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	Earth's lateries				

UGG XXIV Ge	eneral Assembly July 3, 2007	Perugia, Italy	
Abbreviations			
	International Associati		
		onomy	
		Colonna	
		Oceans	
		the Earth's Interior	
	Climate and Chusenber	istry of the Earth's Interior	
	Clobal Energy and Water Exercised		
	Hindu Kuch Himalayan Lou Daimas from	vrational Experimental	
	and Network Data		
	International Association for a logic moo	tanh	
	International Association of Cryospheric Sci		
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	International Commission on Antice Representation Politica		
	International Commission on ontin ota E		
	International Commission of Clin		
	International Commission on the Coupled L	and-Atmosphere System	
	International Commission on the coupled E		
	International Commission on vnami Me		
CGW	International Commission on Groundwater	old gj	
	International Center for International Center for	n Development	
	International Commission on the Middle Atr	nosphere	
CRS	International Celestial Permitting		
CSIH	International Commiss	droloav	
CSW	International Commiss	55	
СТ	International Commiss		
CWQ	International Commiss		
CWRS	International Commiss		
GAC	International Global Atmospheric Chemistry		
GS	International Glaciological Society		
LP	International Lithosphe		
	International Union for		
ON	International Ocean Ne		











IUGG XXIV General Assembly July 13, 2007 Perugia, Italy

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JSS001

Symposium Physics and Chemistry of Earth

Convener : Prof. Ian Jackson

Insight into the physical and chemical beh seismological, geochemical, geodynamic, a provide direct access to the conditions of

slabs, the transition zone, lower mantle and core. Co of materials under extreme conditions difficult to behaviour of mineral surfaces and crystal effer molten materials, are increasingly amenable to be promise a better understanding of the Earl's per aspects of contemporary research into the pays and

e. Considementary the occurring high borat as well as the b estigation. The de erior - from constant nd charge y mar

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a temperature of subducting lithospheric ntary theoretical studies predict behaviour borator experiments. The structure and as the bulk projecties of crystalline and The developments in materials science in control core. Papers are invited on all

1716 - 1731

smology and Physics of the Earth's



































Burtny Pet n Kar

racteristics of rocks are

sotropy of the physical

e causes of the initial

barameters in different

by combined studies of

idies show that in their

with clear-cut elastic-

zons by boring.

The crystalline rocks usually form textures being rsible change mineral matter under the effect of stress fields, temperature and different get al processes. rocks may "flow" at definite PTlh parameters in the deformation process and here 🖥 with es with linear, planar and linearplanar orientation of crystals, grains, detes struct an xtural closely related with elastic anisotropy. Ac Irdin the ract f the parameters of the textured media, the latter m ne classes of symmetry ferentia associated with the history of their formation and by the symmetry of the external effects. After having defined the elastic symmetry of rock it is at difficult to recover the directions of the fields of the external effects. But the difficulty of thes anisotropy of rocks and the inhomogene directions caused by non-structural feature elastic-anisotropic properties of mineral ma ditions. Vp was measured ter ffei in 9 oriented polyhedrons of rocks and that ostatic pressure and the ul 'hy programmed PT-conditions corresponding with rock sampling sites. The elastic anisotropy and the structural and textural features of samples of a core of rocks of the Krivoy Rog ultradeep borehole from zones of destruction and non-destructed eff t of paleotectonic and technological stress in rocks on their dest geologic evolution the rocks (plagiogram luct ires anisotropic characteristics typical of transversely isotropic media. Then, under the effect of recent tectonic processes characteristic of synclinea, and transformations The rock occur that are initial stru detected by petrographic studies as secondary changes of ture or texture and are marked by elastic anisotropy distortions. The character of the elastic anisotropy of the borehole rocks at different pressures and temperatures indicates the orientation of the action of tectonic forces not compensated by lithostatic pressure. Most tion of the relaxation of these forces at an angle of 45 to the borehole shaft. Th estructive effect in plagiogranites when their layering is oriented along the layering), while the stress is mainly relaxed in the stratification plan orientation and micropore localization under the effect of recent tec crushing zone, the rocks most probably experience the effect of ter ers) due to under thrust, which also favour the borehole shaft destru

Keywords: high pressure and temperature, anisotropy







IUGG XXIV General Assembly

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JSS002

Symposium Tsunami: generation and haza

Convener : Dr. Kenji Satake, Dr. Gera Co-Convener: Prof. Efim Pelinovsky,

The 2004 Indian Ocean tsunami, the wo tsunami research, warning and hazard mit Tsunami Commission formed working grou

level data and satellite data. Regional tsunami watch the world with a guidance of UNESCO IO communities world-wide using the state-ofhazard assessments have been introduced i forums for information exchange and discus after the 2004 Indian Ocean tsunami, as we will be carried out as Workshop on Wave an Tsunamis. Following the 2004 Indian Ocea access tools, archive, and analysis techniques of

tsunami event analysis became apparent. The objective of this workshop is to establish standard procedures for data assessment and data archive for tsunami event analysis. The lessons learned from this workshop and subsequent re-analysis

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ompile field survey data, instrumental sea up in basins and oceans in nented in many coastal ēimī techniq es. Probabilistic tsunami he world. This joint session offers ts of tsunami research, particularly A part of this session product development for the availability, quality, and water-level data for

e water level events.

Imamura

cument the 2004 tsunami, IUGG

had a significant impact on

smology and Physics of the Earth's






































tens millions of euros. Unfortunately, there were po during the event. However, we have been efficient 2D numerical model of extreme lo for the period 1989-1999 with focus undertaken in 1997 which included a tria gle gauges installed on the shelf and in bays nle these data to estimate magnitude, speed and direction of travelling atmospheric disturbances and waves, we also consider those cases for noticeable sea level response. Parameters numerical model. Simulated ocean waves are

in the inlet or on the shelf

data to formulate an pecifically, we use data al experiment LAST-97 d eight bottom pressure a islands. We re-examine

associated generated long ocean waves. To understand the generation mechanism of meteotsunami distur ances did not produce a e used as inputs to the bse value. The verified model is _q then used to simulate the 2006 event. Finally, the sea level measurements and numerical computations are used to assess general tsunami issues for the Balearic reion which so prone to tsunamis of seismic origin; a very recent and well-known example is the 20 ligerian ts

working tide gauges

on



Keywords: meteotsunamis_field_measurements, modelling





































Allison Allen, Natalia Donoho, Stapher Eli, Tom Mart Bolin Meyer, Manoj Sama t, Fruert Assinal

The United States National Oceanic and At ric Administra (NOAA) funds several water level observation networks in support of operati These observing systems inam pability include NOAA-funded stations operated by enter, stations operated ersity aw ea Lev by the NOAA Tsunami Warning Centers (T) NOA ART ork, and the stations of the NOAA National Water level Observation Network cribes the NOAA National 0. This Ocean Service (NOS), Center for Operational Oceanographic Products and Services (CO-OPS) contribution to tsunami warning. CO-OPS i ble for operating the NWLON which has supported spon tsunami warning since 1948. NOS continu ulti-purpose data for a variety of applications, including real-tir level trends, habitat restoration, computation of tidal datums er the December 2004 lar Indian Ocean tsunami, CO-OPS was taske expand and upgrade the tsunami warning capabilities of the NWLOI A existing NWLON stations irad with new Data Collection Platforms (DCPs), implement data formats, and fill observation gaps. Work began in 2005 to upgrade 33 existing water level stations and install 16 new stations in priority areas in the Pacific Ocean and the Caribbean . ted upgrades had been targ completed, as well as 15 of the 16 planned 7, CO-OPS plans to have iber completed the upgrade of all coastal NWLO atis tsunami requirements. As of January 2007, a total of 100 NWLON stations are operating in support of the National Weather Service (NWS)s tsunami warning capabilities. Eval though the information If still be transmitted via the Geostationary Operational Environmental Sateline (GOES) both the primary and backup DCPs, for the first time the upgraded DCPs will be transmitting 1-minute averaged water levels every six minutes. 6 and 1-minute data are available to the TWCs directly through GOES, through remote phone dial-in, and through the CO-OPS web page ore 15-second data on a flash drive for post event analyses and modeling. The 15 ally downloaded from the station itself or remotely using the DCPs moder CO-OPS and the TWCs has involved more than just the expansion of but an expansion of a total capability, as evidenced by the CO-OPS su ants for each DART buoy site, so that effective de-tiding can take pla als.

Кеуш
























































































































preliminary conclusions on the possible role of the mapped landslides in generating the 1693 tsunami.






Morocco, by its geographical situation with vo west, and along the Mediterranean coast in risk of the tsunamis coming from the Atlan between Rabat and Casablanca the tsur transoceanic tsunami, report strong deva 1969.02.28 Horseshoe fault (HSF) earthq ake amplitude recorded at Casablanca. Recently, another

moderate magnitude (Mw=6.1) prevented the triggering of a significant tsunami. In this study we present the preliminary results tsunami propagation and inundation results for Rabat and Casablanca areas. Modelling was performed with COI covers the eastern part of the Atlantic Od prone tsunami generation area. Three n 0.0005) are incorporated to obtain a go Ы shore. Results of the numerical simulations maximum velocity. This study was funded by projects

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the Atlantic coast in the s alo estern Africa more exposed to the h. In high populated coastal areas Data on the 1755.11.01, nown_ Agadi the last century the sma sunami, with maximum southern Iberia but its

he simulation domain f Cadiz, from the most tion (0.008, 0.002 and opographic effects near heights, flow depth and and NEAREST STREP 37110 UE.



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part of the Severo-Kurilsk town, up till now non of our country, the height being more than 5 wave heights being up to 15 meters in so coast is not homogeneous and tsunamis are is known that a sequence of earthquakes and its main parameter is the recurrence function

height more then elected "threshold" h0. Analysis of the natural data showed that this function depends from two parameters: $F(h0) = f \exp(-h0/H^*)$. Parameter H* is calibrate (characteristic) tsunami height depended from the coastal point of tsunar varying very slowly along the Pacific coast square method these parameters have be characteristic tsunami heights H* were e Tsunami frequency for this region is f = 0.7parameter allowed to estimate tsunami risk for this region.

an 40 tsuham ers in 7 case nts. <u>Tsunam</u>i equen ent an e d n ch is à

re recorced on the Far East coast d in 1952, 1963, 1969, 1994 the ard distribution along the Russian the So ern Kuril Island coast. It itude eally the Poissonian one quency of tsunamis with

unami frequency, that onstant. Using the least unami data set. Finally, thern Kuril Island coast. h is $\sigma(\ln(f) = 0.13$. These The work was supported by the grants of

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the Russian Foundation of Basic Research 05-05-64733 and FEB RAS I # 06-I-ON3-106.

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A huge tsunami was generated by the south of than 500 people were killed at the village 0 international resort point of Pangangdaran. Kato during July 25th to 29th July on the c west of Pangandaran. The second one was hade near Yogyakarta about 200 kilometers easy of Pang

at Pangandaran where 137 peoples were killed due to the tsunami. The length of our surveyed coast is totally 400kilometers. The run up heights the values of generally 5 to 7 meters and The tsunami magnitude in Imamura-Iida's and found out that the shaking of the earth Most of the people living on the coasts we event was induced by a kind of tsunami ea

htral Java ear e Indian Oce nducted fi m Pan ara ring A st 4 the max

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ake of Juy 17th, 2006, and more oast of Java Island including the The first one made by euk abou 150kilometers Pame from nup east to Parangtritis, n-up height is 7.7 meters

villages on the section of the surveyed coasts keeps of our surveyed coast. erviews to the habitants le i hearest to the epicenter. it is suggested that this

Keywords: tsunami run up height, tsunami of indonesia, tsunami earthquake

these



















IUGG XXIV General Assembly

July 13, 2007

Perugia, Italy

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JSS003

Symposium **Early-Warning Systems**

Convener: Prof. Jochen Zschau

Early warning systems can reduce the ne earthquakes, tsunamis, volcano eruptions and others. They provide timely information

to avoid or reduce their risk and prepare for warning time available may range between years in case of creeping disasters such as technologies are now available to a great ex used to any appreciable extent, showing Integrated earth observation, the development their integration with appropriate facilities for systems and their application to early wa hing remains one of the major challenges of applied earth

scientific/technological developments, projects, programs and best practices in the field of early warning. This will be done for various disaster types, from global to local scale and with special regard to the user needs. Contributions are part early warning.

smology and Physics of the Earth's

1845 - 1892

of extreme events such as floods, extreme space weather duals exposed to a hazard, to take action

se. Dep offectiv າດເ nly ds in ŵ sê dro its. Altho igh eir potential ne user need al-ti ta pro onju n v sciences

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on the type of hazard the hquakes to months and of d information- and early warning e field of disaster mitigation is not e not yet addressed sufficiently. node Ad simulation methods, and rapid information ualiza nanagement, therefore, disast sion will discuss the latest

aspect of multi-hazard






























































an automatic change detection scheme. The index of local (in space and time) change, which is at the basis of the classical RST approach, is here integrated with a differential index. computed by using RST

preion as well, which permits us to ident temporal repetition of the sensor. A po discussed, analysing its actual potential a and efficient early warning system. Moreov polar e.g. NOAA/AVHRR- and geostationa applicability to other present or future sensor data

d

satellite records and on

ht, exploiting the high uch a scheme will be evelopment of a reliable (already applied both to guarantees its complete

















A surface landslide cased by collapse of ab world. The landslides can cause damage to of such territories, and sometimes result in depth up to 80 m in Donetsk city (Ukrai landslides and, therefore, earlywarning ma reliable plans of the mining excavations an which relates to the physical characteris existence of such cavities in abandoned undergroup

combined influence of different rock layers and uniaxial compressive strength of the immediate roof over underground openings on existence sandstones in the rock mass or relatively h contribute to the existence of cavities in a thickness, weak alluvium, argillites and sa the opening and allow collapse of the roc m abandoned subsurface mines defines l bat thicknesses, alteration and types of rock layers, thickness

ned_ mines uildi on th /erlv man life! Cav potentially ese c<u>avities</u> mposs to e ov ng hgs at

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been in many parts of the hinder the economic use urfac in the andoned openings at a erous from the point of view of eded_Unfortunately, there are no enings. Empirical model such a wa roposed to predict the epth. The model shows

_cavities in the openings. Large thicknesses of in overburden.

in the immediate roof d, sandstones of small nt on immediate roof of mapping for cavities in ding on rock strength,













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Destructive earthquakes dated back as far a historiography the following destructive ear (972, M6.5), Ararat (1319,1840,M7.5), Gar 7.0). These earthquakes caused thousands of the Spitak 1988 devastating earthquake mo injured, and 515,000 people were homeles villages of the Northern part of Armenia were alm

and structures in Armenia are designed for ground acceleration values of 0,1-0,2g, that correspond to 7-8 value by MSK-64 intensity scale according to seismic zonation map operating in the territory of Armenia up to 1994. It is obvious that the the designed values of accelerations of destruction of buildings and structures is e Development of Earthquake Early Warning destructive earthquake is getting essential for Yerevan city. The new system will consist of the

assessment and early non-urgent warning (preparation phase): the current seismic regime is evaluated, and in case of relevant hazard the warning bodies, and, if necessary, to public and com rapid information is necessary): the principle of velocities difference between electromagnetic waves propagation and seismic waves. It was focused on the solution of the following tasks: possibility and determination of necessary number of 44 upcoming strong earthquake and others. The first subsystem is dealing with the results of the current seismic hazard which based on the monitoring data of the more than 40 seismological, geophysical and

geochemical, and hydrodynamic, and geo their correlation with the seismic events, The structure of the second subsystem and been developed.

he data of the Armenian ding vin (851-93, M6.5), Talin Arouch zor (1827, M6.5), Spitak (1988M mic losses in Armenia. Only during es, 20,000 people were st the nd str ires in all the cities and h present, the buildings

considerably exceeds d the hazard of great earthquake in Armenia. mation about upcoming he new concept of EEWS two components: 1.Current seismic hazard

at and related governance

uake hit urban area and (ea developed on the base EWLW ich, configuration of EEWS of EEWS crea ein in tim the warning in case of well as on the sizes of anomaly and

> te software has been developed. the special seismic signal have

Key








red geo-sensors structure

bn of the geodynamic

(v) provision of early-

thodology and software

ctromagnetic parameters

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esults can

(test sites) triggered by the earthquakes (F red i The activities that have been accomplished consist ovation for network conditions; (ii) implement and continuous improvement of the real-time monitoring system depending on the landslide test sites condition real time signals processing for pattern recognition s; (iii) in pre disaster and at disaster circumstar precursory parameters related to the land warning against the risk arising from land slid packages have been applied for obtaining l in' tin and to point out their anomalous beha viou geodynamic conditions. Additionally, by combining different data types and analysis techniques, and also by merging electromagnetic parameters with geoelectric tomographic images and with low frequency electric signals occurred prior paradigm, in which is emphasized a earthquakes magnitudes, was carried out. Sus (test site), it was possible to assign the increase of the landslide activity related to the local fault which has been reactivated by the EQs occurre the Vran a zone. In the stage of the system implementation and to wh extent de Sus) site may contribute on understanding such kind of phenomenon in order to provide the information necessary for disaster mitigation.





