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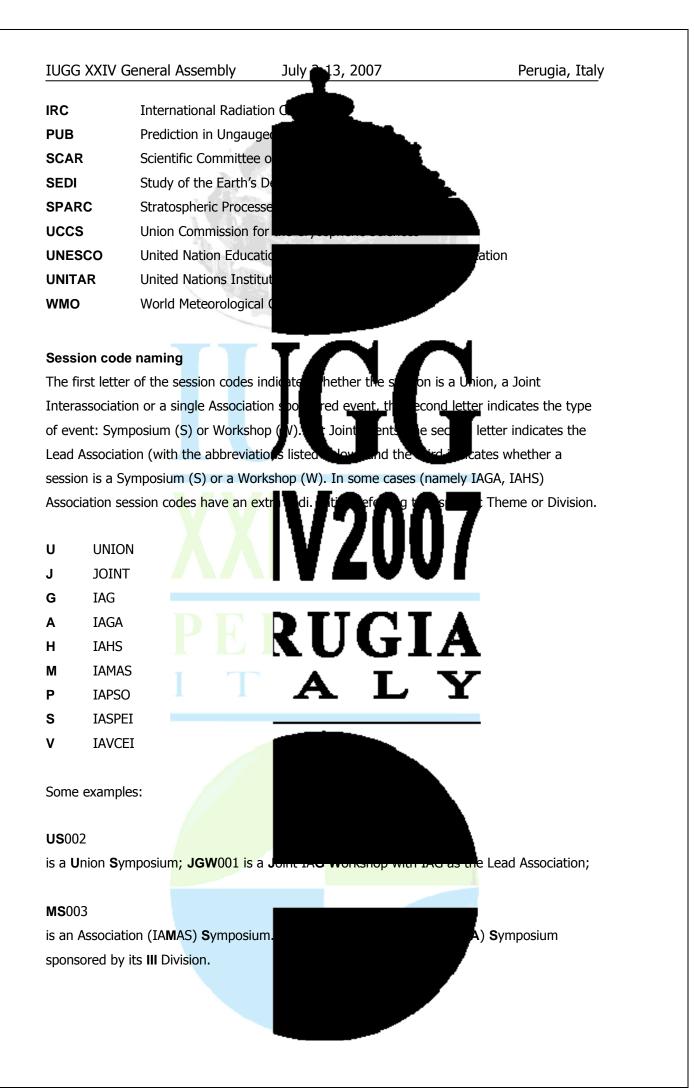
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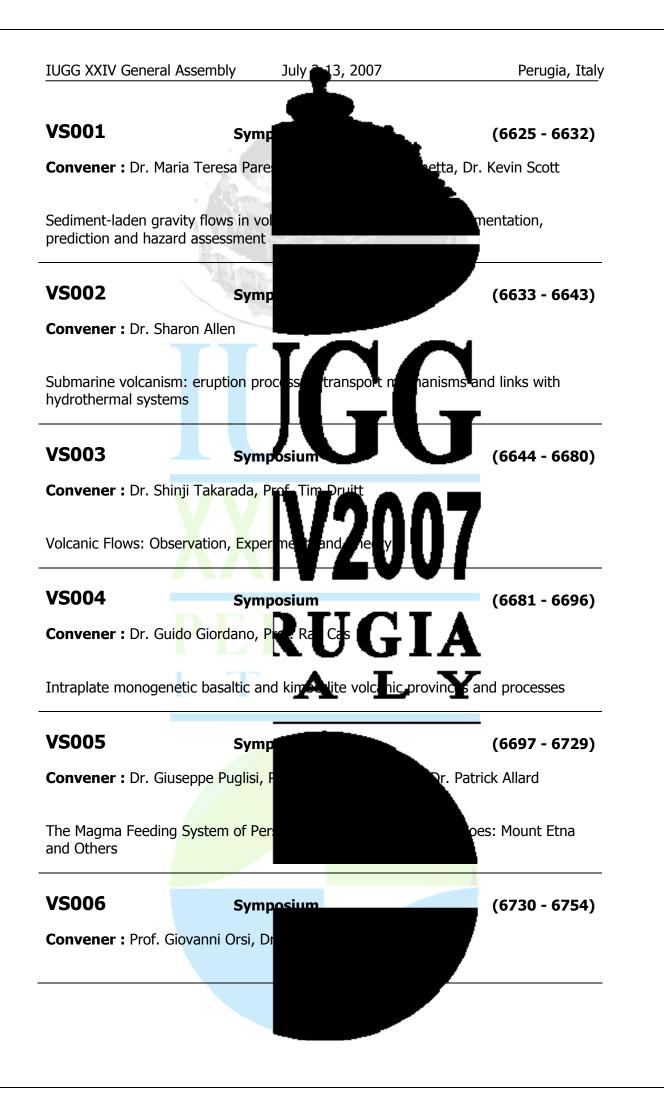
Provincia di Perugia

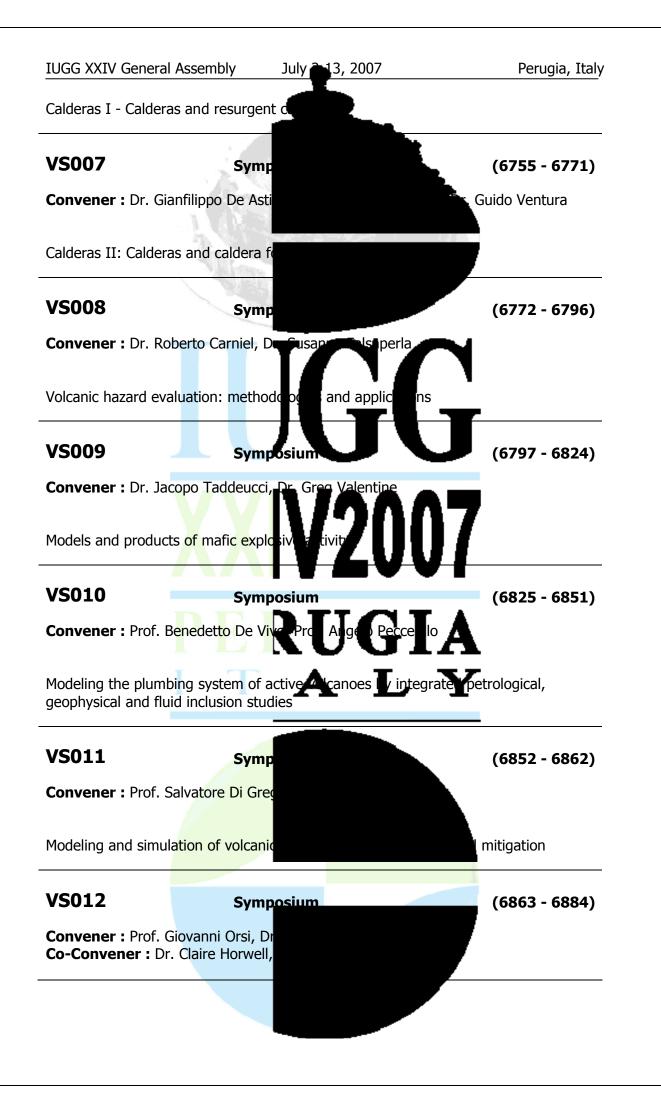
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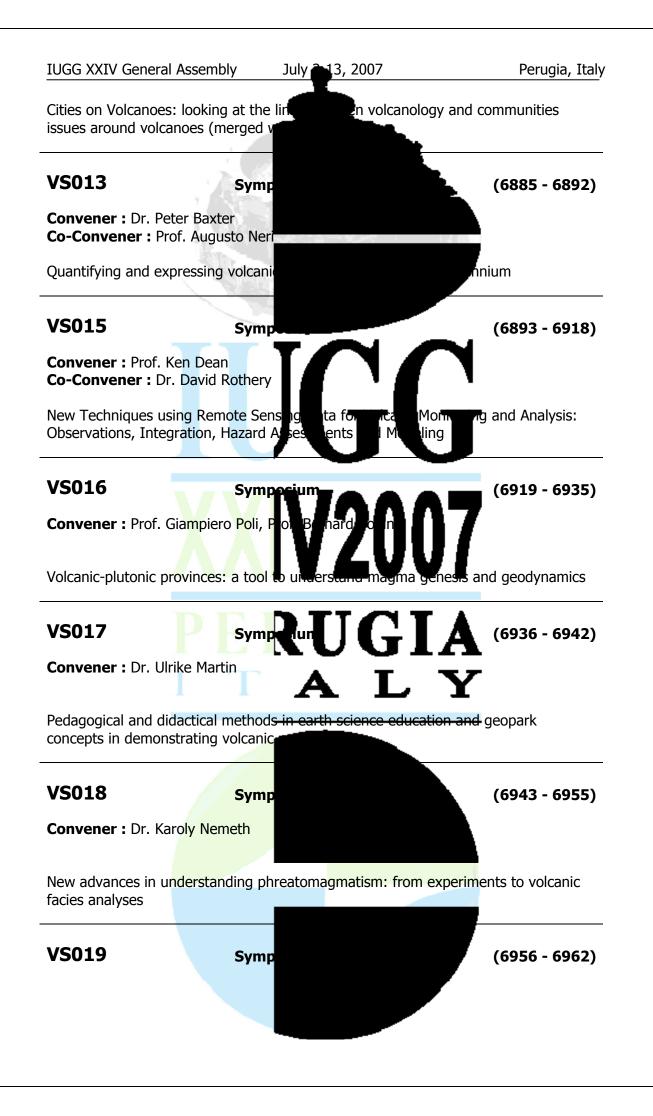
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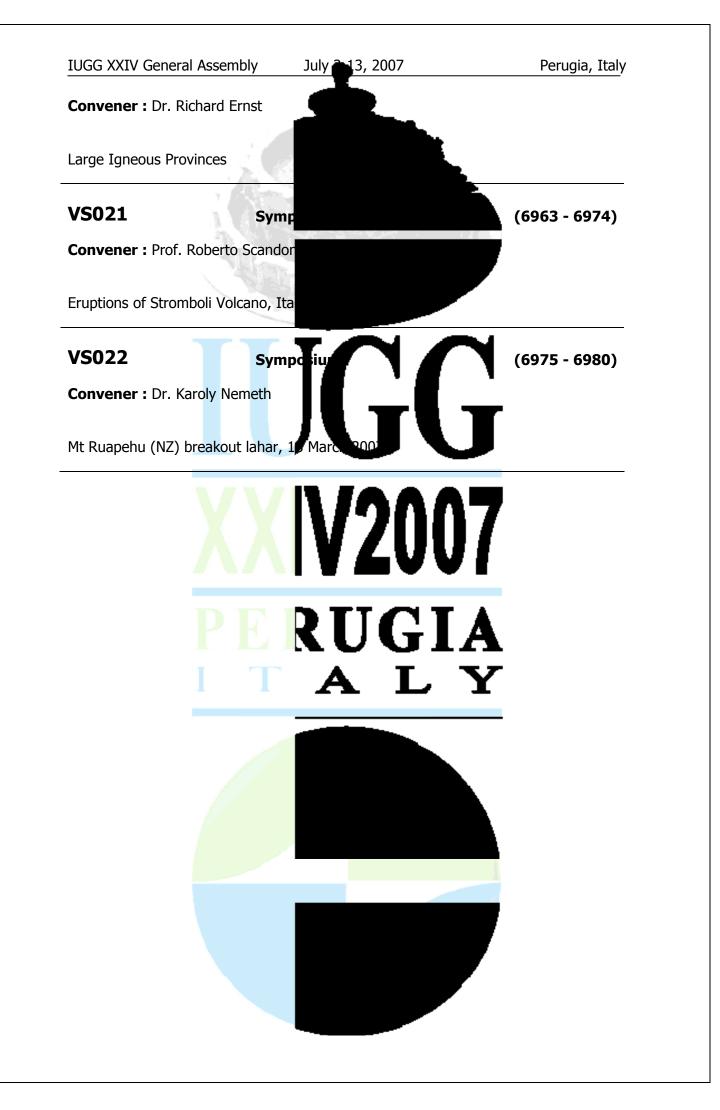
	neral Assembly	July 13, 2007	Perugia, Italy
Abbreviations			
AG	International Associati		
AGA	International Associati		ronomy
AHS	International Associati		
AMAS	International Associati		Sciences
APSO	International Associati		Oceans
ASPEI	International Associati		the Earth's Interior
AVCEI	International Associati		astry of the Earth's Interior
CliC	Climate and Cryospher		
Ev-K2-CNR	Everest-K2 CNR Comm		
GEWEX	Global Energy and Wat	te Expliment	1
HKH-FRIEND	Hindu Kush-Himalayan	Fover egimes from	national Experimental
	and Network Data		T
ABO	International Association	for logic cean	aphy
ACS	International Association	on of Cryospheric Sciences	5
CACGP	International Commiss	propriation Althougheric Shemi	Global Pollution
CASVR	International Commiss	on nu musper -s ill e	etation Relations
CCE	International Commiss	on in continuita Eustra	
CCL	International Commiss	on Clin <b>for V V</b>	<i>,</i> , , , , , , , , , , , , , , , , , ,
CCLAS	International Commiss	ion on the Coupled Land-A	Atmosphere System
ССР	International Commiss	on fiouris and rect	ation
CDM	International Commiss	e on synamic Metroro	gy 🗛
CGW	International Commiss		
CIMOD	International Center fo	r Intranted Mourain De	velophent
СМА		ion on the Middle Atmospl	
CRS	International Celestial	Pofer	
CSIH	International Commiss	dr	ology
CSW	International Commiss		
СТ	International Commiss		
CWQ	International Commiss		
CWRS	International Commiss		
GAC	International Global At	mospheric Chemistry	
GS	International Glaciolog		
LP	International Lithosphe		
NQUA	International Union for		
	International Ocean Ne		













(V) - IAVCEI - International Associa

**VS001** 

## Symposium Sediment-laden gravity flows i sedimentation, prediction and

Convener : Dr. Maria Teresa Pareschi

This session will focus on state of the ar transport and sedimentation mechanisms settings. Methods for prediction and hazar occurrences in vo lcanic settings, where they pres settlements and life. Sediment-laden gravity the most common are intense and/or prolon or glacial ice during an eruption, or seis Volcaniclastic mass flows can be generated centuries later. A common so urce area is volcano. Geological history, morphological hazard as sessment. It is very important to sediment-laden flows in volcanic terrains,

volcaniclastic material. Because rainfall is the most common triggering factor in slope stability, analysis of rainfall intensity and frequency is the flows initiated in this way. Rainfall thresh and to establis h alarm networks in endang hydrologic behavior, including modeling, in assessment. Once flows are triggered, ot sedimentation o f sediment-laden flo ws, inundation. Th ose facto rs include the pro slope profile.

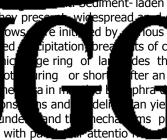
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canology and Chemistry

is concerning the triggering, -laden gravity flows in volcanic comment-laden gravity flows are common



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ontinuing risk to huma n chanisms, among which lous ts of crater lakes, melting of snow des that evolve to debris flows. fter an eruption or some years t o ts downwind from а porta nt constrains for oting th e generation of ms p sion pro cesses in lo ose

most frequently adopted approach to forec asting the debris vil protection purposes ts in our knowledge of lem ents of ha zard ence the transport and potentially subject t 0 the additio n and lo ss o f

sediment and water to the original flow vol ume, respectively, and the effects of longitu dinal change in

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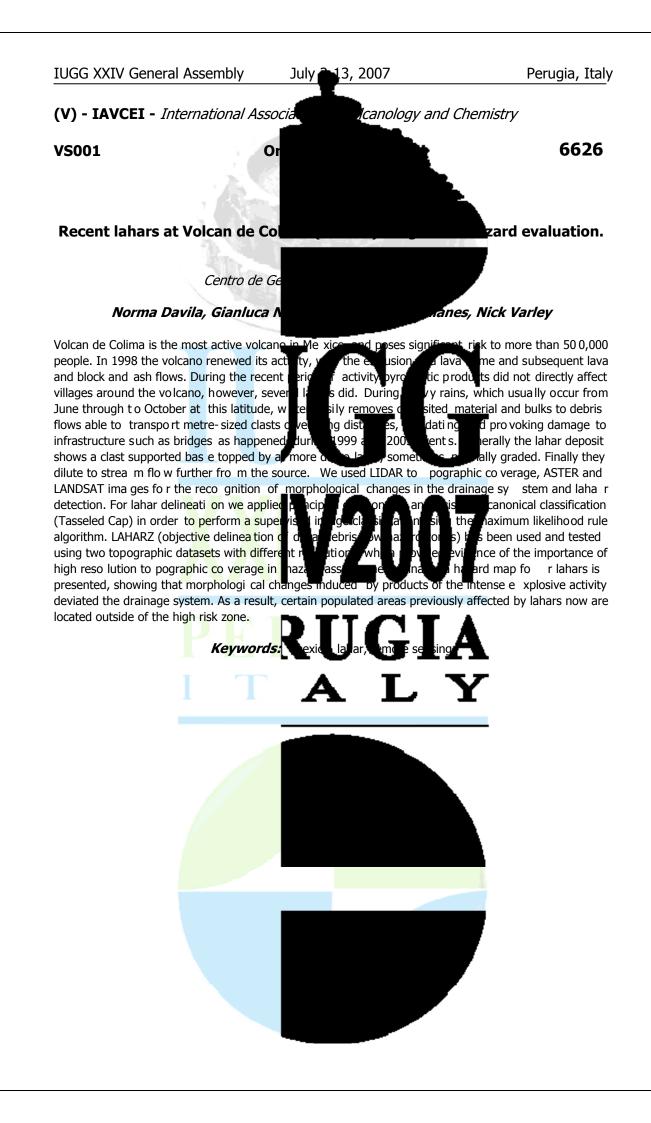
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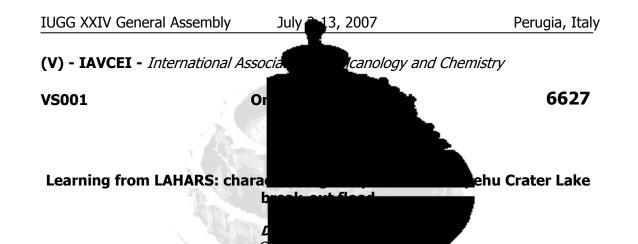
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### Shane J. Cronin, Jon N.

The predicted break-out lahar from the surface opportunity to capture maximum scientific b he to capture as much data as possible fron br energetic for traditional river gauging metho of relief. Ou r proposed sc ientific response complementary components: 1. Instrument tio laboratory flume. Using a diverse range of sensors

conductivity probes, load-cells, pressure transducers, and broadband seismometers we hope to capture time-series data on such key flow parameters as depth, velocity, sediment concentration and pro file, bed aggradation/erosi on, and degree of parameters at a number of key locations v from an initial clear-water flow, to i ts ma attenuation and debulking through downst supplement the planned sensor arrays, ca ptu Collection of t ime-series lah ar samples and visual

Determination of changes in the bed of the Whangaehu River through capture of pre-and post-event, high-resolution topographic and ortho-image using LiDAR te chnology (airborne laser sca first 58 km of flow path and an 18 cm pixel an event will enable changes in ch annel cross-section and profile to be mapped and identification of critical areas of sediment ero sion and deposition. These remove sensing so by differential RTK-GPS ground surveys i n ker loc

with hydraulic data at locations with sensor arrays. 6. State-of-the-art numerical models of lahar flow at Ruapehu will be calibrated against the real newly refined models will be run over the future risk assessments and mitigation combination of resources and skills betwee globally unique data on break -out lahar m mitigation and planning approaches for pro debris flows.

offers a n unp aralleled ete laha event. The challenge is gely unpredictable in timing, to o 5 km of river channel with 2530 m

cole, Hilary K. Mcmillan

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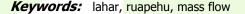
up of a number of ere the bed of a giant , acoustic flow monitors,

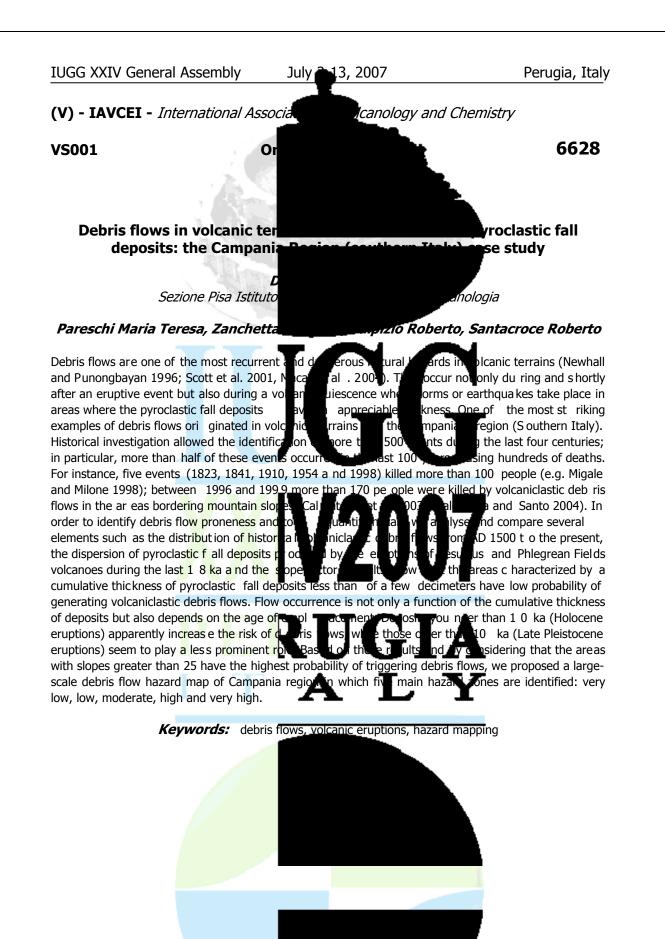
By measuring these

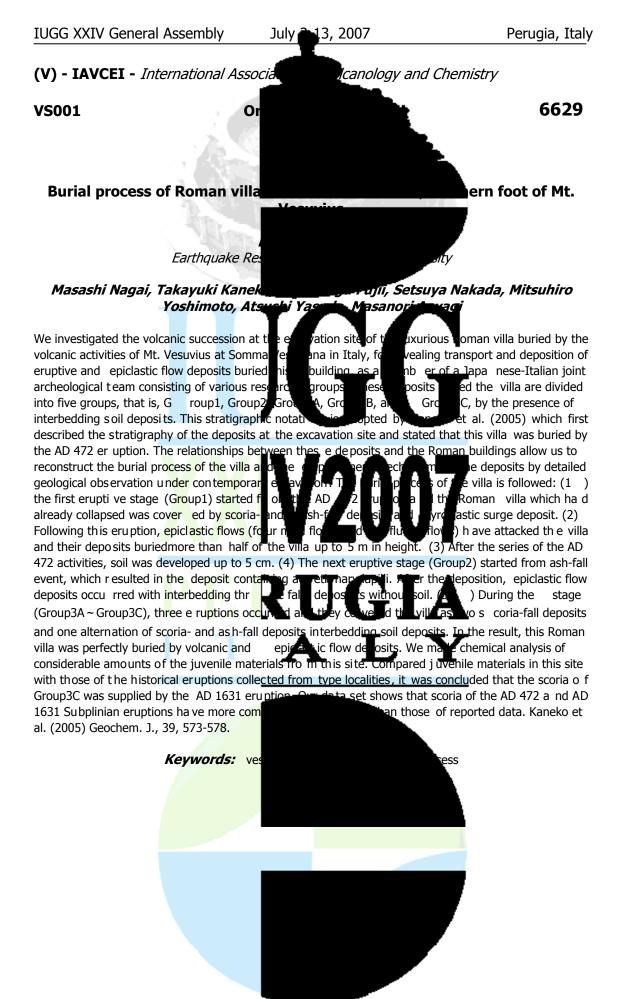
h evolution of the la har nd then its subsequent II and video cameras will eak and lahar e vent. 3. records by observers stationed at key sites. 4.

haselready b een underta ken resolu tion DEM of the ub-m . A repeat survey after air 🖊 not ways will be supplemented event g and survey s of the lahar deposits will allow rec onstruction of flow pa rameters at no n-instrumented sites and cross-referencing from the predicted break-out event. The m the LiDAR surveys to improve er cone volcanoes. Using а

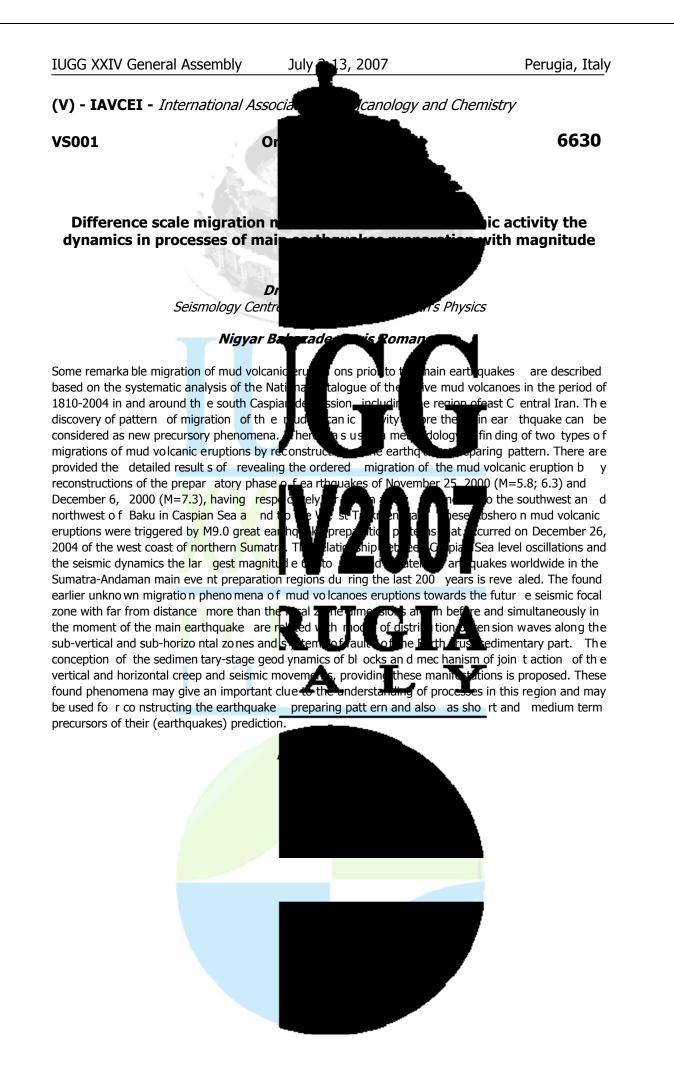
> e the o pportunity to collect will in turn help improve o Icanic and non-volcanic

