves if the prinitive equations on the various it ealized physic all tts-Miller convection scheme. This lvin waves on properties of the ne much produce s convectively the variability within the tropics.

baroclinic mode structure than observations. By varying a convection scheme parameter that increases the fraction of large scale versus convective precipitation, we show that the waves increase in strength,

propagate more slowly, and move to large occurs, the Kelvin wave disap pears, and th relate the decrease in speed to the gross increased larg e scale precipitation. We add model simulations, which indicate that simi



aves increase in strength, e sca le precipitation form-like variability. We which is reduced with full general circulation isive GCM's as well.

Reywords: moisture, tropical, variability PERUGIA ALY















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is also observed that the larger scale feature for example the "envelope" of the MJO is oft Rossby gravity waves, and these in turn are necessarily organized into "waves". This interactions in the organizati on of tropical occurrence of smaller scale, higher freq ue modulation are not yet _____fully understood. While p transport have been reason ably well elucidated,

of diabatic hea ting on the lar ge scale. Sin composed of an entirely dif ferent suite o upscale interactions can be enabled by a w these scale interactions appears to be a cru disturbances in models. A potential aid to the a certain degree of "scale i nvariance" in observed organized tropical convection, from the mesoscale on up to the planetary scale structure of the MJO.

Convectively coupled disturbances uni ver moisture, vertical velocity and diabatic head lower trop osphere ah ead of the wave, with convective region. Low level moisture and thus CAPE and moist static energy is high ahead of the waves, and drying occurs fir st at low levels while diabatic heating precedes deep convective heating, over cooling. These dynami cal signals are consistent with the observation that the waves show a progression from a domin ance of sh allow to <u>deep</u> convection, and then stratifor m precipitation, regardless of scale or propagation directi

evolution of meso scale convective comple also exists at a certain level on the scale of simulation of convectively coupled waves. their rainfall spectra corresponding to the waves identified in models thus far have c and therefore phase speeds that are too f depth, a f act which perhaps provides clue

quatorial waves, so that estward-nertio gravity, and mixed ectrum of mesoscale features not of bo<u>th upsc</u>ale and downscal e he MJO modulates the ms responsible for this convective momentum the importan ce of oth er aspects of th e u pscale

interaction provided by the higher frequency are not obvious, such as those due to heterogeneous fields

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similar ways can be eatures, evidently the hding the precise role of imulation of equatorial

s is the fact that there i s gro ss features of the dynamical structures of

> s in their wind, temperature, se di rbances d isplay a war m ar mann troposp here within the shund the wave. Low level ioist aloft per tropospheric heating a anal of ble fact t hat the temporal and spatial ack to microphysical arguments, ions have implicatio ns for the Is appear to have peaks in

iness. How ever, all of the are universally too deep nd the same equivalent meterizations involved.

Simple modeling and cloud resolving studies are beginning to provide some realistic results, and will no doubt provide useful testbeds for the development of improved parameterizations in next generation GCMs.

Keyw












context, the depth of the mixed layer is a temperature and the amount of energy ava the ISV amplitude tends to maximize for an this basic process, the SST may also be st absorption of solar radiation by the upper ocean

the mixed layer temperature while the noo modify the air-sea exchanges and the bou as the wind increases in association with a at large-scale t hus increase t he amplitude ISV events. To address the role of warm-l tropics is pr oduced using C OARE 3 Fairain algori thm for ced by ERA-40 reanalysis. The results ar



nines the reactivity of the surface . Theoretical studie s showed that meters. In addition t o scale because of the urnal vind conditio ns. These

diurnal warm-layers may form during the convectively suppressed phase of an intrase asonal event and increase even the diurnal average surface temperature. This is because the nighttimes SST is close to



This higher SST wil I vers disappear as soon Warm-layers organised on and may even trigge r diur nal amplitude in the е

validated using the global SVP drifter datas et and em piric relations from earlier studies. An important result is that warm layers tend to develop rong confirming their potential amp role in the ISV amplitude. For moderate to s de > 1K), the horizontal ement with the typical extension is be tween 1000 and 4000 km du ŋg /s ij size of ISV perturbations. These results finally vali date the use of the C OARE 3 algor ithm forced wit h surface fluxes at a horizon tal scale typical of atmospheric ge method gives an estimate of the warm layer death and temper nonodel. I n addition, this era I circula. bysical time scale, making ture at each it usable as a parameterization. The imple mentation of this warm layer parameterization in the LMD GCM is further tested and its impact on the representation of the ISV in the model will be presented.















simulated, although its propa gation speed similar structure to Gill (1980). The selective pr

compared in the active convective regions.

heights (roughly 4 7 km) than the advection

and large horizontal diffusion is considered

used as that of wave-CISK. By using the si

both amplifying convective modes and dam

in the eastwa rd propagating convective systems.

investigate the eastward propagating featu_ res in the large-scale cloud system. Equatorial beta plane obtained by NICAM is g mode is preferentially IICAM. This mode has a the combination effect of

term is greater in some

spheric instability works

hich the moisture is

Keywords: dynamics of convectively, coupled equatorial waves, madden julian oscillation

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excluded and the diabatic heating term is a n ordinary or positive-only wave-CISK type, is proposed to

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heat flux and latent heat flux variability were the amin contributenet one heat flux changes. The observations suggest that the phasing bet were the wind certurbation and convective perturbation is important, as previously suggested by (Du to and Vialarda 000). One must releed h ave both strong winds (and evaporation) and clouds to make the net the at flux negative (ocean losing heat to the atmosphere).

Keywords: vasco cirene, intraseasonal perturbations, downwelling radiation













Madden--Julian (MJ) event as either primary which does immediately follow a preceding e Precursor features of the primary events c convective anomalies are observed to grow to the start of a primary MJ event there. An to destabilise t he atmosphere, leading to t

events appear to be thermodynamically triggered by a previous dry period. Other theor ies of the MJO predict that bo undary layer convergence and humidity, free tro pospheric humidity, pro pagation of dynamical stru ctures right a round the eq extratropical transients may all be imp ort signals from all of these mec hanisms are the successive MJ events, they are all abse precursor signals are actually part of the p generation of the MJO. Additionally, the usual a

no imme iate Abo ut 40% n am<u>biquous</u>ly ay in 5ν ed mid DOS e ger he act

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eceding IJ event, or successive, MJ events are primary event s. ributed to that event. Suppresse d cean, several days prior ature anomaly then acts t. Henc e, primary MJ

hd lateral forcing b y er, although precursor te observational data in e, it appears that thes e a role in the spontaneous that the MJO starts over the Indian Ocean is

tested using t he primary M J events, who se starting lo cation can be unam biguously determined . % of v ents), over half of the Although the most frequent starting locati primary MJ events start elsewhere, from the vestern Pacific.

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In the eastern portion of each of the five of summer, driving a cool ocean current off the w along-shore equatorward winds act to maint and coastal upwelling, favoring the formation of the cool SSTs. Thus the presence of a local unsolved is how the surface anticyclone itse heating has be en suggested as the possibly contrast obser ved in summer between the relative

landmass to t he east. Our model experiments indicate that it is the shallow cooling/heating couplet associated with this therm al contrast, rather than deep convective heating, that primarily and directly forces a subtr opical high in summer. Altho observed three-dimensional structure of th hypothesized that a local feedback loop thr be triggered by increasing insolation over and early fall thro ugh its intera ctions with th un



east. The associated by enhalcing surface evaporation high albedo also acts to maintain este<u>d , but a</u>n important issue still h monsoonal convective ed near-surface thermal d a heat ed subtropical

btropical high intensifies in

newhat weaker, the in our model. It is thus ced and maintained can e loop is prolonged into cool SST. As actually

observed, a wave-activity flux for stationary Rossby waves is distinctively upward and diverging toward downstream in the upper troposphere above those surface highs. Therefore, the thermal forcing acting on the highs is also important in the formation e planetary waves in summer. In winter, most of the subtropical highs weaker from da ana that the high in each of formed and maintained the North and South Pacific and over the rth | lan ore Hig 7 IS mainly as a downstream structure of planetary wave s generated in mid-latitudes, while the Mascarene high in the South Indian Ocean is under the influence of plan ary waves or the tropics. In addition, the transient eddy forcing is found to effect ivery ce the A regrand Mase rene highs, the latter of which moves in winter to the western portion of the basin in the vici nity of the storm track core. The fact that the core is anchored by the sharp SST gradient across the major oceanic frontal zone around the Antarctic Circumpolar Current suggest kind of air- sea interaction is involved in forcing the wintertime Mascarene high, the co unterpart acting o n the

Keywords: subtropical hi

summertime subtropical highs.

netary waves



VOCALS (VAMOS Ocean-Cloud-Atmosphere goal of which is to develop a nd promote scientific acti simulations, and predictions of the SEP interannual timescales. VOCALS is ultimatel coupled climate in both the SEP and over

shown that the SEP climate is a tightly coup

clouds, aerosols, marine boundary layer (MB coastal currents and upwelling, large-scale, lbs the Andes mountain range. We will give a Frief ov (VOCALS-REx), an international field campaign that

central to the SEP climate system. The field program_ enhanced long-term observations. Modelin the context and will directly benefit from t coupled o cean-atmosphere models exhibit str coordination through VOCALS of observation at which field data can be used to improve



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oved understanding, model ystem, o n diurnal t o fe-lar proved model simulations of the cs. In this present ation, it will be y un<u>derstoo</u>d interactions between n dyn is and thermodynamics, iurn a ulations, to the west of LS Regional Experiment

will provide detailed observations of processe s

is carefully designed to comple ment a suite o f OCALS, have provided kn own region where rature and r ainfall. The acceleration in the rate opical climate variability.

Keywords: vocals, stratocumulus, vamos



The five s ubtropical highs have some diffe sides of the s ubtropical ocean basins, but not alway merge with adjacent continental highs ove Áfric dipolar: have maxima on the east and west in summer, but no t all. The South Pacific differences are worth noting since it is co m Each high can be viewed as part of the local The talk will ill ustrate the primary difference purpose of this observational discussion is of clarify

mechanisms have been proposed for their maintenance, both local and remote. Since the Hadley cell can be viewed as being forced in part by subtropical convection equatorward and e forcing and local forcing (such as upwelle observational discu ssion is to sh ow evid maintaining the highs. The talk will show kinds of statistical tests to ass ess some of with bootstrap resampling and 1-point correlations. The rocus will be upon the two highs in the Pacific.

cell w on imilar bet lahs að tropical convection, several mechanisms invoke tropical or ler

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high fo r example, tends to an Ocea n high is often út h T end to be highest central pressure loc al sp ring, for example. These e highs are strongest in summer. nal av je is strongest in winter. n the subtropical highs. One e over time. A variety of

sonality. Most tend to favor the easte rn

sms stress midlatitude n other purpose of the th eories proposed for we shall show two basic these highs: compositing

The South Pacific high has high monthly persistence (though there is much daily v ariation) and has strong, clear links to remote events: conve ads a strength ened high while Midlati de fr convective activity over Indo nesia lags a st al cyclones have strong the lattine For the forth ranfic high strong low-pass at from fron tal cyclones. Con trary to one theory, links to the South and North Pacific high, es tiall filtering is nee ded to see any signal other than th strengthened North Pacific high leads temp erature changes to and there is little evidence for a connection to topped convect the east (including over North America) ime permitting we shall also discuss North Atlantic high statistical associations with remote weather.





Jean-Louis Brenguier, Olivi

Anthropogenic aerosols have a noticeable i pact efficiency, being therefore likely to al so sign however not feasible to do cument such i r generally correspond to different air masses and the accur acy of temper ature and hur di impacts from thermodynamics modulations, dynamics of boundary layer cloud systems have the

models, that now include detailed parameterizati entrainment, surface fluxes, droplet sedimentatio simulations are contradictory. Originally, it result in an increase of the liquid water opposite, because of enhanced entrainmer a

These simulations though were mainly focu this study we consider the s pecific case aerosol impact s interact wit h the diurnal variat ions of the cloud system dynamic

polluted ones with higher cloud condensation aerosol properties are modified during the case. When the sun raises however, the poli due to both an enhanced entrainment flux and a de

the process is robust, though with varying polluted case. This numerical study demo systems to the diurnal cycle and to change processes can be strongly coupled. Differen due to different fluxes at the interfaces, b the decoupling processes induced by solar

diatiy clou perturb ł e lit ro m o bserva to di<u>fferent v</u> ot sul eas ur hts

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and their precipitation cle of boundary layer clouds. It is Indeed, different aero sol types al profiles of moisture and stability, nt to decipher aerosol action tween aerosols and the igh resolution numerica l

ons of dyn amics, radiative tran sfer, clou d top n and drizzle precipitation. The results of the

Thouron, Valery Masson



tion e fficiency would howe ver suggest the e ntrainment efficiency. most often nocturnal. In tocumulus, during which s. Large Eddy

Simulations of a stratocumulus diurnal cyc le are prod uced by assuming either a pristine cloud, or

ulations reveal that, when se sin initi increases in a pollute d <u>r</u> pa ted y than the pristine one. d d s e' ore ap During the second night, the polluted c loud, unlike the pristine one, is unable to rest ore its original liquid water path. Detailed analysis of the simulations show that the no n-new sibility of the process is bupling (be bo undar layer in the polluted case that involves a competition between droplet and drizzle evaporation and short wave heating of the subcloud layer. Sensitivity test on the intensity of the subsidence and the large scale advection confirm that

fferences between the pristine and the e ti mes of stratocumulus cloud e comparable, so that the two polluted cas e are not only of the boundary layer t o ng sedimentation.





As in many contemporary coupled atmosp surface temperature (SST) warm bias in Environmental Prediction (NCEP) coupled Cl the formation of this warm bias, its associa and its impact o n the model's time-mean s include: (1) Where does the warm bias st develop? (2) How does the warm bias relate to th

what is the contribution of the associated excessive surface radiation fluxes? (3) What is the role of oceanic dynamics in the evolution of the warm bias? (4) What are the impacts of the warm bias on the mean state of the model and on its intera outputs from a su ite of CFS integrations, realistic atmospheric and oceanic initial cor decade free si mulations, (c) experimental SEP area [30S-0; 90W-68W] corrected, an (d the CFS with and without mean surface radiation

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lels, there exists a se a E P) in the Nathional Centers for CFS) model. This s tudy examines f model-produced stratus clouds, ability estions we will address pad nd how quickly does it ced stratus clouds, and

> addr essed based o n easonal forecasts from 04, (b) long-term multiadiation errors over the d oceanic component of



















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Well-defined spatial variability in the physical proper. evidence for t he importance of aero sol-clo and recent in situ and remote sensing obser our current understanding of these interacti features o bserved in mari ne strato cumulus ultra clean air masses. Although ship tracks and more reflective and solid clouds, recent reflective clouds associated with marine ultra clear

stratocumulus as observed from satellite and ship-based remote sensing and probed by aircraft. Unlike ship tracks, which are produced by anthro maintained naturally by drizzle cleansing. in the production of MUCA are discussed a marine stratocumulus (via cloud seeding) t

Keywords. clouds, aerosols, precipitation

arin e s l-drj inter of mari e observation cks a<u>nd bro</u>k strond le ndica ne e) wher

nulus clouds provide clear oud cture. A review of pa st cumulus clouds is used to assess idence reviewed will focus on two oud structures e mbedded within pr the between increased CCN nce d oken, drizzling, and less centrations fall below 10

cm-3. Such areas have been associated wi th Pockets of Open Cells (POCs) and rift areas in marine

ni sti

to be generated and drizzle process involved reasing giant nuclei in demonstrated.










simulating stra tus clouds and thus predict Ina interannual variability of upper ocea n in t h using ocean general circulation model (OGC Coordinate Ocean Model (HYCOM) for the tr daily surface fluxes based on the NCEP rear radiation for the period of 1979-2004. Grid ded sur

on WHOI IMET buoy measur ements at 85W, 20S. Als o, the OGCM is a ble to well rep roduce observed interannual SST variations in this regi on. Additional model experiments were designed to examine the relative importance of ocean dyna mics an results indicate that upper ocean dynamics of S ST n orth of 20S in the stratus cloud provide useful info rmation f or upco ming Study (VOCALS).