For updated seismic monitoring and prepar Ines waveform data exchange with other seismi related research Geophysical Institute of Is Network (ISN) and the ISN acquisition syste Nanometrics Co. The ISNAS is based on sa time automatic waveform analysis. The proj based acquisition to the continuous waveform re-

stations to the new ones; 3) Changing of the seismic network configuration; 4) Modernization of the existing software for processing the continuous data. The ISNAS structure comprises four networks: 1) Site network (between data loggers and sa data center hub; 3) Data center acquisition servers and connection between the mai network. The ISNAS includes several type relay communications based on the SeedLi kа real-time detection, picking and location as a part of the

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as well as for improved d for the progressive earthquake It modernization of Israel Seismic the contract between the GII and cation and close-to-real comr 1) Tr formation of the triggerthe part of the existing

n the stations and the tween the hub and the sis network internal GII ellite, broad-band frame cludes new algorithms of JSTAR analysis system, developed in GII.



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Peter Foden, Jaff Pug hilib Wood

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The Asian tsunami of December 2004 was the the Intergovernmental Oceanographic Comr Level Observing System (GLOSS) where new GLOSS guidelines were amended so that i network include tsunami monitoring capab ty a and intervals (e.g. 1 minute sea level samples

development and implementation of "tsunami enabled" tide gauges at Proudman Oceanographic Laboratory, which allow real time data to based around off the shelf components w platform, which performs the data logging OTT Hydrometry Kalesto radar and PS1 Telemetry is through the Inmarsat Broadband directional broadband connection over ethernel transmission is via SMS messages which contain 1 minute values from 3 sensors and are returned over Inmarsat's private network every 5 minutes.

devastating i held a meetin nmendati ons wł gaud ransi sea data ery 15 return

dern history. Following this event, der the auspices of the Global Sea for made tide gauge installations. are part of the GLOSS suitably high frequency Here we present the

d from almost anywhere on Earth. The system is wer, embedded Linux ed so far have been the over a serial interface. stem which allows a bition when required. Data

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JSS004

Symposium

Non-instrumental seismometry earthquakes: balancing the geological, instorical and cont records

Convener: Dr. Gianluca Valensise

General scopes related to Symposia on Non-instrumental seismometry. During the past 20 years, the seismic hazard assessment practice has seismicity based on non-instrumental disci seismology. Over the years, these diciplines more quantitative. Quantification involves n recent and more distant earthquakes (complexity), but also the rate of earthqu development of segmentation models, the triggering. The symposium intends to act lowled

seismology in recent years and particularly since the mig-1990s, emphasize their mutual relationships and show how they integrate with conventional instrumental methods. Due to its long-standing tradition in the analysis of the historical earthqui methodologies in earthquake geology, Ita leading role in this symposium. This is t JSS007. Below the scopes of JSS004. assessment practice is the comparison evidence for ongoing tectonic strain. The c geologic and tectonic parameters on the other h earthquake, the size of impending earthquakes, and the expected rate of earthquake production in any given region. Significant over- or under-estimations

diverse conditions as source complexity du q dynamic fault interaction, failed identification ∩f intends to draw on scientists from different ne disciplinary boundaries and compare their approaches, results and residual uncertainties. We especially welcome contributions from these areas, or contributions schemes: - quantification of the historical earlingua including field studies of cumulative tectonic strain and the development of fault segmentation schemes; partitioning of geodetically-derived strain onto individually identified or areal active tectonic structures.

Papers are expected to emphasize the impa of i) the understanding of the seismic cycl estimates.

1893 - 1904

d future mporary strain

ion of past and future onics, 1 eoseismology, historical herely descriptive to progressively eters of the earthquake source for lenath. coseismic slip, rupture rrence perties of a fault, the ions leading to dynamic ric co chieved in this area of oare

> the development of ate country to take a \$004, JS\$005, J\$\$006, modern seismic hazard geologic and geodetic s on the one hand and of

the basis for assessing the maximum credible potential may derive from such pog assessment of fault size, mic creep. This session to cross conventional

at combine em into unconventional quantification of the geologic record, results or approaches in the improvement

ii) mid- to long-term seismic hazard



Historical earthquake catalogues in the India in the News Papers, writings of Courts investigations. Many new damaging earthquake scale were added in the catalogue based or earthquakes has brought out discrepancies location. During the early instrumental era, 1905 in western Himalaya based on limiter inference about two earthquakes one in Kungra ar

Region nerally b ind larcha lorian anging in seis ove. Howeve time_ of occ re-inte tat pgical a h near

he felt reports published nd Paleo seismological 5gicaľ intensity from VIII to XII on MM ser examination of some of these ce and in some cases even their f the t Kangra earthquake of y questions, due to the ised Further look at the data

has reaffirmed the occurrence of only one earthquake of 1905 in Kangra. The other inferred earthquake near Dehradun was only a consequence of site response called secondary meizoseismal area similar to that observed during great earthquakes d historical earthquakes based on limited dat only if old the seismograms are preserve was initiated in India Meteorological Depar Program on Historical seismograms and earthq their preservation has now led to the modern method of digital scanning of the record. New results



d Mexico (1985). The hich can be undertaken f historical earthquakes he first phase of IASPEI. odology and problems in





















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Italy is earthquake country, but investigatin hidden or positively blind. Geological strai seismogenic areas, and seismic hazard stud GPS networks are still rather sparse and the the country. Our work attempts to bridge th may be used to assess strain-rates in most seismogenic areas. We computed the stran-rate in

faults, GPS measurements, stress and strain indicators, and tectonic regime. Different ways to assess strain rates have obvious advantages and disadvantages, due to the different spatial coverage of data, different temporal scale, and probably dif rates. Using different classes of observatio and show the inconsistencies where the d observations relate to the future strain integration guarantees a more robust result that error procedure, we build a series of 3D dynamical models of and surrounding regions, compute the

velocities, stress, and strain in the thin-shell approximation by using the finite element code SHELLS (Bird, 1999, Computers & Geosciences, 23 available observations. We test many f conditions, fault friction, rheology, etc., and deviate less from the data. We build the structure of the models based on the most recent information on the crustal structure, faults, and rheology. The shear tractions are adopted from the literature. To use regional published datasets and compute the deviation of each model with each dataset. Last, we

signal that cannot be modelled by our app with data (we do not reproduce 17-33% of of 26-30 degrees, and the RMS of velocitie rate of 10-16-10-14 s-1, compatible with t pattern that accounts for the different data observations leads to very different result appear to be more robust.

if not most of them are ently as essed only in very few earthquake rates from catalogues. ed by the peculiar configuration of geodynamic model that tructir pssibl identify as yet unknown : seismic events, active

> ct to the future strain f the separate datasets face how the different hips not yet clear, the . By means of a trial-and-

odel predictions with the

es, including boundary

dels whose predictions

s and the ideas of basal

with the observations, we

combine the single deviations assigning a weight to each dataset based on a possible estimate of the r model predictions satisfactorily agree ns, the predicted azimuth deviates active areas, we find a strainwith somewhat a different hat using only one class of ata, whereas our results

Keywords: strain rates, italy, numerical models





Seismicity rates in eastern Canada (east completeness threshold) per year and beta of a size that had it occurred in the top of t needs to be reduced by about a factor of through the seismogenic crust and also ~2 years (when knowledge of M6 events is fairly we know of just one: the 1989 Ungava su undiscovered surface ruptures. Furthermole a rate

since eastern Canada was deglaciated, ~100+ surface ruptures should have been formed. Arguably we know none, though some candidates are beginning to emerge, and evidence of earthquake shaking events is accumulating. If the Canadian ice ~50,000 years as in Fennoscandia, we sh the ice margin of the day, i.e. 5 times mor yet. Excuses might be made for the Cana craton due to the scale of its glaciation, or rapidly destroyed (perhaps as in southern Sweden?) remains profound.

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7 M>=4 (approximate >=6 events/year, each but C duced a surface rupture. This rate br distribution of events in depth surface ruptures. For the past 100 /e had *100/10 = -1 ruptures; w otł earthquakes might have h the circa 10,000 years

ssing the seismicity for events occurring near have been confirmed as from the Scandinavian hg under the ice and thus gap between the predicted and the known



















Keywords: fault, segment, parameter











the well-known scaling relations; Mo is p saturation of fault displacement in case of of segment size, which is probably governed by the thickness of seismogenic layer of the crust.

> Keywords: fault segmentatio lation

and L for L>10W. The

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Keywords: fault rupture scalin, width limed rupture, v zealand



















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IUGG XXIV General Assembly

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Symposium

Non-instrumental seismometry Earthquake Recurrence in the 21st Co

Convener: Dr. David Schwartz, Dr. G Co-Convener: Dr. Daniela Pantosti,

For the General scopes see introduction established discipline whose major goal geological record. It provides the fundame seismogenic faults in time and space an assessment. Paleoseismlogy has evolved du the field moves into the 21st century there chronologies, reduce uncertainties in recur recurrence on sources that do not not e (trenching). We invite contributions on top recurrence data, particularly for correlating the of identify paleoearthquakes on blind or remote (ie, logistically difficult settings such as urban environments; -technologies that increase the length of the paleoseismic record and paleoearthquake

techniques of paleoearthquakes; -approac of paleo-earthquake magnitudes

smology and Physics of the Earth's

1924 - 1951

seismology and

Gutirrez

Paleoseismology is now a wellmeation and dating of past earthquakes in the ping ar hding of the behavior of tion for seismic hazard inforn d is now practiced worldwide. As ed to: develop longer earthquake better knowledge of earthquake a paleoseismic analysis tradi tatisti analyses of paleoseimic faults ; -approaches to ots a auduction zone, earthquake sources and in

at investigation sites on major faults; -dating es for better estimates









The western Betic Cordilleras have experie nd partly strong earthquakes and veral modera earthquake-related hazards (landslides and ast 2000 wears. The ruins of the nis) <u>duri</u> Roman village of Baelo Claudia (Tarifa) yiel thquake damage on the nce fo storic Iberian Peninsula. Roman settlement starte I Cer licts from the V-IV Cent. AD. lin t We have found indications for two earthquakes, estroyed During the I. Cent. AD, probably an earthquake occurred, the village was restored and rebuilt (40-60 AD; Silva et al., 2005). Ground Penetrating Radar and geo-electric tudie were carried out in the ruins, across fault zones to map and mirror fossilized and active faults ruins of Baelo Claudia are often badly preserved, we encount evidence for coseismic ical avement, pull-ups, and deformation, i.e. high-energy events, e.g joints in the flagstones of the Decuman indicators is systematic, Drie pointing to a shock from the SW, and bldir n tio hel sis temple area is partly excavated. Drums of fallen columns, wall and pillar collapses are directed in S to SW direction, and testify to coseismic building deformation. A crude stratigraphy based on Roman pottery allows us to date the collapse event in the IV. Century nt. AD suffered not only ne I. earthquake and/or landsliding deformation cracks in the walls and inclined walls are interpreted as generated b e other and, big fallen blocks of lov Srm the tiers are attributed to coseismic damage. The eastern aqueduct outside the city walls crosses a little creek. The western part of the aqueduct collars d downhill and some he arcs show rotational displacement around a horizontal axis, this fight be inter ed as a w deformational feature originating probably from small creek-parallel landslides. The city wall surrounds the village, and was built for representative and not defensive purposes. The walls are inclined up to 10 with varying directions. Keystones of arcs are subsided walls are partly displaced up to 17 cm, and/or rotated against each other. During s of a former city wall has been encountered, this wall is topped by a "dem of wall boulders. This horizon may correspond to the 40-60 AD earthqu 5). Silva, P.G. et al., 2005. Archaeoseismic Record at the ancient Rom th Spain). Tectonophysics 408: 129-146. Acknowledgements This v panish-German Acciones Integradas Program HA2004-0098. The au he Archeological Site of Baelo Claudia, Angel Muoz Vicente for facili

Keywords: archeoseismology, paleoseismology, roman remains





Keywords:	marine pa	rbidite





The microscopic fabric of clastic sedimentamelect We characterize the fabric of clastic dikes an sei susceptibility (AMS). We assume that different manifest in different magnetic fabrics. We tu fluvial incision into the lacustrine 70-15 ka sa vertical, up to 30-40 m high, and up to 0.4 m spanning a sector of 60 degrees with projected s salt diapir. Field relations and AMS analyses show

flect the composition of the services layer by month or the services of the service over 250 Ho month of the services of the s

experiment, and deformation history. Buring the anisotropy of magnetic have different microfabrics, which he clastic dikes, exposed by deep ead Sea basin. Typically, they are ey append in a radial arrangement it a structural dome above a rising fillen from above have vertical

Kmin directions, compatible with sedimentary features. Conversely, horizontal to sub-horizontal Kmin directions occur in dikes that show segmentation typical of horizontal propagation of the fractures and

lateral material transport. Vertical zoning c regarded as evidence of multiple injection diapir induced a local stress perturbation, Subsequently the fractures were filled by l vibrations. Sedimentation from above occ remained open. ation typical of horizontal propa and clauration to support the propagation with the rigger of the arriential of rizonal initiation of the law cs recember and to could the bridge

e center of the dikes is rd indentation of a salt radial set of fractures. ressurized by earthquake reached the surface and































approach will help to gain important informations on the recurrence of great earthquakes along the Algerian margin in the last 10,000 years and on their actual imprint on the seafloor.






