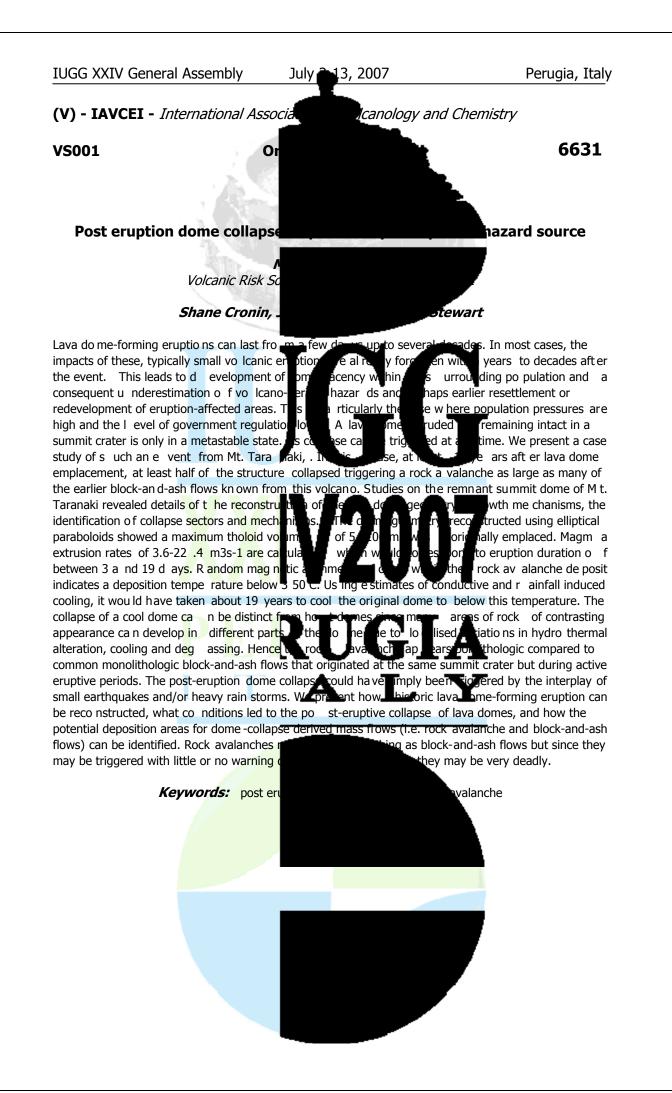


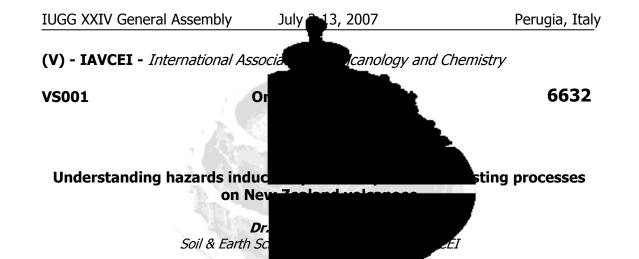




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Rapid mobilisa tion and recurrent redistribut contributing to the morphological evolution plain associations are represented in New from debris avalanches, lahars, and fluvially layers and pal eosols. At T ongariro, in tern tectonic activity; magma intr usion and/or er sudden destab ilisation of a s aturated upp resulted in the catastrophic emplacement of a volca

clay-rich o r co hesive debris flo w. B y contrast , no n-cohesive lahars do minate th e po st-Taupo volcaniclastic landscape at Ruapehu. Here, debris flows and hyperconcentrated flows were and are still generated by (1) expulsion of acidic water and-ice cap by hot pyroclastic material,

eruptions). Key lithological sections from the ring plain suggest however, that large blu the evolution of the volcanic edifice. We ill stra studies conducted around volcanoes of the To ngariro National Park, and in particular Mt Ruapehu,

where flank collapses and la har activity have coexisted for more than 150,000 yr. In addition to field based studies, we have dev eloped a range understand the physical prop erties and flo clay-rich debris flows. Saturated clay+sand+ understanding of the processes that lead identification of potential da mages and mi

ed volcanoge ediments, interbedded with tephra ition<u>s ( e.a</u> ress <u>ure in hydrothermal system;</u> bt are th ht ive in difice. DC of prot 'avalañ

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melting of the snowvered slopes (1995-96 hedial/distal reaches of events occurred early in ion of results from recent

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sequences that include deposits

rounding lowlands. Ring

ed c.60,000 yr ago, the

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ments in order to better exper haraq se the e mplacement of hat ntor tegentent va rying between ture wi at 2.5-15%vol, were used on a 3m long experimental set-up for that purpose. Experimental challenges and preliminary results are discussed. The objective on this dual pproach is **5** fold: (1) in crease our tote but re ntflank estabilisation events on andesitic stratovolcanoes, and (2) refine our current practice used to determine volcaniclastic mass flow hazard in New Zealand ( delineation o f inundated area s; calculatio n of vo lumes a nd frequencies ; This paper is a contributionto the New Zealand Foundation for Research, Science, oject MAUX 0401.

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Keywords: debris avala

g modelling

#### IUGG XXIV General Assembly Perugia, Italy July 13, 2007

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**VS002** 

## Symposium Submarine volcanism: eruption with hydrothermal systems

**Convener :** Dr. Sharon Allen

Recent detailed volcan ological, geophysica on the seafloor h ave provided us with ne eruptions. It is also clear t hat there are

hydrothermal activity. Co mplementary studies on dimensional analysis of the facies architectul sedimentation. We in vite contributions t sedimentology of modern s eafloor vol cand eruptions; (3) transportand depositional character, setting, and temporal and spatial systems; and (4) links between magma res hydrothermal activity.

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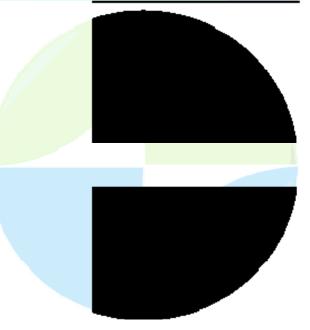
en seafloor volcan oes and ore-forming ucressions provide a 3ocesses involved during itiona ati ons of the volcanology and explosive and effusive seaflo or

sm s and products of submarine

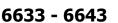
cts of su bmarine eru ptions; (3) seafloor h ydrothermal uptive processes, a nd

studies of modern volcanoes





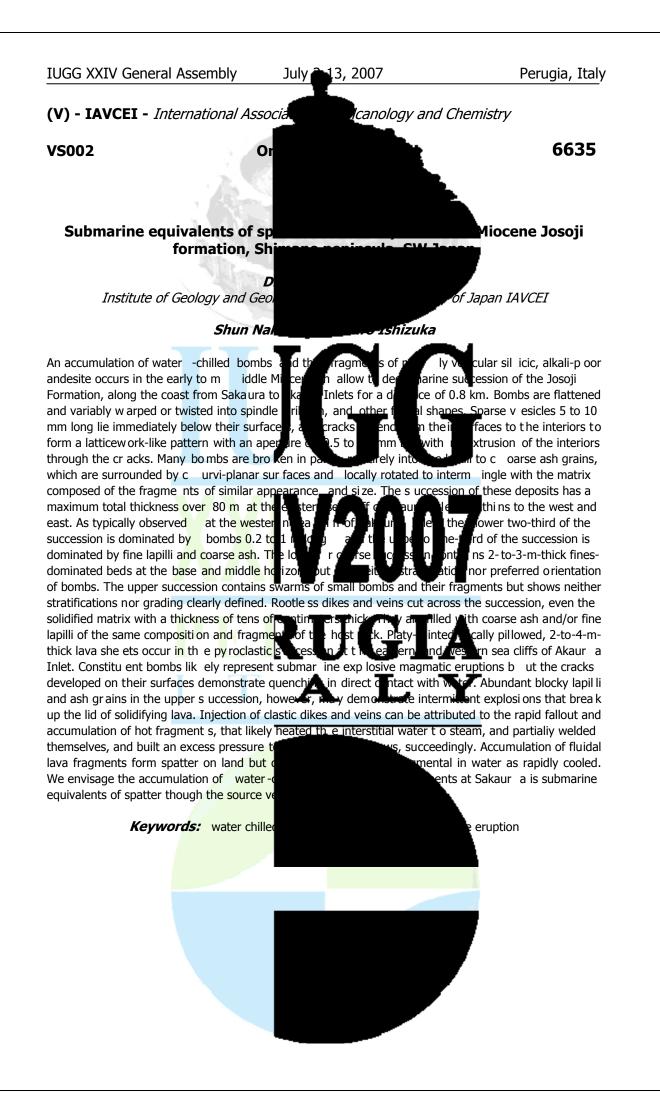
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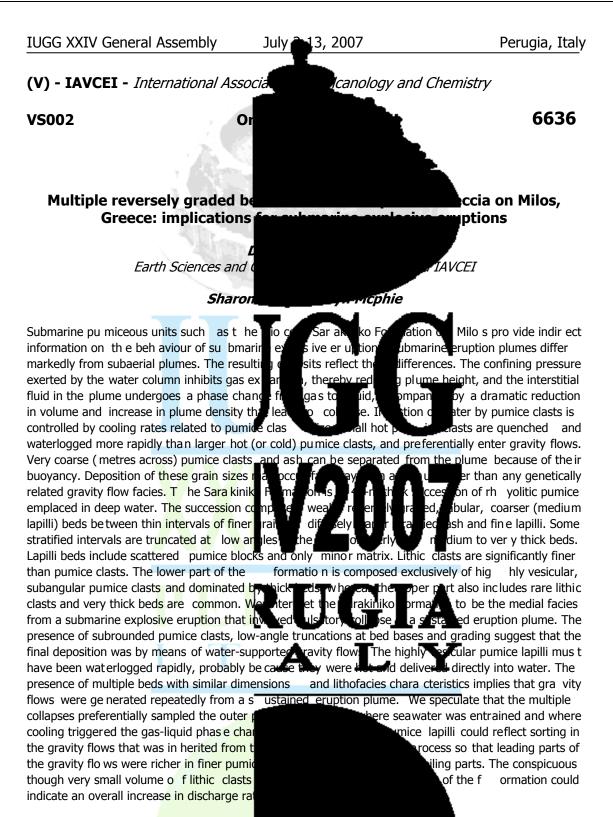


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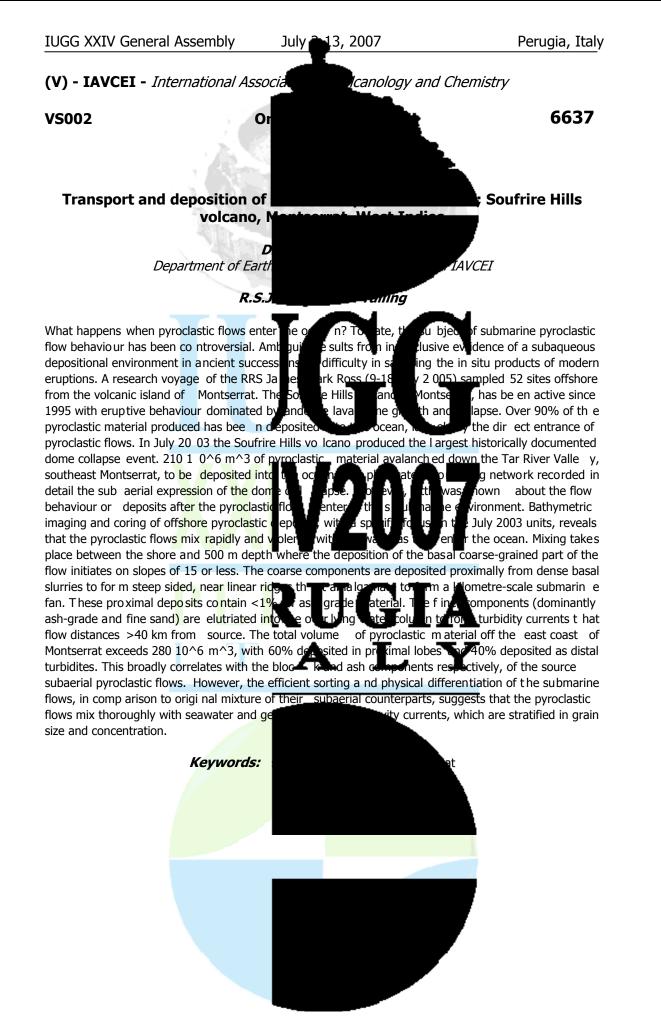




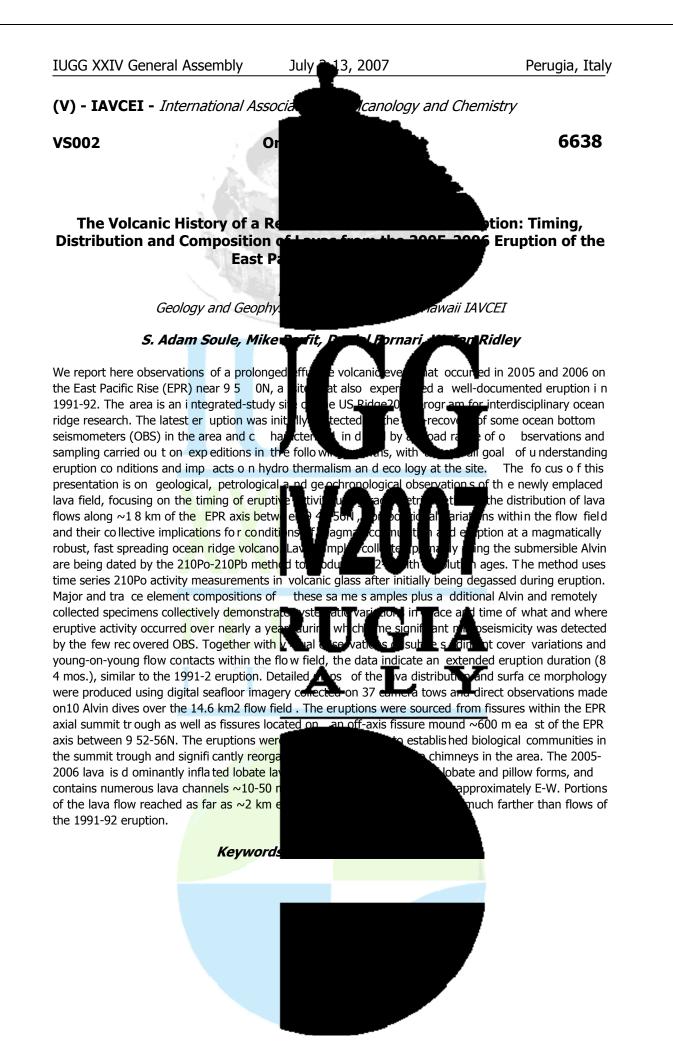


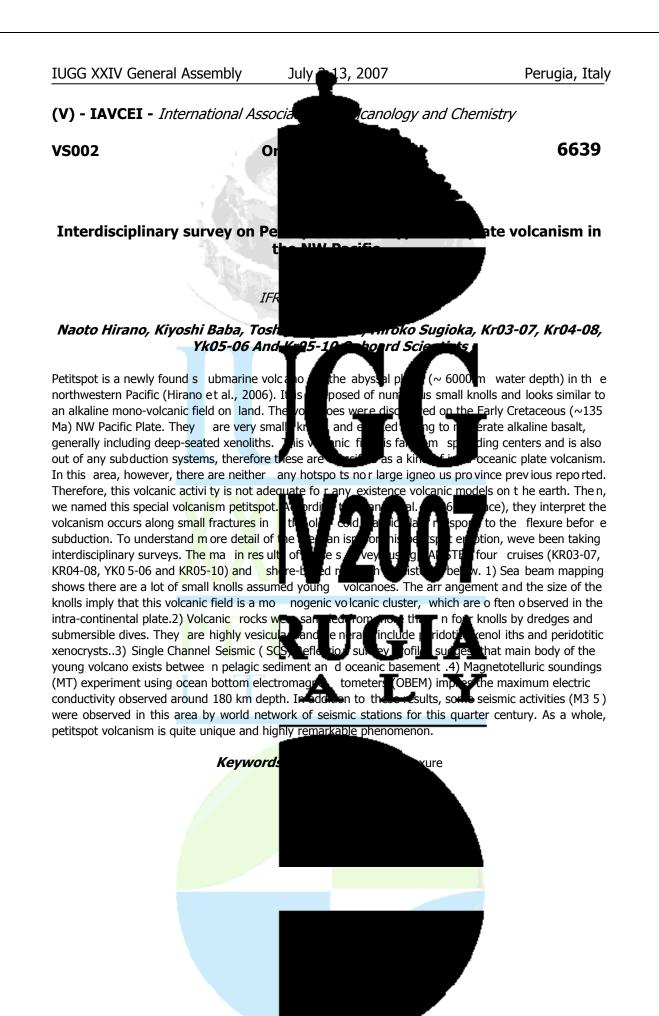


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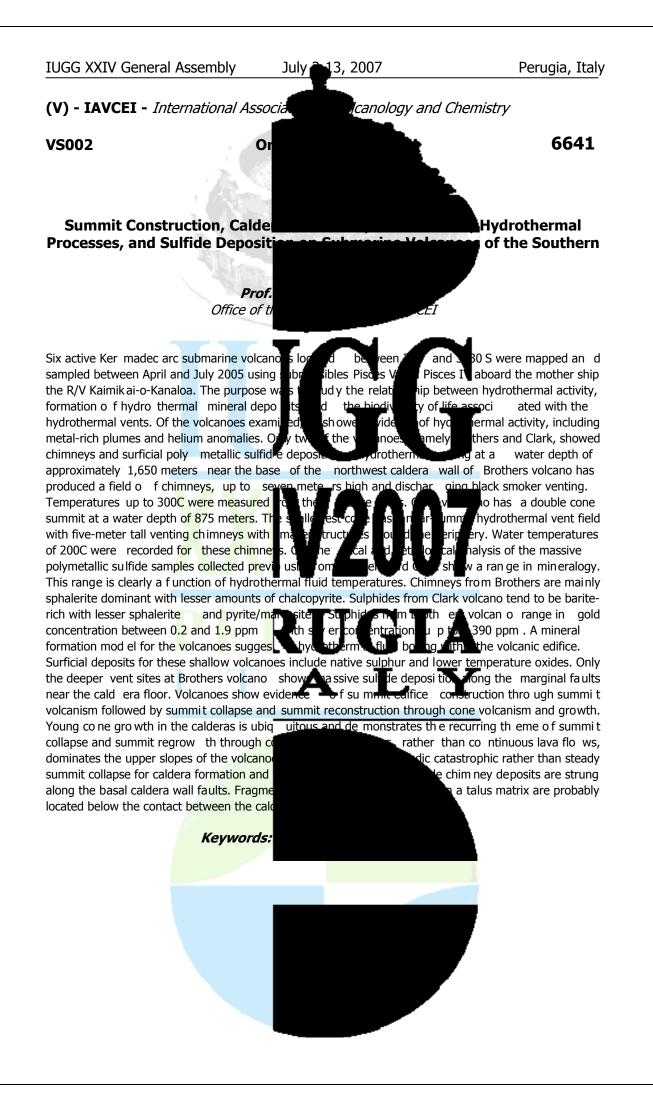


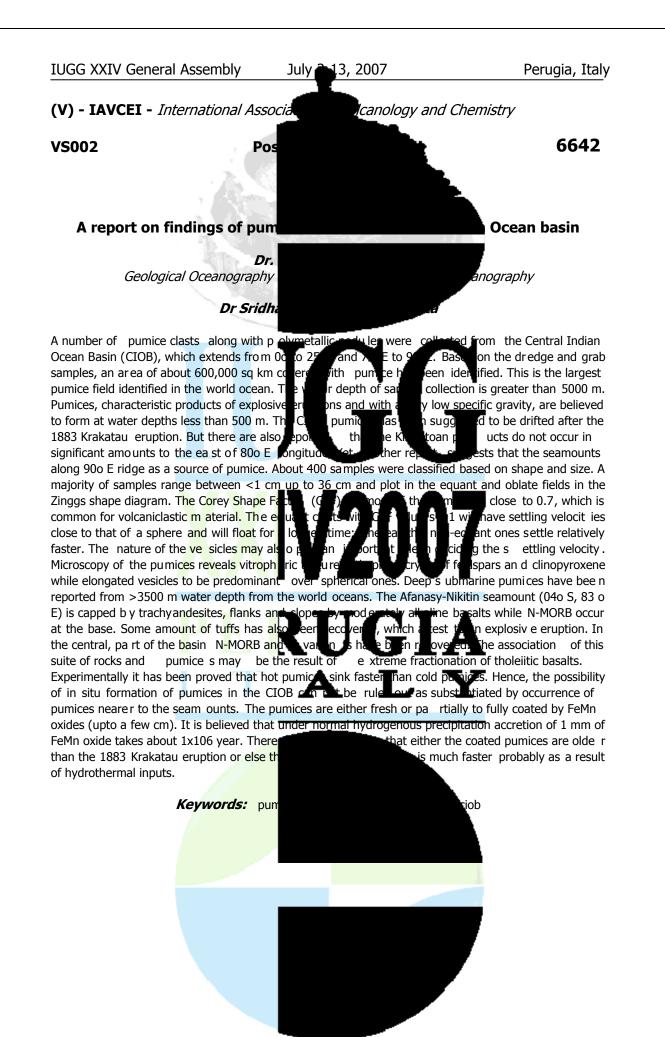
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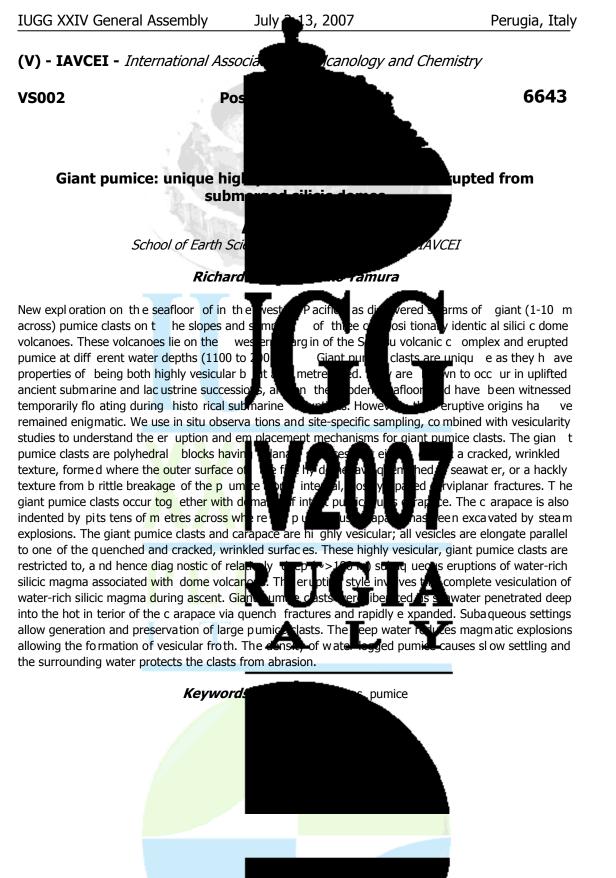