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imate. This study investigates d its relation with SST variability CM u<u>sed in t</u>his study is the Hybrid el was fir st forced with hortwave and longwave e well with those base d

> ual SST variation. The he interannual variation berimen ts wil I h opefully -Cloud-Atmosphere-Land

Keywords: upper, ocean, process



































NPL since 1997. Seasonal variation in OZ autumn seasons and minimum in monsoon shows daytime in situ photochemical build large number of days the surface ozone va for ozone (hourly average 8 0 ppb), which occurrences of hourly ozone more than 80 ppb

the above the regular measurements of surface ozone on round the clock basis are being carried out in is

vels in summer and h ozone concentrations is of data shows that on bient air quality standard s concern. Number of

was round 83, 39, 113, 158, 112, 111, 173 a nd 215 2003, 2004 a nd 2005, res pectively. The surface ozon e during 1997, 1998, 1999, 2001, 2002, measurements were also carried out at other cluding at high altitude stations like Leh / Hanle (one of the highest observat ign n e. The pattern of diurnal in d ent like New Delhi. The variation at Hanle is altogether different from tha envj fb ai оп nighttime increase of surface ozone has been observed under stable boundary layer conditions in nights and also d uring thunderstorms. Since NW side on the pre sen Inced by ag ricultural areas site is sur N (IARI, PUSA Campus), a preliminary evaluation sible dar gerto crop vi d by O3 has been carried out using exposure plant response index (AOT 40) and found O3 exposures are higher than the critical level of O3 and suggest that the present level of O3 may have impact on reduction in crop yields.









Andrew R. Jacobson, Sara E. Mik

Spatial and te mporal variations in fluxes of CO2 at pattern in the atmospheric CO2 concentrat directly, inverse methods have been used w Such atmo spheric CO2 inversions are no to small region s are n ot in dependent of on propagate to another, i.e. if one is able to b the estimates decreases in other regions atmospheric-ocean joint inve rsion method

Ocean lead to substantially improved esti particular, our new estimates of the net air a much smaller sink than previously assum outgassing of natural CO2 and the uptak suggest a rel atively larger sink for atmost Hemisphere (44S-18S), causing a large e Hemisphere. When this information and its erro



te distinct spatiotemporal s canno t be measured :02 fi hese athospheric CO2 gradients. meaning that flux estimates for n formation from one region to in or gion, the uncertainty i n a rece ntly developed as provements in the

Gloor, Jorge L. Sarmiento

estimation and understanding of the pattern and magn itude of the air-se a CO2 fluxes in the Southern mates of the magnitude of the tropical land fluxes. In

b



south of 44S) indicate ation between the he time, our estimates gions of the Southern hic sinks in the Southern ucture is combined with atmospheric CO 2

concentration data through the joint invers ion approach, the tropical and so uthern hemisphere lan d regions emerge as strong source of car a 77 probability that their aggregate source size exceeds 1 Pg C yr? his simil**a**r n gnitu to estimates of fluxes i n of a large tropical CO2 the tropics due to land-use change alone. Atting into que the k ist include a large tropical CO fertilization sink. Systematic errors in the atmosphe ric transport models that we employ to infer the the tropics due to land-use change alone the x ist source of uncertainty for our results. We sources and sinks of CO2 at the Earths surface remain a major a large suite of models. If attempted to incorporate this uncertainty by up ertainty ing our i argions using our results hold up to further scrutiny, the absence of a major CO2 fertilization sink in the tropics would have major implications for the future evolution of the global carbon cycle. References: Jacobson, A. R., et al. (2007a,b), A joint atmosphere-ocea face fluxes of carbon dioxide. Globa l Biogeochemical Cycles, in press.

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emphasizing recent observations.

















In the Tokyo metropolitan area, the extend prevails on summer sunny d ays with weak s familiar to peo ple living in the metropoli tar vertical wind fields associated with the sea b Tokyo on summer days, using a ground-bag Institute of Information and Communications Tech

numerical simulations by use of the advanced rese arch version of Weather Research and Forecasting (WRF) model. Results of the meteorologic and the numerical simulations with a horiz southerly sea breeze on the early afternod Imaging Spectroradiometer (MODIS) show the sea breeze front. The Doppler lidar de sea breeze front passed the observation site. A ceilom eter and a wind profiler installed in the lidar observation sit e also observe d the clouds and u pdrafts, respectively. A fter the passage of the sea

reeze orlaina wind conditi To in<u>vestiaa</u>t rculat we ent D er l ICT).

from the Pacific Ocean frequently The sea breeze is, therefore, quite characte ristics of horizontal and erved daytime wind fields over syste eveloped at the National a nalyzed results of the

Doppler lidar measurements performed on August 10, 2006 through meteorological analyses a nd

> ed. d t

e observation network land penetration of the he Moderate Resolution d band extending alo ng trong updrafts when the

breeze front, vertical wind fields over Toky tructure. The multi-layered wind structure was composed of 1) the sea broken v 0.8 above mean sea level tory return flow layer at (MSL), 2) a layer of weak winds at altitudes om ens the altitudes of 1.22 km MSL, and 4) a prevailing synoptic wind at altitudes above 2 km MSL. This wind structure, which was no t clearly sho wn by m o bservatio s by use vioi of piballs and radiosondes, might influence the movement and de elopmen undercle as that frequently appear over Tokyo on summer days.











As is known, mountain systems are source regime of rivers (glacier-snow, snow-glac ier snow (SAI) in river basin (RB) a nd determines mountains strongly varies from topographica direction of air carry etc.). At the same time variation of SAI takes place with other mountain Kyzylcha RB (Tashkent province,

of free er and requ leper lin etc ics of R s se es of RB (ave cases are ol in g eð an) ha sho

river flow (RF). Inflow accumulations intensity on sn h al distribution. In turn, SAI in height, arrangement concerning a ed, when inside of RB the stron g is the w co vering's maps in windy mountain slopes vind". One of the main

reasons of it is an aerodynamic of RB or its orientation concerning the basic air direction. In particular, at presence of the snow air flow (snowfall) there is an additional snows loss, and the most favorable

conditions are o bserved for snow carry. watershed par ty, one of the empirical forn transferred through watershed (Ms) is fun tio certain meaning (5 m/s), and time (t), du ng speed - 5 m/s and more): Ms = f (Ws, t). Cor the snow covering's maps have shown their satisfactory convergence. The basic conclusions of research consist that the aero dynamics laws allo w to assume that (with other things being equal) : a) under-

nat ur have always much more snow in comparison with which nts

s transf erred to other ula the snow volume for h is equal or more than rved (at the given wind carry via w atershed and

at snew weather (snowfall) an windy slopes have more snow in comparing increase of wind speed results in increase he unde wind ope; c) at orientation of RB towards to the air flow (RB "is open" to flo ow blur e s buld to be accumulated e al v) b ic according to along a channel of the river, as a whole); d) at orientation of RB against an air flow (RB "is closed" to the air flow) basic snow volume should be accurated in the top part of RB. Depending He period various variant s of SAI are on topography of RB and wind directionca rr, du ng sno w possible in different parts of RB. Knowledge of the aerodynamics laws and their application concerning a snowfall in mo untains allows receive a qualitative es timation of the snow a ccumulations intensity in mountains.





Abstract This study attempts to use a full meteorological satellite to examine the chara of- the-art velodrome that stood out at the r October, 2003, two hours to the closing implication of such damage. Deive and statis in the study. With the aid of meteorologi Meteorological office, it was observed th at the

part of Nigeria. Synoptically, features obser_ved also from the NWP products(i.e. chart) show that the storm was initiated due to large thermal fo the ocean observed both at 10m and 900m turning created around Abuja at 850Hpa ch defined organized and propagating type du Also, Inter-tropical discontinuity (ITD) or Intersouthward with the area relatively stable. The reflective surfaces provided by the rocky hills that outcrops most of the r egion. The presence of these

hills energized the system. The stadium is (Zuma and Aso), hence the storm may h maximum gustiness of the storm over the

cumulonimbus clouds in the early ho urs of afternoon over the rocky high ground areas of the central hrt' to

nsed observations fro m fot el that wreaked havoc on the stateplex, Abuja, Nigeria, on Saturday, an G<u>ames a</u>nd socio-economic to illur te the features involved ined from the Nigerian jer y d small in-situ cells of

ergence of winds from values due to cyclonic he storm was not a well (AEJ) during the period. CZ) was already dropping main source of en ergy came from insolation, high

r valey of two prominent hills cends these hills. The as it. ler nots. The absence of tha 6 adequate outlets at the velodrome couple d with th is gustiness may have been resp onsible for its destruction; as vortex was likely to have been created inside the velodrome. The rainstorm did not only ly-built s destroyed property and facilities but also the velocity me of the dium being the worst hit. The 8000 seating capacity bowl was shredded into pieces of rags by the heavy winds that accompanied

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Generally speaking, the intensity of tropical cyclones heat flux and larger surface roughness including, rapid weakening of tropical cyclones. Howev landed and traverse the Indochina do not di or the Bay of Bengal. In this paper, TCs land 2005 are studied to extract characteristics of of small dissipation rate of their intensity. W using GMS, TRMM, rain gauge, JRA-25 (JNA/CRIE

data. The I ndochina is one of the ideal experiment because: 1) It is located in tropics and the<u>refore effects of tranformation from tropical to extra-tropical</u> cyclones are n egligible. 2) M any T Cs of v west-north-westward throughout the rainy 1000 km in the zonal direction, which is probably on g enough to examin e atten u tioı landed Indochina from 1979 to 2005, we bun

rapidly de er la raphy yasu (2005) I they reach nd p<u>assing o</u>v ed tr bance re the tica CO' reana

ver land . Smaller latent e over ocean cause this rted that a certain number of TCs outh Asia (ex. the northern India) ne n <u>orthern</u> Indochina in 1979 to TCs and to find causes d attenuating processes MA Typ hoon Best Track

al fields to study attenuation processes of TCs

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decreases over Annam mountain range just after their ranging. However, it is also shown that a certain

nd mov e westward or land area is mo re than Cs m ovements, and is analyis of 98 typ hoons t of all ty phoons largely number of ty phoons are af fected rather weakly by the Annam range and live longtime over the

Indochina land mass. We divided all TCs lan lived group and a short-lived (normal life) group depend ed on their S. compared areas and atio**h** ra background humidity of typical cases of b he longlived members hd f ind, arc па accompanied wider areas of water vapour flux circulat ion and also move into more rich vapour areas. ver the mid leof Indochina. Noticeably some of longlived members increased eir vorticity












The backward reflection of a stationary Gr in the linear viscous case and for large Reynolds numb critical level at the gro und because the men wi number at the surface (J) is below 0.25, the GV inviscid limit Re~infinite. The GW is a litt le when both the dissipation and J in crease. Reynolds number, with a general tendency large ground reflection favors the downst decreases when J increases explains why the mo

indicates that there are some corresponde <u>nces between the dynamics</u> of trapped lee-waves and the dynamics of K elvin-Helmholtz instabilities. some among the neutral modes for Kelvin-J<0.25, can also be stationnary tra pped-w limit Re~infinite. When Re is finite, these s and decay in the downstream direction. dissipation and J increase, as does the GW abso

appagating toward the gro undis examined Re In this c is nu here lection b the ed, when Re l>0.25, the G а GW 5tic velop t of le bo

e stationary GW presents a mean flow Richardson n tous boundary layer is total in the hite, and the reflection decreases bsorbed for all values of the de cr when J increases. As a ped le the fact that it ave, ers favo r the o nset o f

mountain lee waves. The fa ct that the G W reflection depends strongly on the Rich ardson number



mple, it isshown that unbounded flow when bund and in the inviscid on in the boundary layer ases when both the











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One of the requirements of existing paramet orography is a specification of the surface this part of drag is based on the calculations results of flow over 2D ridges (Smith 1979) importance of the effect of vertical w ind s extensively recognized; ho wever it is no complex wind profiles, in order to devise si mple

assuming that the wind varies relatively slowly approximation must be extended to secor variation of the incoming wind to have an considering a no n-rotating, inviscid fluid expressions for the drag normalized by its detailed shape of the orography. The no involving the first and second derivatives of the wind velocity at the surface and the static stability, as

Teixeira et al. (2004) and Teixeira and axisymmetric mountains. For simple wind with height at a co nstant rate maintaining profile parameters is reduced to a dependence on the Richardson number (Ri). The facts that the drag decreases as Ri decreases for a linear wind profit but generally increases as Ri decre ases for a wind that turns with height, are explained by this simple

behaviour is asymptotically c orrect for relatively high Ri, down to values of Ri of order 1. The drag is found to vary with Ri more rapidly when mountain than when it is parallel. But whe of the mountain, the component of the dra of the mountain.

Keywords: gravity wave

ons าดเม ag. In the Lo hillips (1984) ntain<u>s with a</u> the d s of al fe culate le ti ations

n stratified flo w over ler paralet trization, for example, ch extend previous linear theory ptica<u>l horizo</u>ntal cross section. The surface drag has been drag a nalytically for nt study addresses the

effects of wind shear on the wave drage exerted by a stratified flow on an elliptical mountain by



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odel. Alt

height, using a WKB approximati on. This meter for the vertical atic approximation, and clo sed-form analytica are independent of the mensionless paramet ers

well as on the aspect ratio of the mountain. This model can be considered an extension of the models of ed, respectively, for 2D and

> nd, or a wind that turns ry ing len e o he drag on the wind

ng n some d artures from exact linear theory (where closed-form analytical drag formulae do not exist) occur for sufficiently low Ri, the model erpendicular to the major axis of the

45 degree angle to the main axes Ri is that along the major axis

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Strong and variable surface wind condition cause dangerous flying conditions, significar snow accumulation, may bl ock road s and communities are located, strong and variable and sea ice conditions, affecting again impo impacts there are concerns about changing environmental change. It is t herefore the q that, based only on available surface data is capa

at two locations on southern Baffin Island, and to establish linkages with the larger-scale overlying flow. Due to the high stability of the boundary-layer strati fication, high-latitude low-level winds over complex terrain are strongly affected by blocking communities in the Canadian Archipelago, winds for the most part are from two d iar terrain. Shifts between the two pr evailin geostrophic wind directions within a well actual and geostrophic surface wind directions sugge '

terrain to the overlying quasi-geostr ophic flow. To quantitatively investigate the role of large-s cale pressure gradients in the for cing of orogram winds and sudden wind shifts, a simple dyba is derived from the basic e quations of m directions is described as a function of the geostrophic wind. Based on this stability function, it is shown that, consistent with the observed increase in surf flow, the conditions for the stability of station positive accele ration. Actual surface winds adjust to a new stable direction as dete dynamical system within 3 - 6 h. Over these time-s cales and longer it is therefore advantageous to determine the relatively slow evolution of highly variable surface winds directly. The climate simulations, and for d etermining fu

it does not describe local thermal effects, lo within valleys, or to land sea breezes at the

harards in the Arctic. They wing snow and, through 5n in the dast, where most Arctic e a significant impact on sea state el routes. Due to t hese hazardous nditio the context o f glo bal velop imple dy namical model h surface wind statistics

> ly, at many low-lying sidere d here, surface g the or ientation of the nd are associated with connection between the

sts a coupling of the now below the top of the ia for the occurrence of strong ition , channelled surface winds St Illo ry along- channel win d

al with the conditions for

rmined by the sure dis tribution, instead of modelling s a tool for downscaling of global ing wind conditions. However, tic or katabatic wind systems

d during exampled periods of channelled





A heuristic mo del is used to study the in cyclogenesis. The model is a Semi-Geost rephic version adapted by Smith (1984) to study lee cyclog hesis a mountain massive exert a force on the lar (Shutts 1995). This force is taken into accou anomalies in the mid-troposphere. First, we the orograph ic varian ce is well separated absence of tro popause, the PV produced compared to the surface response due to the large

aloft and cyclonic downstream. Second, w sensitivity tests. We change the specification also repeat some experiments by including approximation. Finally, we consider the case Alps.