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whole surface, which is often difficult du

surrounding the summit. This problem coul a new approach, which combines /ercom/ analysis of both ground and satellite data. I s d on t of the elation similitude, which het eolog enables to fill the gaps in the EM data usin oth and geoph al data, correlating with the resistivity values. In particular, the latter ones onic fragmentation of the sist from rocks obtained by analysis of the satellite photographs of the surface using the method developed in stimation of the coefficient of the tectonic fragmentation (Nechaev, 1999). In turn, it is based on th (CTF) from the specific lineaments' length the studied area. The approach mentioned above was used in or el of the Elbrus volcano ma (northern Caucasus) by MT data measure bssing its summit. Twodimensional inversion of the MT data res same nodes of the grid, Ite where the CTF was determined. A special eur s used in order to extract the subset of these nodes, which supports the biggest correlation ratio between these two parameters. At the next step an artificial neural network was taught to correspondence of these parameters determined only in the selected nodes for istivity distribution in the the re whole area from the CTF values. The anal del resulted in detection of the relatively conductive body at the dep 0.0 S/ dimensions in vertical axis, latitude and longitude being equal to 20 km, 35km and 15km, accordingly), which can be treated as a magma chamber. The model of the Elbrus yaano can be used for sole h of two important tasks related to the monitoring of its activity: optimization f the EM vey netwo and determination of the background level for detection of the time and spatial variations of the manifestations of the tectonic activity leading to the eruptions.

relief surfa

nd inaccessibility of some zones







Keywords: multi scale tomography, hydrothermal sources movment, complex wavelet transform











algorithms carry out the morphological examination of time series, and identify pre-supposed signals in successive segments of the EM records. Boin, a sample or a real record or some ideal pattern formulated by the expert is used to formulate the recognition procedure. The application of algorithms allows to find out anomalies in electric and magnetic data, and to discriminate between anomalies of different type corresponding to diverse plant the second of automation and for revealing characteristic morphological sequences in h

Keywords: fuzzy logic

ic activity





more than 150 nT in about two months with a valuation rate of pour 0 nT/ray. At the end of January 2007 the phenomenon was still in program, even if the variation rate of stically decreased to 0.5 nT/week. The large anomaly could be of the come beac of the rest by available of the rate of station and stopped at about 30 m from the magnetic sensor.

Keywords: geomagnetic changes, mt etna





Institute of Seismology and Volcar

Hokkaido University IAGA

Yasuo Ogawa, Shinichi Takakura, DVU Yusu miya, Mitsuru Utsuqi, Tetsuji Koike, Hiroshi Hasegawa Hir ii Ich leyuk Satoh, Toru Mogi ara.

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1. Introduction * Usu Volcano, southwesterr eruption in the 20th century in March, 2000 m due to a shallow intrusion of magma bene th volcano. Many geophysical investigations eruption. None of them, however, have yet achieved electrical structure, Akita and Shibata (2003) conducted the shallow resistivity survey by audio frequency magnetotellurics. They reported that the therefore, it seemed that the investigation lower frequency band. We thus planned a MT sites aligned in the NE-SW direction center, were interfilled with several audioas the vertical reach. * 2. Electrical resistiv ty s resistivity structure by using the inversion process developed by Ogawa and Uchida (1996). As seen in the previous survey, the general resistivity of this area is quite low (0.1 to 10 Ohm-m). This feature well explains the poor variation of surface self-s shows about 10 Ohm-m, corresponding to Ohm-m) underlies with a thickness of some beneath the upheaval center, while it does from 500 to 1000 m deep in the northern and southern side of the cross-section, just like an umbrella-shape tructure. It is probable in mineral (montmorillonite) immersed in high samity huid is re-VLR then possibly corresponds to the thermal transition (about 200 C) from montmorillonite to other minerals such as illite of higher resistivity. The bump of the VLR below the upheaval center may be related to the isotherm due to a magmatic VLR, though it is not imaged as an isolate geomagnetic changes * Some of the au measurements over the upheaval area sir ended, anticipating the subsequent geo Surprisingly, it has been revealed that mai on. The change looks quite linear with re after the eruption calmed down. The ma

the ground upheaval of about 80 ishiyai the NW piedmont of the this a during and after the ruded magma. As for the shallow part of this area is very conductive, and trusion depth requires

experienced the 4th (and the final)

in 2006. Six wide-band crossing the upheaval fizontal resolution as well es were inverted to a 2D

esistivity (VLR: 0.1 to 1 had at 200 to 400 m deep

a highly conductive clay

VLR. The bottom of this

et al 2007). Surface resistivity

magma should be situated below this oss-section. * 3. Implication from nave started magnetic repeat the surface manifestation surface thermal activity. 50 nT/yr was still going auing even seven years he south, while it has

decreased to the north of the upheaval center, suggesting the increasing magnetization at 400 m deep. Such change is normally interpreted as cooling at the source region; namely, the rock is getting more magnetized as the magnetic minerals in the rock freeze themselves to the present geomagnetic field.

adr

This mechanism requires that the source temperature enough to produce the effec subsequent cooling phase. Such an extension the discussion in the resistivity section. magnetization (TVM) in a lower temperatu

in advance to a high al magnetization in the hlikely, taking account of sms is the themo-viscous osed to an external field may

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be gradually magnetized with time along the a range far below the Curie temperature, in considerable amount of TVM acquisit consistent with the temperature informat magnetization process. * It is worth co magnetization of opposite sense to the pr such a situation, magnetic field change demagnetization (TD). TVM effect can, how an open question whether the magnetiza However, the latter is probable since the the reversed Tertiary volcanic basement area (Okuma et al., 2002) and a marked (Satoh et al., 2002; Hashimoto et al., 2007 and conclusions * We conducted MT resist of Usu Volcano. An umbrella-shaped very beneath the upheaval center at some hundr to the intrusion. Montmorillonite is a plausik correspond to the isotherm (c.a. 200 C) of from the geomagnetic monitoring in the pa deep is acquiring the magnetization to the

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magnetic field. Heating of a rock body, even in crease of the relaxation time, resulting The TVM therefore seems more netotellurics than simple cooling a rock with initial remanent rsely magnetized body). In e one due to thermal y in a low temperature range, since it rotates the reversed component in place of simple erasing as in the case of TD. It is still is normal or reversed. gical boundary between romagnetic survey of this al stage of the 2000 eruption al. The ongoing rapid and persistent changes in the total field may be accounted for such effective TVM at low temperature. * 4. Summary magr on of the 2000 eruption Ístivity m-m) layer was found LR to 1 brella structure is probably related ers deep. this idate for the while the bottom of the VLR may ral. Meanwhile, results n to e cla that ck body at about 400 m on sta ua aeor etic It ore consistent with the

resistivity results to interpret this ongoing change is up to the TVM at a low temperature range rather than considering the simple thermal demagnetization.

d

pres

Keywords: usu vol 20 UGIA





several hundred millivolts to more than one recovers to the level measured on the Consequently, the entire SP profile along a and reaching the foot on the opposite side by Ishido (2004) showed that the primary the electrokinetic drag current associated underlying saturated layers and the presence of a

shallow conductor contacts a deep conductive layer, this conductive structure provides a current path between the low-potential shallow and high-potential deep regions, resulting in increase in SP around the summit. Assuming a plausible value current, the terrain-related SP on the perip the magnitudes observed at a number of with periodic groundwater recharge, which distribution is sometimes not symmetrical a summit. In addition to heterogeneous resistivity and coupling-coefficient distributions, local increase or

*i*olt one. volcan а ine starting f s th<u>e shape</u> the ha do е ard onduct

ne volcano, then rapidly summer crater is approached. the foot, passing near the summit letter "W" Numerical simulations SP dis¹ tion is a combination of id flo n the unsaturated and volcano summit. If the

> dependency of drag V/m, which is typical of ed profile is stable even ions. The "W"-shaped SP obvious near the volcano

decrease in liquid-phase saturation associated with smaller or larger permeabilities or fumarole activities is thought to be responsible for the local S de of high SP around the amplij summit crater is sensitive to the conducti ght j hange over time due to volcanic activities such as magma ascent nal on ction, etc. Evolution of othe high SP near the summit crater with (electrically conductive) magma ascent is expected to be largely the near surface and deep affected by the continuity of the pre-existing corrective structure between regions. These topics will be discussed on the ba of axi fetrical 3 numerical simulations of electrokinetic potential produced by subsurface fluid flow.

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Disaster Prevention Division, Tokyo Me

Intion Specialist (Chief) IAVCEI

Mitsuru Utsugi, Eisuke Fujjj , Mak Uveshima aues Zlotnicki

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Several remarkable magnetic changes were obj jima volcano, such as the precursory cha formation of a sinkhole in the summit, tho during the caldera formation and so on significant change in the total intensity was ot d time of the magma intrusion started on June 26. O

magnetometer operated by NIED at a station on the western slope of the volcano (MKA) recorded a large amount of changes up to 200 nT. Ued be explained by the piezomagnetic effect beneath around MKA. This is the 3rd exam the cases of 1986 Izu-Oshima (Sasai et al. These three observations indicate that m gn dike. Ueda et al.'s piezomagnetic calculation is and GPS data (Ueda et al., 2005). This model consists of three intrusive dikes and a contractive one, which reproduced the deformation data in three time periods from 18:30, June 26 to 6:00, June 27. The magnetic changes at MKA were ascribed other hand, Fujita et al. (2002) searched f fit the tilt change at each one hour interval in E-W oriented dikes intruded in the southern part of the volcano toward eastern side of the island, which were not well represented by Ueda et al.'s (2005 model. Our roblem is the station, located at the southern coast, gradually de contrast to other stations which showed no significant variations. However, even a dike closest to TAR proposed by Fujita et al. can not properly explain the observation there in terms of the thermalmagnetic nor piezomagnetic effect. The or

the shrinkage of the magma reservoir whi western sea. Del Negro et al., EPSL, 2004. 2002. Sasai et al., J. Geomag. Geoelectr., 161, 891-906, 2005. Ueda et al., EPSL, 245 ing the 2000 eruption of Miyake-

prior to and associated with the events, the enormous ones 2003). However, any Ited on the island at the a 3-components flux-gate

(2006) showed that the said magnetic changes could SE direction emerged to dike intrusion, i.e. in Del Negro et al., 2003). position of an intrusive el inferred from tiltmeter e period (19:00-01:00). On the

king Mogi source to best

une 27 from 21h to 24h, mostly

The total intensity at TAR June 26 to uly 10 by about 10 nT in n was the piezomagnetic effect due to gma to the dike intruded into the Inst., Univ. Tokyo, 77, 67-75, 777, 2002. Ueda et al., GJI, 5, 139-154, 2003.

Keywords: dyke intrusi

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In the 2000 eruption of Miyake-jima volcan of the volcano. It started from a sudden c caldera on July 8. The new sinkhole of 900 diameter and 500 m deep until the August tilt-step event took place once or twice a da which was a sudden step-like inflation of the volcan

hours (Ukawa et al., 2000). The velocity waveform of the ground motion was a single sinusoidal wave of 50 seconds duration. The magnetic field observed by proton magnetometers at se minute (Sasai et al., 2002). The electric fi line (a few km distance) SP measurement ys seconds. Unlike magnetic case, the electric fie velocity waveform (Sasai et al., 2002). Recen electric and magnetic variations, namely the short-span (150 m) multi-channel SP measurement with 2 seconds sampling on the southwestern side (Zlotnicki et al., 2003), and 3 components magnetic data by two flux-gate magnetometers with 1 second volcano. All these EM data are compiled tectonic phenomenon, which should be esse 2002). Since the time-dependent behavior is different between magnetic and electric signals, the generation source must be different. The manetic varia on is most piezomagnetism of rocks due to stress changes, whi the groundwater movement induced by rapid stress changes and/or fluid injection from the pressure source. Fujita et al., 2002, Cyclic jerky opening of magma sheet and caldera formation during the 2000 Miyakejima volcano eruption, Geophys. Magnetic and electric field observations du Earth Planet. Sci. Lett., 203, 769-777. U deformation and earthquakes observed b Space, 52, xix-xxvi. Zlotnicki et al., 2003, activity: The July 8, 2000 Miyake-jima erup

aldera was formed on the summit area in the existing Hatcho-Taira in depth enlarged up to 1.6 km in pid gra deformation called the tion p d (July 8 to August 18), aual shrinkage for several

a step-like_change_within_one minute, which was surement interval of 1 ected by the long basea sampling interval of 10 ation very similar to the have been found for the

outh-western side of the

of this unique volcano-

obably ascribed to the one to the ectrokinetic effect due to 29/2001GL013848. Sasai et al., 2002, iyake-jima volcano, central Japan, Miyakejima eruption: Crustal ion network, Earth Planets s associated with volcanic 205, 139-154.

a formation process (Fujita et al.,



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interface is responsible for the divergence of the drag current and it is located at the casing pipe.

Keywords: geyser, resistivity, self-potential





measurements from 2003 and 2006 arou summit areas of Ontake volcano. As the result of SP measurement torus-shape positive SP ar th eisr anomaly has been detected at the eastern p of maly is located between a recent active clusters of earthquakes and near the ground uplift detected by Kimata et al. [2004]. They suggest that the uplift is associated with a region of low resistivity [Kasaya et al. 2002] and anomalous increases in chemical compositions of haging potential shallow 031, hydrothermal activity. Generally, an upfld ction produces positive current sources in the direction of flow. The ensitiate maly supports potential om shallow hydrothermal activity. Recently, we established a continuous SP observation network with the stribution. This network aim of monitoring the hydrotheral activity by reasence to the obtained S uses metallic telephone lines for measuring SP with ing. We also 1 sec sa found the positive sense anomalies up to about 2V p-p around the northern part of summit area. This large anomaly may be not irrelevant to recent eruption vents. Younger vents are located near the positive anomalies. In this presentation, we will report a detail of SF outline of continuous SP observation





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Symposium Progress in electromagnetic st Electromagnetic fields associated with a

Convener : Dr. Malcolm Johnston

Electromagnetic fields are both expected Furthermore, as a consequence of this rul motion into the atmosphere and ionosphel processes driving these tectonic events Unfortunately, not all aspects of these mea understood. This session will focus on the f and magnetic fields near and during activ around seismic and aseismic rupture region and observations from natural laboratories Theoretical considerations regarding source gene quality, identification, separation and removal or smology and Physics of the Earth's

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olcanoes artiquakes and active faulting

smic and aseismic fault rupture.

oposed b explain them, are well

ation: 1) Measurements of electric

neity in electromagnetic structure

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aboratory observations

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leary fields are generated by coupling of ground

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measurements (strain, tilt, pore pressure, displacement, etc) together with EM measurements that can place better constraints on the physics of s e pr<u>e</u>c s befor , during and after earthquakes. UGIA




















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Clarifying heterogeneous subsurface structu es in understanding mechanisms of the intrapla how/where the fault ruptures. In this paper district, Northern Japan. Hokkaido locates in Amurian plate are colliding with each other Dne where there had been occurring 12 large earthqua

Teshikaga region is located in a volcanic belt formed along the northern margin of the fore-arc sliver of the Kurile arc. The other one is focal area near the boundary between the Okhotsk crustal structure, we performed wide-band preliminary study, several 2-D resistivity m were compared with gravity data and intraplate earthquakes and crustal structures. body is imaged at 0-5 km in depth with a horizontal width of 10-20 km, in the focal area of most of the

1938-1969 earthquakes. Comparisons with borehole (NEDO, 1985) and density structure indicate that the high resistivity zone seems to represent which are regarded as Quaternary-Pliocer suggests that the earthquakes occurred are stress concentration to the heterogeneity. The 2-D models_also clarified that the fault of the 1938 earthquake corresponds to the wall of the Kuton caldera Three profile were obtained around the focal area of the 2004 comprised of two layers: upper conductive layer, existing surface to 3-5 km in depth, and lower resistive

layer. Comparisons of the resistivity image with the surface geology and drilling data indicate that the upper conductive layer and the lower resi and older igneous rocks, situating as its b we found a clear upheaval structure in one steep variation in rigidity around the foca strain, which probably triggered the eart anticline structure observed in surface geo of the fault. In order to clarify above constructed three dimensional resistivity st

ke areastis an important element where the stress concentrates and arthquake areas in Hokkaido late 🕯 the Okhotsk, Pacific and ig, wh hqual reas is Teshikaga region and 1967 (Hirota, 1969).