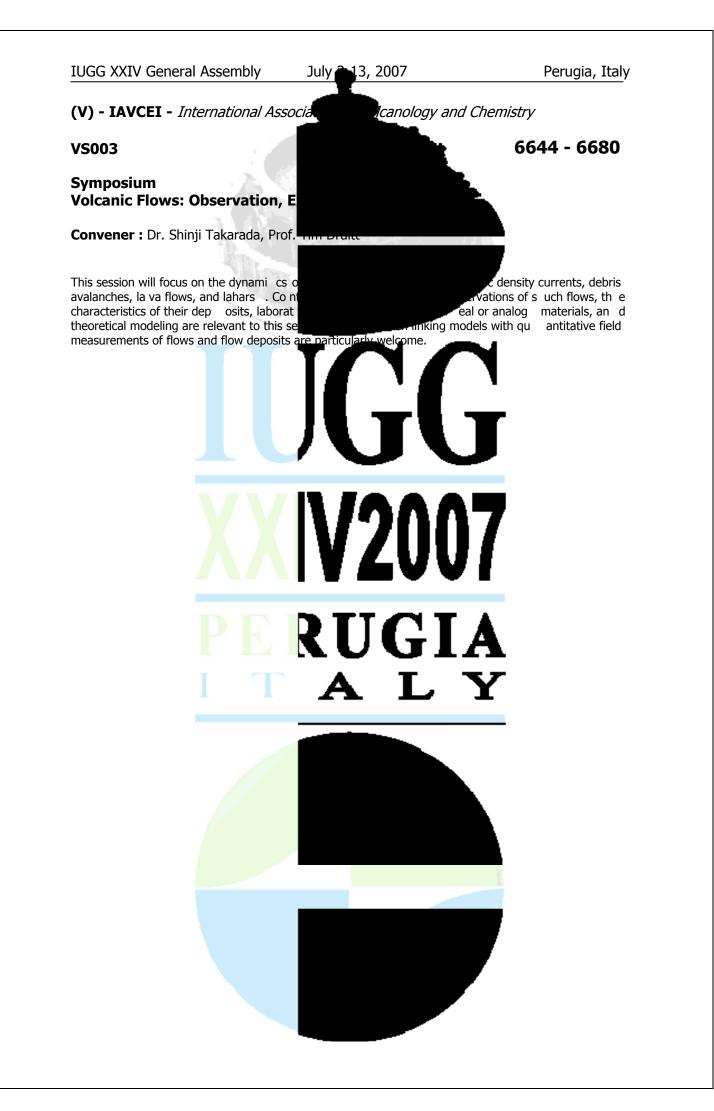


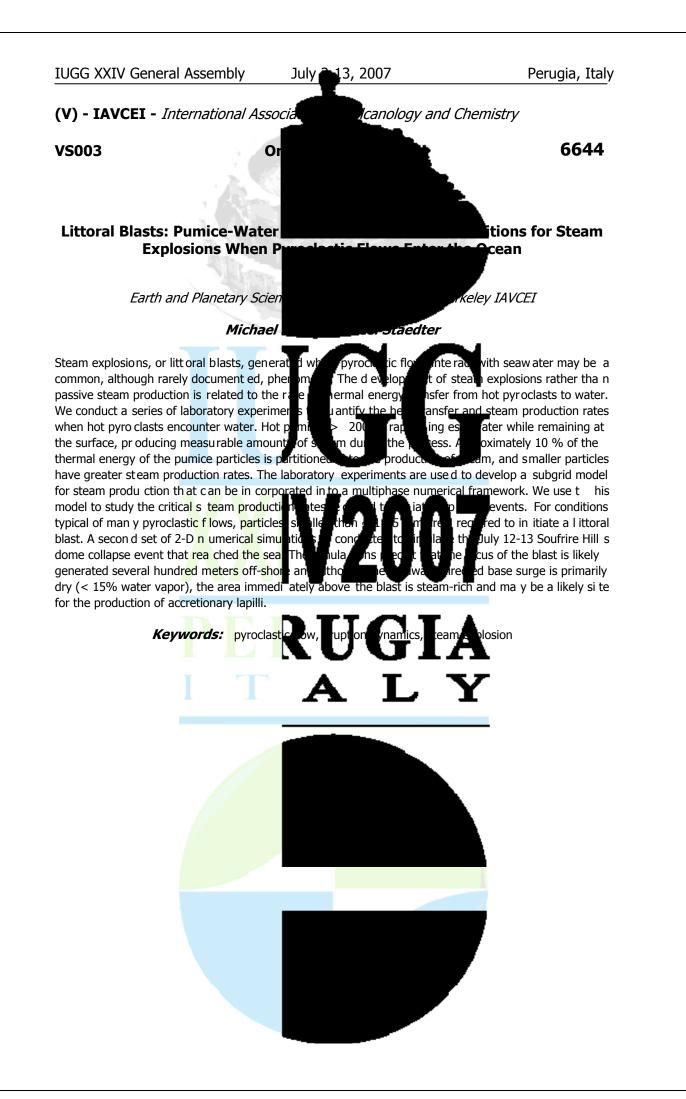


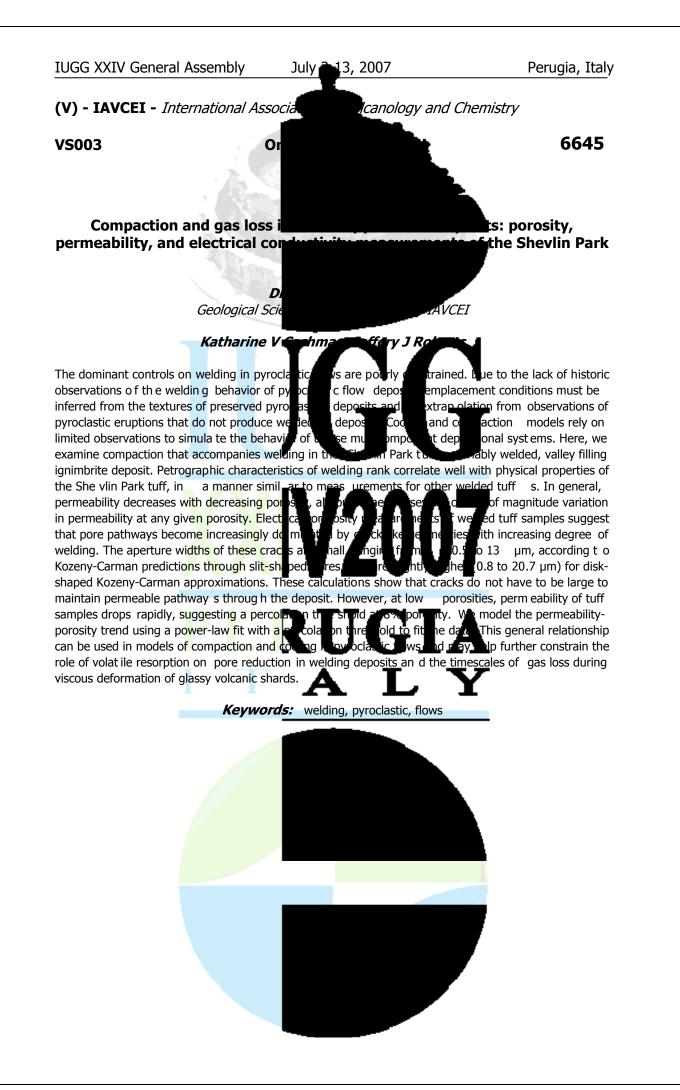


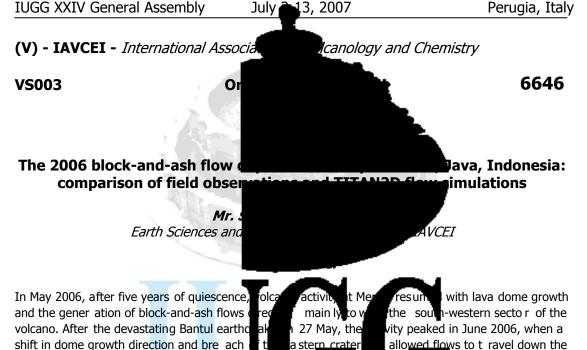












shift in dome growth direction and bre ach volcano's southern and south-eastern flanks century. On 14 J une, the largest block-and the Gendol river valley, causing two fataliti

Associated de posits form about ten < 8 0 m wide overlapping lobes, which represent a record of successive flows generated during and after the major event on 14 June 2006. Both, single pulse (post-14 June events) and multiple-pulse pyrocla are recognised and three ty pes of depos deposits in the main Gendol river valley, spread laterally onto ridges and interfluves valleys and (3) overlying, dilute ash cloud s urae

compared wit h the valley-confined deposite contents within surface part icle assembla

multiple-collapse events of the unstable east

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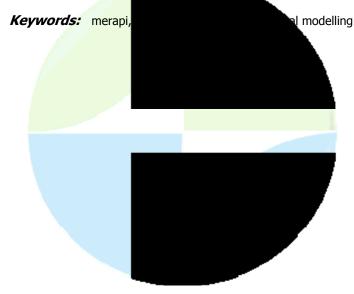
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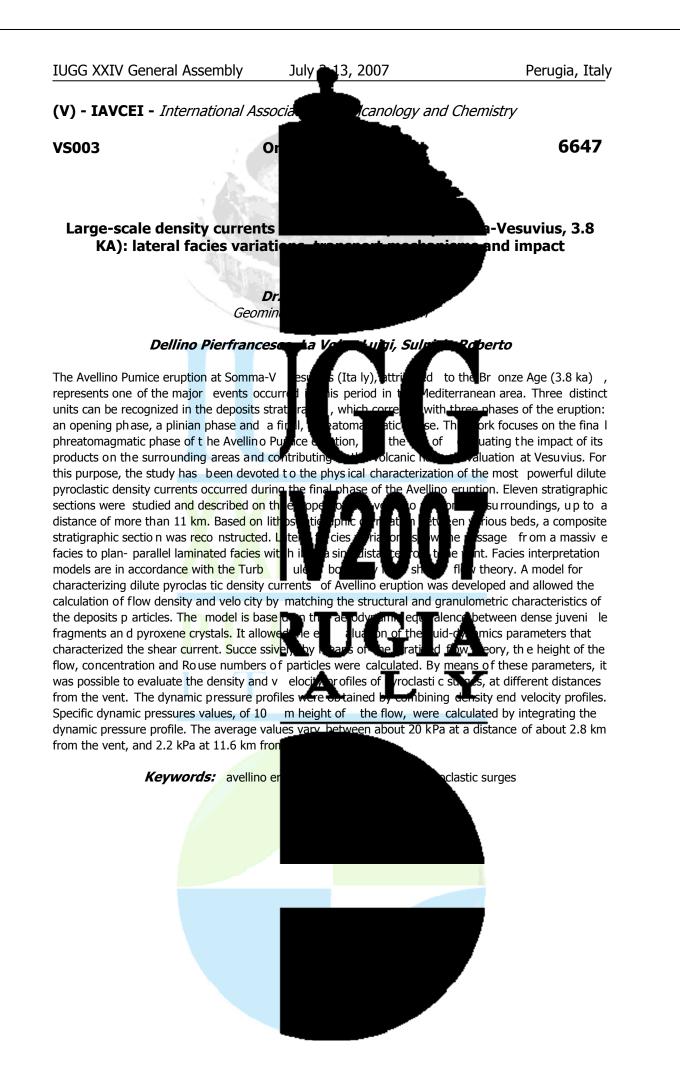
allowed flows to t ravel down the ic flows for more than a pyrod es o f km fro m the summit in e village of Kaliadem.

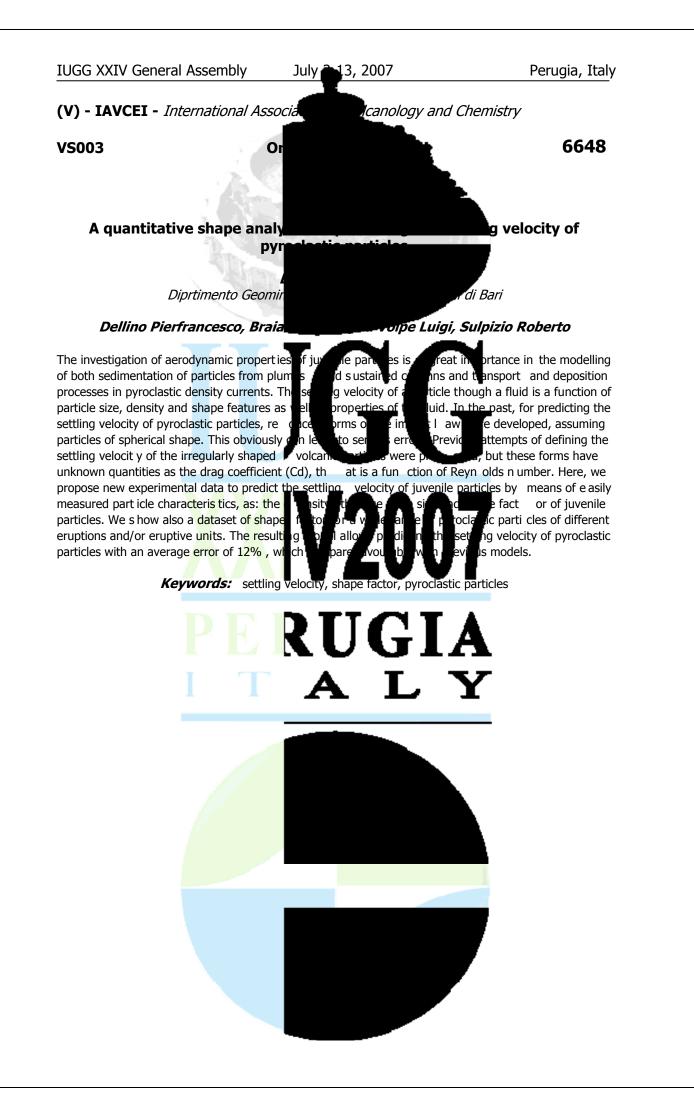
e collapses on 14 June fined ba sal avala nche of the b asal avalanche nto the surrounding river face particle assemblage analyses on different lobe deposits have revealed variations in the abundances of the main lithological

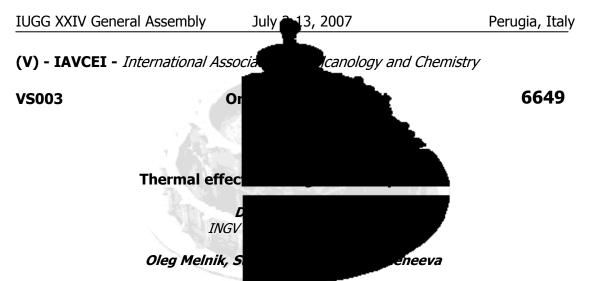
components from proximal to distal reaches. The overbank deposits are enriched in scoriaceous clasts ly altered and oxid ised clas t or flows produced by rks occur on all sides of npa it n

large dense blocks within the deposits and signif icant erosional features were observed at valley margins at distances > 5 km from the summit. T te indicate that the 200 block-and-ash flows can be frictional contacts during transport. considered as unsteady, rapidly agitated granular f s with e Numerical simulations of these flows using the TITAN2D model developed at the Uni versity of Buffalo, USA, stay within the extent of the mapped deposits. Insights gained from these preliminary results are considered invaluable in guiding future haz clastic flows at Merapi.









Viscosity of silicate melts is strongly affected momentum equations in mag ma flows are stratification due to temperature gradients in energy equations for magm a flow inside surrounding host rocks imposing local far fi we neglected magma solidification and melt dissipation and heat loss to the conduit wall distinguish three regimes: i) a conductive-neat-los

iii) a viscous-h eating-dominated regime. When viscous dissipation effects are negligib le heat loss t o conduit walls is responsible of a significant local increase in magma viscosity near the boundary and as a consequence of an increase in the condu of the conduit. Maximum velocity is 5 to 6 profile the ratio between the maximum an are dominant, there is a local temperature viscosity which strongly affect both temper type, with a uniform temperature distribution at the

d by tem ure variation ouple Fond noro a flows, 🐙 so m ical conduit dition<u>s for ro</u>c pst ro ion res a pi d reaii

effects of local viscosity lo stu humerically mass, momentum and he at conduction equation in the mperature. For sak e of simplicity show that both viscous a dy namics allowing to termediate regime, and

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at implies that energy an d

zed in the central part euille parabolic velocity scous dissipation effects sequent decrease of the evolve from a Poiseuille to a plug type with a hotte r layer near the

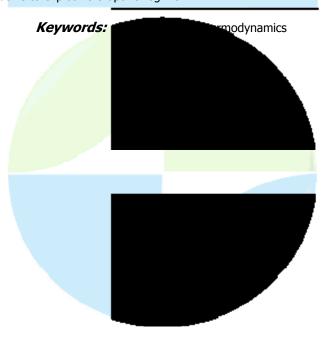
walls. The maximum velocity is just slightly larger than the average in this case. Fully 2D computational results were used for better parameterising e used in simplified 1D models. At the end-me mber cases some analytical boseq alcu lations h ave direct ns are D eno n that can be linked to implications for interpretation of magmatic ndu floy s. d I temperature variations include re-melting a nd erosion of wall rocks due to temperature increase and instabilities, textural differentiation due to s train calization, escription an d/or formation cal crystal due to thermal layering. Finally a possible role ous diss tion effects in leading to a transition between extrusive/effusive to explosive eruptive regime.

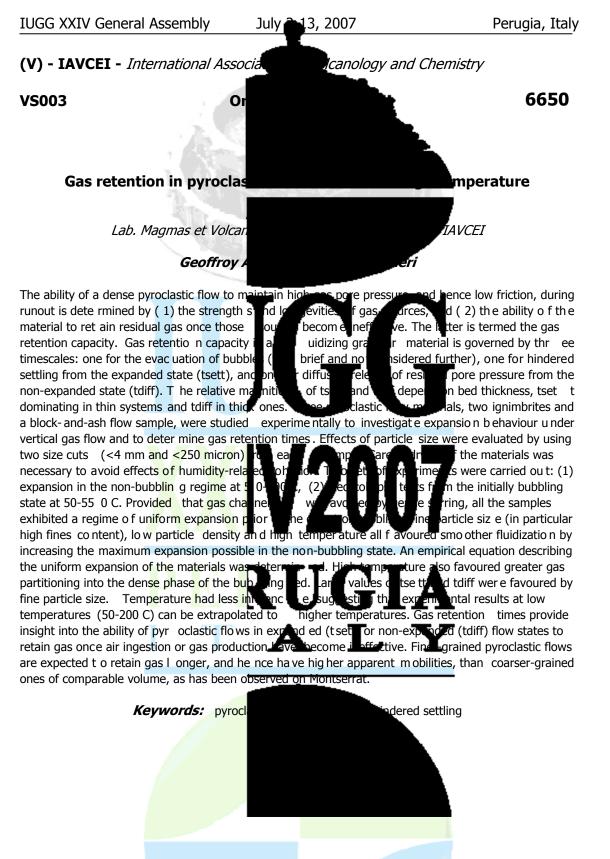
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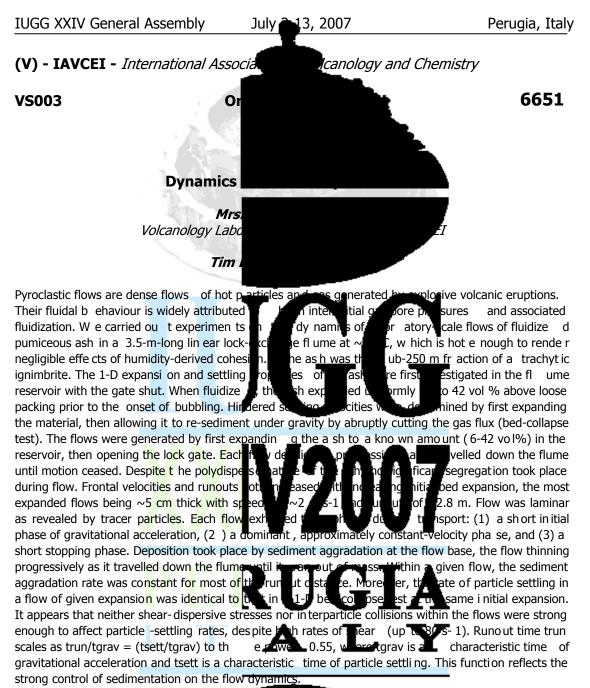
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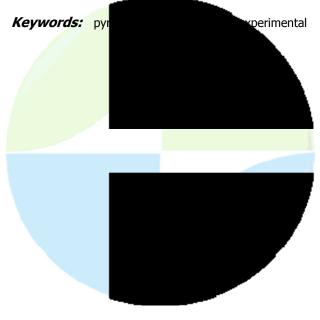
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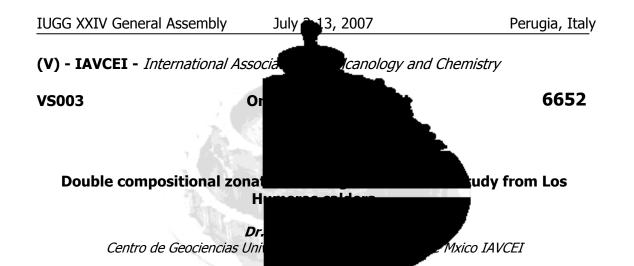
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Vertical compositional zo ning is a relative changing discharge with time. It shows an u evolved silicic pumice at the base of the igni pattern is commonly inferred to record D stratified magma-chamber, initially tapping I chamber and then progressively deeper, de this pattern may be common in multiple floy units showing this feature. One is the Zaragoza ign

in central Mexico approximately 60 100 ky interesting double (normal-to-reverse) pattern of chemical zoning, marked by differing compositions of both the juvenile and the accidental comp up via a mixed zone into a central andesiti upper rhyodaci tic (67.6-69 % SiO2) zone populations. We infer that py roclastic fou granular fluid-based pyroclastic de nsity cu current then gradually changed, first to an desite and then back to rhyodacite. Inverse grading at the

to less-ev blve e eruptive w nse\_more\_s ore ma bar ther r ago. The layer exhibits in a section of 1 6 m thick an

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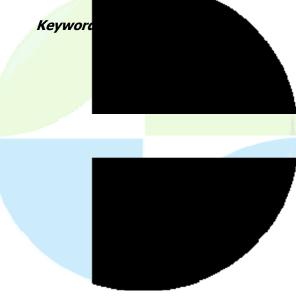
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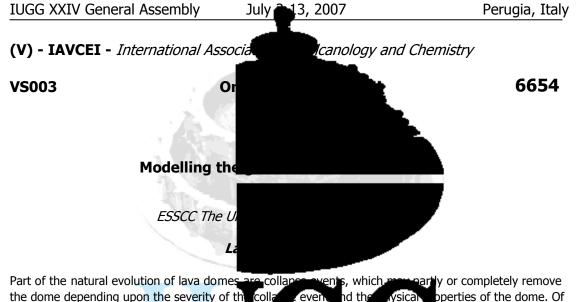
and typically records nponents from relatively enile` ore mafic pumice at the top. This rawal from the top of a density uppermost parts of the magm a the c ber with time. Although few e ples of individual flow om Los Potreros caldera

> .% SiO2) zone grades turn, grades up into an variatio ns in lithic clast ce clasts to a sustained ce clasts supplied to the

base of the massive layer may reflect initial waxing flow competence. The pumice concentration at the ted dering waning stages of the top of the massive layer is entirely rhyoda current, when the supply of a ndesitic pumie tur rhyodacitic compos ition may have been the result of eru ption-con dif oll ar se L os Potreros caldera, it n' ati ' dι ing of hydrothermally alt ered lithic b locks, and/or a marked in the ignimbrite by a widespread influx decrease in draw-up depth from a comp ositional stratified m gma chamb The eruptive mass flux waned. The m assive layer of ignimb rite thins local toless -2 m, yeart still shows the double zonation. Correlation of the zoning suggests that the thin massive layer is stratigraphically condensed, and aggraded relatively slowly during the same time-inter val as did much thicker (  $\leq$  50 m) massive layer.







the dome depending upon the severity of the colla fundamental importance to volcanology is collapse. Collapse events commonly result in and ash, which travel at hurricane speeds ar hazardous events are those that remove suf loss) of magma either within active domes, Thermally, a lava dome can be divided in

surface with temperatures generally less than 100 deg C. As material is extruded into the dome interior, the solidified surface is forced to break apart genera ting debris which is deposited at the flow front and forms a talus that eventually enshrouds the commonly follow a trend of initial talus col This suggests t hat the tal us is critical to the computationally model the growth and e neglected component of dome growth. We of the lava do me. The talus is mode llead to develo p tr om lava solidification via vola tile loss which

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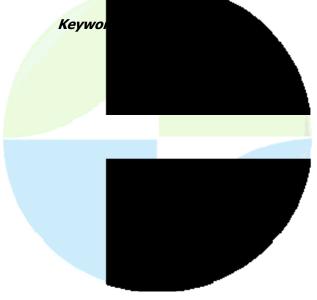
pperties of the dome. Of d conditions are for a dome to astic flows, avalanches of hot rock i wh<u>en ente</u>ring the sea. The most id vesic ulation (volati le els in a volcanic edifice . re) and a cooler outer

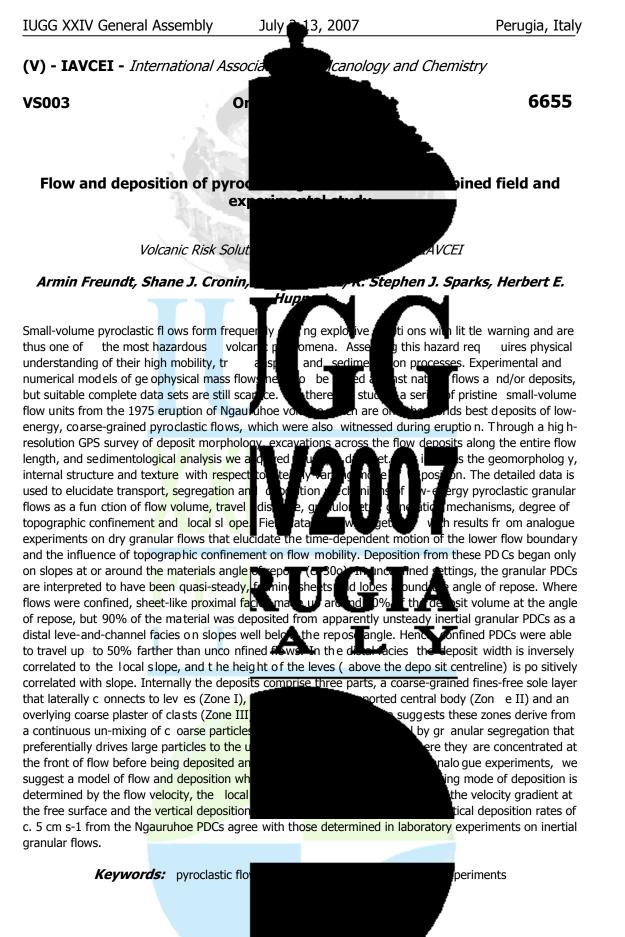
dome collapse events en the core is exposed. va and for stability. We the talus, a c ommonly EM) to model the growth promotes crystallization as observed in inter mediate composition lava flows. We model the dome core

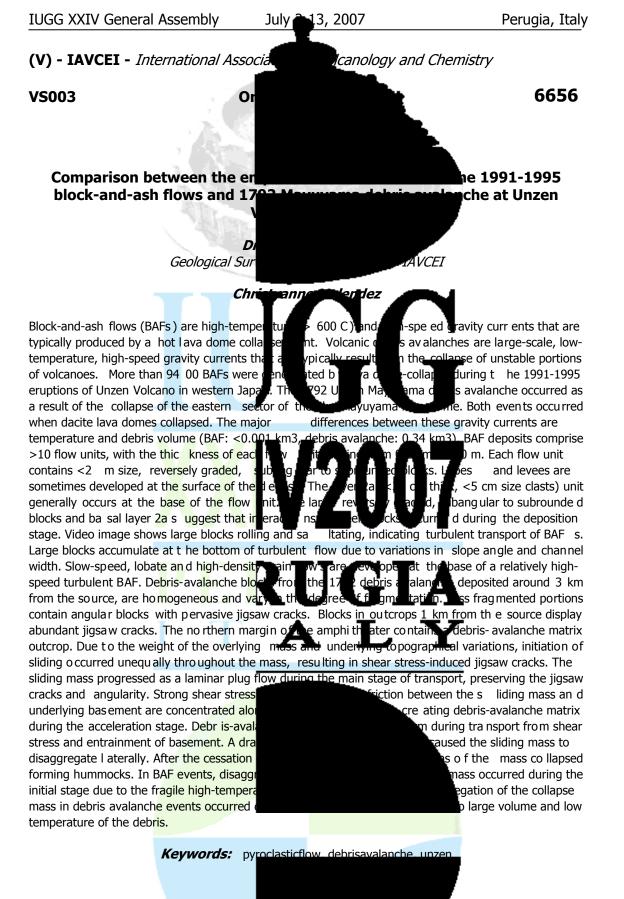
lava as a Newto nian fluid with a constant use the lava cor e will maintain higher temperatures due to the low thermal e talı modelled as a Drucker-Prager material with the interface between þ be olidus transition. The ken LLIG free-surface and core/talus interface are modelled using the level-set method, a te chnique to trace sommetric and provides interfaces and flow fronts with out distort ing the mesh . Ou mode I is a lur le of coll se required before the information on the extent and growth of the and the collapse scar will reach the dome core. The extrusion rate of the lava, solidus transition and the physical properties of the talus gover n the extent of the <u>dome and talus/core interface</u>. Model results show how talus is predominantly generated at the top d deposited at the toe of the flow.