Gravity Wa ves (GWs) breaking o n leeof heEadv this del. th flow when the our model, a r th<u>e case o</u>f small e //sfo has the m

el for baroclinic instabili ty d by the small-scales of Ws fo ncounted directional critical levels produces Potential Vorticity (PV) dealized mountain range such that e contributions. In the pact that is significant r a cold front, the GW s

force produces a trough over the mountain and a la rger amplitude ridge immediately downstream. It opposes somehow to the response due to the large-scales of the mountain range, which is anticyclonic

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result, by a series of e back ground flow. We t he Quasi-Geostr ophic um representative of the

Keywords: gravity waves, lee cyclogenesis, turning critical levels

















Using the new NSF/NCAR Gulfstream V and Sierra Nevada mountain waves entering the The goals were to identify the influence of the from wave perturbations. Three of the diagr In the stratosphere, wind layering was fo un to 10m/s and vertical scales of 100 to 200m Bernoulli function and cross-flow speed, suggesting

discussed. GPS altitude-corr ected static p_ressure was used to comp_ confirming the Eliassen-Palm relation betw or EF was de tected at the tropopause. tropopause, indicating partial wave reflect momentum and energy fluxes were found related to secondary generation near the c



rch aircraft, six cases of ed with 26 cr oss-mountain legs. inguish. background wind layering e new GPS altitude measurements. es, with magnitudes up n the ayering correlated with vap a c hemical-dynamical

layered texture arising from horizontal interleaving. Methods for distinguishing layers from waves are



ute the vertical energy flux U*MF). No jump in MF KE) jumps across the stema tically reversed explanation is proposed

Keywords: mountain waves, stratosphere











troposphere and lower stratosphere up to h of a spatially staggered, semi- implicit, Sem standard atmosphere temper ature distributi stratospheric temperature lapse rates are The modelled reference wind fields are line

generation an iso lated mo untain ridge, p erpendicular to main wind direction, is a pplied. The m ost prominent wave propagation feature is the wave re_ tropospheric wave-guide. The wave-guide positive wind s hear in the troposphere will guide which may reach thousands of km-s effects like a wave-train wiggling can appe shear and disappears with the wind-shear partial refraction of waves into the stratosphere with propagation of some fraction of wave en ergy in stratosphere and formation of an elongated, obli quely down-stream propagating stratospheric wave-

estimated for different model conditions.

0 35 km. Mod ngian<u>, two-t</u>ir a mo vitł the t t r and hidire

g bases on the stationary solution evel wave equation. The US 197 6 const ropospheric and zer o lther stratification examples. wind profiles. For wave

flection on the tro popause and formation of a s already. However, a

ve-train occurs in the ion is absent. Secondary occurs at the weak wind hation is accompanied by

train. The tropospheric wave-train amplitude i ntensity of the wind shear. In this respect, the linear and hyperbolic wind lar ha approximately the same eaual S effect on wave guide formation. Impact the wave-drag size is

> Kevwords: buovance s, wave

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The vegetation type and density is mainly influenced besides other climatic fact ors like temperature. fluctuations as well as envi ronmental cha groundwater quality and quantity ar e inter variations. The northeastern region of India, average a nnual rainfall is 2450 mm, varyi rendered the whole hydrological set-up in t prevalence of shifting cultivation in an annual area

land vegetation cover is bare minimum, results in ab out an annual loss of 88.3 million tones of soil and 225 thousand tones of nutrients from the region. This has caused denudation of hill slopes and, silting of river chann els and floods in the plains. burning of forest material, and more and m growing population, a declining trend in ra over land and other resource depleting ant rod investigation is in progress in northeastern reg cover and rainfall on the sediment yield, so il moisture regime, groundwater recharge a nd runoff. The

land cover (uses) are ; livest ock based (maize for fo dder, oats, rice-bean, pea, guinea grass, tapioca etc.), forestry (Alder exbucklandia, Albizia grains, ginger, beans etc.), agro-forestry pear, lemon and vegetable crops) and shiftin cul 416.2 mm due to the interactional effects of land co covered with grasses and fodder crops. Sig moisture regime and runoff were observed of the rainfall received.

Keywords.

are linked with climat ic preciditation, vegetation and linkage has spatial and temporal 0 km2. is predominantly hilly. The nent of rainwater has isman This rther aggravated by the e practice, in which the

> forestation and in-situ practice to sustain fa st

ease in vegetation cover -disciplinery, long term

eractional effects of land

etc.), griculture (rice, maize for c ulture (guava, citrus, al cr), h rop 0. The sediment yield varied bn (e of from 9 to120 tkm2 in livestock based lan d use to $1\overline{300}$ and $\overline{4580}$ tkm2 in shifting cultivation, when the annual rainfall was 1992 mm and 2705 mm, respective ly. The annual run aried from 15.7 mm to r and r intation. Minimum soil moisture up to 20cm soil layer, was 5.6% and maximum, 27.2% in October due to various land covers and amounts of rainfall. The g roundwater recharge varied from 149 mm in shifting cul tivation to 1328 mm in land sediment yield, gr oundwater recharge,

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ent land use systems and amount





Varadero beach is located in the Hicacos P biggest island in Caribbean Sea. Varadero is one of the has around 20 km in length and around 1-2 km in The physic-geographic conditions in Varade geographic position, its exposure to sea-leve winds, hydrology, its fragile ecosystems, the climate change. The a ntropogenization pr urbanization more the tourist ic exploitation vulnerable space. Its human settlement, it



istic places i n Caribbean. It 5 10 m above sea level. etwe ice to environm ental change. Its hanges in temperatures, sunshine, pects of its natural vulnerability to e analyzed trough of d herability to this natural ep er e Varadero a vulnerable

place to global change. A sight to the climate change analysis in Vara dero allow know a bout variability in the climate in last decade in Varadero, future climate scenario s, d eterminate vulnerabilities and

impacts in to urism, the main economic a impacts. This work, analy ze the physicperspectives, the climate vari ability and f evolution with the antropogenization proce information develop adaptation measures a

vuln erabilities and agnost ic, the future atural vulnerability, its te change, and with the in Varadero.

Keywords: antropogenic vulnerability, impacts, adaptation







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Now there is no expert who would have (ATI). It is proved that XX century is the warmest decade of XX cent ury. The disput consequences for human ability to live. In (CCMs), ATI is expected in limits 1.4-5.80 external and easily proved factors it is poss destruction of the Antarctics shelf ice, and f and Tang-Shan mountains in Central Asia (CA).

therefore questions of ATI influence on wat er dynamics and other factors connected to water use are rather important. The Aral Sea Basin (ASB) is a part of CA. The ASB surface water resources (WRs) are equaled 116.5 km3/year for the multi-yea km3, and in the Amudarya R B 79.3 km3 a essential WRs change will not take place in 2-4 %, and for Syrdarya RB it is increased flow's measurements). H owever, for the observed. E.g., in the Chirchik-Akhang aran RB (CARB; watershed area - about 22000 km2), wich is

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temperatures increase nnium, a nd 1990s are concerning sizes of ATI and its vario us Climate Changes Models with <u>1990. In</u> this context, to the servable for last de cade scale ion of glaciers' area in Pamir harp water deficiency,

> river ba sin (RB) 37.2 uring nearest 25 yea rs uction is expected about imits of mistakes of river hanges of WRs can be

located on the territory of Kazakhstan, Kyrgyzstan, Uzbekistan, increase of river flow (RF) can make up an river) during the vegetation to 6-8 % (Chirchik river), an d reduction of period (April-S eptember). In this connecti erning consequences o f ns co climate change in CARB are r epresented as tte glacier areas on 1 5-20 ser uct on oŋ % to 2030; b) reduction of RF's glacier part on 3 -5 %; c) Increase of RF's rain part on 7-10 %; d) reduction of the snow covering; some other A ccording, RF's se as Junit dynamic will change considerably, and the repeatability of the high water phenom 🖬 ill be raie a. At the same time, such consequences require critical consideration: a) growth of the air-drought in the summ er months; b) change of vegetation of irrig ated plants; c) change of gro wth rate and efficiency of agr arian cultures; some o thers, which depen d o n weather ility o f the weather parameters (air temperature, humidity, etc.) is much more ons, in comparis on with changes, which are predicted by experts on climate.











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The climate s ystem respond s in complex variations in the magnitude of solar radiatio (e.g., changing atmospheric concentrations surface (e.g., through more intense and hi climate itself, for example, th rough change surface-atmosphere exchanges of water var urban areas are very sensitive to local, regi attributes are crucial to understanding the dynam

Knowledge on biophysical properties of forest vege tation retrieved from satellite images enables t o improve monitoring changes due to anthropogenic and climatic influences. The study covers a forested area situated in the north-eastern part of B the changes in biophysical parameters f IKONOS, Landsat TM and +ETM, MODIS a from time traj ectories of sat ellite data is a nev sensing at coarse resolutions. This approa h is alternatives. Further, it oper ates on biophysical va riables and thus shoul d be more rob ust than more data depende nt techniques. The input data are

Vegetation Index (NDVI). Associated with used to estimate surface temperature. The mean growing- season near- infrared radia parameters for the composite NDVI and surface temp erature data. The units recognized are broad lifeform vegetation classes, such as evergree n need leaf forest evergreen b They are compared to a ground truth map.Classification and the comparison method as well as f

climatic effects as well as for forest bioma forecast, and mitigate the risks of climatic adjacent environment areas and to provide derived from satellite data regarding atmos

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may be n atural (e.g., here) or human-induced atmo hate-induced changes at the land s) may in turn fee d back on the tion, radiative characteristics, and heighbo urhood of large ed in ange. dies of their biophysical der human disturbance.

> h aims at establishing image s provided by

hysical v ariables derived land cover via remot e impler than conventional composite values of the Normalized Differe nce

e thermal bands that are pts p n growing-season NDVI, s unac temper ature as input alleaf fore st, shrubs, etc. are varia , depending on the class unction of season of the year. Our analysis indicates a

potentially application of threshold techniques to land-cover classification and changes analysis due to cific aim of this paper is to assess and its biodiversity as well as on e basis of spectral information gradation .

Keywords: climatic cha

sensing



Biomass burning due to land cover change in Indo emissions at a global scale. The most severe purp which due to draining, have reduced ability susceptible to deep sub-surface combustion quantified the relationship be tween drough regions in Indonesia. We also considered th be reconstructed using a no vel interpretation data were obtained from the Global Fire Ernssions

satellite-detected hotspots with a biophysical model to estimate emissions of aerosols a nd trace gases. Land Precipitation dataset (PRECL), the NAS dataset (GPCP) and the NCE P moisture ba statistical model, we determined that seve le ENSO phases. The biggest emissions sour 420 Tg emitted during t he September pea extremely well fit ($r_2 = 0.93$), and used to determin

when the 3-month total rainf all was less than 460mm. Over eastern Kalimantan ($r_2 = 0.84$), the most severe burning was associated with a 3-mm (r2 = 0.95), with a 4-month total of less the unidentified as a major emissions source, th rainfall of less then 540mm . Importantly, severe associated with any of the moisture indicators and the use of satellite-der ived GPCP rainfall pr ovided

haze.Furthermore, we deter mined that the strongly correlated with the atmospheric ex The C emissions in southern Kalimantan, fo extinction coefficient in the city of Palangk September of 1982 and a 76 Tg event in S 146 Tg. The drought conditions under whic threshold estimated from the GFED data.

presents ated v s ass h soil moleture an omalo evere emissi ions e to w optic cal re eor GFED)

ingularly larg e source of C ral activity in peatlands, agric a result these large C pools a re dry periods. In this stu dy, w e from 1997 to 2005 for five key ates prior to 1997 could

ds. Monthly C emissions bines several sources of

Monthly in dicators of dr ought were obtain ed from <u>3</u> different sources: the NCEP gaug e-based Global n Climatology Project

non -linear piecewise d occur during positive represented 74% of the s region, the model was ed that severe em issions events have occurred

m, and over southern Sum atra (r2 = 0.99), previously ed with a 5-month total DCI emissions in central Sumatra were not closely

tade over the gauge-based estimates, clear adv whereas the more sophisticat ed soil moisture model was not useful. This has important implications in the developm ent of operational drought-monitoring__systems used to evaluate the risk of sever e stimates for several key regions are d from visible range observations.

upled from the ocal climate. In general,

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> elated $(r^2 = 0.96)$ with the to identify a 94Tg event in e September 1997 peak of was consistent with the

Keywords: biomass burning, emissions, indonesia

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In February 2 002, the lock gates of the new Algu closed, establishing a water reservoir that w artificial lake. The impacts of the Alqueva La model MesoNH (Lafore et al., 1998), were st The MesoNH model, like most me soscale representation of the evolution of lake temp period. An im proved represen tation of th understand water balance in the Mediterratean clin

DWD and currently operational within the <u>DWD</u> Lokal Model), have been tested and cal ibrated for the Algueva reservoir. Ini tial conditions, atmo been provided by in-situ observations. Mod of the surface temperature will be compare by remote sensing. Some numerical experin lake model have been performed in order local atmospheric boundary layer structure

amin the Gu cov 250 ń are udied with the depe nd critid athe<u>r</u> forecas whic physic OC ing 1

na Diver in Portugal were be the largest European . It v of simulations of the mesoscale on the water surface temperature. de <u>ls, does</u>nt include an explicit ned d ant over the integration d in order to bette r is ne lels, as LAKE (currently

being developed by Lykossov and Stepanenko at Moscow Univ) and FL AKE (developed by Mironov at ed parameters have

pperature profiles and e temperature obtained model coupled with the lake temperature on the





















It has long been known that the land surfa governed by complex topogr aphy or heterogeneitie

frequently a nalyzed using high reso lution vegetation-atmosphere transfer (SVAT) sche balance of soil, vegetation and sn ow, gener turbulent exchange with the overly ing atm climate models is lake, even though it is we be significantly different in the presence of capacity, roughness length, and albedo of water c

differences in the vertical transfer of heat in the water column compared with that on land. In this study, the impact of boreal lakes on the regional _ climate_is evaluated though a series of multi-year regional climate model experiments over central Ca a cascading sequence of simulations in ord surface tempe ratures and ic e cover are s NOAA/University of Miamis 4 km resolution on the local b oreal climate is estimated by been replaced by soil and v egetation consistent with the hearby la not surface. Because most bo real

development of a 1-dimensional lake parameter known as the Canadian Land Surface Scher computes the surface energy balance a no tile in the land surface mosaic of CLASS.