2004_Rumoi-nanbu_earthquake (MJMA 6.1), located e heterogeneity of the earthquake areas. As a these resistivity profiles ationships between the esistivity (> 300 ohm-m)

the resistive body. This

the chitrary, conductive zones,

ock and non-rigid rocks by

of 2-D resistivity images

All images are roughly

to Cretaceous-Tertiary sediment rocks the basis of this correspondence, al area. This structure implies gests local accumulation of da et al., 2007) and the nds along strike direction ng strike variation, we magnetotelluric data at

Teshikaga and Rumoi area, respectively. The resistivity modeling was operated using the 3-D forward modeling code developed by Fomenko and Mogi (2002). Around these seismogenic zones, both models

lake

show heterogeneous structures, where resistive body surrounded by conductive zones. The 3-D analyses also resolve incoherencies in 2-D discuss lower crustal heterogeneity under t of existence of shallower conductive layers

Keywords: m

ing. This enabled us to fficult to detect because











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interference. • A short duration anomaly that that appears to be co-volcanic (with dig representing an EQ precursor requires con rock physics of Freund, 2002 coupled wit 2004 indicate lithosphere-ionosphere cou However, here, we have only reported station. Several automated terrestrial stati of the signal would provide a much better underground construction. Keynote lecture 90/8 pp 18-48. IE Aust. Canberra. Freund Geodynamics, 33, 545-572. Freund, F. T., N.A. and Ouzounov, D. 2007. Stimulated eEarth, 2, 1-10. (www.electronic-earth type of seismoelectromagnetic emission. France. Pulinets, S. and Boyachuk, K. 1995 GermanyP.315. NASA/CIT/JPL, 2007. http://ssd.jpl.nasa.gov/ Rowe, A.J. and Gray Australia. V.165. No. 10/11. Warwick, J.W., with rock fracture: Possible application to t Res. 87:2851-2859.

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Keywords: vhf pulsed emission, p h pulses, stress

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Perugia, Italy

e an EQ and a relatively long duration anomaly The issue of a P-H pulse anomaly as not been attempted here. The iven by Pulinets and Boyarchuk, h electromagnetic emissions. single, manually operated synchronised vectoring m). References Brown, E.T. and Windsor, C.R., 1990. Near surface in situ stresses in Australia and their influence on ey. Nat. Conf. Pub. No. propagation in rock. J. aseer, A., Fu, C.C., Bryant, assessing a stress indicator. ence Australia, 2007. Australian earthquake database at www.ga.gov.au Punnets, S. and Hollis-Watts, P. 2003. P-H Pulses - the new 07035 G Joint Assembly, Nice f Earth akes. Springer: Berlin : neric F System Ephemeris Program nics at 1996. Shake le and roll. The Medical Journal of C. ar lio emission associated 19 Chile arť ike of / 22, 1960. J. Geophys.













Kengo Tanimoto, Yusuke Yamaya Hashimoto, Venera оvц vama, Doblica, Takanori Kajiwara, b Ogawa resh coto

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The 2003 Tokachi-oki earthquake (Mw=8 determined at 80 km off of the Erimo area, region, including the asperity and aftership ťκ direction over the upper plane of the Pacific late the Erimo area at a depth of 50km. The Pacific plate

direction beneath Hokkaido along the Japan-Kuril trench, a depth of about 6000m, situated at 100km off from the land. We carried out wide bank to 0.8x10-4 Hz, at 25 sites in the Erimo ar sites were arrayed in a grid in about 50x5 dimensional resistivity model. The survey a area were seriously affected by a sea eff ct. including sea to become reality. The resignivit part of the area and its complex shape in this area.

dge cting at magentotelluric survey, covered frequency range over 320

aide

September. The hypocenter was Northern Japan. The source m wide in NW and NE out 1 ie sol region reached beneath out 8 cm/year to the NW

> the source region. The s us to construct threeata obtained at the coast ensional resistivity model ducting plate at deeper
























































dimensionality in the long period (1-1000s TM mode, phase in TE mode and tipper were used for two-dime shows inhomogeneity in lsic long the fault plane, we the middle and lower crust. Additionally itv carried out MT survey along the Atotsugawa fault. A seismic gap and a creep-like crustal movement were observed along the Atotsugawa fault. Obtained preliminary inversion result shows lateral inhomogeneity correlated with heterogener which suggested that the ault seismic gap on the Atotsugawa fault plane ive block. We will report GPS data. outline of the both MT surveys and also disc

Keywords: magnetotellura, atotsugava fault, tectorio zone













Keywords: seismic electric signals, electric field transmission, magnetic field variations



















Identification of electromagnetic (EM) precu to the seismic activity is still under 0 arameters rela scientific debate and requires new reliable ir interrelation with changes of on ab ssible electrical conductivity occurred prior to the amic p phasizes the anomalous paper eo behaviour of the EM parameters as possible the d mstances of the specific ory s ls u brer geotectonic characteristics of the Vrancea zone diate micity. In this respect, measurements of geomagnetic field have been performed since 2001 year and recording network has consisted of two high sensitive geomagn Surlari National Geophysical SV tems placed at the Observatory and Provita de Sus Geodyna stem consists of data logger with 6 channels and A/D conver magnetic field sensor (frequency range: DC- 1kHz) and a lapt processing. One of the horizontal components of the three-axis n entated perpendicular to agri the geological strike in order to record its tim hat a large-scale regional ar conductivity anomaly causes a regional amplification of the vertical magnetic component Bz as well as spatial changes of the horizontal magnetic component perpendicular to strike (Bper.). Subsequently, a specific approach regarding the electromage (Bzr Bz/Bper. and $\rho n = \rho / \rho z$, where ρ is resistivity parallel to strike an cording to the temporal ecte invariability criterion for a 2D geoelectric st diti in, king into consideration nor their daily mean distribution versus intermediate_depth seismic events recorded simultaneously, was elaborated. These changes of electrical conducative inside of the Vrance seismogenic slab and its surroundings, before the earthquakes to occur, as the lithosp eric conductivity changes equenc produced maybe by the dehydration of the rocks associated with rupturing processes and fluid migration through faulting systems are reflected by the anomalous behaviour of the Bzn and pn parameters and, finally, several conclusion interrelation within a span of 6 years interval are inferred. We claim that this sp with more complete approach of EM phenomena can improve the seismic ha

Keywords: electrom ismicity















the observed neutral current data and c d observed at Kakioka magnetic observatory and try to show the ield and neutral current rel quantitatively. We calculate transfer function (for input) and neutral ee ēon current data (for output). Instead of the con e have used continuous wavelet transform (Harada et al., 2004). Although the transfer functions for short-period geomagnetic changes were not clearly determined because of artificial noises of neutral current data, the transfer functions for long-period geomagnetic cha e determined with small W errors. By this result, we can discriminate s from observed neutral current data and it would make it easy to pssibly associated with earthquakes from noises.

Keywords: neutral current, geomagnetic induced current, transfer function





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Electromagnetic phenomena have been earthquake prediction (Hayakawa, 2005). T signal observed in a wide frequency range radiation during experiments of rock fract comes directly from seismic source, consi suddenly due to high stress in crust. It see fracturing process of rocks in crust. So a magnetic signal, such as stress variation effect in

effect, and electromagnetic emission associated with microscopic cracking in rocks, have been brought forward. Comparing to field observation, none of these mechanisms can explain all abnormal behavior of seismic electric magnetic signal, but el relatively some reasonable and with cathol to characteristic time of minutes or hours f or piezomagnetic effect need large numb main viewpoints about the mechanism of e is surface charging mechanism, suggesting that the surface charge on crack walls in rock is caused by

appearances of excited hole and electron trapping centers (point defects) on a newly created surface of rocks when fault asperities are sheared. T caused by emission of electrons in crack quadrupole model is proposed by Guo Zig cracking in rocks and frequency spectral bandwidth gotten as 0.5 1.0MHz. The emission of electrons in crack tip end and the model of compressed atomic are supported by the third pinion, and otherwise, a capacitor model is proposed as a supplement suc happens in the extending process of microscopic cracking and the lower or higher frequencies are explained well.After all, the radiation mechanism of seismo electromagnetic signal is still not known

well. It seems much difficult for us to ch around ground. There is still much uncert and direction and so on. The uncertaint convincing evidence of seismic electromage there found abundant cracks in rocks. Once system would get into a nonlinear phase fracture thus macroscopical seismic observed.Considering that seismological G

dicate for the short-term oncer about electromagnetic rthquakes occur. Electromagnetic at seismo electromagnetic signal occur when rock body fractures nal is closely related to agneti med isms of seismic electric ctricity or piezomagnetic

> ic cracking in rocks is s variation corresponds tion, and piezoelectricity n crust. There are three cking. The first viewpoint

d atoms. Even a electric

process of microscopic

discharging in crack tip

seisma electromagnetic signal is

ore reasonable by observation in field al, such as amplitude, time series mic radiation process. Another ing process of rocks is that in value, the seismological dimension scope would generated be and th power law, a typical

fractal phenomena, much similar to the behavior of fragmentation in mining and eruption, as well as tectonic distributing. Some authors regard power law as a characteristic of self- similar seismological

system and others suggest that it is a Gutenberg - Richter Equation is statistically used for a certain region and for a longer crust rocks with different fractal dimension dimension, the radiation mechanism caus

phenomenon. Although of time, it is still often is used to tiny cracks in Suction of furcated fractal is analyzed related to the









multi-parametric study utilizes ancillary i

crossing. In addition, this ing digital topographic data


























(IAGA/IASPEI/IAVCEI) Working Group of Electrop (EMSEV) was established in 2001, with th cooperation in the new field of research now interest to understand earthquakes and vol often using different languages, namely fro science and solid state physics, to name a f is of crucial importance. During the past meeting in Perugia, the second IUGG General As

symposium entitled Progress in electromagnetic studies on earthquakes and volcanoes, consisting of the following sessions. JSS007: Volcanic structure and activities. Convenors: V. V. Spichak; J. Zlotnicki, Y. Sasai, D. Patella and C. Del Negro JSSO active faulting. Convenors: M. Johnston; earthquake precursors. Convenors: P. F electromagnetic studies using space tec Ouzounov, and V. Tramutoli Although separat discussions among participants with different backgrounds inter-connected.

pu of motii Seismo-el ctro tal but their scie nd based de efore] tua ears, SEV EMSE

Studies

be made

tinquakes and Volcanoes hal and interdisciplinary fterna netics. Nembers share a common c background is extremely diverse sics to ionospheric physics, space lerstar g for the common cause enjo high activities. For this e planned an integrated

with earthquakes and Crustal instabilities and Nagao JSS010:Seismo-Pulinets, M. Parrot, D. hoped that active fruitful since the subject matters are all



nol









Panayiotis Kyriazis, Antonis Kyriagopouk E Simon Ana Chiadis, Dimos Triantis, Filipos Illiana os

Pressure stimulated current (PSC) effects en studied arious materials. In a number of previous presentations the applied mechani with the emitted current. In ٩I ss ha lated the present work current emissions are st aterials like marble and lie a se na brittl amphibolite as well as on composite man-m de l rials cem paste ecifically, the stress was nong with applied on the referred samples close to mechanica PSC measurements. The recordings manifest that, dynamic phenomena, like macro-crack propagation and failure plane creation, result in current emissions. In these experidespite the fact that the stress level was maintained ents practically constant in the vicinity of failur ration were observed. The emitted PSC can be attributed to cha change of the sample ۱ar structure while new cracks are formed and both the existence and interpretation of the PSC are consistent vith ons (MCD) theory which relates the emitted current to crack forma on p d 'e hsequent strain variation. Strain recordings in the range near fracture support these findings since strain variations have been recorded without any change of stress. A deep and fast PSC reduction has always been recorded before fracture predicting the inevitable failure. p coexisting causes: The to t former is related to the lack of any obviou he bulk of the material, rain that could lead to charge rearrangement s ress is localized at the st hop ed ' edges of the main macro-crack that guides the failure plane. The latter is related with the observation that the failure plane limits significantly the availa the charge move across the sample e paths fo bulk which makes the detection of any emitted SC Eult. en more

Keywords: pressure stimulated currents, brittle materials fracture





1-2 hours during 5-10 days before the ev since the day before until a few days after an they find that it occurs at day-time in peri dbe the season, on 11-years Solar variations circle, and on geomagnetic variations. In the work we use the

data recorded in time intervals with weak heliomagnetic disturbances (Wolf number below 100 and Σ Kp< 30). The amplitude of the modification effect amounts to more than 0.95. No eff than 80 km. Also, no effect is found for we foF2 variability with the modifications of the electrical conductivity could be connected with an injection of radon into the lower atmosphere.

al behavior of the effect,

hal variability depends on

uakes with depths more

revealed change in the

d the statistical reliability of the

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ear

6.0 Th

Keywords: ionosphere, earthquakes, f layer









mathematical accuracy of the signal. Further application is required to accurate events. These facts demonstrate a possibility of monitoring the state ctivity with using the SS and PCA.

Keywords: singular spectral analysis, principal component analysis, ulf geomagnetic data





Keywords: geochimical precu

drothermal model





recordings from an 1Hz acquisition system analyzed using wavelets and classified acco ng separated into three sub-periods: One consisting the Kythira (36.21 North, 23.41 East) earthquake and two without any significant seismic activity from the indicators of CEP signals primarily in electrical field ordinas



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IUGG XXIV General Assembly July

Perugia, Italy

(S) - IASPEI - International Associa Interior

JSS010

Symposium

Progress in electromagnetic st Seismo-electromagnetic studies using space

Convener: Dr. Valerio Tramutoli, Pro Co-Convener : Dr. Michel Parrot, Dr.