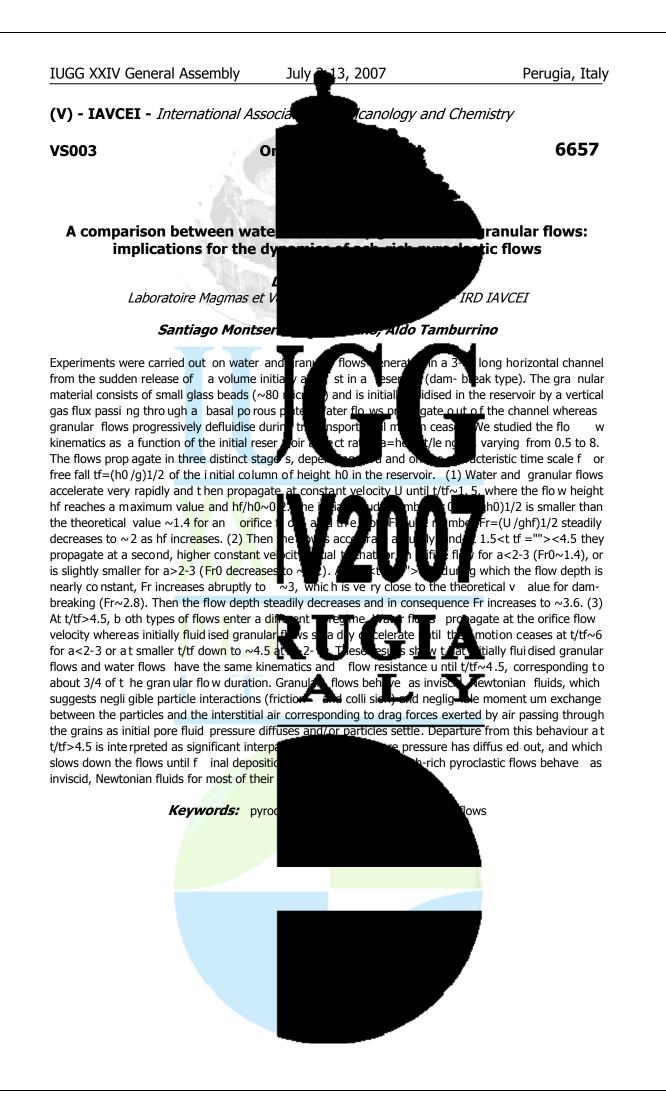
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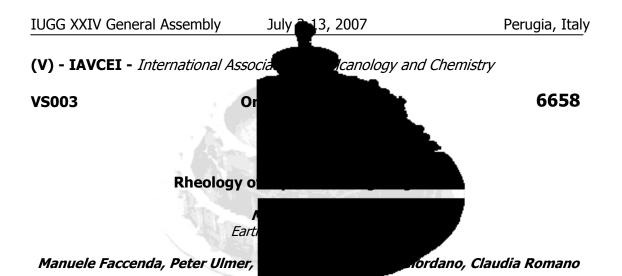
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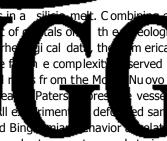






Magmas are s uspension of gas and crysta is in a simulations we focus on the physic al effe suspensions. While the exp eriments provide understand the physical processes responsib bearing hydrous rhyolitic sam ples and natur Italy) have been deformed in an internally crystallinity (50-80 vol. %) and strain rates. they exhibit a tendency to non-Newtonian and Bind

account for the observed shear thinning consequently, the no n-Newtonian effects r numerical simulations have been designed identical stress and strain rate conditions numerical simulations, reveal the generation to the direction of flow. The melt enrichment induces localization of the strain in these regions and leads



experiments and numerical beh avior of magmatic m erical mulations are aimed to served in the experiments.Crystal-Nuovo eruption (Phlegrean Fields, overing a wide range of s lack yield strength but high (>10-4 s-1) strain

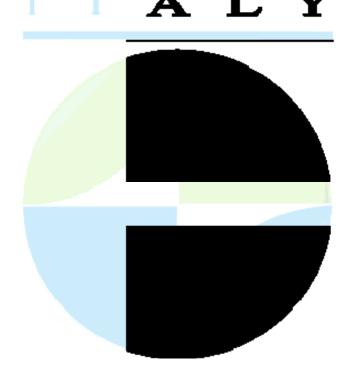
rates. The experiments have been performed under temperature and strain rate conditions at which the melt phase behaves Newtonian. Moreover, in the strain rate range applied, viscous heating is unable to reasing strain rate); Is in the magma. The

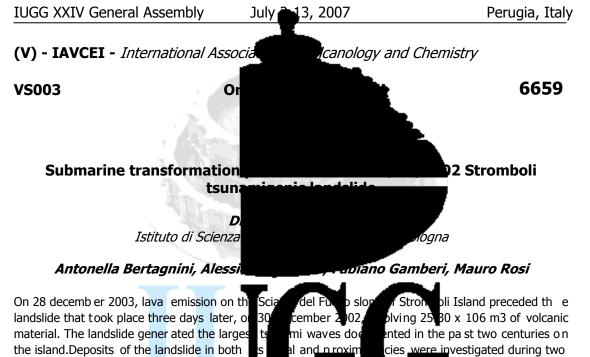
ments and co nsequently from experiments and t around 30 with respect

to a geometrical redistribution of melt and particles that c ould account for the rheological complexity observed.

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depo indicate that the proximal, c oarse-grained, nds ver a kter water depths between 1600 and 2 000 m at a dist to 8 kn consists of several discrete, mainly chaotic assemblages of fresh cobble-sized scoriae and lav a flow

the S

volcaniclastic s and often arranged in ripple surrounding seafloor in being completely d sediment textures indicate that the proxim debris flow processes. Distally, box coring del Fuoco, in a site located on the right sic capped by a 2-3 cm-thick layer. In the s

high resolution marine s urveys conducted o

clasts within a coarse sand matrix. Down-slone and laterally, the coarse-grained deposit grades to black a erfe e of าค

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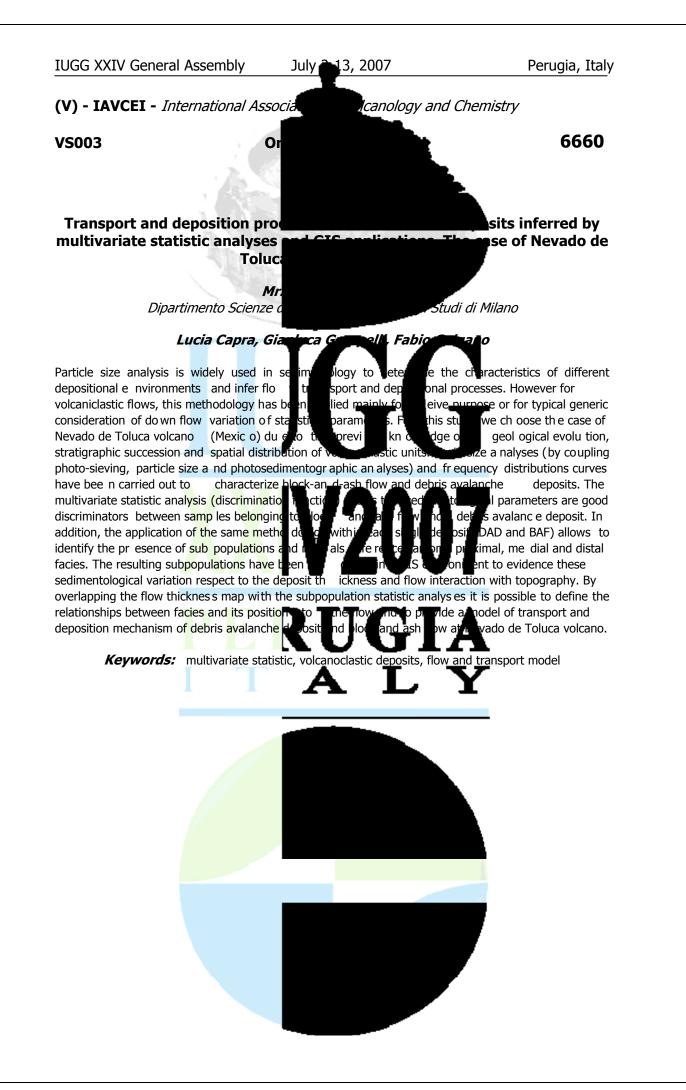
and July 2005. Results W elongated area at shoreline. The depos it

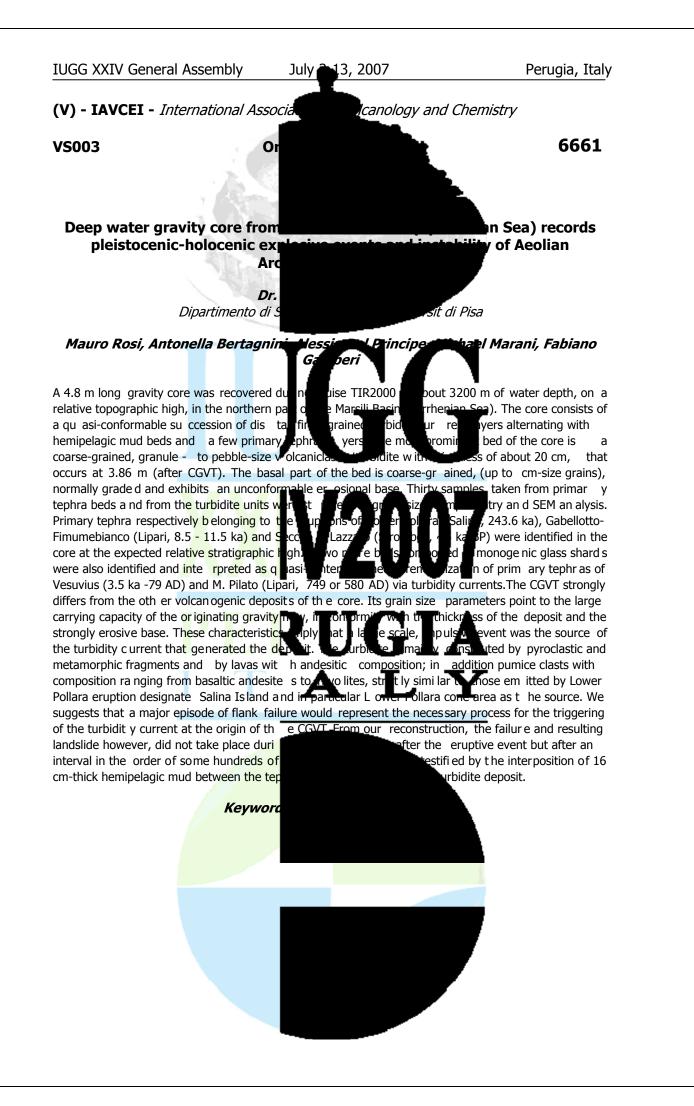
cont rasts with the Initial interpretation of low-coherence, granular he shoreline of the Sciara ed a sedi ment sequence ame sa mpling site, the top most la yer was not present in

September 2002 when a previous gravity core was al so collected. Grain size and componentry of the and glass compositions matching layer indicate that it is a volcanoclastic sand the character of the material of the 2002 la nd c ositional features of the lical layer are consi stent with an origin from a ent, og etic to the debris flow ani c ť pid ' cu generated by the 2002 land slide. Our study stresse s the potential of the marine e nvironment in recording landslide events b oth as the coa rser mined products of debris and a s the cogenetic turbidites that register the distal deposition of edian 📕 arained derived oducts.

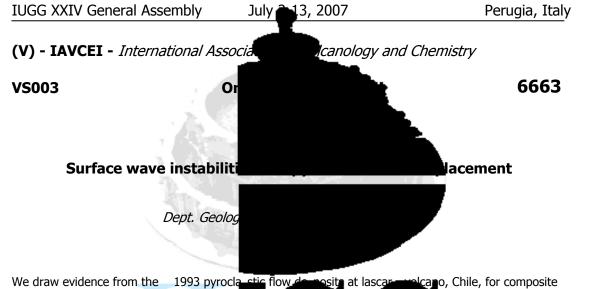
Keywords: stromboli, volcanic landslide, submarine debris flow











pyroclastic deposit fans having been emplayed by their flow sur faces. The instabilities play unsteadiness and the development of succes discrete onlapping units now identifiable with by the unique perspective provided by grou remarkable potential in unraveling transport instabilities are already known from other types of

versus flow-derived current unsteadiness, currents. This work will use the convincin the role surface wave instabilities may have advocate using geophysical tools, such as as detailed field relationships.

anula ows t integratiole Ρđ ves, which in deposit lobes. et rati ād esses eop d exp

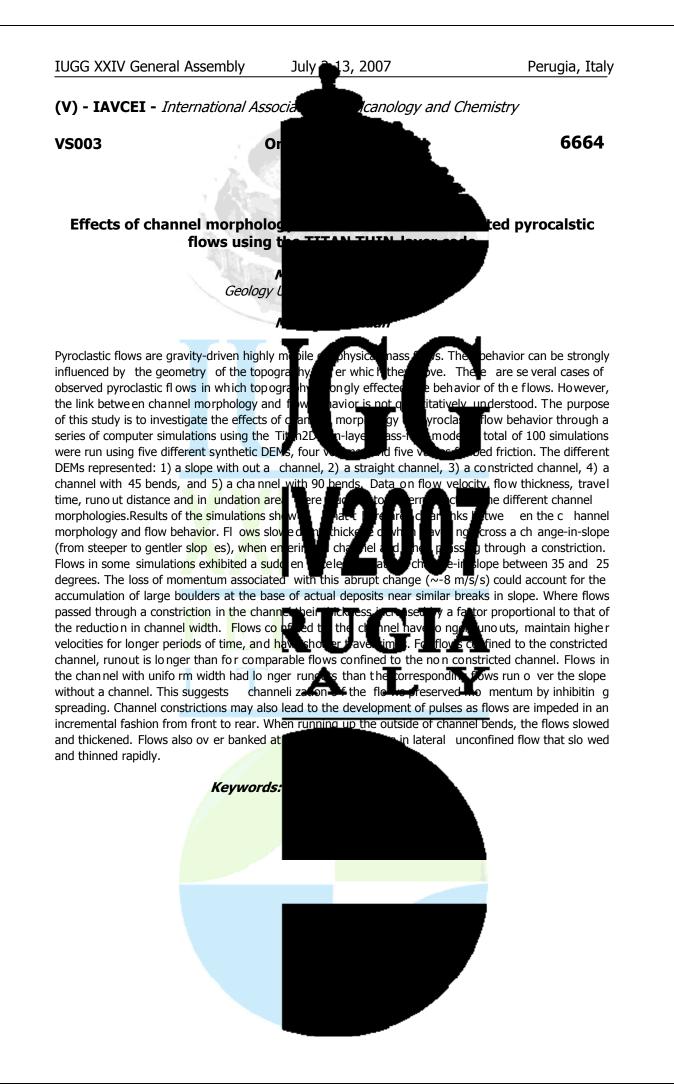
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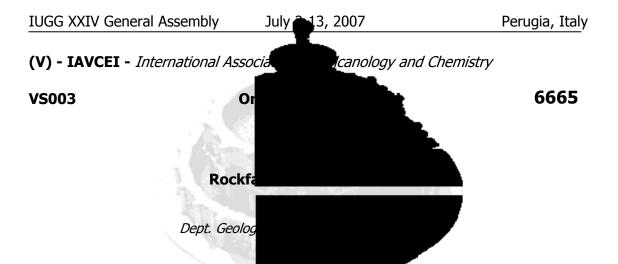
d inertial instabilities on éx hir ow emplacement, generating fina I stages of runout, emplaced alization of these features is aided PR), a that w e believe has cal ma lows . Surface wave ran ular flows, and on a

theoretical basis are anticipated within the context of volcanic granular avalanches. Their existence is relevant to a series of on-going debates including the significance of flow units, interpretation of sourceigh pyroclastic density

a case for considering r sites, and 2) strongly g of deposit-wide as well

Keywords: pyroclastic, flow, instabilities





Rockfalls, are anything but, the benign as being. Monitor ing rockfall ch aracteristics can potentially allow the collapse. Rockfalls are important as they re Probability density functio ns fo r natural overlooking small-tail end- member events. magnitude and timing o f major collapses relationships are best approximated by game other natural hazards. In thi s regard, there the mass wasting processes from growing ava do

predict power-law frequency magnitude and frequency-repose relationships, although experimental and natural avalanches commonly show more complex behaviour. A critical factor is how, a process such as sandpile collapses, ordinarily governed by t space and magnitude), behave when addit well as other external and/or second orde energy, max amplitude, and repose interv extrusion-rate dependant. Extrusio n rate, frequencies when rates are relatively high, whereas other ractors become increasingly significant when

rates are low. In addition, it is seen that rockfall energy (J/Kg) increases exponentially in the 48-hour period before some major collapses (e.g. dome pressurization cycles, normally associate inflation/deflation cycles. An interpretation rockfall counts during the days prior to the 20 May 2006 collapse. Thus, increases in rockfalls, especially with a 4-12 hour periodicity might be the most se Alternatively, the processes leading to colla pse any the result of carapace adjust ment to deep-seated vi scous creep before incipient brittle failure. Despite efforts towards understanding the physical mechanisms behind dome collapse at Soufriere Hills Volcano, Montserrat, and elsewhere, the propensity remains a seri ous issue that still jeopardiz

observe and better underst and the ca u precursors to lava dome collapses, and tac instrumental role in this.

cruptions they are o ften given credit for

eser part inte need to cons brk illustrates o lat<u>ion of r</u>o but ar pa а for epti ing ar lpile a

se histor y of the dome. he cð the mplete inventory not e i mportance of considering the alls. <u>Ro ckfall</u> frequency- magnitude h, to tha t observed for ery probably modelling, oretical sandpile models

ion of incipie nt lava dome

ity (in terms of time, of lava dome growth as seismic signal duration, ccurrence frequency is in determining rockfall

vbrið ea akes and crater rim portra cycles in long period ้รเ s **5** identify these cycles. reliable mea these ro ckfalls might be iv initiated a lapse with little or no apparent warning w, is to develop meth odologies to

is may be a manifestation of a

ifically temporal controls and itoring rockfalls can play an

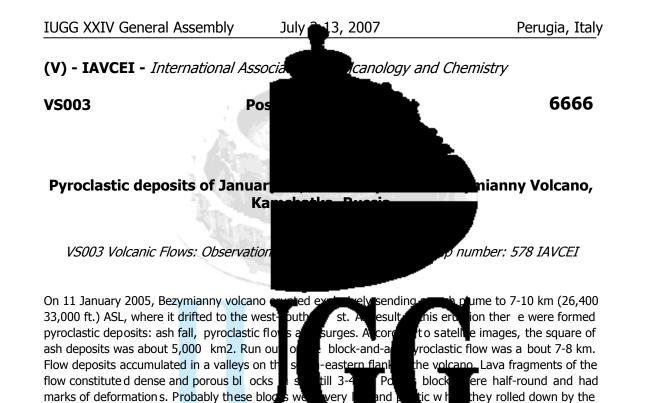
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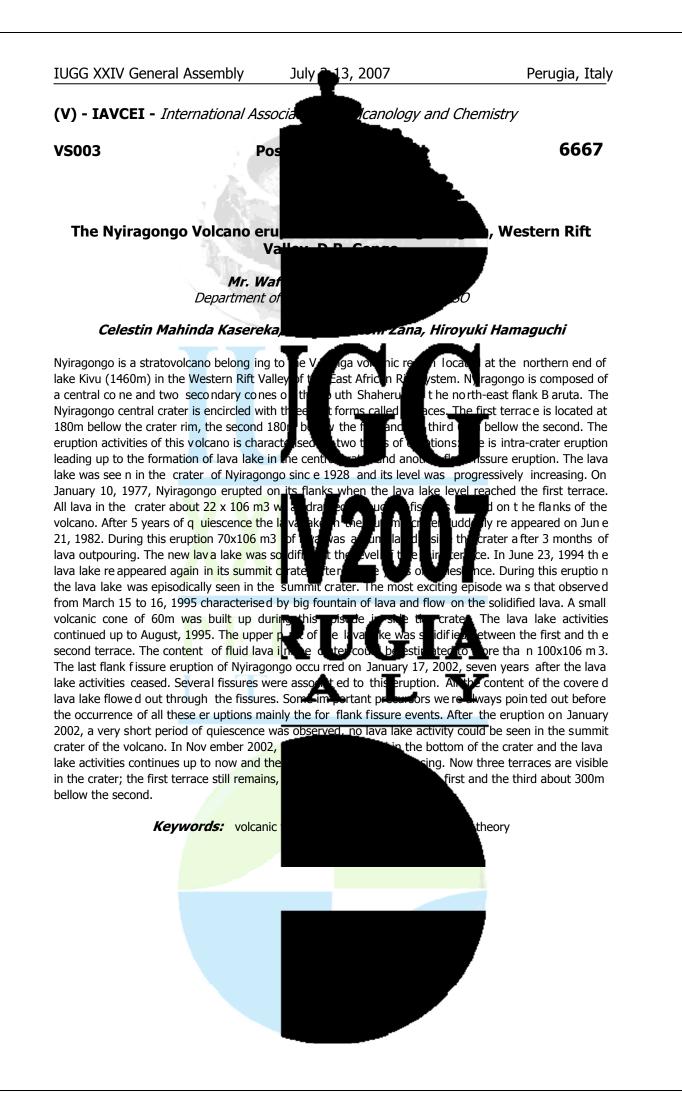
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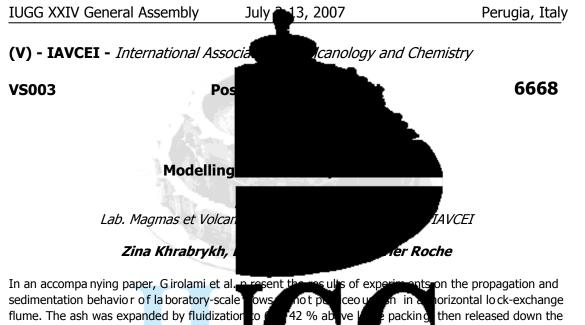


flow. Content of prevailing fraction of pyroclastic constrained by soits mater (0.2000,5 mm) was ~ 30 %. Pyroclastic surge deposits occurred in association with the block-and-ash flow deposits in the valley, and are also found separately on the south-western flank of the volcano. The run-out of surges deposits was about 10 km. The thickness of pyroclastic surge deposits are deposited to be about the purposite of the volcano was about 30-40 cm. Content of prevailing fraction of the rocease surges explicits metrix (0.25-0.5 mm) was ~ 50 % at a top part of volcanic slopes and ~ (0.% - in 0-7 time results). There was noted a gradual change of lava fragments content a volcanic surge are a compositen) interfurge deposits from a top part of the volcanic edifice till the is front. For orderic content volcanic transported gradually to pyroclastic flow on the south-western flank of the volca no. The volume of eruptive products of January 11, 2005,

flow on the south-western flank of the volca no. The volume of eruptive p roducts of January 11, 2005, Bezymianny eruption was about 0.07 km3, and VEI ~2.







flume as thin ( <10 cm), but fast-moving (1 impermeable, the ash defluidi zed progressiv slip conditions at the flume walls. Despite th place during flow. Deposition occurred by p hindered settling velocities in the flows we

bed-collapse tests at given v alues of expansion. We modeled the ash flows using 2-D depth-averaged equations of motion that included terms for deposition by hindered settling and a resistance stress due to material rheology. Wall and air drag we rheology of the expanded ash-gas mixture rheological s tudies of fluidized powder Dimensionalizing the equations revealed th acceleration; tsett, the time for hindered s

m/s) shear flo motion ceas ispers the a tυ /e sec nt e indist withir

Since the floor of the flume was low was lami nar in all cases, with o size segregation took the flow base. Particle those measured in 1-D

irst approximation the

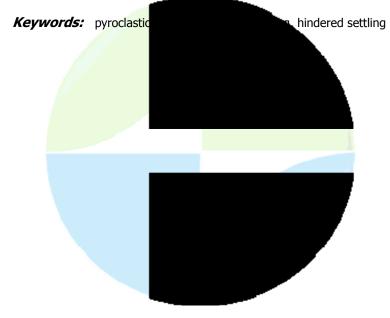
as justified by previous mbers of ash flows.