In the global climate system. Processes

surface veg mode feai l clim Such s d lem rporating a s One surfac ו that at, m to the ter s is th nea

on and soil properties are vith elaborate soilcouple odel the moisture and energy ticated treatment of radiative and

e that is normally disregarded in fre, and momentum can ge diff erence in heat d vegetation, as well as

h) is achieved through

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In the model grid. Lake eriments based on the pross impact of the lakes ns where the lakes have

lakes are sub-grid scale in current generati on climate and numerical weather prediction models, the SVAT used in Canadian models, SCUS This model (iteratively) uatio n u nder specified he at e

atmospheric forcing. Turbulent mixing is parameterized based on an integrated turbulent kinetic energy approach developed in the 19 80s for modeling la and ocea c mixed la sThe model represents a

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IUGG XXIV General Assembly

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Symposium **Extreme Weather and Climate** Likelihoods (ICCL)

Convener : Dr. Lisa Alexander Co-Convener : Dr. Xuebin Zhang

The vulnerability of society to climate varia intensity and frequency of extreme weather This is the case because extremes usually mean condition can cause a large change in the likelih that document past oc currences and futu especially as a result of ch anges in the cl changes, comparisons with model-simulated leading to extremes, and quantification of ur

teorology and Atmospheric Sciences

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and Future

pend more on changes in the on changes in the mean climate. sets on socie ty and a small change in the

n extrem is symposium invites papers extre we er and climate events, P apen reporting on observe d uture extremes, physical processes fact s, projections ties are invite

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Societal impacts of changing environmentainelude, moisture, loss of life, cr op and building damage illness) and, higher order or systemic effect progression of vulnerability to changing eve extent, and duration of the physical hazard risk e.g. population, property; and, (3) adjust and, (4) the c apacity to respond and reco increased calls for decision support and services ris



i creased/decreased soil effects (displacement, 'in di environmental degradation). The (1) the timing, magnitude, spatial r time: (2) exposure in regions of e over time e.g. levees; ture of risk has led to nal axes of assessments i.e. characterizing the integrated physical and social system. The expectation is that increasing the rate

at which po licy makers a nd reso urce managers interactions will result in improvem ents in thresholds are crossed or immediately afte Much recent work has shown that this ex high, uncertainty is great, techno logies distributions of burdens a nd benefits. More is Learning is of strategic importance in the variability and in identifying lessons from

policy windows of opportunity and adaptive in the context of changing baselines, and shown that long-term environmental problem steady public attention and visibility. However, a Latin America and hurricane and drought impacts

logical and be consensus-ba sed but might entry points through these b arriers can be as well as by changes in government and s offering opportunities for including lo ng-t after particularly severe or visi ble events s (1993), the 2003 heat wave i n Europe, or optimally a priori because of uncertainties risk perception .We il lustrate the benefits

_aca_uire knowledge about environment-society

sions before particular ough land-use change. hen decision stakes are and there are uneq ual opulation and affl uence. decades-long process of adapting to climatic change a nd past and current practices. T wo complementary concepts,

Entificial as facilitators of learning rnance. E xperience has orm c th by si agle discrete actions or on alt policies but respond only to continuing, sustained efforts at learning from event to event, supported by n diverse contexts such NSO-related impacts in in th S, the G ula tive short-term risk adjustments to smaller events can actually produce greater vulnerability to future events e.g. through induced development along levees. Thus, the evolution of responses in the short-term can appear

in

g-term risk. Policy win dows that offer ocusing events, suc h as disasters, indows have been identified as reduction plans, especially ina, the Mississippi floods these can be designed as land use ten ure and pts such as adaptive

management and policy windows in responding to cl imatic extremes e specially as the statistics of drivers of vulnerability change over time. We show that moving beyond more broad-based assessments requires approaches that are not only communicated in context, but that are also action-oriented and

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reflectively int eractive among participants private institutions. In addition, responding through climate change) must be placed w (as lessons are learned and incorporated and across diverse administrative units. Ca

of various public and s in the base state (e.g. mate acr oss time-scales ariability through change) variations in extremes such





This presentat ion discusses the large scale extraordinary weather events affecting the event types have been studie d, but these 3 heavy rain, extreme heat wave onset. Data significant parts of variou s meteorological f and geopotential prior to the event have lar region having extrao rdinary weather; can b theoretical dynamics concept of downstream deve

be the prominent role played by a rid <u>ge(in he_ight</u> and <u>thermal_fields) downstream over t</u> southeastern prior to heat waves and har are preceded by a ridge either over Alaska waves). While the Alaska n ridge passes a California requires the prese nce of the so compositing a dozen target event periods are resolvable by most climate models and provid results.

atterns that at entr alifor alley emphalized sa tellite e ot s<u>urnrising</u> lional rs back trad me bst su

been associated with USA on and beyond. Various his presentation: severe freezes, e analyzed to identify statistically e up<u>stream</u> fields of temperature The p n: is far larger than the mes s uggestive of the is so le features are intuitive

and known to local forecasters (and that provides a check upon our methodology). Less apparent may he

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dinary weather types st west of Alaska (heat ordinary event to affect ical analysis consists of uch larg e scale patterns alternative app roach for downscaling model





important agricultural area strong ly impact d during winter months with a frequency from **B**0 North and Western part of the basin. In ord variance effect, estimates of the mean warm outputs (DMO) using climate model simulati between simulation of the climate of the

obtained using a downscaling technique dev eloped for daily temperature extremes (Tmin and Tmax) in the region and applied to the same IPCC temperature series estimate of frost occurr benefits of using a statistical downscaling r emissions scenarios on frost occurrences a evaluate the importance of reproducing the order to infer f uture changes for climate extrem

ost. Fur herr the mountai tudy <u>the com</u>l minim ires ar 'np AR4 d nived ie I 0th ce of the

fr ost occurrences are observe d South-East to less than 1% in the effect of the mean warming and rived from Direct model base, us ing differences ry forced with different

emission scenarios. The mean anomalies are applied to observed series and compared with results

MO and downscaled pared to evaluate the thermore, the impact of downscaled series to on of local temperature in the contribution of the mean and variance

effect across the range of climate simulations is asse ssed and the importance of the variance effect is evaluated. Keywords:

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Given the complexity and global nature of the climate and interdisciplinary programs are indispensible f climate events usually have strong impacts cause a large change in the likelihood of a Climate Change Detection, Monitoring and I on extreme events using free ly available so the ET by the Climate Resear ch Branch of E can be fre ely exchanged within the inter national

indices, RCl imDex software and users gui de are av ailable from http://cccma.seos.uvic.ca/ETCCDMI. RClimDex creates 27 core indices after performing several QC checks on the raw data. The information provided by the indices not only includes h statistical distribution of the data changed trends in extremes. We have run RClimDex We selected the same data period in order provide an overall picture of climate varia index, relative to the period 1971-2000. The re

have increased about 6 days per decade. Growing season length has increased over effect on summer agricultural pro ducts but minimum and minimum of m inimum temperature have incr warm nights have been increasing all over aut

intensity index have been increasing in mo declined in 30 stations located in the Aege days have been increasing e specially in the extreme flood events. The maximum one eastern Marmara and south Anatolia regio Karapınar, Cey lanpınar a nd I ğdır which a increasing trends in Marmara, Aegean and especially in eastern parts of the Marmara

em coopera and a dictin ionite ety and a sma me. A joint W has <u>defined</u> 2 RClim ped a ent C es ha ommi

ctivities within international limate change. Extreme ange in the mean condition can CI/CLIVAR Expert Team (ET) on re climate indices mainly focusing naintained on behalf of he adva ntage that they bmplete list of the 27

time but how the

information about the riod from 1971 to 2004. same climatic period. To verage trends for every

nights have be en increasing all over Turkey whil e ice days and frost days decreasing . Summer days cally gignificant at the 5% level. ions , is will be have a positive coastal r me ffe 5 w b e exp ienced by orchards for ea tive example which rely on co ld conditions. Maximum of maximum, minimum of maximum, maximum of ased at most stations. Warm days and ool days an cool nights have been w while decreasing. Warm spells have increased while cold spells have decreased. Diurnal temperature ran ge has increased in most inland stations while it has de_creased along coastal areas. Trends in simple daily even mean annual total precipitation

sults snow that numbers of summer days and tropical

be number of heavy precipitations an regions a nd usually cause ave also i ncreased exce pt eased especially in Konya, un fortunately there are et days have increased p, Sivas, Rize, Kilis and

Mus while decreasing in the Aegean and Konya. In summary, in general there are large coherent patterns of warming across in the country affecting both maximum and minimum tem peratures but there is a much more mixed pattern of change in precipitation.

Keyword
















investigating the impact of land-atmosphere coupling for intra-annual in addition to inter-annual climate variability; 2) global climate model experim Dutch/German/Swiss ESSENCE project, wh interactions for climate change. In addition discuss the extent to which soil moisture-a heavy precipitation events R eference: Sen atmosphere coupling and climate change in Euro



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region al clim ate simu lations It in the context of the ole of land-atmosphere variability, we will also ecipitation variability and

nd C. Sch r, 2006: Land-

ate interactions Keywords: heatwaves









Water related disasters such as flood and landelide economic and human loss. Therefore, there is n precipitation a ccompanied by global warmin (AMP) is focused as an extreme index and it which is developed by MRI-JMA (Meteorolog Phenomena such as t hose for heavy precipit horizontal resolution. However, from the res km-grid AGCM is directly comparable with rain ga

20th a nd 21st Century, res pectively. Both simula_tions are forced by the a dequate SST and the greenhouse gas concentrations. The PDF (AMP are analyzed and estimated by L- mo which is a scatter diagram of L-skewness pareto distribution) in many regions, both shows that the observed AMP generally ob ys the simulated PDF of the AMP may insufficient in kurtosis. Future change of L-location increases in most part of land area. However, high-order moments such as L-CV, L-skewness and L-kurtosis show strong



netimes lead to seriou s re changes of extrem e h to fl ual maximum daily precipitation stigated using a 20-km-grid AGCM nd Japan Meteorological Agency). ie AM very sensitive to model dy, th MP simulated by the 20ain areas. Two 20- year

runs, the present and the future, are conducted by the model. The target periods are at the end of the

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he return value of the moments ratio diagram obeys GPA (generalized hwhile, p revious studies distribution). Therefore,

western Atlantic Ocean, western

e PDF does not largely

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ch areas as those listed

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spatial variability and no noticeable change Indian Ocean, western Pacific Ocean and se change, just shifts toward large value, in m above, upper t ail of the PDF extends more toward la rge value, which is mainly caused by intensified tropical cyclone. Return value of the AMP estimated by L-moments increase which indicates the magnitude and frequency of hea

Keywords: climate model, future change, I moment

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the southern Australian margin. Maximum es that wave events at on s ind Cape Sorell lag those at Rottnest by approxim telv B(), uedic by approximately lou ad at pe, le v 7 hours. The location and intensity of generating cyclones for the N-largest wave events of each year of record have been tracked fr om the NCEP-NCAP -analysis MSLP, a not being lationship with wave properties (height, period and direct ion) investigate Statistic btain ed poi out a we ak variability of peak directions (almost all wave events c oming from the same 10 sector), regardless of cyclon e position, with highly energetic, strongly seasonal. extreme events. Interannual variability of the storm count is large, and relationships to climate stigated.







Recent changes in hurricane activity in the Nerth Atla greenhouse-related warming on tropical cycone the observational record is d ifficult because observational record and the role of interd resolve the detail of convective processes in to analyse po ssible changes in tropical cyc maximum potential intensity of tropical cycl model results from the WCRP/CMIP3 archive to as

ity in ean l bf lack of consi variability. Lo al cyc<u>lones a</u>n activit en sed ir njun hes mpacts

erest in possible impacts of d the world. Analysis of ĥs a rð observiting technologies over the berio d climate simulatio ns canno t erna<u>tive me</u>thods must be applied ced gi ho use conditions. The with halysis d ata and climate change tropical cyclone

activity. MPI correlates with cyclone activity in at least some basins and has been shown to characterise the cyclone season in a number of location s. Two_algorithms for computation of MPI are applied to a

range of climat e models for the AMIP2 pe cyclone season and on the basis of pote characteristics of individual storms, MPI is the season and the peak (potential) intensi modelled thermodynamic potential intensity. Ba output, the MPI algorithms are applied to future



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the basis of length of assertio ns abo ut th e ikin on, in terms of length of etween GCMs in terms of and availability of mode 1

CMIP3 archive. The pre dictions of thermod ynamic potential intensity computations from these climate





sub-basins, projections of disastro us extremes in discharge using highimportant for local risk assessment.

varied am ong rivers and resolution GCMs will beco me











The possible relationship between anthrop activity has been a topic of g reat interest and discussion Workshop on Tropical Cyclones, a consensus state attending hurricane researchers. Among the evidence both for and against the existence climate record to date, no firm conclusion c other conclusions of this group will be revie results. As a step toward using models to fu framework for simulating Atl antic hurricare activity

hydrostatic regional model, run over observed specified SSTs and nudged toward observed time-varying large-scale atmospheric conditions (Atlantic domain wavenumbers 0-2) derived from NCEP Reanalyses. Using this perfect large-scale model appro find that the model successfully reproduce hurricanes and several other tropical cycle simulated vers us observed annual hurrid hurricanes intensities of up to category 3 scales, the model reproduces the observed

variability in the Atlantic (e.g., shear vs. p useful simulati ons/projections of pre-1980 reliability of these projections will depend of conditions from sources external to the model.

ange and hurricane or tropical cyclone In Novembe nt wa oduo sus condusio to tectable anth hade <u>on this</u> p of m ng wit uþ infor nis d e uced.

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6 at the Sixth International on input from over 120 draw as the for owing: Though there is genic signal in the tropical cyclone The basis for this conclusion and ecent relevant research w dyna mical modelin g is an 18-km grid non-

> ons (1980-2005), we in numbers of Atlantic mple, th e correlation of The model simulates b). On interann ual time

ENSO-Atlantic humcane covariation reasonably well. We conclude that the model appears to be a feasible to ol for exploring mechanisms of hurricane frequency he model may potentially make urr iq activity. However, the Atlahtic tmospheric and SST











Meteorological Agency (JMA). The statistical pe seasons. The relation betwe en Dec. 2005 heavy sn owfall and cloud-top heights is averaged LNB in 2005 is higher than 700 h other years. The appe arance rate of LNB i

Dec. 2005 was caused by the environmenta ond

1 km, CRM). The initial and bo undary conditions of which the mixing ratios of clo ud and ice cl

are performed 4 times a day by the CRM, precipitation distribution predicted by the C precipitation (R-A), although the pr ecipita brought from t he underestimation of R-A. CRM are high er over plain areas (about 6 mountainous areas, because clouds forme of appearance rate of predicted cloud-top

clouds caused the heavy sn ow fall in Dec_2005_Th estimated from the results of CRM, well co 2005 over the sea, the appearance rate of Jan. 2006. The comparison of heights bet appears with about an half frequency of L the sea are very similar t o those of cloud

comparatively examined from the horizontal distributions of averaged LNB in Dec. 20 05 and the other years. The g 50 J Pa higher than that in the refore, h eavy rainfall in ner mutoning i not only easily form, und

but also develop higher. The consistency b etween cloud-top heights and LNB is examined using the predicted results of a cloud-resolving model (JMAnnhydrostanc model with the horizontal resolution of ERM are oduced from the 12- hour forecasts of JMA nonhydrostatic model with the horizontal resolution of 5 km (its initial and boundary conditions are produced from the RANAL). The precipitation is cale ulated using a bulk -type microphysics scheme in

> raupel are predicted. 9-hour forecasts data are used in this study. The A Radar-Raing auge analyzed is overestimation could be 2005 predicted by the relativ ely lower ove r ed. The vertical profiles 05 app ears exceeding

two times over a 600-hP a level more frequently than in Jan. 2006, and the vertical level with the maximum freq uency is exce eding 50 hPa higher. In other words, the higher de velopment of sno w <u>e vertical profiles of app</u>earance rate of LN B,

> . In the profile of Dec. ably higher than that of s that cloud-top heights e arance rates of LNB over indicates th at snow c louds





Instrumental observations of European ten the nineteenth century (Houghton, 2001). warmer than normal and a lower number describe the variability of occurrence of wa site and day s pecific way on the basis of th mean and standard deviation were calculate day was numbered among c old (warm) wh deviation lower (higher) from the reference period

length of cold (warm) spell. Records of the longest spells in seasons and in t he year together with records of seasonal and annual frequencie variability was analysed using different me square metho d with Studen t's t-test, and minimum and maximum temperatures from Pulawy, Siedlce, Zakopane and Sniezk a) in the patterns accompanying the occurrence of warm and cold spells were distinguished. To do this the daily

NCEP/NCAR reanalysis were used. The set established. The day was defined as cold these sets the composite maps for all geop a 014222References:Houghton, J. T., Y. Ding, D. J. ThirdAssessment Report of the

CambridgeUniversityPress, 944 pp.Sen onKendalls tau. Amer. Stat. Assoc. J 63: 13 ature ave rey uld ex his ' fing š than r cold spells. V rence<u>period</u> ive da ind emp e re d le. Suc

arming since the end of reater number of day s s bv ys. The aim of t his paper is t o and cold spells a re defined on a -199<u>0. For e</u>ach ca lendar day the entere h this calendar day. The s da s at least 1.5 standar d was performed for daily

minimum and daily maximum temperature. The number of consecutive cold (warm) days was called the



ed and their long-term estimated by the least Sen. 1968).The daily e, Kalisz, Lodz, Poznan, Additionally the synoptic

gridded geopotential heights from levels: 850, 700 and 500 hPa and s ea level press ure values from Is lasting at least five days were

least two stations. For were prepared in each ntia lev pre su season separately and analysed. The work was supported by EC under project STATME-MTKD-CT-2004-🌆 der Linden, and iggs, M Noguer, P. S. D Δ Xiaosu, Eds., 2001: Climate Change 200 1: The cie fic Basis **c**ribution • Working Group I to the Intergovernmental Panel on ClimateChange (IPCC).