During the last decade, the possibilities oceans, such as earthquakes, volcanic active through yet unresolved process the lithout was postulated from ground-based observat using multi satellite sensors have been concentration (TEC) in the ionosphere, grou emission and/or cloud cover associated wit epicentral region. Latest observations from t Frequency (VLF) electrical and magnetic sig are still insufficient and the physical under surface/atmospheric/ionospheric precursory events

session include: 1) Case studies of satellite observation related to seismo-electromagnetic observations comparison with ground-based observation between the process in the earth crus earthquakes; 3) Thermal infrared (TIR) er related to major earthquakes, and volcan based relevant data, such as meteorologica of space technology in tsunami early warni

smology and Physics of the Earth's

technology

physic

2067 - 2092

olcanoes -

ge events in the solid earth and marms, may affect the atmosphere and ionosphere



13, 2007

Coupling. LAI coupling he recent investigations icant changes in total electron anomaly by thermal infrared (TIR) kes have been observed over the ae evidence of Very Low vever, convincing results lid earth processes and opics to be discussed in this

al mechanism of the connection bmena prior to main and

EC anomalies, possibly mparisons with grounda; 4) Possible usefulness











MODIS, onboard NASAs Terra and Aqua; on A)/AA RR; (4) geosynchronous A weather satellites (GOES and METEOSAT); (5) DEM e have analyzed: surface r; (6) GS emissivity; sea and land surface temperature (LST); emitted earth radiation (OLR); air temperature; surface latent heat flux (SLHF); Total Elec TEC <u>(GPS</u> /IFC) quasi-continuous electrical ron Content fields; and thermal plasma parameters.A r year data (LST, SLHF, OLR, and VLF) by systematically compa ervations to determine O meaningful statistics that can be applied to quake. Our rationale for lari using this complement of observations is the emporal coverage of any at t ar one of these pre-cursor signals on the global so bach is to enable multiple of ap and previously validated physical measurements to be integrated into one framework with the latest theoretical models of seismo-electomagnetic generation and propagation and to provide feedback on data gaps that may then be acquired in the gnificance of our satellite The based multi-sensor approach was define 2006) worldwide strong earthquakes and applying the techniques ed an mails. This joint approach provides an opportunity for a comprehensive study of Earth electromagnetic environment, and can be used to understand the relationship between sei -tectonic rocesses in solid Earth and surfaceatmosphere-ionosphere variability.








































as earrhquake precursors and to develop a new nethod of eachquine prediction. One of the possible ways for such separation can be implemented by nears manalysis of the neutrons' flux data and the seismo-electromagnetic data obtained by since a issions. In this appendix the results of cross-section analysis of the neutrons' data and the charged particles' precipitation under the Earth's radiation belts in order to find the characteristics of geodynamic variations of the neutrons' flux are presented.

Keywords: neutron, precursor, seismo electromagnetic













IUGG XXIV General Assembly

July 13, 2007

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Perugia, Italy

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JSS011

Symposium Earth Structure and Geodynam

Convener: Prof. Thorne Lay

Integrated theoretical, observational and geomagnetism are essential for advances problems such as the driving mechanisms

water in the transition zone, the existence and geochemical reservoirs in the mantle, the layer, and the chemical and physical multidisciplinary approaches. These topics w constraints on the structure and amplitude elastic anisotropy using many imaging and r thermo-chemical nature of mantle convection transitions, time evolution, and deformation bn i experiments. These two fields interact closely,

geochemistry and geomagnetism to develop quantitative understanding of how the planet works. Both disciplinary and multidisciplinary papers drawing from these fields are invited for this symposium. This symposium will be subdivided into 5 m structure and processes (L) 2) subduction (P) 4) transition zone structure (T) 5) D" a

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eismology, geodynamics and amics of our planet. Fundamental the fate of subducted slabs, the role of umes, the isolation of rmost mantle boundary the nd mantle all require core symposium. Seismology provides tic properties and density and on rovides insights on the dyna tion of flow with phase or, int odeling and laboratory erica s from mineral physics,

> e and asthenosphere nant/ upwellings and plumes





Keywords: intraplate earthquake, local earthquake velocity tom, crustal mafic intrusive













short of active dynamic process in deep. The er are where mantle convection and its impact are weak. This is consiste activity in these areas. tect Beneath the Baikal Rift Zone exists an upw ical to the low-velocity h is ide le ' anomalies originated from the 670km discontinuity imaged by seismic tomography. The Hangay-Hővsgől Plateau is underlain by mantle upwellings, resulting in extension stress regime A complicated pattern of mantle flow and convection-generated stress western ibited in ongolia. The consistency between the surface tectonic features and the mantle flow pattern, as well as the resulted stress field, suggests that the small-scale convection within upper mantle may be one of the main mechanisms which control the regional tectonics in Mon





APM direction. We should consider a. The direction of Gondwana rection. In general, the fast ss. We consider that the an-African orogen event

Keywords: seismic anisotropy, upper mantle, antarctica

that it is reasonable that the structure

asthenospheric mantle flow. The upper lay

the anisotropic structure which is past tec

continent break up was NW-SE. This is p

polarization directions are consistent with

anisotropy of upper layers is caused by

(~500Ma).















Keywords: tectospheresandmantleflows, blackseaopening, tyrrheniansubduction



















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Earthquakes are not uniformly distribute especially uneven distribution is present all a The zones in which the deeper earthquakes examined. Abandoning the traditional 2-D se the help of 3-D plotting on larger scale, which characteristic inhomogeneous pattern of hy the Italian region as well under Mediterra global catalogues of relocated earthquakes, filame

or spoon-like patterns. These filaments taper downwards, resembling the shapes of trees, columns, smoke from chimneys, and leading to the idea of an origin in a narrow region of disturbance. Because very hardly a subductive process can interpretation of the Wadati-Benioff zones resulting global tectonics framework invo expansion, rifting, isostasy, surfaceward phase changes. The associated model of e plut an isostatically uprising mantle column which segments slowly overcome a solidus-solidus limit of the

temperature-pressure phase diagram. The outpouring of the exceeding material drives the gravitational nappes to overthrust the sediments of the emulate the subduction process, but withe between uplifting material and down-pushe mixing, migmization, upward transport of fragments of the buried lithosphere etc. are possible. The mere existence of the earthquakes in the brittle partion of the hosphere (h depth) is at odd with the existence of the Earthquakes are the more important circumstantial evidence of local storing and releasing of deviatoric

fragment could be mechanical product of g of kilometres should be considered. This tomographically revealed P-wave and S-w most orogens and arcs, and the obtainable increase that are associated to the main i upward movement of mantle materials ca orogens and to the widespread observat between two lithospheric fragments is a de

arcs or in depth. An veen Africa and Eurasia. irgin l heir regional and global context is he trench-arc-backarc zones, with extent of a Wadati-Benioff zone, a d sub ion zones is revealed in ctive rgins. Using the recent hizable instead of planar

> distributions, a new zones is proposed. The es deriving from global spreading, and mantle to the volume increase of

70 km. At the boundary

Wew tens of kilometres of 🖬 channe a low viscosity channel. stress, which can be cause of local overpressure. Then the possibility that lenses-like HP-UHP exhumed rrence at depth not exceeding few tens old belt is in agreement with the derlying with different slopes nt with the values of volume this view, a discontinuous tinuous evolution of the ally, the rate of rifting of the orogen toward a

ng the on a burial path which

herome on like metamorphism,

true fold belt (low rifting rate) or in a continuously enlarging depression (high rifting rate), leading to a true marine and oceanic sea-floor generation. Indeed, some zones like Tonga-Kermadec-New Zealand-Macquarie seems to suggest all these aligned different zones trench and expanding ridge, mature fold belt, oceanic ridge respectively as different

Keywords: deep ear

uilding


























Keywords: anisotropy, lithosphere plume interaction, upwelling



















short station spacing enables us to measure veltim without spacial aliasing. The pattern PKPg df PKP ntial di earthquakes in South America shows a clear sy radient example 0.5 s difference occurs over a 300 km distance for PKPbc-df, suggesting a short scale variation of P-wave velocity of the top 300 km of the inner core. The core phases propagate the inner corein the east-west directionbeneath Pacific whe heterogeneity in addition to the well docu the scatter of the differential traveltime regional structural variation of the innerre obsevationalconstraint forthe style of conve

eriod b waves from waveforms eltime from deep focus Japanese islands. For

> ted. Rich short-scale inner core may hide in scale and size of the eismic array will provide

Keywords: traveltime, seismic core phases, inner core



























bod correlation between the slip im principal axis and the plate bern area, it is estimated that stronger than the northern outheast strike and exists structure obtained from a northern part of Hyugathe spatial variation of

Keywords: stress field seismic counling byuga nada

boundary. Although slip distribution has r

subducting plate in the southern region of

region. On the boundary region, Kyushu-I

on the PHS plate, subducts beneath the Ky

seismic tomography, there is high Vp/Vs

nada. The geographical features, and/or th

the state of stress in Hyuga-nada.







Keywords: subduct

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In the Apennines subduction (Italy), earth chain axis. The events concentrate in the u solutions indicate normal faulting. In the fo the crust. Focal solutions indicate prevailin faulting in the southern one. The deepeni foreland follows the deepening of the Moh associated with uplift of the chain axial zone, vo

pushing of the asthenospheric mantle and the mantle-derived, CO2-rich fluids trapped within the crust below the chain axis causes this seismicity. All these features indicate that the axial zone of Apennines is affected by early rifting processes. In foreland reflects active accretion processe faulting and the lack of reverse focal so present. In our interpretation of the Aper plate is due to the dynamics (uprising and

therm he nicity and gh CÔ2

rriding plate, along the bove the mantle wedge, and focal olume affects the upper 35 km of northern foreland and strike-slip lume the chain axis to the ve the mantle wedge is extension. The upward

> eper seismicity in the erved dextral strike-slip esses are not active at micity of the overriding heric wedge.

Keywords: apennines crustal seismicity, rifting, subduction





other by the closest image are connected with nicity is most brightly shown on boundaries of plates, and seismic belts (zones) Indarie nd configurations of the ermin ese plates. On the other hand it is enough sure, le p neter glob hospł rotation are determined (Greep, Gordon, 1990; Argus, Gordon, 1991; Ban ometrical, kinematic and . The dynamic regularities of plate motion have been established (Barkin, 2000). The maximal tension at sliding of lithosphere and, accordingly, the ened accumulation of elastic energy and the most trength active displays of seismic activity should t f rotating lithosphere. Therefore we had the right to expect, that al rd ation of lithosphere, the pole Pm of the angular moment of relative a planetary (most active) no seismic belt should have a close positions t research have confirmed ea the made assumption. Thus, the global rotation anetary scale) and directs 'litl hni in I seismic events, and its equator of rotation is set actually the position of a planetary seismic belts (zones) of the most active seismic events, as it is observed actually. The dynamic model of lithosphere plates and lithosphere with various powers area (Barkin, 2000) has been ner used for analysis. Only surface displacement accordance with known node for n, 1991). The special kinematical theories NNR-1 and HS-2 (G An s, bn.' computer method of axography (Ferrandez Garcia et al., 2002) has been used for determination of the most active poles (Ps) of latitudinal and longitudinal alignment in positions of epicenters of large earthquakes in 20th century (coordinates in degree : 1) 65 60.5 E. (e analysis of 112 largest earthquakes in 20th century); 2) 53.5 S, 45.5 E (392 earthquakes with M>7); 3) 54.5 S, 41.5 E (112 earthquakes); 4) 52.5 S, 56.5 E (392 earthquakes) etc. Various indexes of longitudinal (1), 2)), latitudinal (3)) and equatorial ordering et al., 2000) have been used. The obtained coordinates of a pole of a seismic ach other and with corresponding coordinates of the following poles (Barkin of angular velocity of global rotation of lithosphere; pole 45.4 S, 57.6 E ative motion of lithospheres plates under theory HS2-NUVEL1; pole 48 ntum of global rotation of lithosphere. It is shown, that the vector of inertia Lд of lithosphere (together with the Earth) is located in an ending node of seismic belt Ws (longitude 145.4 E). The centre of ates 41.0 N, 36.0 E is

located near to the big seismic belt. The pole Ps is located on the other big seismic belt covering big northern arch of Pacific ocean, with ascending unit on equator (a longitude 105.0 E and an inclination about 55.0 degrees), up to seismic zones of South America. With poles of a vector of the moment of inertia forces of lithosphere, caused by th connected. They form of the spirals loca clockwise) and 270-360 E (are twirled cou north - south. Their origin can be connected lithosphere. The moments of these forces

of active seismicity are 6 E (spirals are twirled ntation in a direction the of inertia for non-spherical of plates result in additional







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The observed geoid anomaly shows very bear circum-Pacific trench, when the longest-way and 3) are subtracted (Hager, 1984). Our wavelength geoid anomaly is significantly mantle, i.e., stiff (high viscous) subducting plate-tectonic mechanism, by the use of the 2-D

examined possible effects of such LVVs on the long-wavelength geoid by using 3-D spherical shell models. In contrast with a traditional propagator matrix method by Haner, our new numerical approach can treat the mantle flow including LVVs. basic equations governing the instantaned more actual global density models compared 2004), we have used a model coupled with (1) in the upper mantle, and (2) the S-wave glob lower mantle. The radial viscosity variation is layered; the lithosphere, the upper mantle, the transition zones, the lower mantle, and the bottom boundary layer. The low viscous asthenosphere is also considered. The reference viscosity is fixed

between the lithosphere and the mantle is the lithosphere. The viscosity of the plate in (Kreemer et al., 2003). First we have calculated the geoid anomaly by using a no-LVV model, in which the stiff subducting slabs and the weak plate makins are not considered. highs over the subduction zones arise only men mantle and the lower mantle (RLM) is around 103. This value seems to be one order larger than the viscosity contrast suggested by the post-glacial rebound analysis (e.g., Peltier, 1998). We have next

imposed the stiff subducting slabs only in t the lithosphere, on the no-LVV model. The the subduction zones, especially, the Ja significantly high, 104. This is because the the mechanically strong coupling between imposed weak plate margins on this model systematically examined the effects of the 100 (i.e., no viscosity contrast) and 104.5

luction zenes, especially over the spherical harmonic degree are 2 2004) has shown that the longariations (LVVs) in the viscosi pus) į margins related to the In this study, we have

> r the discretization of viscosity. To construct a et al., 2001; Yoshida, el based on the seismicity and Boschi, 2002) in the

ual effective viscosity of

sing the lobal strain-rate model result shows that geoid scosity co rast between the upper iscosity of which is the same as that of nally strong negative pattern over renches, even when RLM is ons strongly depress due to ing slabs. When we have nchanged. Here we have s in the range between Ve have confirmed that

r marele. The viscosity contrast

when the viscosity contrast of the subducting slabs is around 101 to 102, the geoid anomaly over the subduction zones becomes positive pattern over such regions, if RLM is around 103. Imposing weak plate margins on this model reproduces the broadly positive anomaly which explains the observation. If

RLM is lower than 102, the geoid anoma These results indicate that the viscosity lithosphere. When the low viscous layer (Cserepes and Yuen, 2000) is included, emerge when RLM is 101.5 to 102, which

hains broadly negative. weaker than that of e second asthenosphere ver the subduction zones sity contrast derived from the
















































but not just its structure. That allows incoducing y corré he mentioned expected reaction of the medium and, furthermore, to other issues related to creation of seismic and geophysical

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earthquakes, explosions

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ess changes from 1 to 15

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first that an increase in

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Intensity of microseisms

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networks, monitoring of natural events, finding out possible earthquakes precursors and so. Moreover these two most important parameters read taken into account in order to solve variou or microseisms registrations. For instance, seismic noise intensity distribution showed km the noises level changes from seve al h measurements. But for seismic zones the noises level much lower and changes from 15-20 to 85-90

units. The opposite situation is observed during registration of strong signals: earthquakes and explosions. The analysis of the number of sediment thickness leads to a decrease quantity of registered events in a seismic zo aseismic zone. The next results we have got analyzing influences of stress state of the medium to registered information. As the level of stress st permanently decreases. Surprisingly in quiet-area approximately in 4-12 times. From that point of view becomes understandable a physical meaning of so called seismic gap appearing during the period of earthquakes sources activation. Thus, the specificity of seismic noise intensity change revealed

to judge the level of stress state of the me



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Seyed Mojtaba Mirmajidi, Dr.Ahm

One of the best ways to study deformation anisotropy, or the dependence of seismic v an anisotropic medium, shear wave split int different velocities and write characteristic e wave's trains. In this paper, we present the using aftershocks of the 28th May 2004 Balader

in the direction of microcracks, and coincide tectonic stress in the region. These effects crust in the region. It seems that, this she through stress-aligned fluid-saturated micr parallel to the crack surface, the velocity is polarized perpendicular to the crack surfac



tion and interpretation of seismic ation and polarization direction. In ogonal polarizations that travel at ree components seismic es int ting i e crust of central Alborz shear-wave splitting was

observed on the records of the selected aftershocks. The polarization of the faster shear wave is aligned

with the direction of the maximum horizontal compressive uni

hisotropy of the upper caused by propagation shear wave is polarized clearly if shear wave is

Ima

Keywords: crustal anisotropy, shear wave splitting, fast wave direction









Deformation processes in the Caucasian reg of strong Earthquakes are considered. For earthquakes for the period with 1970 on 199 of speed of deformation essentially depen components of tensor average speed of def separate areas chosen for consideration. average speed of deformation for separate element

lead the detailed analysis of the most intensive areas of deformations in the Caucasian region. Quantitative characteristics of dominating given. Orientations a component of tensor are intensive vertical movements. In a so mainly in a horizontal plane, but their dir horizontal making speed of deformation in its orientation in the central part of region in structures adjoining to coast of Caspian sea.