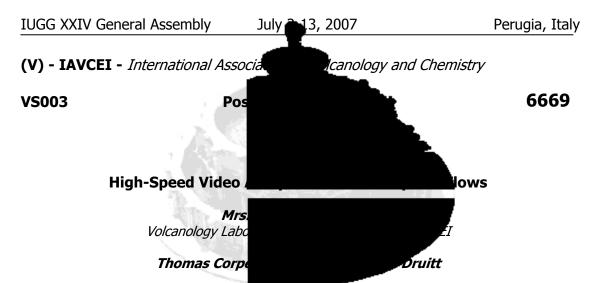
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e time for gravitational leration by macroviscous stresses. Three non-dimensional parameters governed the system: a, t he aspect ratio of the initially

fluidized bed in the reservoir, the ratio tsett/tgrav, and the ratio tgrav/tvisc. The equations were solved using a numer ical code capable of reprodu solutions to a hi gh degree of accuracy. Hindered settling velocities were d a fu ion of expansion from 1-D bed-collaps e tests in the e reservoir wi the ee variable was the g macroviscosity of the fluidized ash. Distance-time curves for flows over the entire expansion range were fitted by a single value of kinematic viscosity (1.4 alont to dynamic viscosities 10-4 m2/s This is equ. Lethe range of previously of 0.09 0.13 Pas (values decreasing with increasing expa nsi י which lie published values for fluidized powders. The study yielded simple scaling laws for the runout distances and durations of 2-D flows of fluidized ash.

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In an accompanying paper, Girolami et al. sedimentation behavior of laboratory-scale lows flume. The ash was expanded by fluidization to the flume as thin (<10 cm), but fast-movine impermeable, the ash defluidized progressiv deposition occurred by progr essive sedime velocities in the flows were indistinguishable at given values of expansion. In order to investi

experiment several times, we were able to <u>o</u>bserve the s edimentation behaviour both as function of time and with distance down the flume in because the experiments were reproducib Using specially designed image analysis so measure sedi mentation agg radation rates collapse tests. Analysis of particle velocity evolution of velocity gradients.

resent t uls of expe lot p ceou h in 42 vol% aboy 3 m/s) shear motion cease d atior he v bas rror f easur tho w dvi

orizontal lock-exchange bose packing, then released down s. Since the floor of the flume was key feature of the flows was that article hindered settling h 1-D be d-collapse tests ore detail, we filmed a

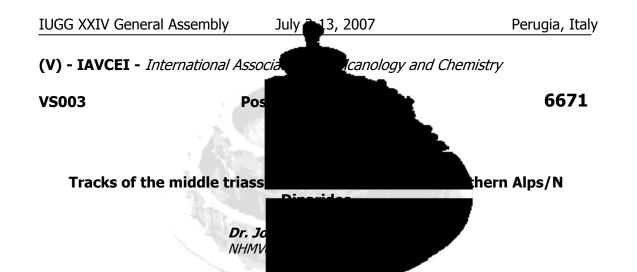
nts on the propagation and

number of experiments using a high-speed (1000 frames per second) video camera. By repeating each

ws. This was possible ill be presented her e. e the video footage t o etermin ed in 1- D be dvs temp oral and spatial







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Diagenetically compacted subaqueous pyrople deposited in non-volcanic, ba sinal environm deposits of 2 major explosive events, an A boundary (determined within ranges o f 40 ignimbrites several m thick. The Kammleiter indicating tropical climate during the Triassi are indicating a paleo-relief a nd erosion. F known (Julian Alps, Italy; Karavanke Mts., Slovenia

caldera are basalts and intermediate igneous rocks (slightly alkaline to calcalkaline) indicating volcanic activity before the caldera f ormation. Shallow marine carbonate clasts\_ missing of a central volcanic edifice in the event happened in a shallow marine enviro of subaerial settings. Uncharacterized volc The caldera er uptions happened at least p beds. The thic knesses (20-30 m) of ash t during the eruption. Co-ignimbrite ash falls seem to have remained partially in situ. The separated fine

ash had been responsible for producing relative crystal-rich spfs. A retro-diagenetical decompaction of flattened pumices (chlorite-smectite patch units up to 3 to 4 times as a conservative this ca. 200 km long segment of the Sout pyroclastic ro cks indicate a rather huge c aldera, which had been co mpared with the Campania n Ignimbrite (CI) caldera in the past (4) accor ding pyroclast dated yet as fossils are scare. New outcrops in the relative to the Tolmezzo area suggest an e ven larger diameter than the CI caldera (ca. 20 km). It is

remarkable that the non-com pacted spfs of the Mino an eruption, Thera (5), are only ca. 80 m thick in proximal facies. The non-compacted CI de Vesuvius ca. 100 m (7) with out remarkab represent one pulse within less than 2 m.a events preserved as pyroclastic/volcaniclas

up to 130 m thick ar e of) and ash he W sec tions comprise Twod L) event which bracket the A/L T he Carn ic Alps (Austria) h ost erize<u>d by bauxitic alteration on top</u> edbe d d er cluded

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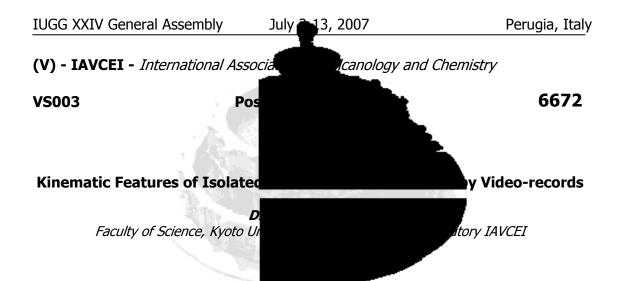
en conglomerates which welded ignimbrites are bsits of t he Pietra verde at the base of the flows, the lusion that the caldera pted, oxidized products he river Po plain, Italy .

> illis are known from ash in dicate heavy rain f alls

Furable thicknesses of individua l l and uctural reconstruction of a. The v arious lithified, l r iss orphology. Various sections are not na) at unkr wn tectonic displacement

dera rim is ca. 30 m thick, offshore erved 2 Triassic ignimbrites might a. covering ca. 40 explos ive the W Southern Alps (6).

Keyword



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Analyses of o bserved vo lcanic clouds using 1D mo. dynamics of eruptions an d some propertie entrainment hypothesis assuming that the velocity has played a central role in these a relation that was obtained from a simple din to volcanic clouds. If this Scorer's relation i conservation law, we can cal culate the loca time in a simple way as far as the conversion of the

cloud, i.e., the relationships among the ascent velocities, radius and height of the clouds. In this study, we examined the applicability of the entral in nature with their charact eristic constar obtained at As ama and Miyakejima volcan treated as an isolated thermal, i.e. a sph kinematic analysis reveals simple relationsh clouds. The empirical dimension nless constants

ned w ie and nsity chanid apply the above simple mod els to real volcanic clouds, we must check the kinemati c feature of the

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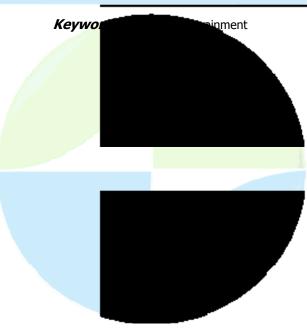
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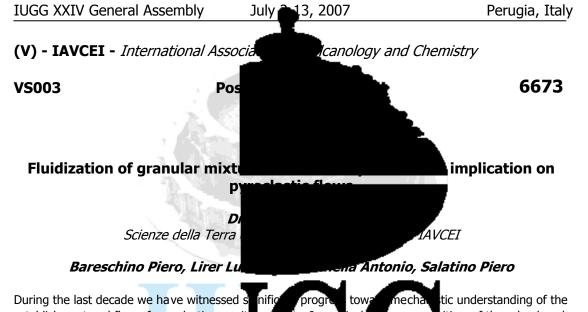
useful info rmation on the ld be noted that the t ŝ air is proportional to the ascent e entrainment hypothesis, another rer (<u>1957) m</u>ay be usefully applied inmer pothesis and the mas s cloud as a function of vol it can be neglected. To

tion to volcanic clouds bhs of volcanic clouds of the volcanic cloud is ly by its buoyancy. Our us and height of volcanic characterizing the entra inment hypo thesis k and the

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eir mean valu es are about 0.36 and 0. 6, Scorers relation C are more or less scattered and th respectively. The wide dispersions of the are probably c aused by mor e irregular conditions of the fluid ejection com e in lab atory These empirical relations can be used t o evaluate kin ematic feature ppr ithout thermodynam ic ds imą can сď егу consideration. Our results suggest, however, that C may change with the height, probably reflecting the Cor the volcanic cloud s effect the ambient density stratification in part. rthermore the values o iments. So y tend to be le ss than those obtained by the lab atory ex have to examine th applicability of Scorers relation more carefully.





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establishment and flow of pyroclasti c gravit by fluidization in the emplace ment of dense recently received additional support. N pyroclastic flows is still poor. The current published liter ature on rapid granular flows mostly addressed steady flows of granula materia

typical of pyroclastic density currents. Due to this limitation, most studies miss the complex dynamics of the frontal zone, which is likely to play an important relative extent of frictional, collisional, stre the rheology of the flow, does not even ap of granular currents moving over the grou big rotary dru m, 1800 mm inner diamete horizontal axis by a controlled step motor the experiments was a narrow cut of FCC, 40m diameter, 1580kg/m3 particle density, belonging to the

dynamics of the frontal zone of the current Factors that p romote air ent rainment in t assessed.Experiments have been carried du corresponding to peripheral velocities in phenomenology was observed across this range,

solid/plastic to free-flowing nearly inviscid at line the inherent instability of the rapid granula whose collaps e promote ext ensive air ent

completed by the extremely slow deaeratio established, is preserved over time interva avalanching and slow deaeration rate is the rheological behaviour of the current at larg

ts. In pallicul as tic flows, h ndin<u>g, quant</u>i ers tan tal th t t ha clines a

h e recognition of the role played hesized since the la te sixties, has ass<u>essmen</u>t of fluidization i despite the extensive estigations in this fiel d s far s maller than th ose

t <u>rol</u>e in <u>the</u> on <u>set</u> o<u>f flu id</u>ization. Moreover, the resses, which dictates tic flows. The behaviour ments carried out with a er is rotated around its granular material used in

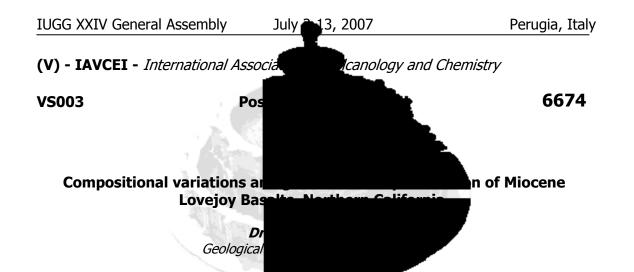
A-group of the Geldart classification of po wders. Experiments were focused on the structure and the urro unding medium (air). the s bli ent of fluidization were etween 5 and 30rpm, elocities the range 0.47-2.9 m/s . A pronounced change of the abrupt charge of the flow pattern from

th an almo relocitie ween 1.5 nd 2 (c orresponding to canonical Froude number of about 2.5). The change of the flow regime is apparently related to the onset of avalanching, at rotational speed corresponding to Froude number of about 2.5, as a consequence of entually develop into plunging breakers

of the current. The scenario is that the fluidized state, once e scale. The combination of and of the nearly inviscid

Keywords: self





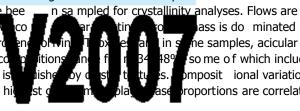
Lovejoy Basalt lava flows of northern California County . Flows traveled nea rly 250 km w Vacaville, California . Isotopi c ag es of Love content, but a whole rock age is reported indicate that t he flows were erupted within thousand years (Coe, 2003). Lavas are thole from 50-52% SiO2 but are also note worthy Ba and low Nb, Lovejoy Basalts have good i sotop

erupt ents loc t, a the alt lavas are c 68 My (Garri lative<u>ly short</u> rema I٧ th' igh B nte

hompson Peak in Lassen Val to Putn am Peak , ne ar ficult because of low phenocryst 200 4). Paleomagnetic signatures eframe. a few h undred to a few ositio homogeneous, ranging ppm). In spite of high ons with Columbia River

Basalts which are contemporaneous. A sequence of eight lava flows were emplaced at Red Clover Creek

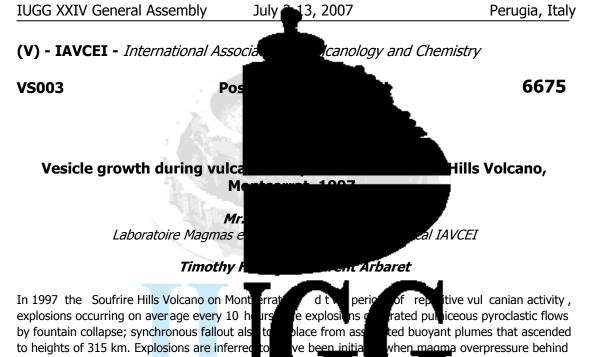
(approximately 30 km from source) and have bee approximately 10- 20 m thick and several h plagioclase wit h lesser amoun ts of clin or apatite. Gro undmass plagio clase ano rthite post-emplacement cooling, which can be between flows at this locality are minor, b with highest Al2O3 composition.



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ass is do minated by he samples, acicular so me of which include ional variations oportions are correlated





from the explosion s reveal h istories of bu bble an d gi breadcrust bombs are dominated by a pop ulation of small vesicles that must post- date fragmentation,

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since the glass y rinds of the bombs are vesicle-free. These vesicles have diameters of 10-80m, with a mean of 30m, and ~60 % of them are nor from associated pyroclastic flows, each co 10-80 m population very similar to that population, and (3) a large, 500-3000m p pul almost all vesicles are connected, probably for the conduit during slow magma ascent between explosions. By analogy with the breadcrust bombs, the small population is interpreted to be syn-eruptive . How ever, unlike in the breadcrust bombs, this

a highly viscous, degassed plug exceeded a

to a depth of 1-2 km. High-resolution vesicle

population nucleated prior to magma fr fallout pumices preserving t abular forms population is tentatively explained by the interview of t the conduit. Nucleation was triggered immediately after eruptive onset by the rapid propagation down the conduit of a decompression wave. Decompre followed by fragmentation, so that vesicles in the

At deeper levels, the time interval between a rrivals of the decompression and fragmentation waves was long enough t hat syn-eruptive vesicle growth was well advanced prior to fragmentation. Hence the smallest vesicles in pumices are syn-erup probably took place largely in the conduit pumices experienced different thermal hist

Keywords: vesicle size

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ion emptied the conduit age ex n pur s and breadcrust bombs e vesicular in teriors of

> e fallout and the other opulations: (1) a sma II, termediate, 100-300m te populations, in which coalescence at depth in

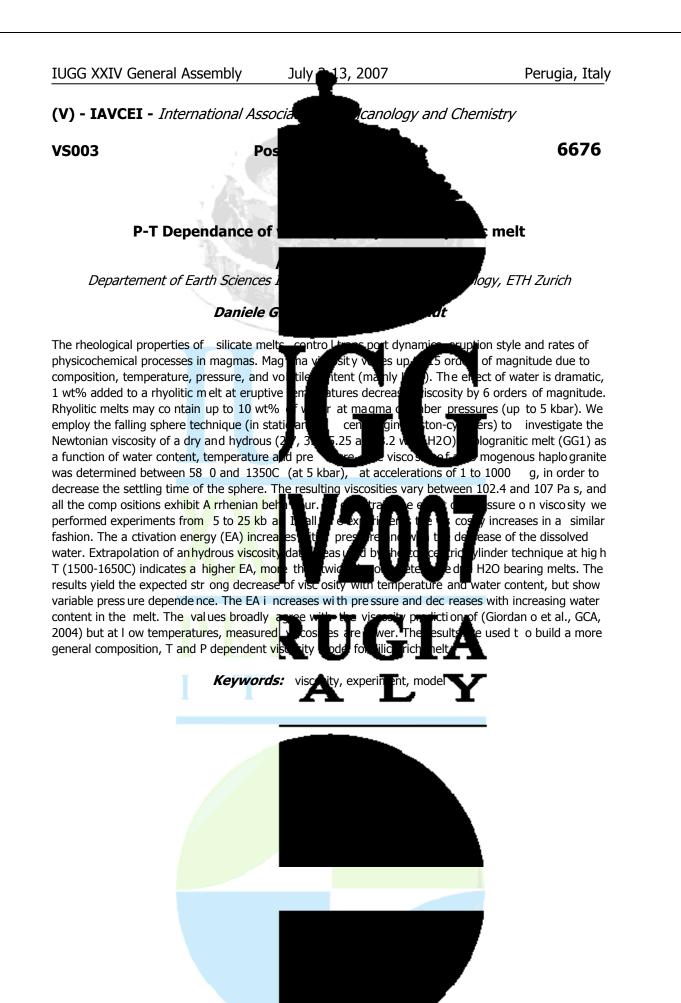
ritygradients are present in in of the small vesi cle sion and tragmentation waves in

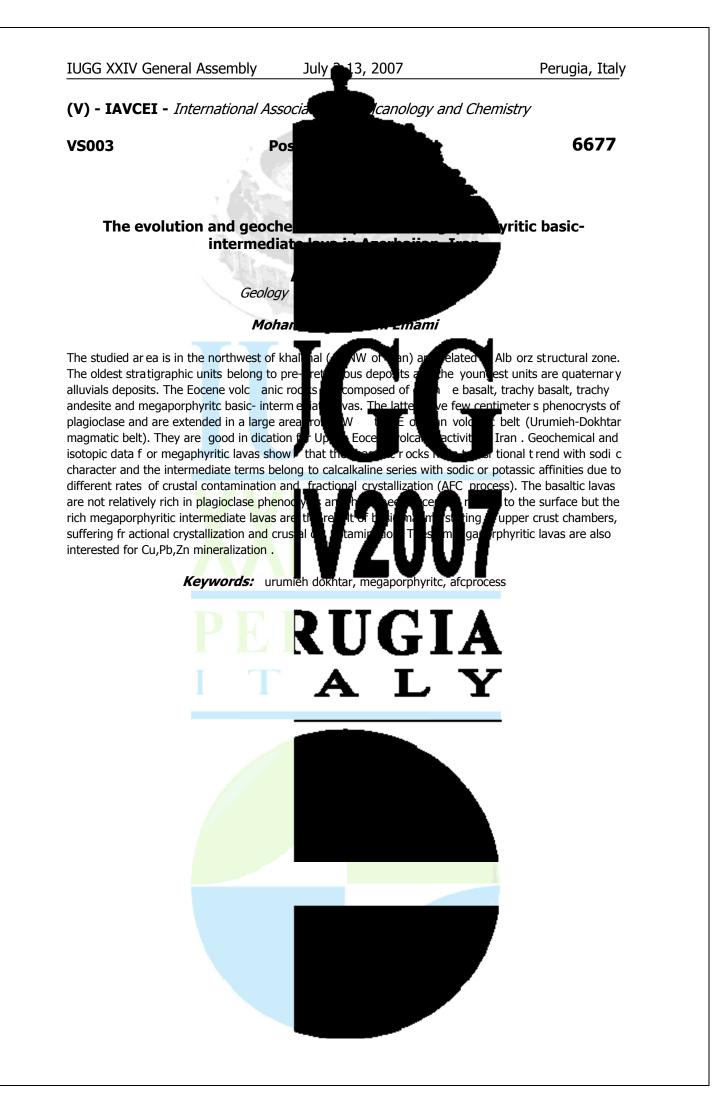
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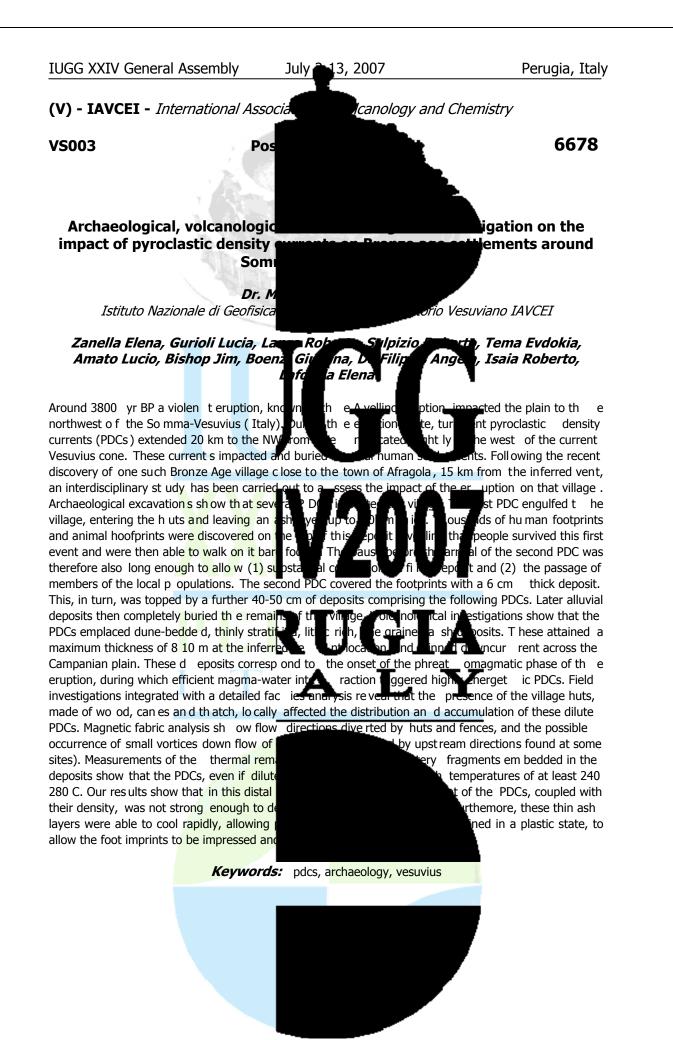
rely after f ragmentation.

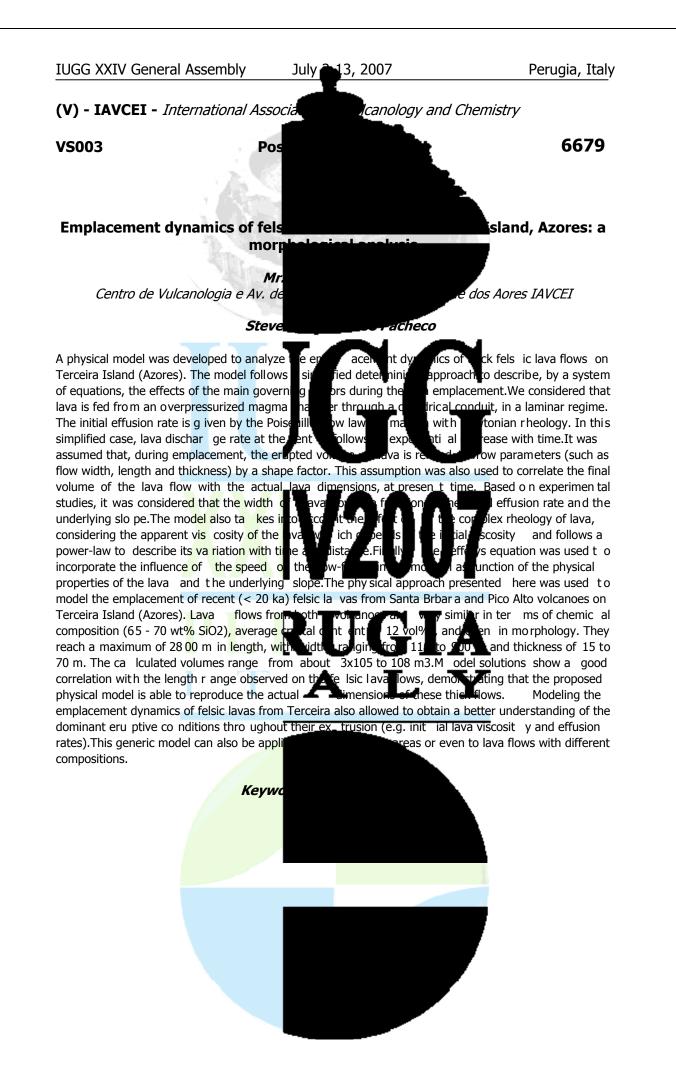
entation. Syn- eruptive vesicle gro wth ct with air, since fallout and flow t have the same VSDs.

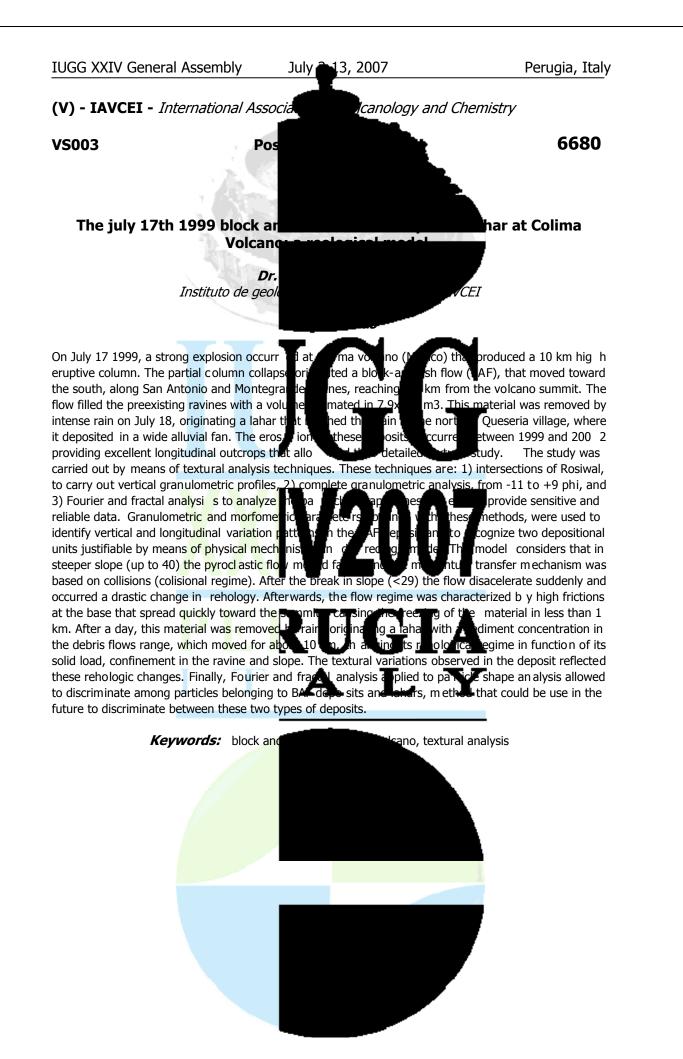
oufrire hills











## Perugia, Italy IUGG XXIV General Assembly 13, 2007 July

Icanology and Chemistry

(V) - IAVCEI - International Associa

**VS004** 

Symposium Intraplate monogenetic basalt processes

Convener : Dr. Guido Giordano, Prof.