PK, <u>1968, esti</u> mates of the regression coefficient based









In order to control water reso urces in the face of c present a serious threat to human life an duatu being required more freque ntly by policyclimate change use general circulatio n mod spatial resolutions (those with grid spacing particularly limit in representing the ter extremes in s ystems such a s the Asian mo estimates are available to planners. In the se days

the changes of extreme pre cipitation during the Asia n summer mons oon. Results from time-slice ensemble exp eriments using a high-resol monsoon resulting from climate change. M warming over land tha n over the ocean, a and an increase in mean precipitation durir þ tf precipitation events increased significantly. to greater atmospheric moisture content. In cor of mean precipitation.

flood, ap ht. ns, p cosy Hydrolduical ak Ms) in t he m oximately 300 water/ ĝу ere 'ery expei

le o sion, which frequently d risk assessments a re ctions dicti ons in at acco unt for global Despite numerous studies, coarse and <u>uncertain</u> physical processes ac tioi nd the variability and ricted ount of regional-scale relatively hi- resolution

(T106) GCM become to be p ossible in the virtue of improving computer's ability. This study investigate



So uth Asian summer ditions suggested more ric mon soon circulation, number of extreme daily cipitation were attributed rather limited the intensification



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Keywords: indices of climate extremes, regional climate model, era40 sres scenario

















The most frequently used statistical approaches appl events are those based on modelling either Te up its time series as a whole. In the former cas (usually leading to the Gener alized Extreme Poisson process model and the Generaliz ed examined variable is generated by means frequently employed for daily temperature) simulated artificial series. The present study y focus

unusually high daily temperatures observed in specific attention is devoted to the setting delineate extremes, and t he minimum ser preserved) and estimating the order of the framework yields also estimat es of future of scenarios of the mean temperature change

e**s**timate re tail d ie dis estimation pr distribution) q distri<u>bution)</u> а stocha hf mc abil of Ind mparis

probabilities of extreme ution he examined variable or lure makes use of block maxima aks-over-th reshold (leading to the e latter case, a long series of the (an a egressive model bein g ts are de rived from the applicability of the tw o

approaches in estimating probabilities of r ecent temperature extremes, particularly heat waves and

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central Europe during su mmers of 2003 and 2006. A he ar າລ

(the threshold used to their independence i s b erature. The statistical ity of heat waves under

Keywords: extreme temperature events, heat waves, frequency estimates



The present work deals with the causes (trid environment) of one of the most importan Italy) during the last century . Between the days, a lot of downpours hit almost the enti recorded the maximum value of the century triggered a lot of damage all over the region Department of public works has been acquired an

documents, mainly co ncerning applicatio Department. D ata obtained were organise database. An estimate of economic dam allocated by the Department to peop le for calibrate a procedure for damage estimatio Damage Index, obtained multiplyin g the value of elements

ll) an rainfall eve 1972, a nd th rian<u>region:</u> in fal o Fo e pur S 0 d in o

mage to s ociety and which occurred in Calabria (Sout h ginni ng of 1973, during about 30 vera<u>l rain-gauges</u> the daily rainfall cte ris y high cumulate value, unpub lished archive of work enact the framework of

damage in one of the most s everely hit provinces of Calabria named C atanzaro. From the above said archive, data concerning landslides, floods, and sea_storms were obtain ed checking a huge amount of from people to the

icipality in a specific he amount of funds ed values were used t o and based on a Local nd the percenta ge of damage

damaged a (classified in three levels) affecting it. Keywo

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but less frequent than in the second. Dr oughts and floods appear in all the seasons but floods only Detween 1700 an d 1715. 1717appear with severity during the second half 1793. Increased frequency of events. Drou s prij and fall b ut their lengt h led. decreased. Floods were mor e frequent in this e ignest num ber of years (7) ter. pel bws with both drought and flood. 1794-1850. The number of droughts and floods decrease d drastically for all the seasons. 1850-1900. T his period is charachrized by schal changes, by fled to the progressive abandonment of the rogation ceremonies, and an • crease thiropic presure on t he thalweg. A s consequence only some spring droughts are recorded up to 1875, when their record ce flood record spreads to all the seasons. Thus, seasonal comparison of the frequency of events allows a better characterization of the annual series n argue that the 1550 to 1800 episode was a period of high frequency of events d

supported by the Spa nish Ministry of Scien CGL2005-06458-CO2-01/HID. Fernando D (BES-2003-0482).

knowledgements: This research is cts REN2002-04433-CO2 an d d by a MEC research gran t

ases, but the





Extreme weather events can cause great damage to. of increased frequency of such extreme events in in recent years. There are many difficulties i the inherently low number of observations models. This paper describe s a met hodole extreme events in a probabilistic way throu project ENSEMBLES.Sensitivity surfaces are, differentiated changes in the observed input data

for reaching an a priori dete rmined threshold any given year is then calculated from the results. The threshold can be set generally, such as the probability of reaching the 50-year return period, or be more site-specific, relating to past observed ever the future is then visualized by plotting t models in the same diagram. An advantag variables from different models can be as assumes that the intra- annual variation in

threshold is insensitive to intra-annual variations, our areas were studied; Lake Vnern, Lake Mlaren and Lule River, and the runoff from the entire Baltic Sea basin. The hydrological model was the HBV model, a conceptual rainfall-runoff model which is were selected for Lake Vner n and La ke ML allowed release discharge during a period in the fall/winter of 2000-2001. For Lake Mlaren the threshold was set a swater level b elow a certain level for more than 100 consecutive days. For the was compared. Preliminary resuls indicated the La intra-annual variations of the driving variables. The results for Lake Vnern indicated that only a slight increase in precipitation subst antially increased the probability of reac hing critical water levels in the

lake, thus forcing a maximum outflow for a decided upon. The results c an be used damage to societies and human health.

Keywords:

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er human health. Concerns esearch within this field drive of increases in extremes, such as tion of extremes in global climate rological models ability to model e st ud as conducted within the all mo for a drainage area with bitation). The probability

> k of a certain event in biecti ons from climate e projections of weather rea and threshold. This the models, or that the

atchments. Specific threshold s nern the hr e d was set as maximum cal event that occurred limi ling as in the modelled runoff studies c ha tive to diff erences in the ras most s er

The sholds for the other areas are not yet

to minimize the risks for future





heat waves to be observed in differen t lo ation Peninsula. Heat waves are defined as perio threshold defined by the 95th percentile of First, we analyse the outputs of the model points closer to the different observatories corresponding to observed daily temperatur are applied to makecompara ble the simulated and

processes resulting from the GCM rescaled trajectories are comparatively analyse d with the observ ed ones, regarding both the occurrence and severity characteristics. Alternatively, the GCM trajectories are used as input in the st atistical model pro "downscaled" heat wave process properties best of both approaches we obtain a mediu these extreme events in the scenario SRES Centelles, A. (2007). Modeling a nd fored continental-Mediterranean ar ea. Glob. Planet.

Madr Barce the daily m data, from 1-T63 in the eir dis we c Ъа in 195 000 ved ti

u a, ...) of th e Iberian í. I a um temperature, Tx, exceeds the to August, for period 1971-2000. BM experiment, for the model grid tions with the ones rectio h variability and position Second, the he at wave

> rder to obtain the served ones. Using the e expected behaviour of J., Cebrin, A. C., Asn,

sti central Ebro valley, a doi:10.1016/j.g loplacha.2006.11.005 Beniston , chande, M., Stephenson, D.B. (2004). Extreme climatic even ts and their evolution under c hanging climatic Mare intense, more frequen t conditions. Glob. Planet. Change 44, 19. M and longer lasting heat waves in the 686, 994-997. DOI: del T eso, T. (2004). hnc

10.1126/science.1098704 Prieto, L., Garca Minimum extreme temperatures over Peninsular Spain. Glob. Planet. Change, 44, 59-71.

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in many impact studies. As a result, t wo pw downscaling te chniques, have emerged as a m information based on the large scale GCM at no evaluation of statistical and dynamical down ca scales more appropriate for impact studie . Co temperature data from southern Quebec and down

a muc arger aling applicach to generate eri c circulati output co arison madu atial-so e than is usually needed namely statistical and dynamical al or station-scale climate change ata. This study m akes systematic uc ting mate scenarios at spatial using merved pr ecipitation and calcocclimate models (CGCM2

and HadCM3) and the corresponding values from the latest Canadian Regional Climate Model (CRCM4). The downscaled results are validated over the 1961–1990 baseline period, and climate change scenarios

are computed for the 2041-2070 period.


climate Houel (CRCHT). climate change scenarios hasis t o the statistica l er the r eference period ent seaso ns while both general, the study helps oducing the temperature







A study has been undertake n to investigate the whole Indian region as well as over eastern p data. A downscaled high-resolution future cl part of India as a special em phasis. A war season while the eighties for the postmonso the monsoon rainfall is also noticed. Future using five GC Ms namely H adCM2, CSIRO greenhouse gas forcing as well as sulphate aero so

per oC global mean temperature change is noticed in the composite scenarios, with more warming in the northern India and less warming in the southern India. However, unlike mean temperature, pockets of positive as well as negative changes Downscaled scenarios using a statisti cal te that obtained from the all-India scenarios seasons rainfall indicate a dip ole structure vitl (-5%) in coastal belt. So the projection f



n the 20th century o ver sis and observed station rean ve been constructed over eastern m the fifties in c ase of the winter nificant recent decreasing trend in ge in t eriods 2020s and 2050s M4 bee n derived due to ount of 0.3 - 0.6 0.2oC

parts of the country. C which is higher than rios in the premon soon d and negative changes g of about 0.5-2.0oC by

introducing the global change of GCMs ('2.5oC) for green house gas forcing alone but slightly less when introducing sulphate aerosol forcing with it.







nesting data are not perfect, the RCM simu presentation we present the response of ar R model framework nick-named the Big-Brothe E due to the nes ting process excluding other hod impact of the large-scale nesting data on the small

changes by the RCM. In general, the RCM restores a great part of the driving-model small-scale errors, even if they do not take part in the nesting in regions with important orographic forcin

of the driving model have er rors, the sma that the large scales precondition the small to provide the accurate large-scale circulati can be affect o er<u>rors in</u> ment rors le d hulate g data. The study uses a perfectts to evaluate the errors permits to analyse the M. It is found that the

errors contained in the large scales of the driving model are transmitted to and reproduced with lit le

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en to improve slightly , when the large scales ors as well, suggesting all scales, it is necessary N

Keywords: big brother experiment, lateral boundary errors, small scales





driving wind field from the NCEP/NCAR rean buoys, light ships and platforms in the east e the regional m odels is obtained when corre and instantaneous wind speeds is higher than that

and CL M are bilinearily interpolated to measurem ent locations and st atistically com pared. In the standard hindc ast approach a regional model is in model is force d by the reanalysis at the lat can freely determine the prognostic varia wanted property for small s cale process reanalysis by data assimilation are ignored wit cyclone tracks and the location of pressure sys especially deep inside the modelled area far from the lateral boundaries. On e mechanism to overcome

this shortcoming in the standard regional modelling method is a dynamical downscaling approach called spectral nudging, in which the ho rizontal accept the forc ing reanalysis for large scale regional model (von Storch et al. 2000). regional models do not ha ve an added val ue both in "open ocean" areas and the German Bight. However, in the English Channel, where local tor graphy and associated important, the regional model shows an a dded va wind speed distribution there's a clear indication for r an added value of the regional models in coastal regions, especially for higher wind speed percentiles, while in "open ocean" are reflecting observed distributions. These fin status, meaning whether the measurement the North Sea/North Atlantic are compared the Mediterranean. Literature: von Storch, technique for dynamical do wnscaling pur Ratsimandresy, J. Carretero, A. Bentamy, J year atmospheric hindcast for the Mediter global reanalysis. Climate Dyn. (25), 21923

n Atlar truth. Added value from idered ice w me a d statistical distribution otł fields from NRA, REMO alysis.