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chanisms of the centers deformations tool data of strong 5.4 and <= 7.0 are used. The size welume in this connection ogen as a whole, and for the for reg size e component of tensor a terrestrial surface, is

processes in horizontal and vertical planes are al part of region there processes are focused direction of a vector of aced on near meridional on southwest in the













IUGG XXIV General Assembly

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Perugia, Italy

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JSS012

Symposium Earth Structure and Geodynam

Convener : Prof. Hitoshi Kawakatsu Co-Convener: Dr. Andrea Morelli, Dr

Seismic tomography has revealed the pres associated with the subducted oceanic lit current and past subduction zones. Alth

established, their cause, fate and impact on are not well understood. We solicit contrib eventually answer questions like "Why do occurs when they fall into the deep mantle? of subducted slab (including oceanic crust) pressure experiments to delineate the fate slabs and surface tectonics and volcanisms,

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tle Slabs

malies, which are likely to be me mantle transition zone of the of such "stagnant slabs" is now well olution of the earth system nes of earth science to dis where do they disappear?, "What de: detailed geophysical mapping antle, dynamic modeling and high-

tion between stagnant

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Pacific slab stagnant in the mantle transi Philippin Sea using data from a þr ne be broadband ocean bottom seismograph (BBC hant Slab Project (2004-S) work. the \$ 2008), twelve BBOBSs were deployed on the Philip Sea. We analyzed ninein t twelve months long data recovered by the cruise (c research vessel KAIREI with the to determine the mantle discontinuity depths beneath the northern Philippine Sea and westernmost Pacific. Among them, eight stations have co tinuou ecords longer than nine months, which were used in the further analysis. We employed the functions (Gurrola et al., 1994). We stacked receiver functions e stagnant Pacific slab round the stagnant slab. imaged by a seismic tomography to detern The 410-km and 660-km discontinuity der transition zone are 384 hs th km, 692 km, and 308 km, respectively, velocity correction. As a comparison, we estimated the discontinuity depths beneath a normal Pacific region from three BBOBS stations in the westernmost Pacific Ocean, giving 392 km and 651 km for the 410-km and 660-km discontinuity depths, respectively. The 660ep in the stagnant Pacific ntly q slab. Assuming that the deep 660-km disc mperature environment C.0 of the stagnant slab, the thick mantle transit estimated to be 300 K colder than beneath the normal Pacific Ocean.

Keywords: transition zone, bbobs, philippine sea













Keywords: anelasticity, mantle, bbobs























Keywords: thermal conductivity, thermal diffusivity, high pressure







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smology and Physics of the Earth's

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Symposium The lithosphere

Convener : Prof. Sierd Cloetingh **Co-Convener :** Prof. Hans Thybo

This session will include multi-disciplinary geochemical, and geologic data. The foculithosphere, including its physical proper

processes that have formed and modified the continare welcome. Some of the topics to be addressed seismic models of lithospheric structure; (a) goal structure; (b) studies of mantle xenoliths in the colithospheric stability; (c) models of lithospheric or electrical properties, and other data. (c) di currents according to method, e.g. seismology, s iuccul petrology.

ntine collitrosphere and context of context ormation from a os of context of context ormation from a s of context of cont

ect and direct methods. Studies of the osphere circle the Archean to the present integrated the sal, gravity, electical, and ion methods of lithospheric commentary geophysical results; (4) from antle xenoliths, seismic anisotropy, autore of decisions of the lithosphere elemination of the lithosphere elemination of the lithosphere

athosphere from geophysical,














































this shortening has been interpreted to b odat /ard acci by w block bounded by conjugate dextral and sinistral stri as also been shown that a ault syste change of dextral component of reverse faults to sinistral ones took place at 52 Ma ago. We have used remote sensing data in order to investigate rrelati on bet ineaments pattern and 009 *l*eer Riedel model. The study was focused on D highlands; and Landsat

ETM data were utilized to spot fault lir 2540 lineaments were distinguished. We discuss: 1) how compa abl lineaments is with the Riedel model and 2) how a useful tool is t is r edi uctures that may form in nt zone (PDZ) commonly shallow crustal rocks deformed in a strike slip d forms parallel to the shear couple. Due to the direction of Arabia-Eurasia collision (N010-040), the prevailing structural trend is WNW-ESE; and 115 was taken as the dominant trend. Normal faults (T) should form perpendicular to the direction s. Rie el shears (R) form at an m limi angle of (~15-20) with the PDZ, where ne faulted material (we considered which is typical of the upper crus se of strike-slip motion as the PDZ, and they typically form a suite of en echelon fractures. Conjugate Riedel shears (R') can form at (~60-75) relative to the PDZ. These corrected have he opposite ises of strike-slip motion as the main fault. Finally, folds and thrust faults should form with their axe s or traces, respectively, normal to the main compressive stress (P). Fold axes will initially be oriented at about 45 to the PDZ. Activation of first order phases results in second and third order R, R', T and P systematic fractures. Directions of the lineaments spotted on s ested with directions of the structural elements predicted by Riedel model for systems. Most lineaments are compatible with this model. Exercising with cations for characteristics that fit each lineament.

Keywords:















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The general regularities in the motion of the largest seismic events on the Earth sur by the closest image are connected with boundaries of plates, and seismic belts (zo plates. On the other hand it is enough sure (Greep, Gordon, 1990; Argus, Gordon, 199 dynamic regularities of plate motion have sliding of lithosphere and, accordingly, the strengt

active displays of seismic activity should take place along the inclined equator of rotating lithosphere. Therefore we had the right to expect, that the pole Pw of the axis of global rotation of lithosphere, the pole Pm of the angular moment of relative seismic belt should have a close positions the made assumption. Thus, the global rot seismic events, and its equator of rotation (zones) of the most active seismic events, plates and lithosphere with various powers of oceanic and continental areas (Barkin, 2000) has been

used for analysis. Only surface displacements are considered in this model in accordance with known kinematical theories NNR-1 and HS-2 (computer method of axography (Ferrande most active poles (Ps) of latitudinal and earthquakes in 20th century (coordinates in degrees): 1) 65.5 S, 60.5 E. (the analysis of 112 largest earthquakes in 20th century); 2) 53.5 S, 45.5 E earthquakes); 4) 52.5 S, 56.5 E (392 earthquak latitudinal (3)) and equatorial ordering (4)) (Ferrandez Garcia et al., 2000) have been used. The

coordinates of the following poles (Barkin rotation of lithosphere; pole 45.4 S, 57.6 E plates under theory HS2-NUVEL1; pole 48 lithosphere. It is shown, that the vector of (together with the Earth) is located in an belt Ws (longitude 145.4 E). The centre of located near to the big seismic belt. The northern arch of Pacific ocean, with ascen other. T termine these meters of ala (in, 20 ertain tablis (Ba 2000 umulať

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ribution of epicenters of Plates and earthquakes stigat licity is host brightly shown on indaries and configurations of the thosphere rotation are determined metrical, kinematic and he maximal tension at tic energy and the most

> blanetary (most active) esearch have confirmed hetary scale) and directs planetary seismic belts mic model of lithosphere

273) 54.5 S, 41.5 E (112 of longitudinal (1), 2)), as indexe obtained coordinates of a pole of a seismic belt are coordinated with each other and with corresponding 65.0 E of angular velocity of global of relative motion of lithospheres pmentum of global rotation of of inertia Lд of lithosphere escending node of seismic inates 41.0 N, 36.0 E is eismic belt covering big .0 E and an inclination

Goddon, 1991). The special

for determination of the

of epicenters of large

about 55.0 degrees), up to seismic zones of South America. With poles of a vector of the moment of inertia forces of lithosphere, caused by the Earth rotation, the extended zones of active seismicity are connected. They form of the spirals located in the field of longitudes 90-180 E (spirals are twirled

clockwise) and 270-360 E (are twirled cou north - south. Their origin can be connected lithosphere. The moments of these forces accumulation of elastic energy and influen centers located on the basic seismic belt

ation in a direction the inertia for non-spherical blates result in additional Three ring zones with the 5 N, 120 E; 5 S, 290 E) are

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marked. Ring zones are closely connected indication, that activization of seismicity ar general mechanism. References Barkin, Y Towards an Integrated Global Geodetic Geodesy Symposia. Vol. 120 (Eds. R. R Symposium (October 5-9, 1998). Spring character of the plate motion: implication Observing System (IGGOS). International 231-233. Barkin, Yu.V. (2000) Geometrica and Astrophysical transactions, Vol. 18 regularities in plate motion. Astronomical a Barkin, Yu.V. (2000) Dynamical regulari Transactions, Vol. 19, Issue 1, pp. 1-12. G 17, pp. 1109-1112. Argus D.F., Gordon R. incorporating plate motion model NUVEL-1. 2042. Ferrandez, M.G.; Barkin, Yu.V.; Ferrar the earth-like planets and moons. In: Eart Symposium, 3 - 8 June 2002, ESTEC, Noord 514, Noordwijk: ESA Publications Division, IS

July 13, 2007 Perugia, Italy

s on plate boundaries and it serves as the on of ring zones are determined by the rotation of the lithosphere. In: S). International Association of H. Hornik). IAG Section II Yu.V. (2000) The regular egrated Global Geodetic 120 (Eds. R. Rummel, H. Drewes, W. Bosch, H. Hornik). IAG Section II Symposium (October 5-9, 1998). Springer, Berlin. pp. structure. Astronomical V. (2000) Kinematical 18, Issue 6, pp. 763-778. tronomical and Astrophysical 790) Current plate velocities relative to the hot spots incorporating the NUVEL-I grooar prate motion model. Geophys. Res. Lett., 1990, vol. otatio current plate velocities ol. 18. N 11. pp. 2039hvsical _etters 1. (2002) **D**rd positions of formation centers of lanets and m Proceedings of the 36th ESLAB q, B. Battrick. ESA SPhe N 9092-8 er 200 p. 129 - 135.

> Keywords: lithosphere rotation, se sinc belts positions, regularities

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We present a transversely isotropic shear w the Mediterranean region, obtained by nonvelocity maps result from the regionalization and Rayleigh wave dispersion, carried out techniques. The linear inverse problem is a group velocity reference model, derived fro al. (1997). The implications of different regulariza

damping or smoothing with different criteria) are analyzed and compared. Both in group velocity maps and in shear velocity resulting model we find confirmation of the larger-scale deep geological features known for the region, namely the different North Western Africa and the slow seismi slow anomalies associated with the magma of the Rhine graben. Fast anomalies are dense European seismograph array results

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tle beneath Europe and wave group velocity maps. Group ments of fundamental mode Love and phase-matched filtering ring in the form of a global forma¹ y dataset of Ekstrom et se ve cally equivalent to norm

> of Eastern Europe and e length we image well nd the extensional zone Hellenic Arc. Use of the previously attained.









suggest a high marin level so that the envi

Keywords: mineralogy paleoenvironment, ypresian lutetian, illite kaolinite smectite



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JSS014		2252 - 2327
Symposium Crustal structure and Tectono active continental blocks and	op then boundaries	heric structure in

Convener: Prof. Kevin P Furlong

Although the concept of localized plate bou clear that deformation associated with the oftentimes distributed over a substantially defining the structure of plate boundaries, interaction, the depth extent/distribution of defining the nature of plate (or crustal bloc range of geophysical and geodetic studies, deformational style, and the resulting diffuse

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ation of modern plate tectonics, it is crustal and lithospheric blocks (or plates) is cus on studies aimed at n associated with block forma novative tools or approaches for rage contributions from the broad provide insight into the structure, iry stru e.





Guralp CMG-3T (120 s) seismometer an continuous mode at 50s/s. The stations were align strike of the Himalaya. The stations were closely spaced at ~7-10 km interval from the Himalayan

Frontal Thrust (HFT) in south to the STD i craton. We analysed over 100 teleseismic April 2005- Oct 2006 to generate velocity structure model at individual station is different azimuths with the surface wave include: 1.Northern Indian shield (Arava comparatively lower velocity (<3.4 km/s) in the upper crust relative to the southern Indian shield

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nd are widely spaced in Gangetic plain and Aravalli by the network during The 1-D shear velocity receiver functions from of the important results km thick crust with a

(Dharwar craton) where it is >3.6 km/s. 2.Dip of the Indian Moho is 6-8 deg. beneath the lesser Himalaya with the average Pn and Sn veo erive from the travel times of updip and downdip refracted waves. clearly map the Main res Himalayan Thrust and also the Indian Mohb, han le if their dip to the north of ing MCT. This is also supported by the seismic trend. 4. Presence of low velocity in the depth 20-30 km in the region between MCT and STD is modeled from receiver function series. 5. Uplifted 410 Seismicity is prevalent in discontinuity while the 660 km discontinuity split into 660 and Kin depths. the lower crust while it is nearly absent in the uppermost mantle.