The session will focus on the effects of the the magmatic source on the development The session is aimed at giving researchers volcanic featu res that ch aracterise plains basa strombolian, violent strombolian and phreat mag provinces. In both systems regional distribu of the rising magma with the structural fabr stress field and the presence e or absence ancient systems are welcome, as well as tho provinces. Hazard-related studies are also w

n es and are a c sty olcan oes bno e crust, the o nd or surface khun

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asaltic and kimber lite volcanism. discuss the variety of eruptive styles and erlite provinces. Effusive ally a

provinces and

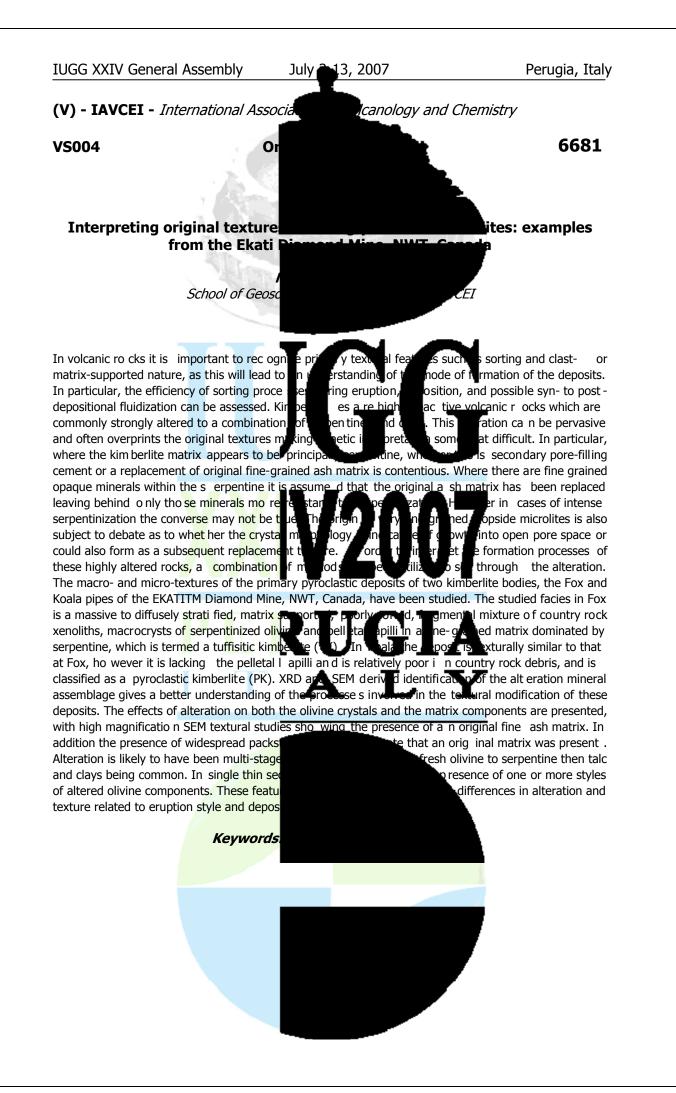
isplayed in plains basalt cano types reflects the interaction ation and intensity of the regional er. Studies on b oth modern and stems of such volcanic plum

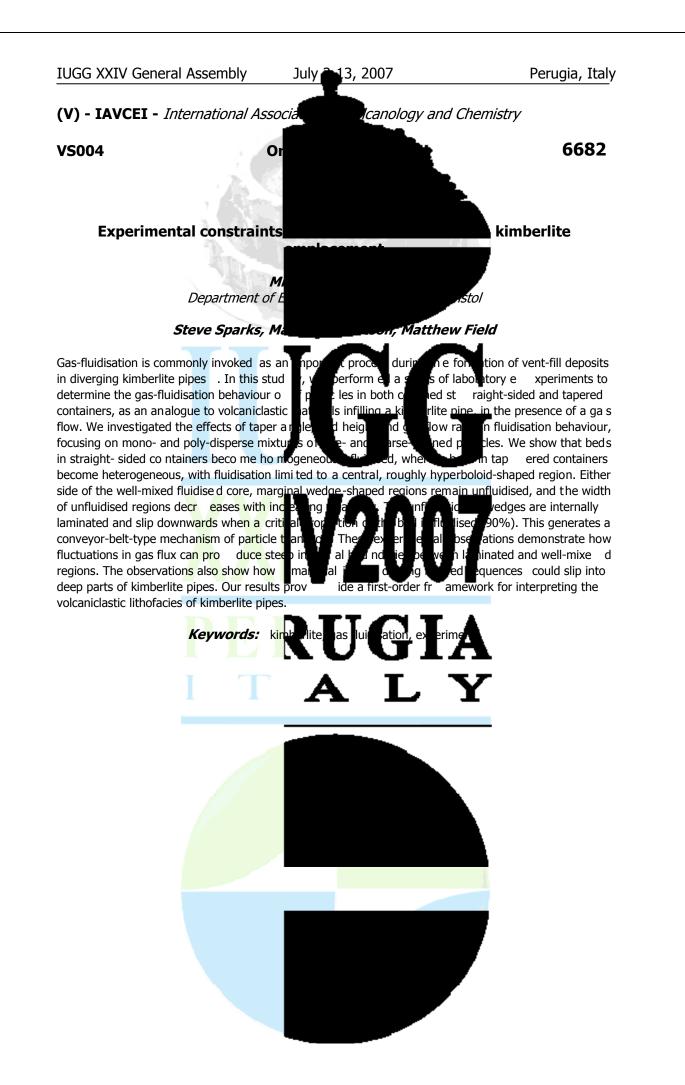
s, other regional factors and

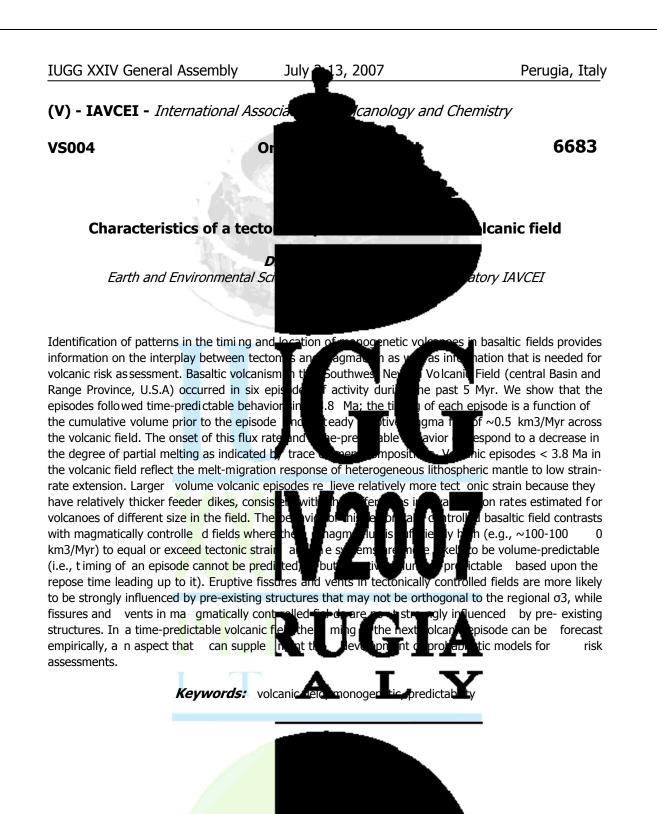
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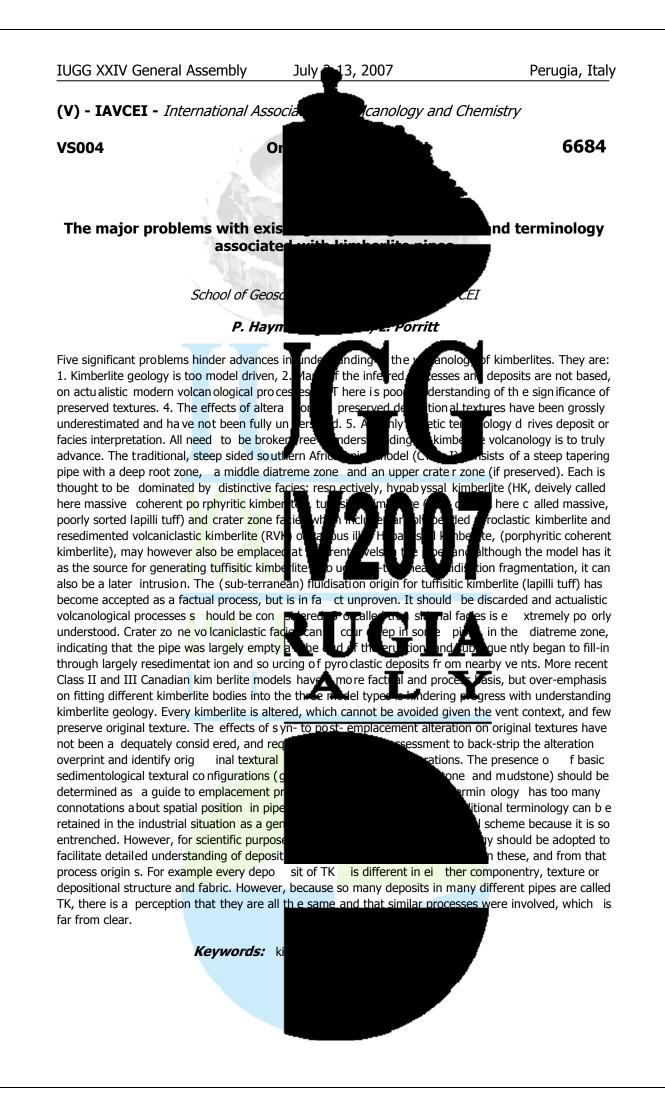


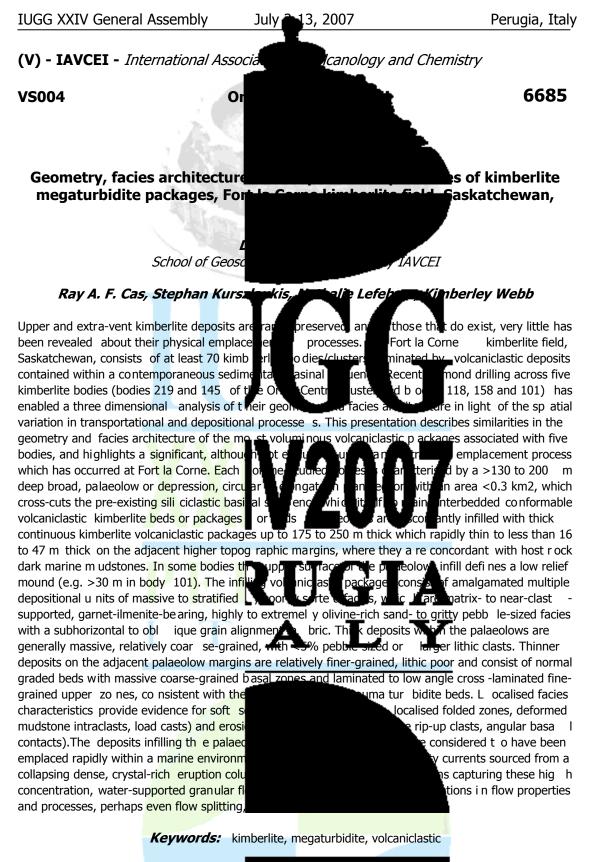


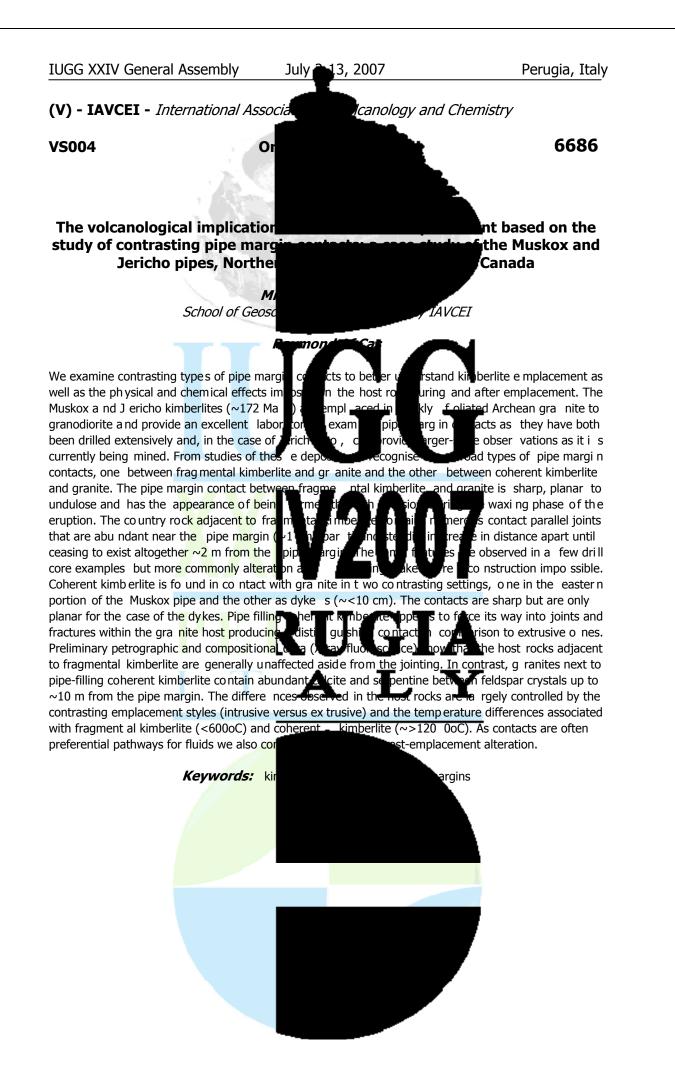


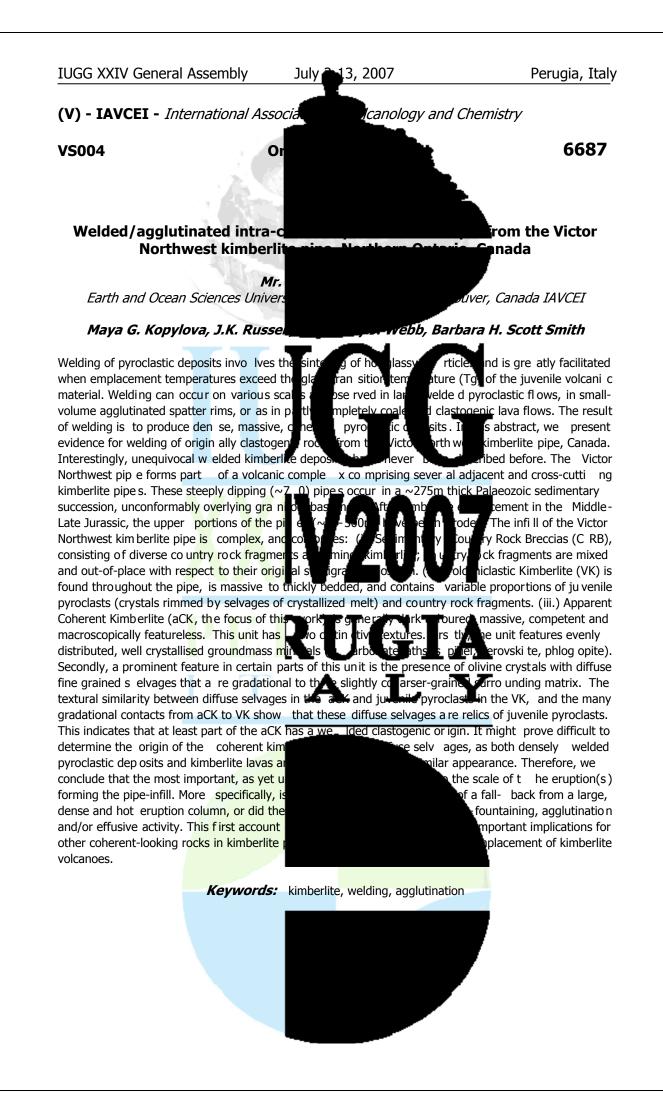


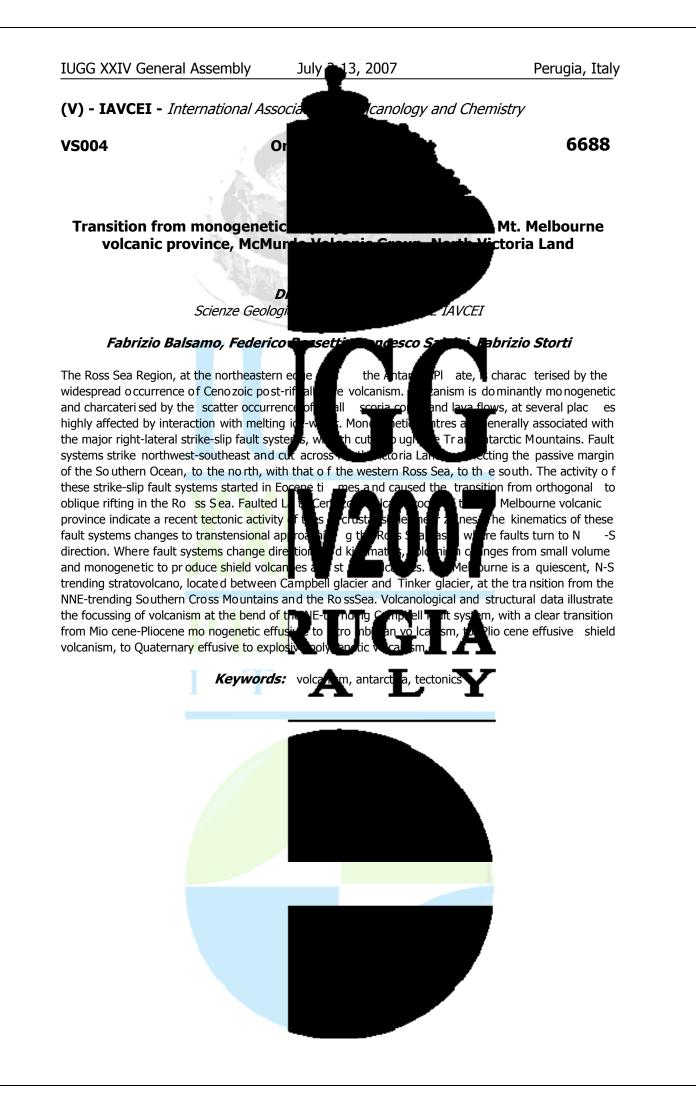


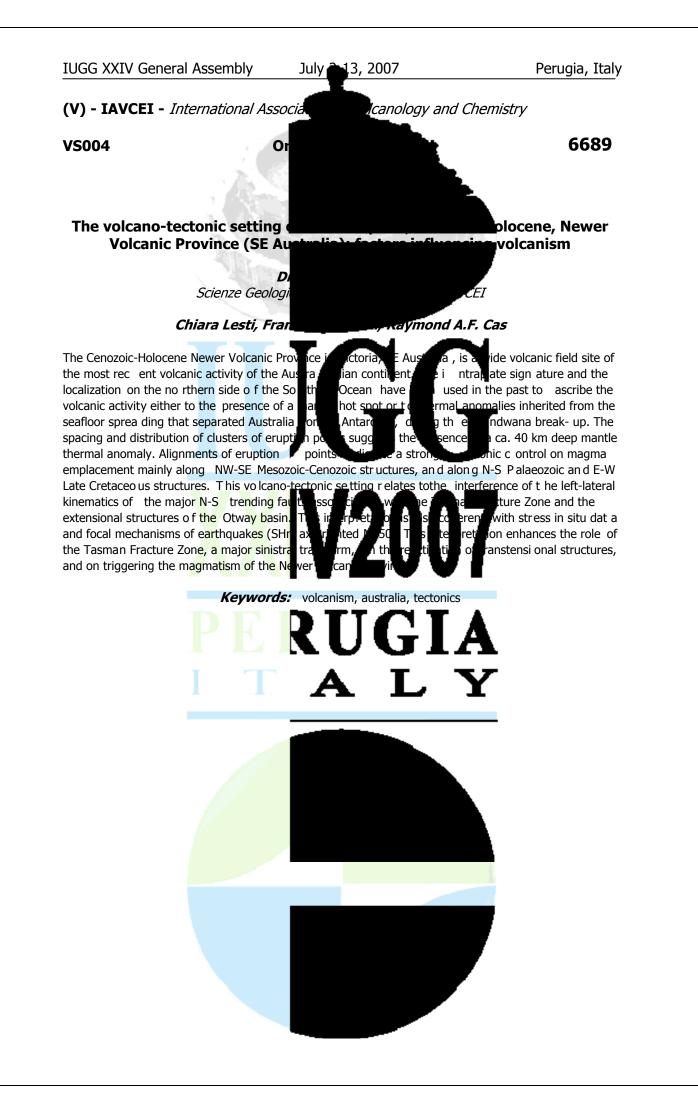




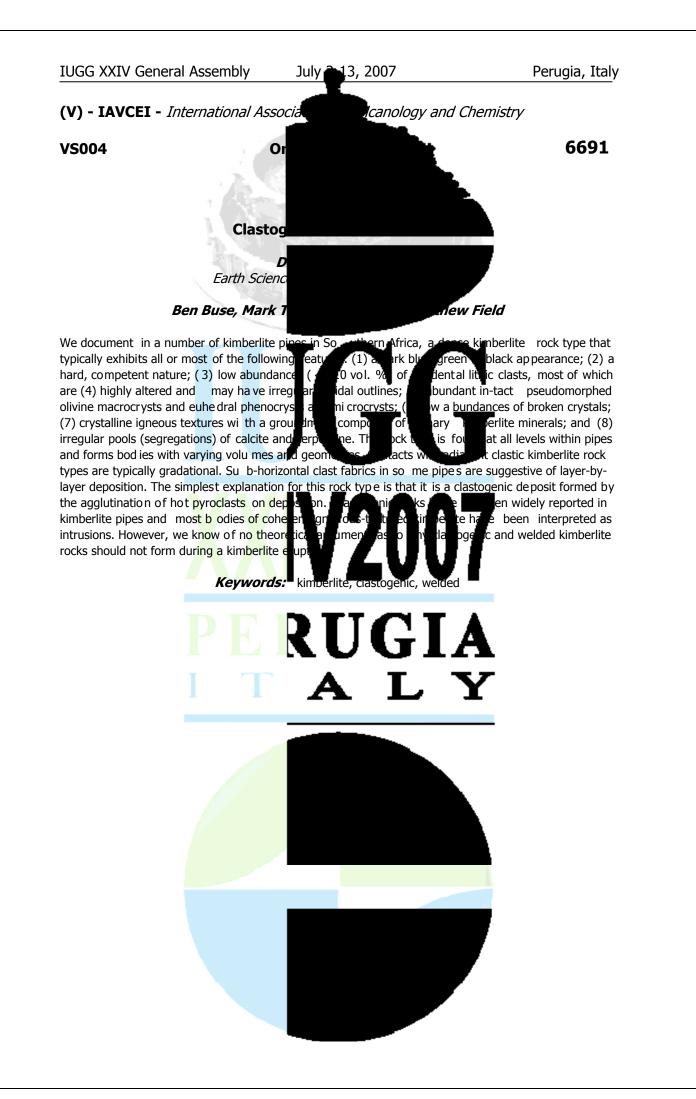


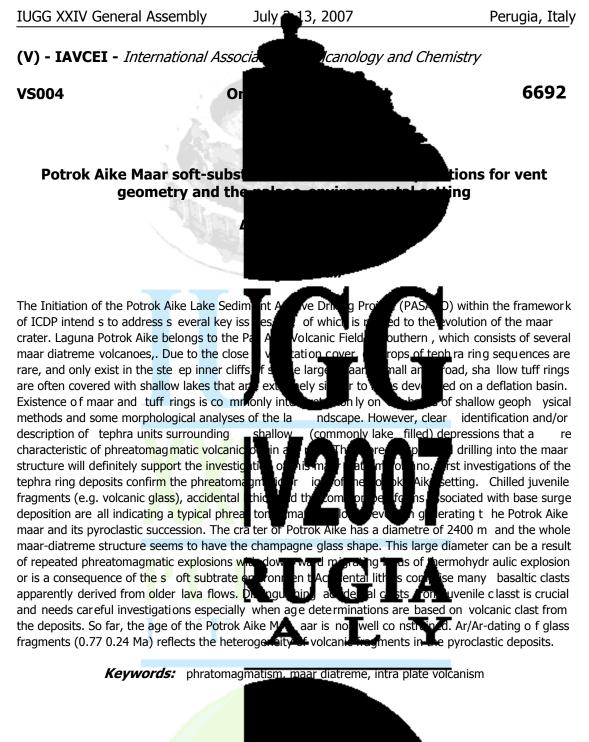


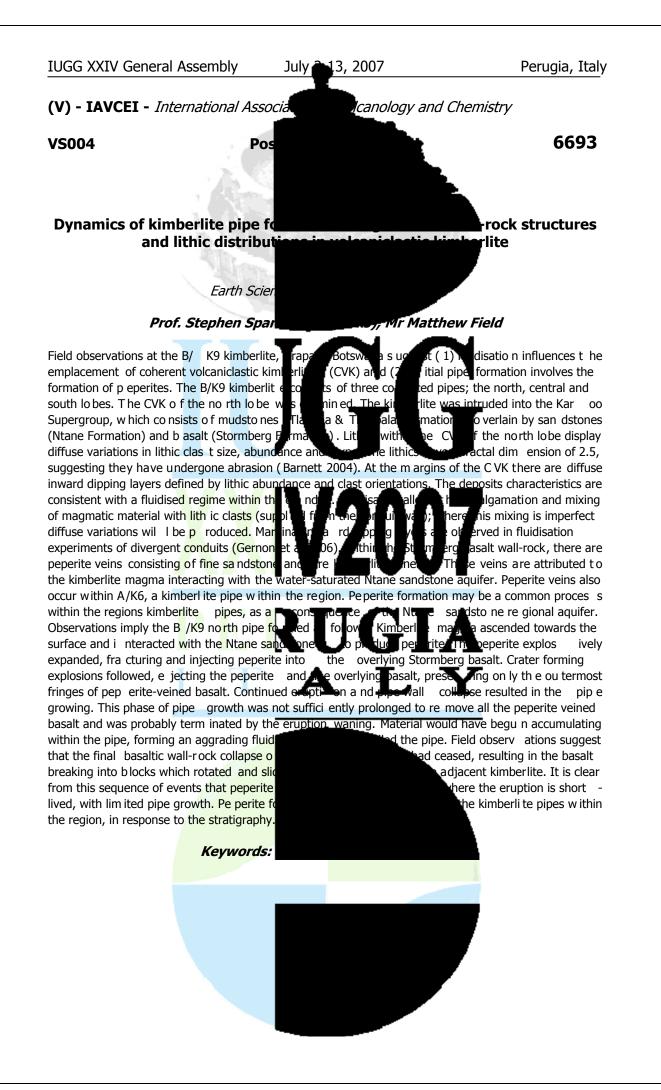


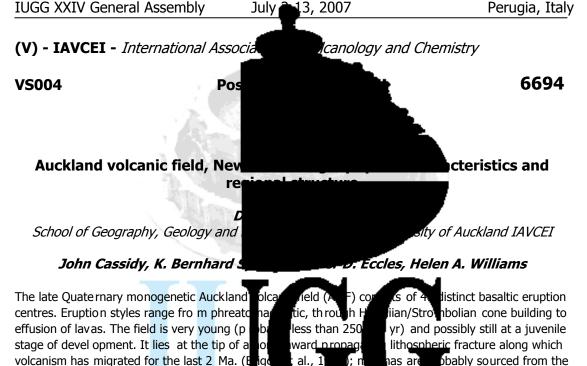












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of crustal complexity within a linear Mesozoi c terra (part of the Maitai terrane). This suture zone, re activated during the C enozoic and now buried, is marked by a NW-SE linear magnetic anomaly known as the Junction Magnetic Anomaly (JMA), which is a key tectonic marker and ex tends throug data show that in the Auckl and region the anomalies across the AVF whi ch are intern Mountain Ophiolite rocks. The AVF occurs i an abrupt narrowing and subtle change in AVF is also located towards the southern end of a large (50km wide) gravity anomaly, interpreted as

lithospheric mantle at depths of about 80 kg

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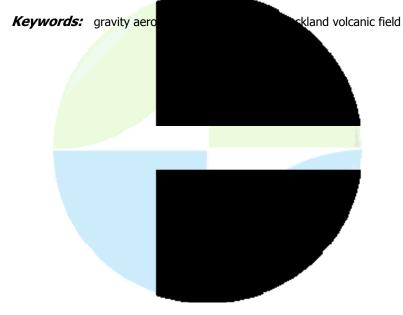
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bably sourced from the coincident with an area Mo untain Ophiolite belt