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itialized by a reanalysis. Furthermore the regional iat

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a the regional model ations. While this is a bly reproduced in the regional model. Thus rom the observed ones,

integration area are forced t o p be dete rmined by the

hstanta yous wind speeds the wind regimes becom e neo us wii speeds. Concerning the as NRA is bett er ntly of the measurements' assimilation

eanalyses or not. These results for ade by Sotillo et al. (2005) for r, 2000: A spectral nudging 36643673. Sotillo, M., A. 05: A high-resolution 44gional i mprovement of

Keywords: downscaling, regional model, observations

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Five Regional Climate Models (RCMs) were i hple the objective of assessing their "transferabili continental scale climates over different reg models are: the Rossby Centre Regional Atm Regional Climate Model (MRCC); the climate Canada (GEM-LAM); the Climate Versio n g Spectral Model (RSM) from t he United States. The

conditions for a five year period spanning fr simulated surface energy budget (surface radiation observations. This aids in un derstanding t from those observed. To gauge the transfe separately for winter and summer seasons, site, 2 mid-latitude continental sites, 2 Arct and 1 equatorial island. We present timediurnal cycle results for these stations, highlighting geographic areas and/or seasons when all the RCMs

perform well or where there are systematic deviations. Most RCMs simulated the annual cycle of surface temperature close to the observations ov largest deviations were observed at high a reasons. Analysis yields insig ht about how determine the cause of a given model err individual modelling group concerned.

gions of the globe with seve fe rer RCMs to mulate the variability of t is, the ability f the world w minimal parameter changes. The c Cli<u>mate M</u>o RCA<u>3) from Sweden; the Canadian</u> al for t model of Environment bei rom (any, and the Regio nal (CĽ 40 and NCEP boundary e pilot

om 2000 to 2004. Whe re useful, an analysis of the



and turbulent fluxes) is made using CEOP flux eratures and h umidity ated their performance h: 1 mid -latitude coastal 1 equatorial land station as estima tes and mean

> r opmand North America. The the tropics for different conve ction or front al

Ide R si the i ming progression in different regions of the world. Where possible we us e extra CEOP observations to hmendation make reco impr ovements for the or an














































IUGG XXIV General Assembly

emissions leads to an increase of the mean driven by an enhanced seco ndary produ period of study and the higher pressures 1960, 1980 and 2000. References Forkel, impact o n ph otooxidant co ncentrations in climate-chemistry model, J. G eophys. Res. and P. J. Adams (2006), Sensitivity of concentrations to climate change, J. Geop emissions on regional air quality in Ca doi:10.1029/2005JD006935. Unger, N., D. Streets (2006), Influences of man-made

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s of pollutants in most regions, which may be e of the temperature increase in the egions observed for summertime Re gional climate change and its ions with a coupled region al JD006748. Racherla, P. N., ine particulate matter 005JD006939. Steiner, A. L., S. Tonse, R. C. Cohen, A. H. Goldstein, and R. A. Harley (2006), Influence of future climate and ., 111, D18 303, h, J. C ofala, and D. G. n tropos pheric ozone, Geophys. Res., 111, D12313,





Perugia, Italy





Islands with significant topography are influenced orographicprecipitation enhancen in current climate hydrolog ical analysis, th possible without appropriate downscaling m lack of sufficie nt observations. In the prese model recently proposed by Smith and Barg Azevedo (1998). Results obtained by the two mode

lara precir the analysis tion of spa itatis<u>tical do</u>w y we è iat 3) an ad (no bared i

es which are highly mate change scenarios, and even distributed precipitation is n ot ling <u>is often</u> not feasible due to a winscaling models, the simp of a m odel proposed by v ers ations in a network of 25

meteorological stations in the Islandof Madeira, and with results from si mulations with two mes oscale

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models, MM5 and WRF, at horizontal resolutions down to 1 km. Results are first evaluated using 2 full years of simulation, one wet and one dry modified model setup and pa rameters. The with resultscompared with observations. current climate control simulation from the change scenarios A2 and B2. The propose in an application to the Azores, with a rather diff

la m

tivity experiments with ateda 15 year period, ownscaling of a he d corresponding climate other Islands, as shown







Jan Erik Haugen, Morten De

In the Norwegian RegClim project (Reg iona Clim has come to an end a fter almost 10 yea i uncertaintyand the thus imp lied risks for dynamicallydownscaling climate projections Norwegian version of the HIRHAM model with A1B and CMIP 2, and the models were from (BCM), and CAM- Oslo co upled to a slap o

ofunforcednatural variations of thephasesof the At lantic Meridional Overtu rning Circ ulation on the climate change in the region. The scenario from CAM-Osloincluded advanced treatment of the aerosol direct and indirect effects. The spread in th because orography and coa stlines increa regimes. We have in particular studied the have also revisited the discussion of the siz of applyinga developed version of the "big-brothe

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Warming), which now was to investigate sources of Euro pe. We have do ne this by ent <u>global cl</u>imate modelsusing the e mis scenarios were A2, B2, Bergen Climate Mod el Hamb study the influence

siderable, in particular in

eferred regional flow of weather events. W e mical downscaling by o duced by Denisand co -

workers (2002).Our HIRHAM-based results suggest, pe rhaps surprisingly, that the do main should be sufficiently large to enable improvements of regional flows and include downwind adjustmentfrom inflow portions of the lateral boundaries.

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Keywords: flow















This study makes primarily use of the satellite derive regional climate model in simulating snow in the the sensitivity of snow simula tions to snowused in the model to determine whether the compare the modeled temperature and prec in Turkey. The regional clima te model is the NCARs RegCM2 model) at the International, surface schem e of RegCM3 is the Biosphere-Atmo

centralized at Turkey, and covers an extensive area including Eastern Europe and most of the Middle East. The spatial resolution is chosen to be <u>24</u> km to make it consistent with the resolution of satellite snow cover data, wh ich are obtained fror System. The model is continuously run for October 1, 20 03 June 1, 20 04. Subseque transition temperatures including 1.8, 1.0, becomes comparable to the satellite snow gove mm at which the surface alb edo is modified. In gene ral, snow cover is well simulated by the model.



ss the performance of a urkey. It also addresses be an e, a threshold temperature that is m of rain or snow. In addition, we tions from meteorological stations whic as developed (based on ysics ted in Italy. The la nd he models domain is

> ow and Ice Mapping 01 - June 1, 2002 and for different snow-rain the modeled snow cover er equivalent is at least 5

Monthly spatia I correlations between modeled and s atellite snow cover for selected regions change

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0.6

1.2.

between 0.5 and 0.8. The modeled snow is temperature. The model generally overesti temperature f or Turkey, however, the mod observations.



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to snow-rain transition stimates maximum tches well with t he

snow, performance Keywords: racin





























been successfully used for decadal simulations over_Europe using ERA40 as lateral and SST boundary conditions. The simulated la ke surface te ditions are compared to observations simulations w m odel PROBE h as been used. Unfortunately, info rmation on lake depth i rge regions. Results will be presented showing how sensitive the resul

s will also be compared to a g how important lakes are

Kevi

and to corresponding RCA

for the regional climate.

simulation where lakes have been replaced









RegCM Regional climate models such as Modeling the climate can detect differences area under study. Variety of simulations hav physical computational processes and sche domain size a nd center of do main. In this different convective schemes including Grell, into two differ ent schemes of Arakawa-Sch Decembers up to February) of 1376 and 1779 wer

simulati re cap etwe real nate done for be ensi ch as radiati he ensiti<u>vity of t</u>ł ed Kud E S) ar itch l of Sti

fferent clima te processes. simulated climate in the mode investigation of the model to the convectio n, land use, reso lution, gC M3 has been measured to the iel. Gr cheme itself has divided pe I (Winters (started from the mod el in producing

precipitation of 151 synoptic stations of Iran and CRU research, the regional appropriate schemes Kuo whit %16.5 error for nor theast, Emai error for southwest, Emanuel with -%10 er and Kuo with -%16 for southwest of Caspi of Iran have been rejected statistically bec use analysis of two years simulati ons, Grell schemes

Keywords: regcm3 convective s



reanalysis data. A ccording to the results of this error for southeast, of Iran, Kuo with %20 %80 error for northeast l and northwestern parts according to mean error

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approximately % 20 errors have good results for winter precipitation simulation of Iran.











Prodromos Zanis, Ioannis Kap

Regional climate simulations have been performed resolution of 60 x 60 Km over the broad months periods using three co nvective para (AS) clo sure, b) the Grell scheme with Frits Comparison of the model re sults with the temperature database indicates a significant closure schemes) which is significantly red uced w

summer when convection is kn own to be more of ten. The temperatur e differences between the two Grell and the Emanuel scheme dominate in the lower troposphere extending up to 700 hPa. As far as it concerns the to tal precipitation no system observed throughout the year for the Euro greater when Emanuel sche me is used. T convective scheme choice indicate that the often than the Grell scheme and is more e The water vapour surplus in the two Grell

which in turn balances the comparison between the Emanuel and Gre lies between the Emanuel and Grell AS compared to t he two Grell s chemes is attr cloud-radiation feedback mechanism: enhanced convec tive precipitation, more efficient drying of the atmosphere, less cloudiness, more incoming so

more convection.

th Re 1 rea n and M dite ations, a) the happ<u>ell (FC)</u> Rese with Gr

cold

model with a horizontal clim ean area or thre e individual 12 -II scheme with Arakawa- Schubert re and c) the Emanuel scheme. CR U) x 0 .5 gridded surface

mrios Melas, Christos Repapis

oth with AS and with FC used, especially in the



hanuel

manuel schemes are of total precipitation is simulat ions due to the bked by the model more s drying the atmosphere.

amount of total prec ipitation. The gratest differences occur in the

he me

le the Grell FC scheme behavior perature in Emanuel ace t ct_ exert on the following t ef ne varmer surfa emperature and ag ain

Keywords: rcms. convection, feedback

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A conceptual study of climate feedbacks ha

assumptions. The term feedback is used prototype usages, stability-alt ering feedback an impulsive forcing, negative when stability terms of the systems stead y-state response diminishing) have been isolated for study. T background of control theory, which provide forcing and w hich is often seen as providing the

relationship between the prot otype climate feedbacks is si mple in the context of the ze ro-dimensional model. Here, the stability-altering and sensitivity-altering feedbacks provided by a given interaction are of the same sign. Also, the sign of the s always coincides with its sign as measure model, however, the sign of the prototype feedback as d efined in cont rol theory. In between the prototype climate feedbacks a that, contrary to the common assumption that these reedbacks are always of the same sign, they can



reduce under simplifying limate literature. Two 'ntľ sys tem asymptotic response to ty-altering feedback (measured in rcing, negative w hen sensitivityk cond are viewed against the e embracing all forms of erspe used in climate. The

> by initial tendencies es. Even in this simple the sign of the systems ty of t he relationship el is lost. Here it is shown

be of opposite sign. Also, the sign of the stability- altering feedback as measured by initial tendencies can be opposite to its sign as measured by dition, it is shown that there is no simple relationship between the sign of mate dbacks in the two-zone se res ults point to the models and the sign of t hese models feed trol k as eor . . need for greater precision and explicitness in the definition and use of the term climate feedback, both to facilitate interdisciplinary dialogue in relation t and st errone ous inferences eedback ar to guard within the climate field. Expli cit definitions of the the prototy climate feedback studie d categories o here are proposed. Reference Bates, J. R. 2007. Some considerations of the concept of climat e feedback. Quart. J. Roy. Met. Soc. , 133 (in press).

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Climate sensitivity, the equilibrium global average sur dioxide, might be the single most important hum to anthro pogenic greenho use gas emissio m. The difficult to nar row down for more than twe type a number of studies that have recently come to variate warming over the last century, climatol occal and temperature response to volcanic eruptions temp data of the last glacial maximum, or combinations

age surface warming for hum a to quantify lo . The range of clim ity ars. An attem of constraints of ar call an pains, a emplicature ation hations are a there is

reaction bling of atmospheric carbon term con ate change in response sensitivit has b een no toriously m ade first to assess the large nate sensitivity using the observed Earths matrix imbalance, the ver threast 700 ye ars, and proxy

guess value of climate sen sitivity of about three degrees Celsius. The lower bound of the range is well constrained, and provides a minimum estimate of climate change we have to adapt to if atmospheric

Keywords: climate, sensitivity

greenhouse g as concentrations are stabiliz difficult to constrain, because e observable recent proposals that the combined constitighter upper bound on climate sensitivity remaining open questions are discussed in



upp er bo und is more mate sensitivity. More of evid ence provides a liction.net ensemble and









diagnostics. This metho d uses Marko v Ch uncertain distributions for variables in the representing unforced variability are estimation (AOGCMs) and thus, have uncertainty and climate system properties are treated as a study has two majo r advanc es. First, we now includ e an uncertain di stribution for t he covariance matrices repre senting unforced variability

estimate the parameter dis tributions for separately. We will also present new results used in the estimation. Specif ically, we indu means as well as using global averages vs. higher re Sensitivity to estimates from multiple AOG day (2a) Sensitivity to record length and s changes in sea-ice area.



s to sample from the

he covariance matrices ral circulatio ns mo dels variable. Likewise, the

h

thod as well. The present while pr evious work used fixed estimates. Second, we

y rather than treating eac

the 🗖 mate change diagnostics tes e th area for the recent 30 sie ges hsa-N years as an additional diagnostic. Also, we explore the dependence of the posterior PDFs on the spatiotemporal patterns of climate change. We explore he use of ing-term trade vs. changes in decad al Each of these changes in olution s a averagin spatio-temporal pattern will be discussed. In sh ort, the major advances include: (1) Updates on unforced variability sampling via MCMC methods. (1a) Limits to information content from AOGCMs. (1b) dating climate observations to presentcs (2b) Additional diagnostic using





(DGF), we have to rely heavily on new instruments (AQUA), ground measurement circulation models acco unting fo r aero so summarise our research findings in this through the ongoing International Polar Year.

(F climate implications area, thei

CALI PSO) and passive th sophisticated climate-This pr esentation will and t he strat egy followed









intermediate-complexity or (more recently) with observations in order to arrive at some climate sensitivity is derived from the weigh The problem with this approach, as noted b arbitrary decisions about how parameters proposedan approach to this problem that makes

nplexity ima r e of th eir re ribut<u>ion of t</u>h et al pled :he ative

odel, simulations are compare d e like lihood, and a distribu tion for nsitivities of the individual models. nat r es are acutely sensitive to al design . Frame et al im a const raints and prior

opinion in a si tuation in whi ch a single parameter do minates how sensitivity varies, along with the likelihood of the fit to the da ta, across the ensemble of models. While it will generally be possible in

principle to generate a "su per-paramete parameter-dependencies in the model, this equivalent to that of Frame et al in the sind likelihood of the most likely model as a fur tio climate sensitivity. We will de monstrate thi ap change with a very simple climate model and applied to simulations of current climatology performed by

ehensive mapping of re practical approach, is to map the relative est, which in this case is recent transient climate

the climateprediction.net experiment. A further concern about climate sensitivity is the possibility that atmospheric feedbacks in a system underge (either a secular trend, volcanic pulse or seasonal cycle) may be different h a s m tha t has regained equilibrium. This point is il lustrated by the cli ensitivities" (estimated de (ctiv au from 1%/year increasing C O2 experiments) and "equ ilibrium climate sensitivities" (estimated, in from 1%/year increasing c c c c c general, from slab experiments) reported in the reported into simple climate Spir We will show how such ent IPCC Fo irth Assessin an effect can be incorporated into simple climated nodel s, d the impact has on our ability to constrain the long-term response to 550ppm stabilisation even if we can constrain the climate sensitivity that is effective today.







current climate change consists of a mix of a li trend is exactly the same for both hemispheres, value is in discrepancy to the widely accepter v discrepancy is that a mutidecadal temperature phase of such a value that this oscillation contri In particular, a superposition of this oscillation r be the reason of the current warming peak duri

oscillation c oincide well with some prominent strongest El Ninos of 1941 a nd 1997, maxima changes in the Caspian sea level variations The oscillation, we speculate that the global clina stage can last two or three decades, simila relative cooling stage is characterized by the m compared with the mean temperature of the N negative NAO, and hig h humidity in the Sa hel present-day level.

an an adecadal oscillation. The qual to 0.2oC/100 years only. This boC /100 year. The reason of this am plitude of about 0.2oC and a bute to be current warming trend. We line warming trend seems to the pair the maxima of this

oscillation c oincide well with some promin ent even ts of the current climate dynamics such as the strongest El Ninos of 1941 a nd 1997, maxima of NAO, m inima in the Sahel-precipitation, and strong

In the orderide con e... In a teached the precedulg boo gota ten crate e teached tern crate e teached tern crate e teached ne. The Caspian sea le vi

Sahel-precipitation, and strong on the urrent phase of the set agr of relative cooling. This are during the 1940-1960s. This uttorn Hemisphere higher as lience of La Nino over El Nino, le vel should be stabilized at its







evolution has been studied extensively in the past, still a subject of much debate. Observational stud much of the southern hemisphere in the past, troposphere and stratosphere plays in this part circulation model (CAM3), a series of the time potential responses of the climate to ozone change of ozone change and the stratosphere and crop change in both the troposphere and the stratopshe

st, the response of the almost of the have a lowed wat the decades the set the oz g trend is no the lunders e climate runs we bee . The shappen d here active ssocia here an wed.

e atmosphere to ozone changes is at t he mate has warmed over the ozone changes in both the understo od yet. Using a general ve been performed to investigate been dow with different scenarios ssocia with subscribed ozone











There have been many studies about climate char only a part of climate change. Anomalies fro have hardly been studied. Climate change time-mean fields and ano maly fields. Thus, particular, that associated with global warmi (historical exp eriment: 1851 ~2000, SRES Research Institute in Japan. They are mont the initial value in each ensemble of the SLES exp