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structure for the source/network vicinity using GA-HYPO (Kim et al., 2006). GA-MHYPO is a genetic algorithm for optimizing hypocentral parameters solution to the 1-D velocity within prese between observed and calculated phase a thickness) in this study is modified de distribution. Preliminary results indicate th Himalaya region are significiant, as has 2006).

and crust. Phases were re-picked using an

reduce reading error. We then jointly solved for ne

Keywords:

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mea es (Park et al., 2004) to and average 1-D velocity

1-D velocity structure by searching for a global travel-time difference ber of layers and their uakes and their depth velocity structure of the (e.g., Monsalve et al.,

















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Giulio Selvaggi, Antonio Avallone niek ni, Eliş P'Anastasio, Sergio Ma nutò

The Adriatic region has always puzzled an la studies of the Alpine-Mediterranean plate hι studies have described the Meso-Cenozoic e oll African plate, seismological and space g bdei indipenden Adriatic microplate. In this study we u

together with the analysis of earthquake slip vectors, to study the crustal motion and boundaries of the Adriatic region. Based on a rigorous statisti which describe the kinematics and the act relative motion between two microplates: single microplate with Apulia resolving the We use a simple block model to illustrate h relative motion. The proposed present-day fragmentation of the Adriatic promontory.

of the researchers involved in the stratigraphic and paleomagnetic rigid promontory of the gion a rt the evidence of an ily su style GPS measurements

of the data, we propose a set of Eulerian poles ranean in terms of the e Ionian region forms a ary between Adria and . nodate the Eurasia-Africa tion to the evolution and















Despite considerable attention directed tow cum-Adriatic bnics in the recent years, the rate and direction of Adria microplate motion an mentati remain unresolved. We al m analyze tide gauge data from the eastern dr tial variation in vertical Sea i de assess tectonics. Tide gauge data are available fro tes ir nia, C ia and Montenegro with eig y, long time series (30-95 years). Correlation among ates that the interannual age reco sealevel variations are common mode. We are thus able to mitigate bias in trend estimates associated with interannual variations by common-mode filteri (Wdowinski et al 1997 · Davis et al., 1999). We find relative vertical crustal rates (with res fluctuate between 0.7 mm/yr which may be indication of active t rates determined from a sparse network of continuous GPS site d) in the region are in general agreement with the tide gauge r analyzed campaign and bul nce continuous GPS data collected along the easter dr o investigate the pattern of horizontal deformation. We find that NE directed horizontal shortening across the Dinarides that varies along the coast of Croatia, revealing apparent rotations that are not well resolved by the present GPS data set, but which nevertheless provide noder configuration and motion of the Adria micoplate(s). Velocity estima with evious models for Adria ompatik microplate motion (Battaglia et al., 2004; Ca tria bay a Northward motion as expected but sites further south along the coast (latitude 44 to 45 degrees N) have a more westward motion, up to 40 degrees off from expected with Adria p les. Motion sites on the Island of tform, respectively, differ Palagruza, Croatia, outermost Dalmatian Island, and ily, Puglia pla Matera; by less than 0.20.3 mm/yr in magnitude and 03 degrees azimuth, which implies that Palagruza and Puglia lie on a common southern Adria microplate - We will also discuss a new project that will involve addition of a number of new continu ditional GPS campaigns, and other measurements.













at the mid crustal (15g the NW trending fault km), and by right-lateral epth (25-38 km). Ground st a fault interaction model,

fault/rupture at depth, one in NE direction and the other in NW direction. The aftershocks along the NE trending fault/rupture show reverse faulting <25 km) as well as at the lower crustal (2 or rupture occurred by pure reverse faultir strike-slip faulting at the shallow crustal observations are compatible with the ab

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which illustrates that the main shock original and it propagated ruptures along the N revealed more detailed picture of the com 2000b). It shows high Vp, low Vs and hig area is a fluid-filled fractured rock - matrix Seismotectonic study of the two recent So 138. Kayal, J.R. and Mukhopadhyay, S. earthquake source area, Bull. Seism. Soc

India and its seismotectonic implications. O P, De R. and Singh, O P. 2002b. The 20 hypocentre and its implications for rupture July 13, 2007 Perugia, Italy

ase of the paleo-rift zone by reverse faulting, e NW directions. Seismic tomography res in this rift-basin (Kayal et al., a, which indicate that the source References Kayal, J.R., 2000. Geol. Soc. India, 55, 123ure of the 1993 Killari R., Ram, S, Srirama, B.V., and Gaonkar, S. G. 2002a. Aftershocks of the January 26, 2001 Bhuj earthquake in the western , J.R., Zhao D, Mishra, vidence for fluids at the (24), 51-54.

> Keywords: fault pl , tectonic model /2007 UGIA















formation of the Dabieshan orogenic belte difference in the peak period and the disc coesite and eclogite from the Dabie oroge reentry from mantle around the depth of evidence of the deep structure and dynamic p explore the 3-D velocity structure of S wave beneath this region by using the inversion method with 4 4 network based on the dispersion effect of Rayleigh wave. Our purpose is to reveal the formation

mechanism of the UHPM zone in the Dabies well as deep dynamic processes of intrap crust and upper mantle. Then we would put dislocation of the Dabieshan and the Tanlu fault zone, and the coupling effects of intraplate subduction.

The results show that, there is a high-velocity subduction s b of tongue ape with underthrusting depth 160km extending from east to west. A new velocity manue hotspot with a 500km wide head is found at depth of 70km in the upper mantle beneath the Qingling-Dabie orogenic belt. The outcrop of coesite and eclogites in the Dabie orogenic_belt_is_the_synthetic consequence of the movement and dynamic effects of deep (upper mantle) ma deep substance and energy exchange, strong strike-slip motion along the Tanlu fa luction.

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has been carried out by colleagues in Japan Contraction Geological Engineering, Co. Ltd.



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Survey C

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Gou Fujie, Kei Murase, Ryuji Kelenta, Filiko Nishiya Tanaka, Kayoko Tsuruga, Shigeharu Miz hata Shige, Kayanura, zusa Nishizawa, Kanta Kaned.

In order to obtain on accurate best crustal it is stru important to make the best fit between the major seismic phases de OB eflection and refraction ed in de-and survey data, and the MCS reflection section loped ew interactive software Re ίv, ave module for the Modeling-Pasteup crustal structu is tool, developed by Fujie et al.(2002), for OBS survey data. The improved interactive modeling module has two major functions: Modeling and Pasteup. Modeling and Paste under X-Window circumstance of Linux OS. can Under such computer circumstances, we by a PC. Modeling is forward analysis software to carry out 1 el time calculation for refracted first and later arrivals, reflected ves, and headwaves, 3) raypath calculations, 4) depth to time co 5) gravity modeling. We /er define the crustal structure by several ver ds the layer interface. When a discontinuous layer boundary is required, we can define two different velocities for the above and the below the particular layer boundary. We can compare traveltime pickings based on observed records and theoretical traveltimes on the Paster vel times of wide-angle ute t CO reflections and superpose on the OBS seis re module can compute of so the two-way travel times of normal incide structure model. The om rus software module also superposes the time contours of the layers on the crustal structure model over the migrated time section of MCS using the Pasteup a hug volume of data, we confirmed that strong reflections from layers in the crust and the the OBS w e-angle reflection records bho seen are fairly consistent to the reflectors in the MCS section. By use of forward analysis, we also evaluate kind of waves for later arrivals. It is great help to the crustal modeling to use full wave information on seismic records. Synthetic waveforms calc directly compared to observed ones. Gravity data are also used to evaluate the t model. Through the interactive analysis, we can obtain the best fit mode above two processes, we can confirm the correctness of the resultant cr result of forward modeling as the input of traveltime inversion to min ne inversion. The average rms misfit is approximately 30-40ms for nce if arrivals are clear enough. In the actual processing, we used in shots for the longest survey line.

Keywords: crustal structure, obs, inteructive analysis
















seismic velocity structurehas beenimaged ershocks recorded by 47 by OBSs and relocated the hypocenters simu ion of aftershocks shows ane landward dipping planar shape with an increased dip angle of the hypocenter distribution at around the distance of 100 km from the trench axis. Furthermore, the hypocenter distribution projected on the parallel plane to the Kuril trench indicates heastward. Such a varied to no dip angle of the hypocenter distributionres epth of boundary of the subducting oceanic plate as evaluated from ng the air gun control ΝĽ .an' sources. The rupture area of the 2003 Tokachi-oki earthquake and its neighbouhoodcan bethus interpreted as a highly deformed zone where that i Japan ard ets the Kuril arc moving Vortheaste eting obliquely under the westward, which issupported by the evidence that e Pacifi te is subt Kuril arc.



















Keywords: gps, northern apennines, geological geodetic strain

















Taiwan is located along a strongly oblig subducts northward beneath the Ryukyu Ar Arc. In order to provide insight into the collis damped stress inversion technique (Hardebe of the data misfit and the model length, to from the Broadband Array in Taiwan for S region was gridded with 0.1 degree space node. We next simultaneously inverted for

degree x 0.3 degree rectangle centered at the grid node, but the grid node with less than 8 earthquakes was abandoned. For the damped inversion corner of the trade-off curve, in the lowe relatively small, and stress orientation und the entire data set. The 95% confidence r solutions closet to the preferred solution. general the direction of compression is con EP but varies with the tectonic units. The direction of compression shows a significant clockwise rotation



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Philippine Sea Plate (PSP) ward beneath the Luzon EP) e -building for Taiwan, we applied a which minimized the weighted sum which are shallower than 35 km om 19 uly to 2006 December. was a ned to the nearest grid all events within a 0.3

we chose the value of the damping parameter near the and data variance are botstrap resamplings of by the 95% of bootstrap version indicates that in h direction of the PSP and

rifting of the Okinawa Trough in northeast the crustal stress field in southeastern significant stress boundary which is corre northwestern tip of the subducting PSP.



00.

Keywords: stressi



































Mineo Kumazawa, Junzo Kasaha kaliro Nakajima, Ken Kunitomo Taka Hasegawa, Yoko Hasada, Toshiki vata be, ipen luan Toshiaki Masuda, Mikio Satomui suyoshi Mia ayashi

We have developed an active monitoring s S (Accurately Controlled Routinely ame<u>d the</u> Operated Signal System) in which a tensor rai funct wit ahly re approach will be the best way to discrimin small ange npoi (material dispersion) and heterogeneous structure ust tensor transfer (Greens) function sampled at finite discrete frequencies in a limited frequency range. From the Toki transmission site we have continuously transmitted circularly (horizontally) polarized seismic waves with modulated frequencies seismic ACROSS transmitter at Mori-mac energy with frequencies from 3.5 Hz to 8 that can cause temporal variations of stre ks. eruptions could be reflected to the scattered sources. The heterogeneity in the lithosphere originated from both stress state and heterogeneous distribution of fluid-bearing rocks can be the scattering sources. Temporal variation of such scattering sources due to the structure sensitivity of fistics of seismogenic regions as well as the active volcanic regions. The a wo detect and clarify such an evolving process uctive solutivity of rocks in the hat nec e s crust and upper mantle. Among many structure sensitive phenomena, probable changes in the reflected or scattered seismic or electromagnetic signals are expected Temporal w anisotropic dispersion of the transmitted signals and likely to 🖌 in the s scattering sources are evolving associated with the movement of the fluid mainly composed of supercritical water in the crust and upper mantle conditions. Recent discoveries of intermittent occurrence of slow slip events and deep r the subduction zone could be one of the most challenging targets to clarify their active monitoring techniques, as well as the dense networks of GPS and several permanent and temporary station

eters). The acquired data by is purposes to discriminate hat the data acquired by at least three possible by nearby structure, (2) heterogeneities of the

rad

media. Factor (1) can easily be eliminated, since they are common for all the receiving sites surrounding the transmitter. Different signatures of frequency dependencies of each tensor component of transfer function would also reflect from combinations of the above causes presumably material dispersion and

heterogeneities of the propagating media. vision (color vision) of the complex struct are developing Electromagnetic (EM-) AC Shizuoka University, in order to physically

signals of the heterogeneous crust from

seismic ACROSS often show highly frequ

causes for it; (1) frequency characteristics

material dispersion of the propagating m

e frequency dependent to Seismic ACROSS, we current dipole source in tterers in the crust.

e error estimation. This

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Observable phenomena

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duction zones where the


























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400-1000 km in various modes of radiation Res structure and monitoring its stress state executed Science are submitted. Estimates of the vibroseismic signals amplitudes and the required vibrating sources power for recording at the teleseism of super-power vibrators, problems of the precise control are analyzed. Variants of c of 1-10 thousand tons for the operation or CO power vibrating sources network for realization research problems of the Earths global hmc

most powerful today 100 - ton seismic vibra

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the upper mantle is executed. The bratin nals at the distances of ic res ches of the Earth's crust the Russian Academy of

are received. The principles of the construction and problems of their sources with the force d. The worldwide superth is proposed. Possible f them is the study the medium movement in internal areas of the Earth and, in particular, rotation of the Earth core.











Keywords: seismic velocity structure, <u>controlled</u> source seismology, philippine sera plate





Dipartimento di Scienze Fisio

derico II, RISSC-Lab

Giuseppe Pasquale, Raffaella Mathi

The Southern Apennines is one of the Italia as evidenced by the seismic catalogues earthquakes (Boschi et al., 1998). The mo producing more than 3000 casualties and e activity is characterized by low-moderate e wave velocity crustal model of Southern Apennines

model we re-localize the earthquakes and then we compute the focal mechanisms in order to estimate the stress field acting in the area. In order from 1988 to 2003 by Istituto Nazionale Seismic Network) with arrival time of af temporary network. Only the earthquakes whi and the final database consist of 1196 ear hqu used the linearized, iterative tomographic ap inversion of local earthquakes first-arrival travel-time to solve simultaneous velocity model parameters

and hypocentral parameters. The iterative technique takes into account the nonlinearity of the problem but in each iteration the method is based pr influence the inversion process. In order to obtained as the mean of the 250 velocity va with different 1D reference models randomly generated in a selected range. Uncertainty associated to the velocity value for each cell has been analyzed and the model resolution as been evaluated through the velocity value for each cell has been analyzering the model resolution

a standard checkerboard test. In order to give an improved re have re-localized the 1196 earthquakes in the 3D final velocity model using a probabilistic, non-linear, global-search earthquake location method (NonLinLoc_code, Lomax et al., 2000). The relocated seismicity is shifted eastward caused by a It is clustered around the system fault between 0-20 km.For the computation of plane fit grid-search algorithm (FPFIT) of the minimum number of first arrival polarity s) and maximum values for ERX (≤ 0.8 k FPFIT algorithm computes the orientation

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most intense geodynamic activity; strucked by several destructive on <u>1980</u> 23 november, M 6.9, Nowadays the seismic a three-dimensional Pocal earthquakes. In this

nchone, Aldo Zollo

a <u>3D</u> velocity model we merged the data collected work (Italian National hquake recorded by a tations were considered arrival time readings.We al. (1996), which allows

hat the starting model will

Lin the D inversions performed

final velocity model was

sentation of the seismicity pattern, we eastern part of the investigated area. wake and have a depth ranging echanisms we used the fault-). Restrictions are place on table RMS residual (≤ 0.5 m) and gap ≤ 180. The prrespondence between ating the stress field in this region, we applied the Michael (1984) procedure to our data set of fault-plane solutions. It

is vertical. This result

on, as evidenced by the

computes the best uniform stress field for the dataset and meaningful confidence regions using a statistical tool known as bootstrap resampling. Stress inversion shows a nearly horizontal NE-SW minimum compressive stress axis (σ 3), w reveals that Southern Apennines is generation fault-plane solutions of major earthquakes

P and T axes and the principal stress orie

Keywords: tom

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Symposium Underwater observatories

Convener: Prof. Barbara Romanowic:

The international earth and ocean sciences in the oceans in order to provide optimally when appropriate, and for the long-term i

local scales. International Ocean Network (LON) disciplines, and to facilitate cooperation i systems, harmonization of those elements developments are required that should prov from the land to the ocean bottom. The s scientific tasks which need joining of interna problems solution to create new generation but with sufficiently high level of parameter, devitimes and take advantage of observatory infrastru-

d for long-term observatories obal scale processes, in real-time dependent processes on the regional and ormed to f synergies among different ht of nents of the observing ical⁻ to achieve these goals, technical e study of investigated parameters or

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concentrated on the discussion of Well as on technological long-term observatories operate stably over long iterfaces to facilitate the

armoniza interchangeability of instruments and the sharing of maintenance tasks is also between sessions interests.