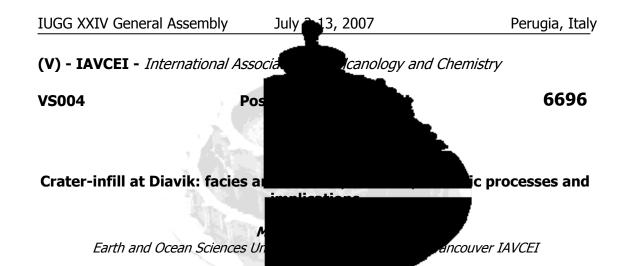
egion al aeromagnetic arallel li near magnetic shear zones in the Dun suture zone, marked by ity studies show that the

mafies may have momentarily within the Maitai terrane. This large comp arrested the northward prop agation of the whic plcanis m has migrated, lor permitting the joint arrays in the complex s to d into and decompress ss t at ť tk t ext lithospheric mantle regions, allowing magma to be mobilised. The distribution of volcanoes in the AVF is complex, with no une quivocal alignment of ver alt houg so me poss b alignments parallel to magnetic d regional Quaternary structural trends are apparent ecent pa a suggest that a number of volcanoes in the A VF may have been active at the same time but t here appears to be no simple structural relationship between these volcanoes

resulting from the largest block of den se material (most pro bably no n-sheared ultr amafics) known







Processes responsible for crat er-facies kimberlite der few well-preserved occurrences globally. New ope in the Lac de Gras region of the Northwest Here, we describe the geom etries, structur 150m of the A154N pipe. We then interpre basis of these properties. The base of th primary massive volcaniclastic deposit (MV) (PK2). The PK2 is overlain by ~65m of re-sediment

overlain by a moderate to well-sor ted pyroclastic kimberlite (~40-50m) that is sourced from another kimberlite volcano (PK1). The PK2-RVK contact is sharp and marked by the onset of bedded grain flows, wedge-shaped volcanic debris, and large, a represents the penultimate deposition of marks a transition to episodi c sedimentation crater. Deposits from post-eruption, resedimen variations in component sizes and types and im grade upward from coarse, kimberlite-rich debris to fine-grained, mud-rich thinly-layered deposits. This

variation is consistent with a change from large volume, high-energy, en-masse debris flows in a subaerial to shallow sub- aqueous environment from suspended fines in deeper water. Late crosses the stratigraphic contact between P

volcanic processes and grade distribution.

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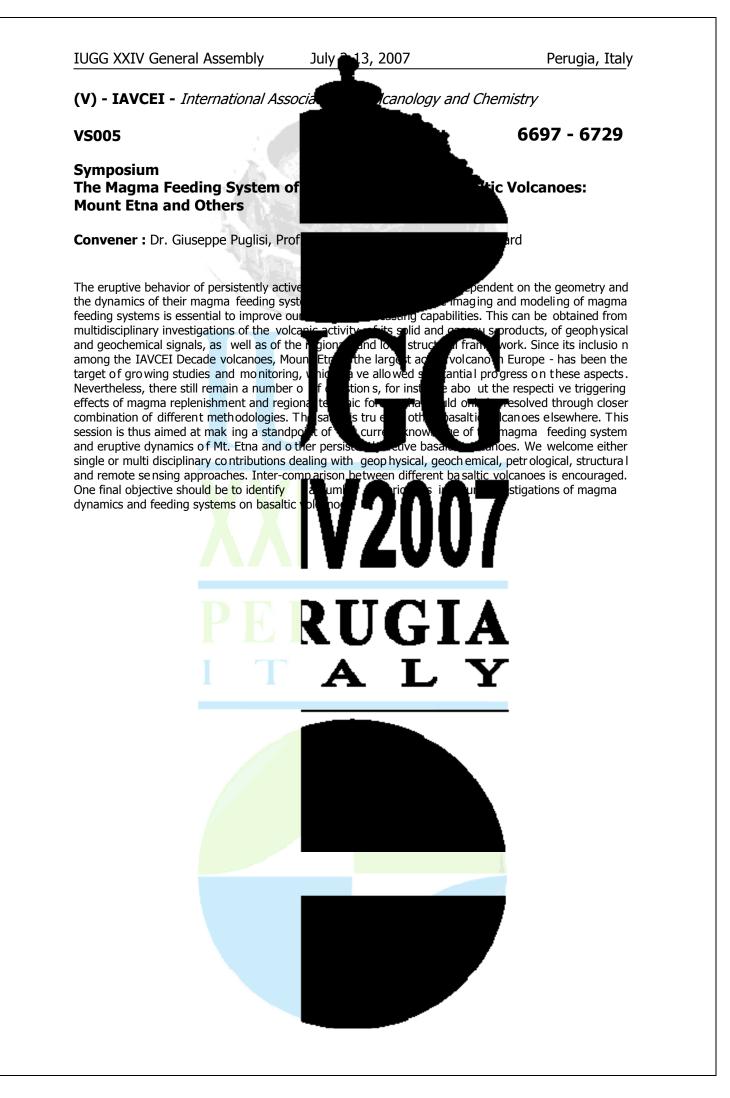
enstood because there are kimberlite pipe at Diavik pres ervel crater-facies deposits. he deposits from the uppermost ble for their emplacement on the -s orted, magnetic, an d a po ettered pyroclastic kimberlite (RVK). The RVK is, itself,

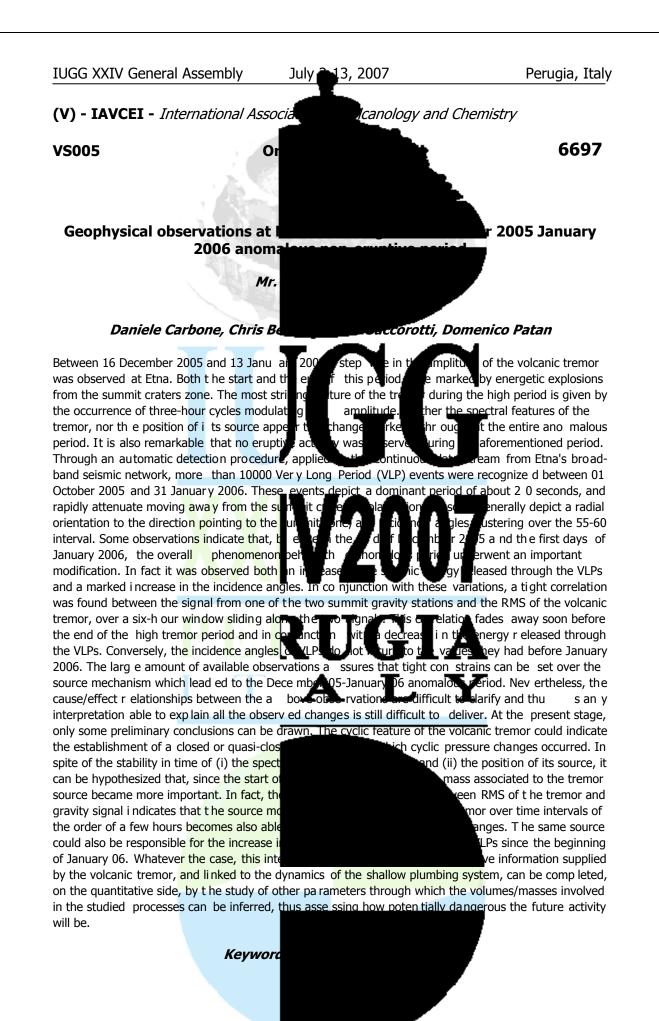
> berlite. T his in terface d of the eruption, and nic edifice into an open RVK) record vertical

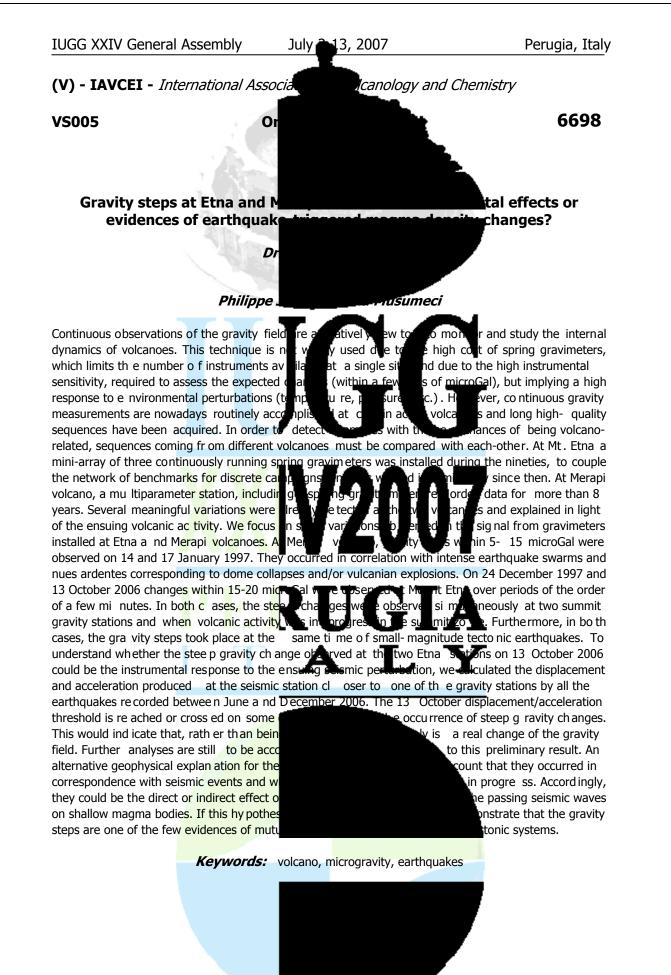
entation style. Sediments

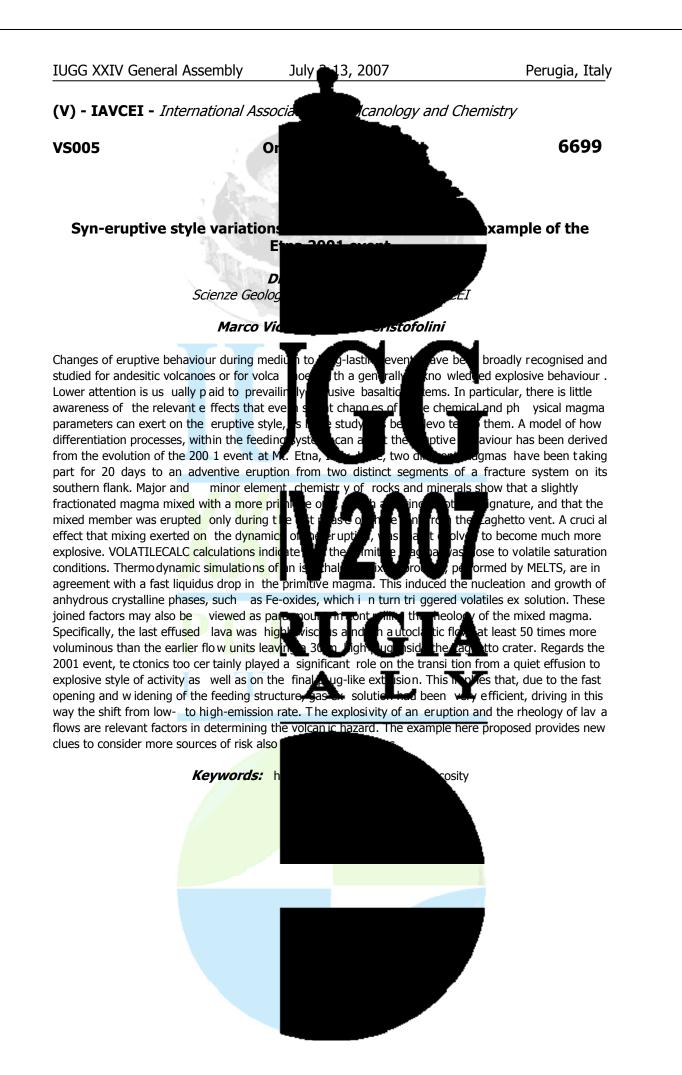
lows and sedimentation grain tion reaction front which benti () marks the top of the kim erlit : (F A154N pipe. The PK1 shows a variety of tex tures on different scales which suggest a pyroclastic origin, including a grain and clast-supported fabric, vesicing ted juven e pyroclasts and the absence of matri x mud. Grading over a 40-50m scale in the deposit, s ⊯p conta with under ng RVK deposits, and the absence of be dding suggests hydraulic sorting of an en -masse deposition of this pyroclastic kimberlite into a deep wa ter column, contributed to the A154N pipe from one of the many adjacent pipes in the Diavik area. The deposit, therefore, repres of A154N by an extra-crater kimberlite deposit from a later- erupting pipe in the o s have implications on kimberlite

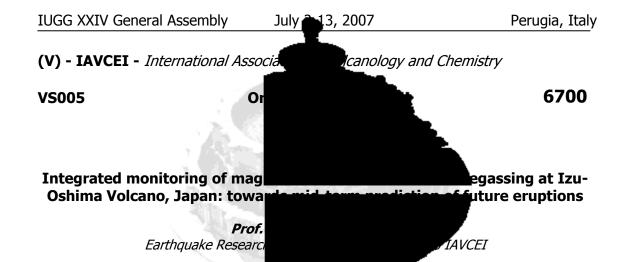
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In order to conduct mid-te rm prediction especially, magma accumulation and the way of r toward eruption. We have detected the sec lar last eruption in 1986-87, and further reveale resulting a net inflation and the accelerated seismic activity in the caldera region. We na supply of m agma from dept h. However, w at is processes causing the deflation; magma drain bac

degassing. If the latter is t relaxation of magma beneath the volcano and closely relates to the way of magma achievement of the conditions to start its rising up toward the deflation cycles might give us an inval uab the degassing of basaltic magma accumula has a low solubility in magm a and separa September 2005, we started continuous minit summit of the central co ne Mt. Mihara of Izu-Oshima volcano. Measured data are

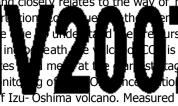
minutes in a logger and accessed via r adio LAN sy stem. All the instruments are powered by sola r battery. We further surve yed the distribut installed another continuous measurement The CO2 concentration data showed tempor features. 1) So il CO2 concentration so metimes increa sed with duration of se veral hours to days. 2) Decrease of 1m-depth temperature followed that is soil CO2 concentration with delay of several hours, suggesting that both the soil CO2 and high temper from depths. 3) There occurred several correlated

seismic activity in the caldera. 4) There occurred peculiar seismic events (with a predominant frequency of about 1Hz) originating from very sh allo concentration. These low freq uency event the summit. We will further elucidate the n Oshima volcano by integrating ground def electrical resistivity and CO2 concentration

Keywords: eruption pr

ecur sory processes: ditions to start rising up the hima volution of since 1989 after the rep eated inflation-deflation cycles panied by the elevation of shallow flation is caused by the /olcan eflati There ar e two possible umulated magma due to

he case, the inflatio n-deflation cycle indic ates the accumulation and



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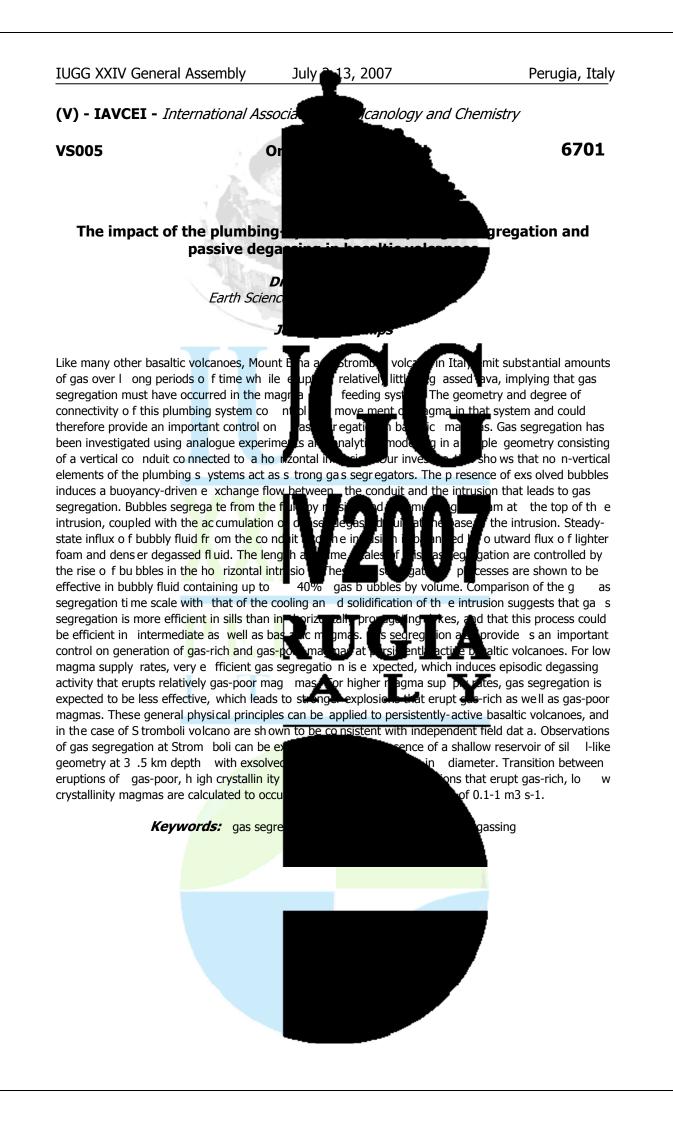
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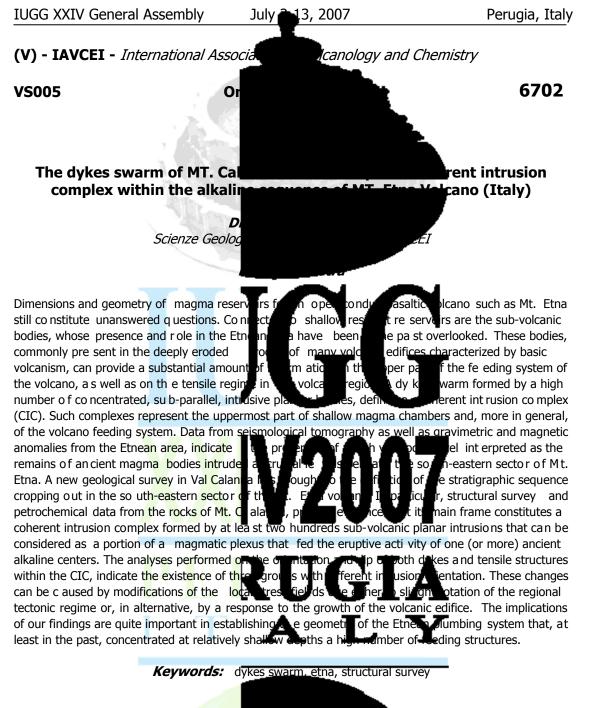
bservation of inflationprocesses. To monitor st helpful because CO2 of accu mulation. On 28 at the eastern part of the stored every 5

around the summit area, and er 2**0**06 the stern part of the summit. f 01-2 vol% and the followin g kes are fed volcanic gas emanating increase of the soil CO2 concent ration and the summit in the period of elevated CO2 pid flow of volcanic gas beneath assing processes beneath Izu-

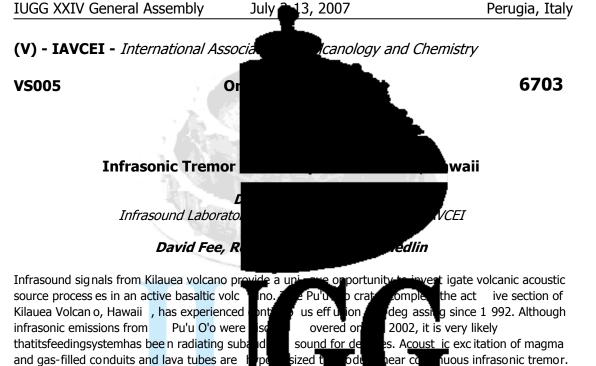
changes in magnetization,

gassing









To capture the sounds from the ongoing ac itν in October of 2006 within a dense tropical I fore meteorological data is trans mitted from t he array in r eal-time, with results being displayed on a dedicated website.An abund ance of acoustic tremor signals has thus far been recorde d. Although the tremor is assumed to be constant at Pu' changes. Some of the variability appears to

atmospheric conditions, particularly diurna representative of regional wind patterns, ld proxy. Harmonic and gliding tremor have a conduit geometry may be pa rtly responsible for

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trem or uous infrasonic tremor. und array was deployed rater. Infrasonic and

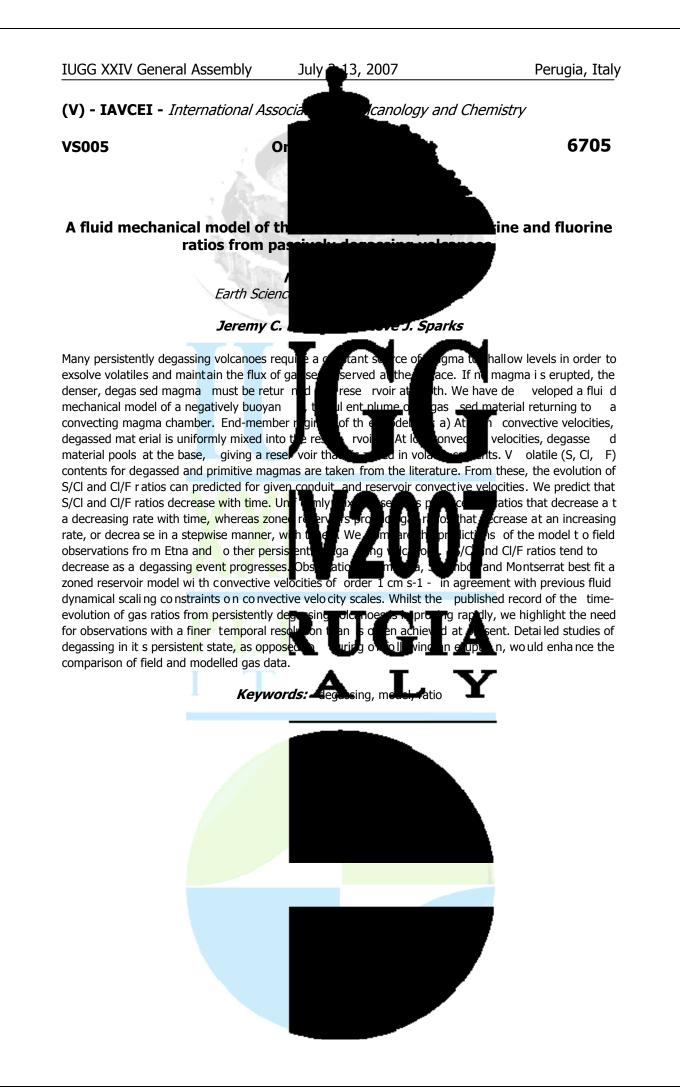
it signi ficant temporal ts created by changing