(1851~2100). The data are a nalyzed every season and ensemble. Anomalies are defined as deviations from trends and averaged for each season the root mean square (RMS) for 30 years. averaged in the high latitude (50N~90N) h latter half of the 20th century. Since glo experiment, this suggests the possibility th other hand, there are no lon g-term changes of the RMS in spring and s ummer. In addition, the RMS

n time-mear tim hean ds (ai grasped ld y considers amin<u>ed.</u> Da ment: north ne vali

However, this describes) also ch ange, but they aly fie ctive by aking account of both ate change in anomaly fields. I n re tw<u>o clima</u>te change experiments ied b y Meteorological les of 3 members. Since in each ensemble of the

historical experiment, these two experiments are combined in 1990. Thus, we can get 250 year dat a

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essed as time series of emperature, the RMS ecially autumn after the 960s in the historica ming are related. On the

averaged in low latitude (20N~50N) also has no trends. Thus, we examined the process of decreasing temperature anomaly in the high latitude i ming temperatur e in the Arctic region is so low that the north-south different re is lard . Hov er, since global warming diff reases. In addition, the is stronger in the Arctic than in the mid-latit enç : u east-west difference in temperature also decreases. That is, the temperature gradient in any directions is reduced. As a result, the exchange of cold and warm iir weak e nerefore anomalies of the gene ral decrease of temperature becomes small in the high latitude tum n. In contra st t temperature anomaly in the high latitude, there are some regions where it increases locally even in high latitude. The reason may be that the temperatur e gra dient there becomes large locally and the interannual variability of the sea ice increase





An analysis of observational r ecords shows by approximately 0.7oC for 100 years. Dur decade in the Northern Hemisphere. This precipitation, a decrease in snow cover atmospheric and oceanic circulation patterns indicator of cli mate change and variability. long-term trends is complicated by strong variability in various climatic ally important parame

the climatic ti me series have a high level computed trends may be complicated sub Northern Hemi spheric SAT time series is respectively. Traditional trend analysis app correlation, and postulate the hypoth esis d they may lead to false conclusions, and sp by employing special metho ds of statist ical trend an alysis. Various statistical es timates like means,

standard deviations, correlations, trends etc. may be described in terms of discrete approximations of generalized stochastic integr als (i.e. an ir function). This approach has been broadly sample means (von Storch and Zwiers, 1999 variance of trends seems to be questionable due to non-unit weight function. We propose a theoretical approach which allows calculations of variance of practicable algorithm has been used for esti makes

for avoiding of that problem. The third and correlation of sample tren d and sample tr that correlation. Using the suggestion about temperature we have co nstructed the der parabolic regression. We have analysed so and Northern Hemisphere air surface temp hypothesis should be rejected at the 2 - 5 level estimated on the base of routine proc theanthropogenic induced global warming concept.

et the surface air. late g th eriod, ng is cl hcu a-ice extent, s (especially l r, it is bor h٧ uenc riab ident i**r**

aure (SAT) has increased 0s wer e the warmest ˈand` with variation ns of land-surface evel rise, and changes in the trends) are frequently used as an ze that identification of o red In pa Ilar, the lo w-frequency nstrumental and proxy

records from the Northern H emisphere and the Arctic. Because of this strong low-frequency variability,



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corr elation and tests of statistical significance o f rrelati on in the mean wo-ye ar time lags ount this intrinsic serial random variables. Thus, erpreting these estimates

an dom function and a weigh t varian ce of statistica l nate quation to estimates of cati of ins trend s u ng a n ar br weight functio n. Thi s statistic nificance climatic tr ends. Another important problem is c onnected to the rand om nature of t he estimated sample variance . It meansan overestimation of statistical significance of observed trends. We suggest a special Student-type statistics problem arises from the sufficient

of lored some ways for avoiding stribution of t he air surface coefficients of linear and time series. For the Arctic we have found that null e the trend significance vided analysis support

Keywords: climate, change, trends






























The variations of UV i _____rradiance during la detrimental biological effects related to the routine UV measurements over Europe have radiation in the past or pres ent (over sites obtained using UV reconstruction models parameters is the total column amount of o ground-based stations (rather limited numb (giving whole globe coverage since Octobe 1978).

ozone field ov er Europe (25 W-35E, 30N-8 0N) sinc e Jan uary 1950 with 1 deg (latitude) x 1.25 d (longitude) res olution. The model, elabor ated within the COST action 726, has been validated by comparisons with the total ozone measurer series since early 1950s). The accuracy of variations of total oz one are perfectly repre-January 1950) using COST 7 26 Total Ozor shown.

special exc έUΫ liatio initiated in the measure with the co otal oz of ations ce é ent a è

sts because of various he gro und-surface. The achin beginning of 1990s. Thus, UV ts a re not available) can be only ent input_data. On e of the input ozone been measured by the 1960 d satellite instruments odel providing the total

> eq and Lerwick (the time 5% and the long-term ed UV time series (since ta base (cloud field) are

Keywords: atmospheric ozone, surface uvb, statistical models

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during the sta gnation stage. For the most e formation of polar vortex and mb ozc h es ozone depressio n o ccur under co ndition of the var ying circulatio n in equato rial strato sphere. In overwhelming majority cases of the 24- and 30-months QBO-wcles starting in December- January TOZ has a relative maximum during the first ozone d relativ iinimum dui g the second ozone hole. Nearly always TOZ shows a relative minimum during the third ozone hole in the 30-months QBO-cycles starting in June-July. As a re sult, when series of the 24-months QBO-cycles are observed in equatorial stratosphere the course of TOZ demonstration of relative maxima and minima. When

the true alternation failures, since ing the 30-months QBO-cycles 79 and after 1993 coincides However, t he very large nths QBO-cycles occurs. months QBO-cycles in Gabis I.P., Troshichev

O.A., Influence o f so lar UV irradiance o n quasi- biennial o scillations in the Earth 's atmosphere // Advances in Space Research, 2004, V. 34 (2), p. 355360.Gabis I.P., Troshi chev O.A., QBO cycle identified by changes in height profile of the zonal winds: new regularities // Journal of the Atmospheric and Solar-Terrestrial Physics, 2005, V. 67(1

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the 24-months series changes for the 30-

the relative TOZ maximum is followed by t

starting in June-July. Minor d escending tre

with mixed 24- and 30-months QBO- cycl

descending trend of TOZ was typical of 19

The tendency is displayed for less amplitu

comparison with the amplitude during the



Polar-night Jet Oscillation (PJO) is a promine and downward propagation of anomalou indicates very clear quasi-periodic nature thr signal, it is interesting to examine how muc it is hoped that tropospheric signal similar to such signal is known to appear with the dou had performed numerical experiments start the PJO appeared very clearly. The results are also

can be well predictable if prediction is started after about 15 December, when the stratospheric sudden warming (SSW) is we II predictable in win associated with the downward propagation the occurrence. In contrast, the predictabil

found to be very short. There results suge with very larg e leading time only wh en the active.



Imindicates slow poleward variability somet imes and the PJO s lon g-lived cohere nt ble. If the PJO is well predictable,) is <u>also we</u>ll predict able, because e PJO. attack this problem, we er of 2 /04 whe n the signal of nts in winter of 2002/03

when the PJO was not very clear. It is found that the occurrence of the PJO until the end of the winter

bccurrence of the AO e than 1-month prior to in winter of 2002/03 is SSW is well predictable or is predicted as very















Springtime ozone over the Arctic changes sig chemical destruction, two me chanisms that phtr to disen tangle those con tributions to an om Here, observed changes between warm and from SBUV-V8, which provid es the 3D str structure are then related to contemporan ECMWF. Relative to warm winters, springting DU. Reflecting the rms change between years, this

been ascribed to chemical deplet ion. A majo r contributor, however, appears conspicuously in the 3 D structure of March ozone. Mixing ratio su rfaces have been driven into coincidence wit h theta surfaces following war m winters but remain deflect ozone structure reflects isen tropic mixing Arctic. In concert with anomalous downwel the observed deficit of springtime ozone o er during March, about half is erased during Ap winters. The Arctic is then opened to isentropic

earlier. The observed reduct ion is consist isentropic mixing. A 3D mo del of dynamic changes of str atospheric dynamical and che cal planetary way es and on the QBO. The integrations

temperature and ozone, which have been composed from the observed records of ECMWF and TOMS. Characterized by a strong an omaly of one sign opposite s ign at subpolar latitudes, each b ears the signature of the re sidual mean circulation. The structure is very similar to that associated with the Arctic Oscillation. nomalous upward EP flux from the troposphere, r epresentative of that observ concert with is entropic mixing by planetar latitudes sharply modifies tot al ozone ove

heterogeneous processes indicate that che ozone between warm and cold winters. Co remaining ~80% follows from anomalous t

ificant vears e win ime one rest u po winters are in foz<u>onemix</u>i an ge ίv over Ard compa

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nomalous transport and zone. Pre vious attempts ase ò e isolation of air over the Arctic. igated in the rep rocessed record atio. Observed ch anges of oz one e in the record fr cal stru om ano usly lean - by some 60 zone changes that have

> ers. T he difference in ozone-rich air in to the counts for at least 2/3 of s. Of the deficit present weakens following co ld нa

mixing by planetary w aves. Although delayed, the ensuing tran sport reduces the oz one an omaly relative to warm w inters from that pre sent a month

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er latitue

nalous Arctic oz one from Pan o to explore interannual en us eir lep dence on tropospheric reproduce the salient features of anomalou

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and a com ratively-weak anomaly of ous dow nwelling of oz one-rich air. I n richment/l eaning at extratro pical nguished by the omission of 0% of the anomaly in Arctic the observed record, the

Keyword
























Gravity waves drive large scale circulation most modern general circulation models. The effects climate models, and the effe cts of non-sta middle atmosphere circulation, including ch Gravity wave parameterizations require d horizontal momentum carried by gra vity v gravity waves include to pography, convection both geographical and temporal variations.

needed information on a global scale. Prog allowed observation of smaller scale gravi generally detected in satell ite observations as te mperature fluctuations. The conversion of measured wave temperature amplitude to momentu horizontal wavelengths and wave propagat space-borne temperature profile data have horizontal sampling of the measurements. AQUA and EO S-AURA satellit es have not o providing a fully resolved thr ee-dimensional view

, and are tr eated via parameterization in ountain wave 6nar impa aves climate nodel nis n formation d om t<u>heir sou</u>r unbala obser ns satelli

are parameterized in most lels see king a realistic it in r at forecast future ozone changes. he spectrum of vertical flux of plobally . The kn own so urces fo r ch are known to exhibit hope of quantifying the trument resolution have

ty wave s and their global properties. Gr avity waves are

all of

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on of the vertical and a momentum flux from harily due to the limited ng of instruments on the h some cases are now

waves from space. Examples from the Atmospheric Infrared Sounder (AIRS) and the High Resolution Dynamics Limb Sounder (HIRDLS) will be presented.

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level is changed. We also investigate the effect of

frontogenesis is accounted for by using the <u>Charon and Manzini (2002) scheme</u>. Our results show that the middle at mosphere climate can b parameterization: the summer mesopause and minimum temperature can change s improved at the expenses of the summe stratospheric warmings, while in general o results illustra te the difficulty of tuning the parameterization.

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h to typical tropospheri c

sources of gr avity waves: convective sou rces ar e introduced by usin g the Beres (2002) scheme;



n the gravity wa ve assumptions, its height cold pole bias can be currence of wintertim e can be improved. Thes e gravity wav e











Simulations of the recent past from thirteen couple of the CCM Validation Activity for SPARC (CCM of oz one projection s made by the same temperature, trace species and ozone in the 1999 but long -term trends are compared for reasonable agreement with observed total of spread in the ozone trends in polar regions, are in reasonable agreement with satellite Ind rad

continued co oling o f the glo bal strato sphere due to increases in gree nhouse gase s. Differences i n stratospheric i norganic chlor ine (Cly) amo differences in simulated ozone hole recover differences in the simulated Cly, with the C 2 ppb to over 3.5 ppb in the CCMs, and th before 2030 to after 2050. There is a corre ozone back to 1980 values.

y-climate nist to pr l) a /alua everal differe ΓM . The core pe exten ded pe ends a Glob SD ٦q

lele (CCMs) participating in ce for the interpretation é gui diagnostics are us ed to evaluate of the evaluation is from 1980 to (1960-2004). Most CCMs show y on a bal scale, but a gre ater eric temperature trends strat nulated ozone evolution

in the 21st century in the C CMs is mainly determined by decreases in halo gen concentrations and

g the inter-mode bst antial quantitative e varying from less than 980 values varying from of recovery of Antarctic









We have examined the 532-nm aerosol b Lidar and Infr ared Pathfinde r Satellite Obsterva stratospheric aerosol. CALIPSO makes obse profile, backscatter coefficient values repor stratospheric aerosol observations was demo Montserrat in the earliest CALIPSO data in e 2006 eruption of Tavurvur (Rabaul). Howe makes routine observations of the stratospheric ad

ati to ~40 km. T d by <u>the obs</u>e e as v ve ry

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fficient meas O) fa éir uð that spall fro serval íς load aer

ents by the Cloud-Aer osol the observation of S to 82 each day and, for each pos sibility of using CALIP SO for on of the 20 May 2006 eruption of showing the 7 October within th e stratospher e pisodic volcanic events.

Nonetheless, we found that averaging a complete days worth of nighttime only data into 5-deg latitude by 1-km v ertical bins reve als the main stratospheric aerosol data center ed near an altitude of 20 km

though the values are clea rly too small and distinctive features of the stratospheric aer small maximum in the lower tropical strates significantly improved by increasing the me 5% suggesting that the current method of likely the source of the low values.

tosphere. In addition, olar v ortices and the data. The data ca n b e erosol) by approximately sphere at 30 km is most















In this work, we examine the upper-air te timeseries (RTT) of all stations in Chines discontinuity in o riginal dat a, we co nduct identified breakpoint and adjustment reveal

The impact of homogenization to the trend 1958-1978, removal detected breakpoints fr in upper- and middle troposphere. Over the enhanced by homogenization, while the way hin averaged from full Chinese r adiosonde network

stratosphere and upper troposphere, while troposphere. However, such inconsistent has been attested to be occurred just since 1980s by a nalysis with various temporal scales. Furthermore free-atmosphere inconsistence is determine diurnal range between day- and nighttime 12UTC and the diurnal range averaged ov cooling trend and larger warming trend th trend of diurn al range decreasing versus altitude from

inved fro m the radiosonde temperature radi os network. In uali 0C) o ntro cant effe t of trary corresp inal <u>series (O</u>l e era, nds

si

toelimin ate error and enization to RTTs. The ĥorr homogen zation to orig inal RTT. ng to pre-and satellite era. During fficiently weaken the cooling trend trend m ORI of stratospher e 00 hF ere weaken. The trend, nt rariety of cooling in

warmin g trend has s ubstantiated in mid- and lower-

ver



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the warming of lower pper. To address the trend profile of 00UTC, e series shows smaller lower tr oposphere. The

400 to 100 hPa and hear a constant for othe r level. The tre nd averaged over all stations is a ffected by sampling frequency in network as w ell. wea ken cooling trend at leve l

osphere.