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destinated for the International Ocean Network and its performances are partially supported by the STCU Project -3165.

Keywords: ocean bottom, geomagnetic observatory, orientation

sented. This study was















M8-class damaging submarine earthquake almost regular intervals at the red repeated subduction plate boundaries around Japar the Philippine Sea plate subduct Pacifi beneath the Japan Islands along ocean tren Trench and the Nankai the K Гrе the Ja Trough, which locate off the east coast of Plate duc cause epetitive occurrences of IJap damaging earthquakes in Japan. Therefore the more seismic in the focal region of the damaging earthquakes is very important for earthquake forecasting. We have developed a new longterm ocean bottom seismometer (LT-OBS) ability to record seismic signals continuously hich over one-year period. The LT-OBS is the n hic activities in the sea by a dense seismic network on the sea flo earthquake data in real time. Ocean bottom cable systems (OBCS) monitoring in the focal region of the damaging earthquakes. W ctive OBCS, jointly with de engineers of various fields such as co nd eal engineering, metrology, electronics, and mechanical engineering. A main specification of a stationary type OBCS is as follows, 1. Landing Station: two place; both stations have power-feeding units for redundancy 2. Sensor: 40 of Seismometers and 3 of Tsunami sensors se observation 4. Cable for d length: 900km 5. Dual Ring structured ne rant and redundancy 6. fault Smaller sized OBCS housing for lower cost livery accuracy of time າar stamp: more than 0.1ms by PTP 8. Operational life: 20 years Authors will discuss new cost effective OBCS in detail in a paper, especially, new network tructure for the OBCS, h precision time stamping, ype OBCS and high reliable analog signal digitizer interface design. Mobi within our system design target, too.





Keywords: seafloor em station sfems, the northwest pacific basin, geomagnetic secular variation



















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Sn are 8.2 km/s and 4.7 km/s, respectively. 3.5% for S-wave. From this result, a seisn fast direction seems to be perpendicular to WP-2 is suggested to originate in a prefe The upper mantle velocity structure is infe and the previous studies of the northwe earthquakes with epicentral distances sma km and a rapid increase of velocity at a

from 30 degrees to 90 degrees, and ha bandpass filter of 8-360 seconds was appli the WP-2 model which is modified lasp91 airgun experiment was constructed. After model, travel times for large amplitudes o discontinuties were read. Averaging the read times, conversion depths of 416 km and 666 km were obtained.

variations are about 5% for P-wave and about ppermost mantle is suggested, and the herefore the anisotropy below the rystals in the uppermost mantle. guakes recorded by the WP-2 ate first arrivals from the ne below a depth of 30 form receiver function analysis, we selected 16 events with a magnitude greater than 6, that have epicentral distance ranging om the WP-2 data. A A velocity model named aced by the results of the e calculated using the WP-2 presponding to 410 km and 660 km

> Keywords: seafloor seism , seismic anisotropy atory, lithosp

> > /2007

UGIA





Kinematic GPS analysis and measure the g tan etwe floor by Acoustic measurements. Next we dete GPS/Acoustic observation has been conducted at the Kumano Basin and Suruga Bay since 2004 and 2005, respectively. For the repeatability of observation a precision of 2-3 cm at horizontal comport three continuous GPS sites on land for Kin on the top of the pillar and dual-frequer records GPS signals continuously with a sa hpľ of velocities are inconsistent with those obtailed two factors. The first factor is the difference of observation period. The second factor is annual and semiannual variations of GPS time series. We neglect annual and semiannual variations when we estimate GPS sites velocities. GPS/Acoustic for GPS site coordinate, we adopt the ave before and after observational days on the а efe

benchmark on the sea e benchmark. Repeated the location of benchmark is determined within et al., 2006). We have ing antenna is mounted second. each receiver ever estimated GPS sites

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y be caused by following ays par one campaign. As which includes one-week inate for Kinematic GPS analysis.

There is a possibility that GPS sites coordinates have some bias, because annual and semiannual variations are contained for about 20 days. In the study, we vestigate ho he determination of GPS sites coordinates affects the estimation of the location of be mark on t sea floor. We use data between November 2005 and October 2006 at the Kumano Basin and estimate GPS sites coordinates using Bernese (Ver 5.0) software. Then we calculate the variation of the location of benchmark according to the bias of GPS sites coordina tic GPS analysis.





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KALACHAND SAIN National Geophysical R The energy-demand in developing countries the world. At present we import more than hydrocarbons have not taken place during our huge population depends heavily on the form of energy. Gas-hydrates seem to be water and low molecular weight hydrocarby low temperature in the outer continental margins

global attention due to their wide-spread occurrences and potential as future major energy resources. Parameters such as bathymetry, seafloor temperature, total organic carbon content etc. indicat margins of . The methane gas stored in t exclusive economic zone is estimated to b our energy requirements for several tens aegis of the Ministry of Petroleum & Natur of scientific plans and technological advancement for the exploration and exploitation of gas-hydrates.

As a member of the national program, we are mainly involved in finding out the prospective zones and evaluating the resource potential using gem identifying an anomalous reflector, known based on its characteristic features. We idea Krishna-Godavari and Mahanadi basins and the Andaman region. The recent drilling expedition at the identified locations of gas-hydrates occurrences, Godavari basin. The Mahanadi basin and the and gas-hydrates. Much of the deep-water regions are yet to be explored. New data sets with suitable parameters are to be acquired to comprehend the lateral/areal extension of gas-hydrates, and assess

their resource potential along the margins the existing technique, generally used for a concerted effort is going on in this direction from below the gas-hydrates can be explo to produce gas from gas-hydrates. There hydrates and evaluate the resource potent like traveltime tomography; waveform inve attributes etc. and will demonstrate their a

rabad - 500 007, India to industrial countries of mpar uirement and major discovery of . Since the economic growth with esperately looking for an alternate are crystalline from of -hydr are f ed at high pressure and hese have attracted the

, sedimentary thicknesses, rate of sedimentation, the huge continental ing free-gas within our which can take care of ates Programs under the ences have initiated a lot

or BSR on seismic data

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und very prospective for

hydrates are mostly detected by

s gas-hydrates are not stable at NTP, cannot be used for exploitation. A gy, it is expected that methane It may take a decade or so prospective zones of gasped several seismic tools quality factors, seismic mic reflection data. We

will also show that studies of quality factor and proxies like pockmark and gas escape features (like faulting or gas-chimney) offer indirect evidences for the identification of gas-hydrates in absence of BSR. As pure gas-hydrates have much higher seismic velocities than that of host sediments, presence of

gas-hydrates increases seismic velocities reduces the seismic velocity. Thus estimat by rock physics modeling or invoking effective gas-hydrates and/or free-gas across a BSR

ates-bearing sediments m tomography followed ellent tool for quantifying








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JSS017		2359 - 2394
Lithosphere thermal state an to models	d	n measurements
Convener : Prof. David Chapman		
The lithosphere is identified as the oute act as a coherent tectonic unit. This syn geophysics that relate to the thermal str petrology, etc) either from the perspect constrained by these data. Papers addre mid plate volcanism, delamination, und are especially welcome.	r la seconda papers from the material second papers from the nates of the lither of th	ntly cool and therefore rigid to nany branches of geology and mology, electromagnetism, sets or from creating models as riting, plate convergence, ence for lithosphere evolution
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The most remarkable feature of the Central characterized by 3.5 km average elevation, flow. Furthermore, below the Altiplano-Puna has been established by a number of indep broad low seismic velocity zones, etc.). Th inť huge concentration of Neogene ignimbrite. (most

Puna Volcanic Complex. On the other hand, the forearc and the foreland basins have lower heat flow, thinner crust, and lower altitude. These features suggest that thermal isostasy could play a role in the compensation of the Altiplano-Puna. The variations in the lithospheric thermal regin variations in rock density in response to continental elevation is difficult to asses, be This study estimates the elevation effect due to adjustment, revealing the thermal and geodynamic effects on elevation. The effects of compositional

and thickness variations within the crust were removed using the crustal density structure obtained for the Central Andes between 19S and 30S f well constrained by a large amount of geom elevation was adjusted for compositional b our 3D gravity model, relative to a reference crustal section (average crustal density: 2850 kg/m3, average mantle density: 3350 kg/m3, crustal thickness: 40 km. GTOPO30 sligital elevation model was used to estimate the actual topography. The heat f values recently published. The thermal isostatic relationship describing the thermal contributions to the

elevation was determined using a reference <u>deotherm</u> corresponding to a surface heat flow of 30 mW/m2 and assigning a lithosphere having adjustments range between 300 and 300 observed that no correlation exists betw values. In contrast, the compositionally ad an increase of around 3000 m elevation I foreland basins areas are characterized by the Altiplano-Puna plateau, the Western (and higher elevation adjustments. Our

sufficient to account for crustal cross section needed in the Puna south of 22S. Other at explain the uniform altitude of the Altipla determined that the above mentioned p between about 22 and 24 S and beld

ho-Puna plateau. This plateau is stal thickness and very high heat melting zone at mid-crustal depth high conductivity zones, extre supp ed by the presence of a melting): the Altiplano-

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ess whereby regional on changes result from thermal contribution to hd thickness can mask it. emoves it by an isostatic

geochemical data. The thickness product from

n this study includes new

Elling The gravity model is very

an elevation of 0 km. Average elevation of approximately 6000 m. It is the corresponding heat flow relation with heat flow, with ones. The forearc and the n adjustments, whereas show higher heat flow nal component of the

Altiplano elevation would be of 1 km, the thermal contribution to the Puna elevation would be of 2 km. Previous works highlighted the fact that the Puna and the Altiplano have uniform average elevation in spite of showing great variation in the amount of structural shortening. Shortening estimates are but are less than that

and crustal flow would esults. Moreover, it was nd in the backarc region Complex area of caldera







































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JSW001

Workshop Subduction zone related volcar

Convener : Prof. Sri Widiyantoro

Many densely populated countries are volcanic hazards may have severe conseq The 2004 Sumatran great earthquake wit

study such areas in order to understand the nature picture to understand its possible effects. S veral countries located along subduction zones. S great damage to infrastructure. The symp mitigation" is to bring together people wor especially in volcanic hazard mitigation. Mitigation strategies for subduction zone r plans to reduce risk in subduction zones 'iv) volcanic hazard knowledge, and v) Integrated ge improve volocanic hazard mitigation would be most welcome.

n zones whose seismic and nic conditions of these countries.

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astating tsunami highlighted the need to esand provide a coherent th century in developing the a large umber of casualties and ne related volcanism and hazard udies along subduction zones and hazard mapping, ii) i) • f integrated emergency opmei focused in improving al stu k geochemical studies to

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The analysis of modern data of geodetic, g phenomena of a drift of outer core of the Ea frequencies relatively to elastic mantle (Ba 2006). Displacements of the core are acc variations of its tension state which in turn f speaking, in all planetary natural processe component of process and by periodic varia have inversion character. Activity of process increa

and activity decrease in opposite hemisphere. It is principial property of considered phenomenon. This theoretical conclusion based on a geodynamical model about forced small relative displacements and rotations of the Earth shells under action celestial bodies, has obtained a set of conf opposite hemispheres of the Earth (Barkin the phenomenon of inversion in activity of Dynamic studies and the satellite data about g core in present epoch (predicted by author) to the north in the direction of a geographical point 70 N, 104 E (Barkin, 1997). In hemispheres N and S, separated by a plane of inclined equator, the subduction

zones and zones of rifting are located asym of compression (of subduction zones) are spreading zones) are located in opposite ha zones of volcanos in northern hemisphere will result in their activization. In southern hemisphere S the changes of volcano activity will have the opposite tendency. The specifie actually. The number of eruptions of volcanos 45 st 5 eruptions per one year (for the period in hundred years). Thus the number of eruptions of volcanos of

known data on volcanic activity for a time method of sliding average the diagrams of period 1800 - 2000 have been constructed of the forced displacements of the core al particular with the decade periods. By rehemispheres N and S will be correspon specified hemispheres. It was actually rezones, falling on 1885, 1950, coincide with

antiphase of zones of activity. These i corresponding zones of subduction and rif intensity of volcanism in the specified peri the Earth core and developed geodynamic The cyclicities of volcanic eruptions reveal

tigns testify a reality of the with a wide spectrum of fillatio Vilke, 204; Barkin, Shuanggen, ormations of the mantle and by of seismic. volcanic and, generally harac ed by the slow secular ;adal more long periods, and rection of the core drift)

> on, the Sun and other of natural processes in drift of the core causes southern hemispheres. the displacement of the

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Thenomenon is observed

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the charges of tension states in

rifting zones on the contrary decreases. This phenomenon was open as a result of the analysis of the 1977 (Mekhtiev, Khalilov, 1987). By the rs of eruptions of volcanos for the 2004). The geodynamic model n of its cyclic oscillations, in ents of the core between volcanic activity in the volcanos of subduction ng zones. Peaks of the

periods of passive volcanism of subduction zones of 1860, 1900, 1910, 1960 will be coordinated to peaks of increase of activity of volcanos of rifting zones. Only to peaks of activity of volcanos of subduction zones 1910, 1930 and to the peak of passivity of 1940 are not strongly pronounced of some imposing of

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nd S. Thus, changes of spective displacement of e Earth shells as a whole. the periods 20, 40 years and

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others (Khain, Khalilov, Ismail-Zade, 2004) about unity of cyclicities of variations of a core oscillations ("a motor", "a heart" or Secular Redistribution of the External and of the Geopotential and Earth?s Rotation. Annales Geophysicae, Supplement 1 to Vo variations of the Earth rotation at diffe Problems (Eds. Steven Dick, Dennis McCa V. 208. Sheridan Books, Chelsia, Mich endogenous activity of planets and satellite Nat. Sciences, Issue 9, December 2002 Shuanggen, J. (2006) Kinematics and dy (Vienna, Austria, 2-7 April 2006). Geoph geodynamics states and possible pulsation sciences, issue 12, pp. 9-16. In Russia geodynamics. Nature (Moscow), No.5, pp. 4

Keywords:

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redicted earlier on the basis of a hypothesis caused by the uniform mechanism of References Barkin Yu.V.(1997) and Their Role in the Variations § (Vienna, 21-25 April 1997). V. (2000) A mechanism of Historical and Scientific pf IAU Colloquium 178 (Cagliari ,Sardinia, Italy, 27-30 September 1999). Astronomical Society of the Pacific conference series, 2002) Explanation of o Zemle, Rus, Acad, of ssian. Barkin, Yu.V. and eres. EGU General Assembly 5, Volume 8, abstract # EGU06-A-01680. Khain V.E., Khalilov E.N., Ismail-zaue T.A. (2004) Periodicity of volcanic activity in various th. ius. P of section of the Earth liev E.N (1987) Volcanoes and SI Russian ion, 12007

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Smedile Alessandra		1947
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Currenti Gilda		1956
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Gvishiani Alexey		1958
Jacques Zlotnicki		1959
Napoli Rosalba		1960
Hashimoto Takesh		1961
Ishido Tsuneo	V V IV/900'	1962
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Sasai Yoichi		1964
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Mogi Toru		1994
Nenovski Petko		1995
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Cutler James		1998
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Fan Xiaoping		2001

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Qian Wei		2007
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Uyeshima Makoto	11/2007	2014
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