Decreasing sampling number of the statio of 300-850 hPa, and weaken cooling trend at







































investigated the altitudinal dependence of dynamical par ameters. In this work , we u November 199 9 to December 2001 with Japan (43.5N, 143.8E) operated by National 1999. Potential temperature and potential y tici displacement, respectively were calculated from NC

model in order to evaluate the contribution of variations of o zone. We found that about 80% of the short-term variations of ozone at four altitudes (20, 22, 24, 26 km) can be explained by th the major fact or, while the role of horiz comparable with that of vertical displacem nt an index of the horizontal transport inste d c horizontal transport has increased more that h tł the results and will discuss about the difference equivalent latitude.

nships be wee data of vert d-ba<u>s ed mill</u>ii e for l on indi of th sis dat

e oz or mixing ratio and the oz one distribu tion obtain ed from r-wa ve radiometer at Rikub etsu, al Stu (NIES) since November rizont ansport and the vertical ed a multiple regression

these two dynamical p rocesses to the short-term

cit

ertical displacement is nd its contribution is ed equivalent latitude as that the contribution of We will present details of the results by using potential vorticity and


















About 50 tonnes of interplanetary dust enters the ea 11 to 72 km s-1. Impact with the atmosphere metals which ablate from meteoroids are the sou that occur globally in the up per mesospher has been the depletion of at omic Ca in the 1L Na. There are two likely explanations for this meteoroids. This theory will be examined us ablation, which includes sputtering and impact ion

to the calcium problem p laces constraints on the distribution is currently the subject of inten se debate in the radar community. The second explanation for the depletion of atomic Ca is that the other metals. This possibility is being exa pertinent ion-molecule and neutral reactio O3, O2, O and H. The results of this study, into a new model of calcium in the upper a

tmosphere (of th és m of the laters wer thermos more than rstly, m ew k linc c ar aporat

aporize completely. The ust to etal atoms such as Na, Fe and Ca re (MLT). A lo ng-running problem ders of magnitude with respect to late m erm d

less efficiently from amic model of meteoric Application of this model

meteoroid geoce ntric veloc ity distribution. This

adically different from of the kinetics of the ric constituents such as model, will be combined

Keywords: meteor ablation, calcium layers, upper mesosphere









Instances of traveling wave trains observe undular bores. They have been observed to occur in and strong wind shears. The bores are thought to within a duct, as for tropospheric bores, not result of trapping due to inversions and bac with the lowes t mode the system can suppo since in addition to the inversion in the mes also under typical co nditions a higher ducti

been identified as mesospheric internal iation with t liated be as Morning Glo winds. Trop the mesosph vicinit in the er numerical study of ducted w aves in the upper ma nd fo

ng airglow walls, inversions generation mechanism รดี henomenon, where ducting is the ere bores are generally associated he ducting can be more comple x, in asso ion with bores, there is . This talk reports on a no sp nditions (e.g., inversion

upper mesospheric and lower thermospheric are found. We find ducted modes that agree with observed phase speeds and wavelengths. These mod with the inversion in the mesopause region more vertically extensive part of the total winds similar to those described by Smith the wave in t he lower stable layer is much stronger more nearly resonantamplification than



stable layer associated an be the stronger and so spheric inversion and 9500, 2003) the part of pper layer, and exhibits





It is considered that gra vity waves are hodograph an alysis for small vertical- scale wind an detemperature flue stratosphere by radiosondes indicate that severa downward. This means that there exist gravely and Sato, 1999). We made intensive radiosend days of each season at Syowa Station in the ln using an aqua-planet model with sufficient l g۲ the analysis of these observatoin and mode lata around the polar vortex in the stratosphere is discu

ent o ens e source in serv ations w rctic <u>in 2002</u>. o lutio res deta e n

tions obser ved in the pola r aves pr opagate energ y a vity olar winter st ratosphere (Yoshiki me intervals of 3 hours over ten eover, we performe d a simulation gravi aves explicitly . Through ravity wave generation nism

















the Arctic middle stratosphere. This EME capture probabilistic character of its phase tran sitions. Also, the Arctic stratosp here at 10 hPa is about 2 K warmer during the easterly phase of the QBO than during the westerly phase. A second EME including a

time series of the solar radio flux at 10.7 temperature in the Arctic stratosphere is sh periods of h igh or lo w so lar activity, resp correlated with the no rthern annular mode stratosphere propagate into the lower strat of about two and fours weeks, respectively. T confined to the lower stratosphere.



s of the QBO and the

en the QBO and the es by about 1 K during bles chosen are highly ano malies in the middle osphere with a time scale stron g tropospheric NAM anomalies is





The SABER instrument onb oard NASAs temperature in the stratosphere, as well as January 2002 to date. The measurements h descending orbits, except for short breaks d makes SABER observations especially well s (FFSM). Although the FFSM has usually bee polar-orbiting satellites, we show that it is straightfor

length by interpolating the observations to a common Salby s-coordinate. This is useful because, given that the orbit of TIMED precesses in local time. it allows resolution of the diurnal tide when at least two yaw cycles are mapped as a single data se construct a continuous climatology of trave We show that, in addition to the mi grating id the presence of fast, planet ary-scale way inertia-gravity waves, and no n-migrating t tles amplitude of these waves and attempt to identify the factors that produce such changes.

MED mesos pl n made almo spacecraft va r map b to of ati general

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glo bal observations o f ower the mosphere (MLT), from ntinuously, on both ascending and ane <u>uvers</u>. This sa mpling pattern bys Fa ourier Sy noptic Method aken ng a single yaw cycle of ta segments of arbitrary

on of FFSM mapping to on tropical oscillations. s strongly influenced by day wa ves, Kelvin and erannual changes in the d

















Fourier coefficients of 3D-decomposition of model simulations have shown that no nimeal Atmospheric (0,0) emission layer results t nightglow. The magnitude of these mod ulations brightness.

itic ้อมร \mathbf{O}^{\dagger} can achieve value

d by gravity waves. The bassing t hrough the O2 brightness of O2 sky of 10% of the background



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decadal modulation of the str atospheric QBO. Involving high-frequency variability (periods shorter than 5 yrs), it too varies with the 11- yr cycle of UV irradiance. The cyclic m odulation of the QBO makes possible a relationship to the tropical trop term average (upon which prior studies ha in their relati onship reveals coherent chan interannual varibility, their relationship mai the Hadley circulation. It operates on the s (TBO). L ike t he QBO' s relation ship to the polar stratosphere, its re

maximizes in the tropics of each hemispher

distribution of an omalous temperature, sug



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and anomalous z onal wind tha t ıre, in cert with the ho rizontal gani convection, as well as a Perhaps related is a

e invisible in the longts for a cyclic variation tropos phere. Involving ciated with the QBO and pheric Biennial Oscillation lationship to t he the tropical troposphere reverses on the time scale of a decade. The systematic swing in the relationship between

the QBO and the Hadley circulation mir stratosphere, where the phase of the Q troposphere, d irectly, throug h the tropica circulation, which couples the stratosphere and troposphere through transfers of mass.

oz che h eating in the upper cts with the tropical gh th e Brewer-Dobson

Kevwords: atosp





The total solar irradiance (TSI) and solar sp nm to 2400 nm have been measured by NAS was launched in January 2003. The Sun is t irradiance and its variability are important in surface, and oceans. This talk will provide function of time and wavelength during the and progressing to the low s olar activity in 200 (minutes-hours), active region evo lution and so la

(months-years). The variations in wavelength ar photospheric emissions, which dominatein the near infrared, visible, and near ultraviolet ranges, vary by about 0.1% over the 11-year solar cycle ar emissions from the solar chromosphere and

ultraviolet and far ultraviolet ranges, and associated with the e volution of the bright coronal emissions, which dominate in the X by factors of 5 to 1000 over the solar cycle

e (6SI) from octral irra 's S n and Radia external prci r studies of view<u>of</u>the s ni ssio art h e dd avs),

nm to 34 nm and from 115 eriment (SORCE), which nate or climate changes, thus the solar vith th oral

ner getics of Earth's atmo sphere, irra diance and its variability as a gh solar activity in 2003 t ions are due to flares ycle magnetic evolution

e dependent on where the emissions arise. Th е

nd bright faculae. The fied in the extreme to a factor of 3 are on the Sun. Finally, the ne ultraviolet range, vary











Andreas Fischer, Stefan Brn

It was hypothesized that duri ng the first ha consisted of the steady increase modulated with General Circulation models driven by th climate response have bee n underestimate the spectral so lar irradiance variability in c century we have carried out transient ense 1900-1960 driven by time e volving solar pectral

we present the solar signal in the atmosphere ex regression analysis technique. We also def to the time evolution of the simulated quar decadal scale solar irradiance variability. The from the transient simulation covering 197 -20 temperature extracted from the observatio

of t 20th decadal cale es in total so ipar i<u>son wit</u>h nd ozd ha with Cher , sea s

of the so lar irradiance tuatio ne Numerous experiments radiance showed that the obtained ervations. To estim ate the role of durin first h alf of the 20th /- Clir Model SOCOL covering perature, sea ice , land

properties, sulfate aerosol loading, greenhouse gase s and ozone destroying subst ances. In the paper

Egorova, Werner Schmutz

tracted from these transient runs using multiple

lar irradiance change s to the long-term and the solar signal obtained al in the land surfac е

Keywords: solar irradiance, climate, ozone

ble



Martin J. Jarvis, Richard

Climate change occurs on various time scale due major source of natural variability is the Sur both influence atmo spheric circulation by ch affect weather systems and climate patterns a quasi-11-year cycle, while t he solar wind field causing changes in at mospheric che irradiance and geomagnetic activity on the atmosp

0.4 K increases in temperature anomalies the stratospheric quasi- biennial oscillation stratosphere, comparable increases in tem with geomagnetic activity. I n these region easterly, or ge omagnetic activity is high, westerly, or so lar irradiance is high. High solar i each others signatures either making the sign

extending the signals to broce order areas, or stratosphere and are 2-5 times larger than the in Ta is asymmetric about the equator, wil high Ap and Fs, suggesting possible disturbances the QBO. the positive temperature ano malies resulting from

rise in comparison to greenhouse gases?



F. Graf, Peter C. Young

hropogenic causes. On e rra diand, and the solar wind ca n ing rates which may subsequently e at<u>mosphe</u>re directly and follows ndirec ia the E arths magnetic es th lative effects of sola r honthly to inter-decad al

time scales. Geomagnetic Ap (Ap) signals are found primarily in the stratosphere, while the solar F10.7cm radio flux (Fs) signals are found in both the stratosphere and trop osphere. In the troposphere, 0.1-

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anals are found when blar region of th е vith solar irradiance and hen eit her the QBO is vhen either the QBO is

rradiance and geomagnetic activity te nd to enhance als stronger and sym metric about the equator or

dominate the middle Arctic lianc low, the signature of Ap rct suctosphere and negative signals at mid-latitudes of the NH strat osphere. Weaker stratospheric QBO signals are associated with he signals o Tand Fs are distinct from eanic er s. These sults will help to addres s a key question of atmospheric research: to what extent solar activity has contributed to the temperature























It becomes cle ar during the l ast years that also energy of the solar particles and the sol problem. Long term variations of the solar v of the satellites observations for period of th density and velocity of the solar wind, full ve field, connected with sector structure of sola of the Sun w ere considered. It was found periodicity of the variations. Full vector of the inter-

12 cycle activit y. Maximum of velocity of the solar wind is observed usually 2-3 years later after the maximum of t he S un spot activity. T he <u>vector</u> of magnetic field conn ected with large-scale weak magnetic field of the Sun has period of ch variations of the so lar wind density. Energy energy, which can be trans mitted into th geoeffective parameters of t he solar wind However it must be taken into account that l tra depends on state of the atmosphere and ground surface. Local ground surface electric conductivity is an important parameter of this mechanism. It is not

parameters in global scale, which is typical of temperature in the Arct ic basin demonst character for large weak magnetic field of the to solar-terrest rial relations could discover other weather and the solar activity in its different form

ity includes e sola This S on i ameters ere 0 years. Tem the i<u>nterplan</u>e and d nir ete hagnet

solar UV irradiance but of such approach to the exam yzed in this report. We used data I variations of such parameters as magnetic field, vector of magnetic y larg ale weak magnetic fiel d the s wind have different

emonstrated strong 11 -

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period is observed in unt of electromagnetic differe nt periodicity of effects on the Earth. solar wind into geospace

d. However, I ong term changes

reasonable to lo ok for periodic ity of atmosph eric

e q to \sim 40 y ears, which i s on/ to ope that such approach allo ways of coupling between the Eart hs climate and

Keywords: solaractivity, solarwind, earthclimate



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cycle for time of Maunder pe riod and for 1720-1850, but in opposite phase at the end of MM (1700-

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simultaneously were raised about twice . The quas

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coincidence for periods of these oscillation

quasifive oscillations are over the root of the

number indexe s. The res ults obtained cor

period the decline in solar activity may hav

1705). In little extent, the similar pattern is seen for Dalton minimum,

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and t he end of MM

phase with Suns 11 yr

is some evidence that

or BWPI so for sunspot

nthropogenic (pre-1850)

erity that re flected in the

close

Th<u>e fine</u> structure of quasifive

is showed the







tree-rings reflect the state of solar magnetic tree-rings, we can trace back the characteris periods. The S un holds several long-term cycle and the 22-year cycle in the polarity variations of the Sun have caused several long-las

(1645-1715 AD), and have br ought cold spells as referr ed as the Little Ice Age. Howeve r, not only the mechanisms of such long-ter m solar variations but_also the mechanisms of so lar influence on climat e are not clarified yet. In order to clarify the climate, we investigated the c hange of the 1200 years, including the grand solar activ (1415-1534 AD) and the Ma under Minimum b annual time resolution. The spectral analys s o the eleven-year variations a nd the slight stretching of the cycle lengt hs during the grand activity

activity maxima. The 22-year polarity reve

of the 11-year cycle. We compared these r

years, and found that the significance of the

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he radiocarbon contents in annual lar cvcles during the pre-historical n to tl 1-year s unspot activity and 208-yea r quasi cyclic the Mau nder Minimum

st and their effect on ar cycle during the last y, the Spoerer Minimum ntent in tree-rings with realed the suppression of minima. On the contrary, slight shortening of the eleven-year cycle was fo und for the grand solar

in an ociation with the change d teŋ ratures for the last 1200 er than hat of the 11-year cycle especially around the gra nd solar minima periods. It suggests that the polarity change of the Sun is taking important role in the multi-decadal climate variations, and that the electromagnetic forcing is UV radiation.


















We examine whether the teleco _____nnectivity of atmo depends on t he phase of the solar cycle. T correlation maps and effects of modes of log-fr precipitation), both in 500 hPa heights. Win classes by the level of solar activity, measur are calculated at each gridpoint with the 50 At each gridpoint, the area with remote calculated in each solar activity class separately.

indicate that at the majority of gridpoints, the teleconnections are more spatially extensive under a high solar activity. The difference in correlation midlatitudes over the Atlantic and Pacific o modes have been detected recently. In variability are calculated in mont hly mea surface tempe rature and pre cipitation at a between a high and low solar activity.

ric circulatio é tel ivitv j con ency variability nths (Decem -10.7<u> solar f</u>ι eight fi corre ns th ar id

the No rthern Hemisphere ed here by one-point aract surface cimate (t emperature and to March) are divided into thre e (ithin<u>each c</u>lass, co rrelation maps orther emisphere north of 20N. r thar ative threshold is eg agreement between the

correlation patterns for diffe rent solar activity cla sses is qua ntified by pattern correlations. Results



axima is largest in the activity in the variability des of low-frequency te that t heir effects on ro pe sig nificantly differ

Keywords: solar activ rculation























representing a spects of Earth's climate, but he w much do they mean? 20th century, the only one for which we have go emphasize the importance of reliable uncertainty best guess is that it has zero likelihood of ben the difference between the patterns (in spac also be presented. This sort of material is outside it. Some comments will also be mad amplification" from GCM simulations.

tional obse iges - wit out rrect. GCM ex e) of <u>climate</u> 8 withi G far he us Foi lab

novel examples for the vill be presented, which /erage n, after al, all we can say about a les which show the importance of ng &<u>those o</u>f climate response will ommu , but less appreciated filteri & on evidence for "solar



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