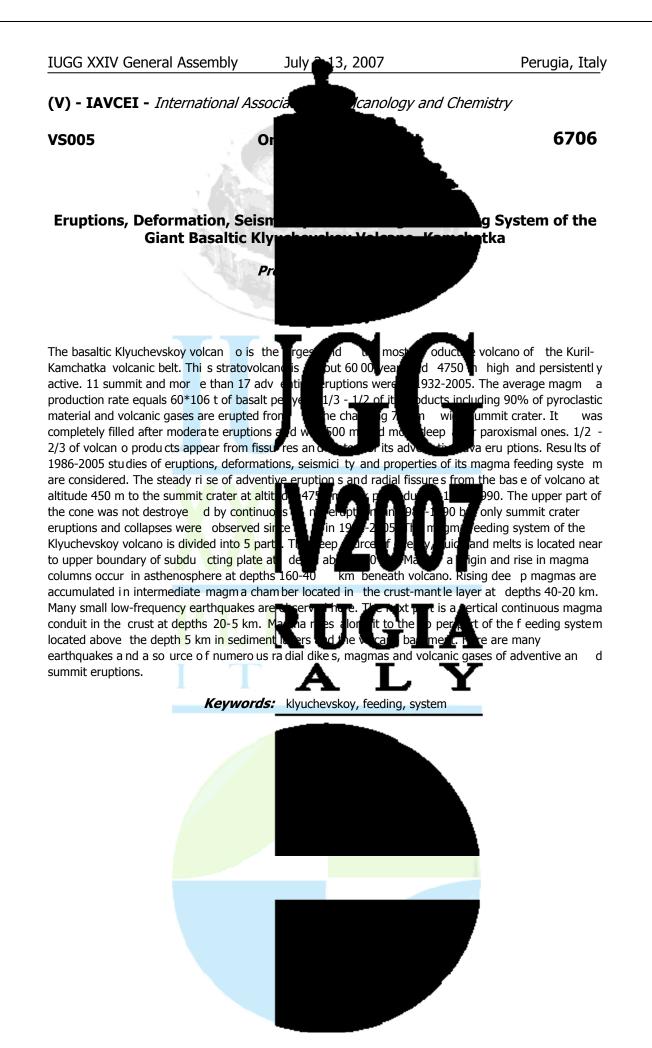
eorolog ical data is not litudes can be used as a content, mass flux, and

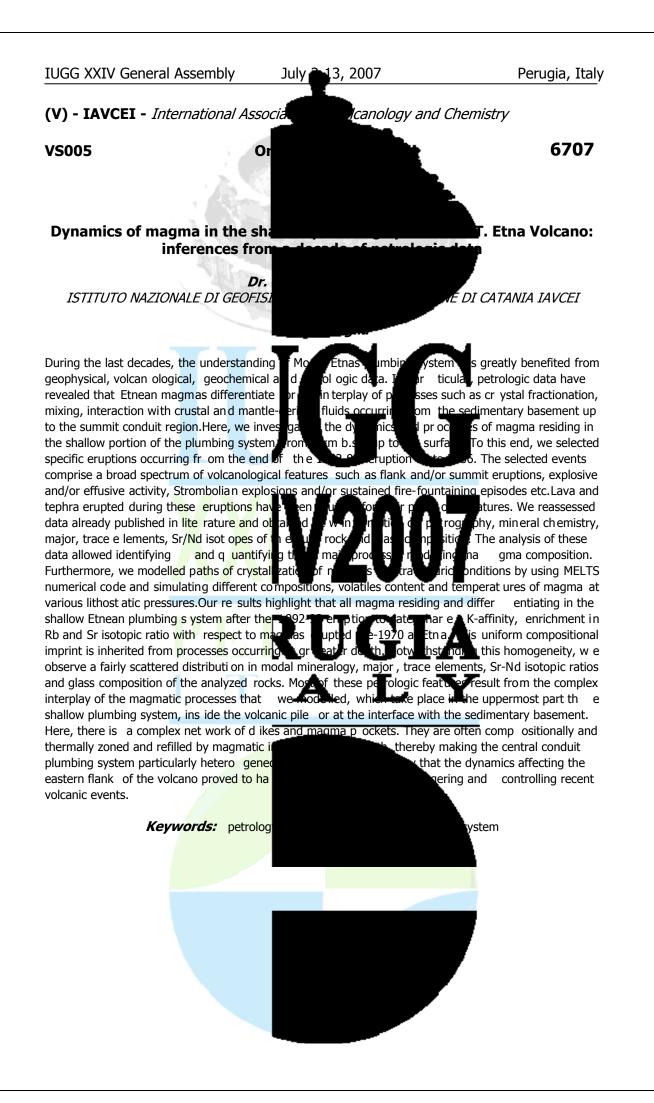
variability. Gas pistoning events are also be potential sources of deep infr asound. To further examine the tremor signals at Pu'u O'o , we want to study the relationship between the infrasour ation and meteorological data . Preliminary comparisons wit h seismic data inflaso c sigr may h ave a differen t source process. An expedition to install an to nd perpendicular to the he y SO епе lava tube syst em is planned for Spring of 2007. These new results may permit a better identification icing infras and characterization of the volcanic processes pro und.

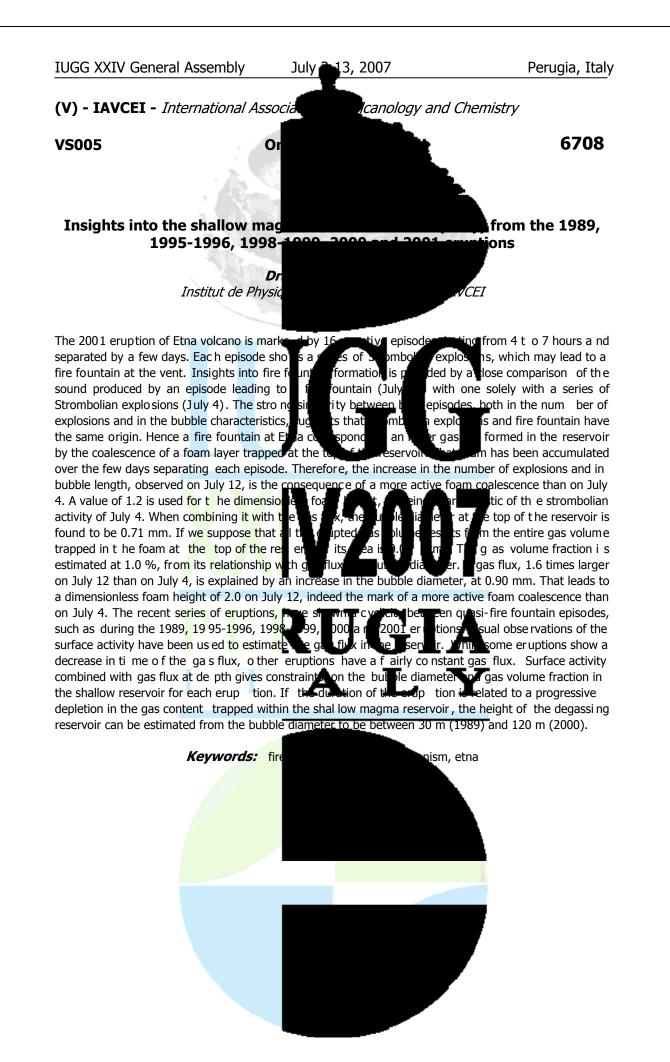
Keywords: infrasound, volcano, hawaii

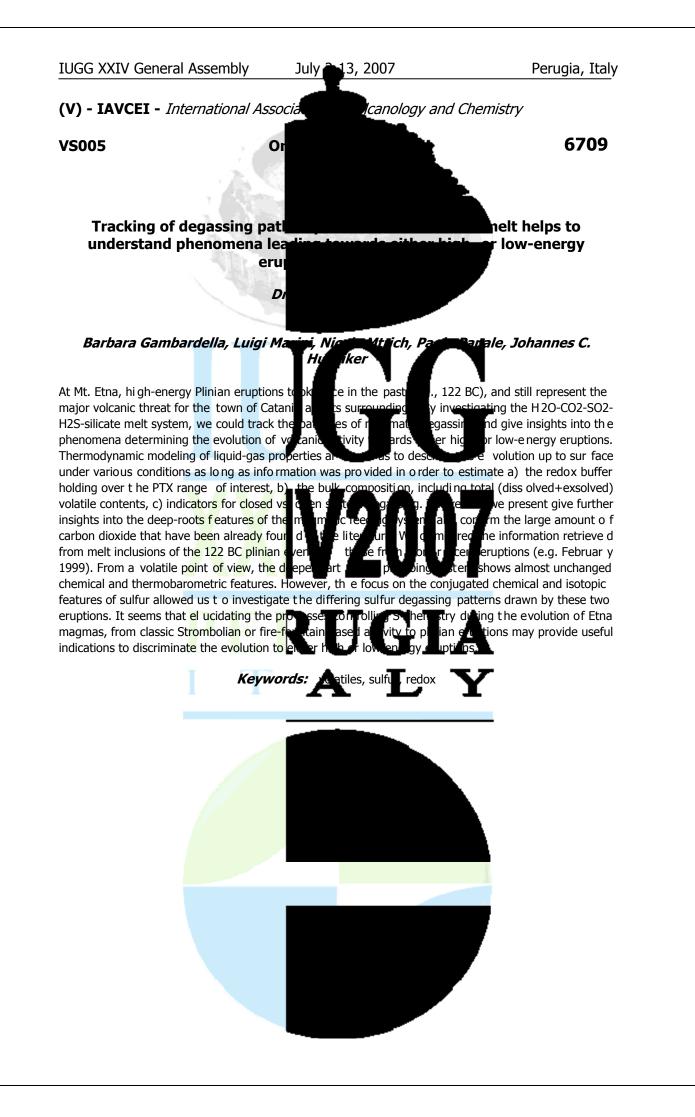


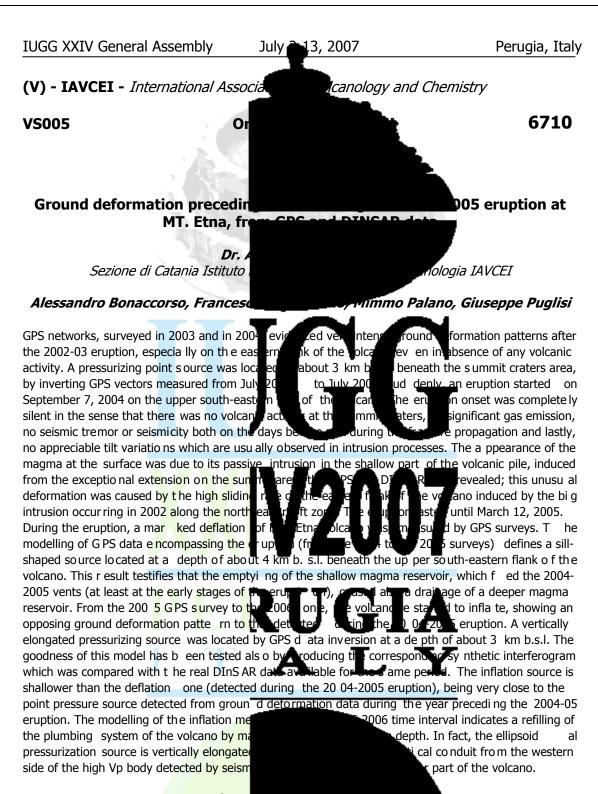




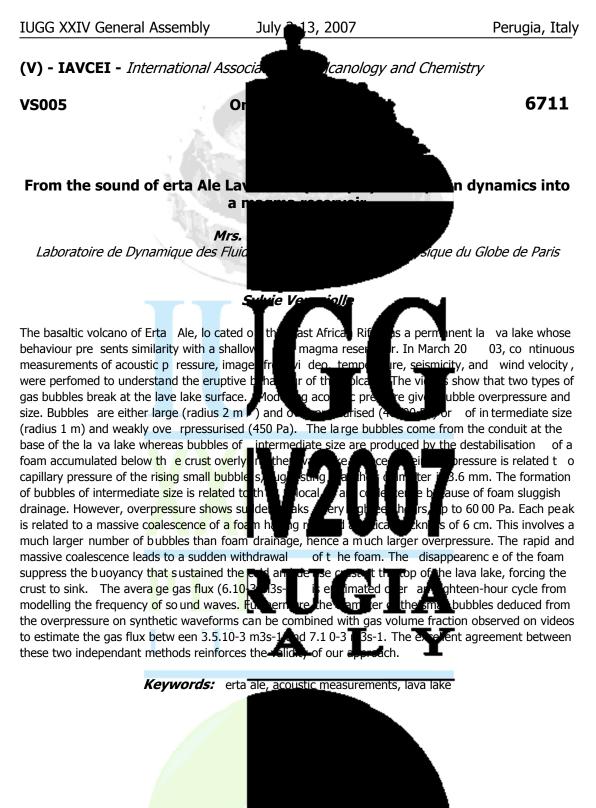


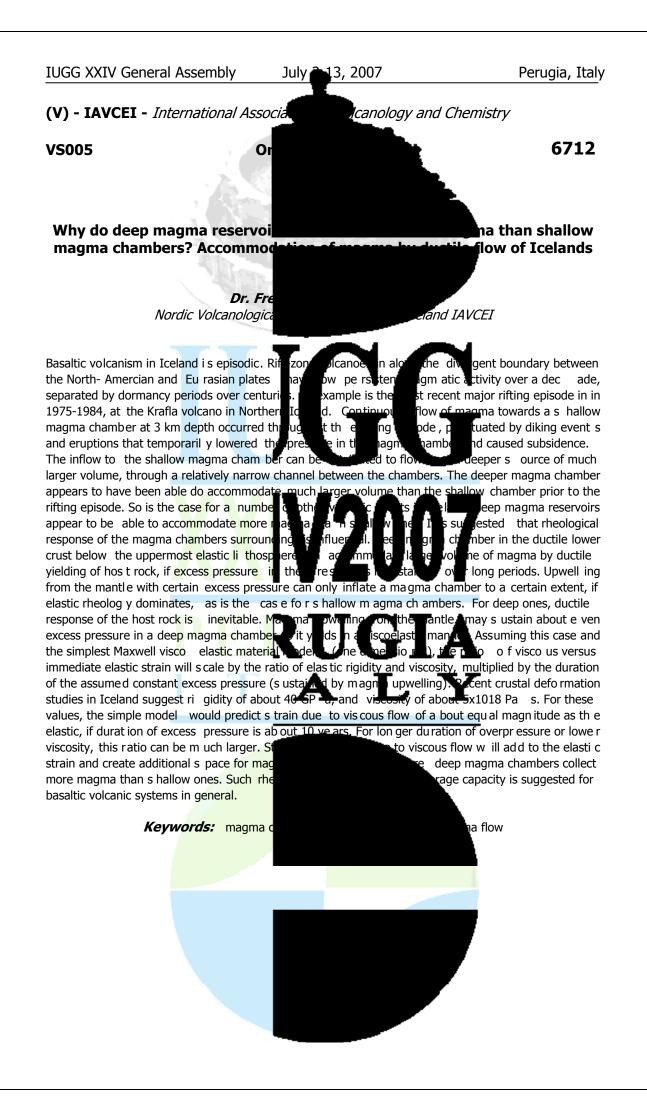


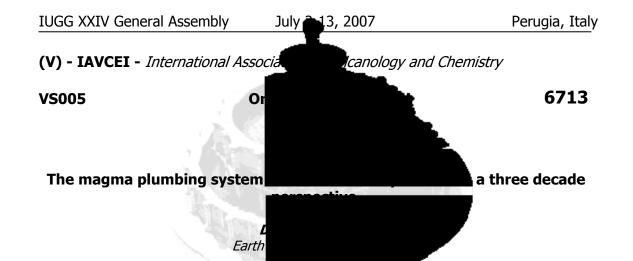




Keywords: e







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We propose an updated ass essment of Eta petrologic, geochemical and geophysical da major event in this period was a mantle-deri magma that i nvaded the i ntermediate to craters) at a mean rate ~3.7 m3 s-1, trigger magma differs from all trachy- basaltic prod content, higher Rb/Th, K2O/Cl and S/Cl ra features are recorded by melt inclusions encrapped

they cannot result from crustal contaminations in the shallower sedimentary basement; instead, they track recent partial melting of a geochemically distinct portion of the mantle source. The new feeding magma could reach the s urface almost ur 1974, then du ring powerful summit lava f highly explosive flank eruptions in 2 001 lateral dyke intrusions that bypassed the system. Otherwise, the new magma grad previously filling the plumbing system, as demonstrated by the spectacular evolution of alkalis/Th ratios

versus Th in lavas erupted since the early seventies. From this mixing trend, the cumulative amounts of erupted lava, and SO2 constraints on the a the overall magma storage capacity of the Molten magma would thus o ccupy a tiny the ~10 km thick sedimenta ry basement. Concordant information from seismic to mography, gro und deformations and crystal melt inclusions suggest two mag km beneath the craters, in coincidence with two ma

Continuous degassing but dis continuous extrusion of the evolving magma mixture has been regulated by the shallowest ponding zone, where partial magna dehydration and crystallisation due to deep CO2flushing o ccurs. This po nding zo ne is als deformation. A t the interface between th depth), a network of sills and dykes, conne important control on magma degassing an

of the cumulated supplied magma were e magma was likely drained back into sub-v accumulation in the plutonic body, as well

ing syst g acco unt of abundant st thre e decades. The here er th lse of hew plumbi ng sy ance<u>d erupt</u>i re cen ti dio Sra Pa(≥Î

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re alkal ne volatile-rich basaltic  $( \leq 15 \text{ km depth} \text{ below summit})$ tivit<u>y and la</u>va extrusion. This new by ha ving higher K2 O B is otope ratios. These g-rich olivine crystals, so

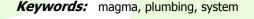
> eccentric eruption in luminously, during two ade pos sible either by fast a scent across this Ikali-poorer trachybasalts

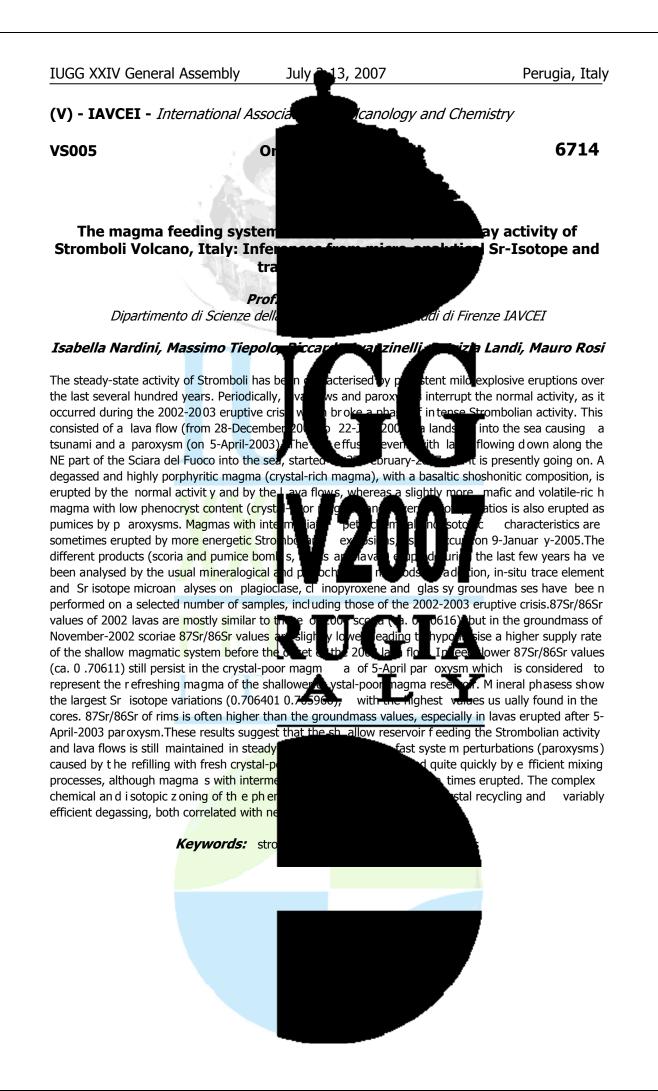
re-evaluate as ~3.5 km3

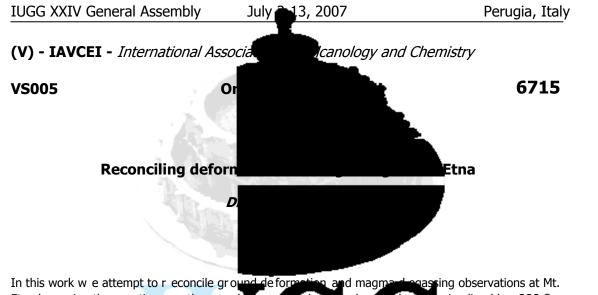
hic hoay hat is e mplaced within

system of Mount Etna.

a ponding 2005 exist at ~132 and ~53 discontinuiti in the crustal basement. f pre- eruptive seismicity and gro und im entary baseme nt (2-2.5 km its ~25 m wide, may exert an Over 30 years, only 25-30% r upted, denser degas sed lid accretio n and stress rn volcano flank.

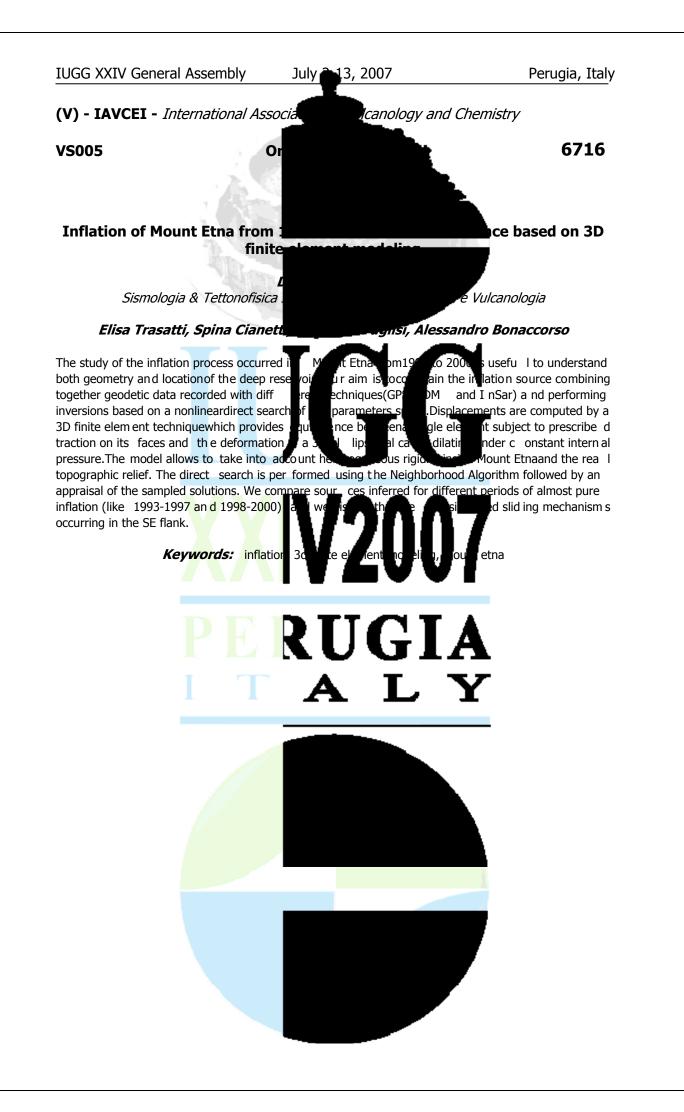


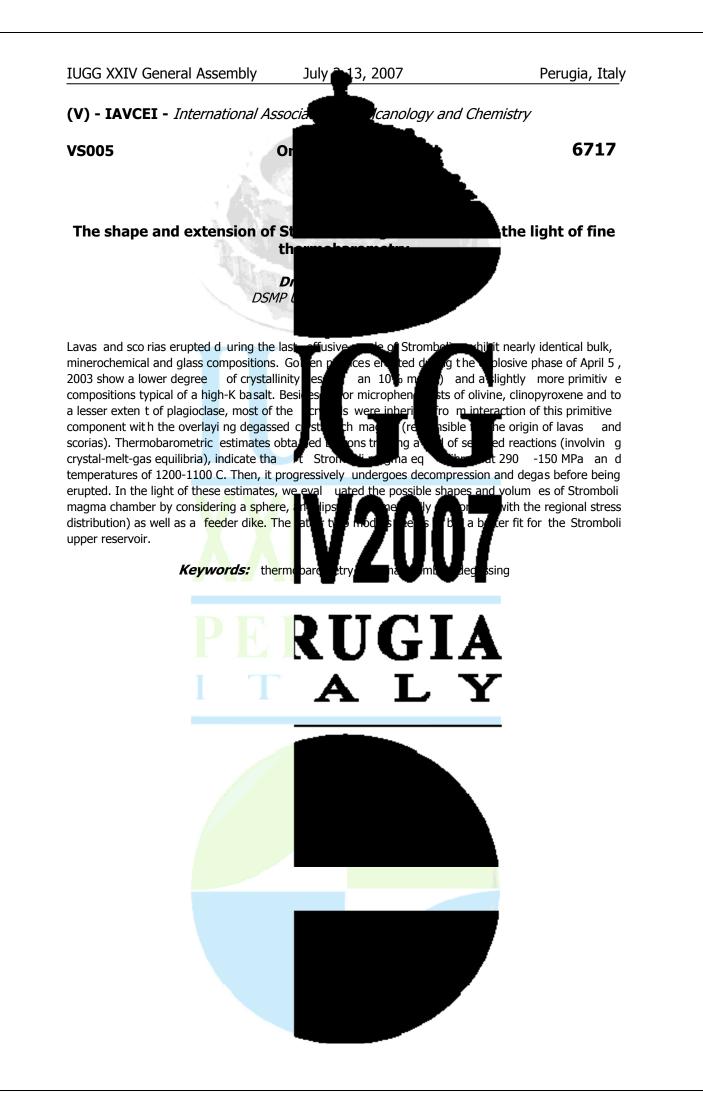


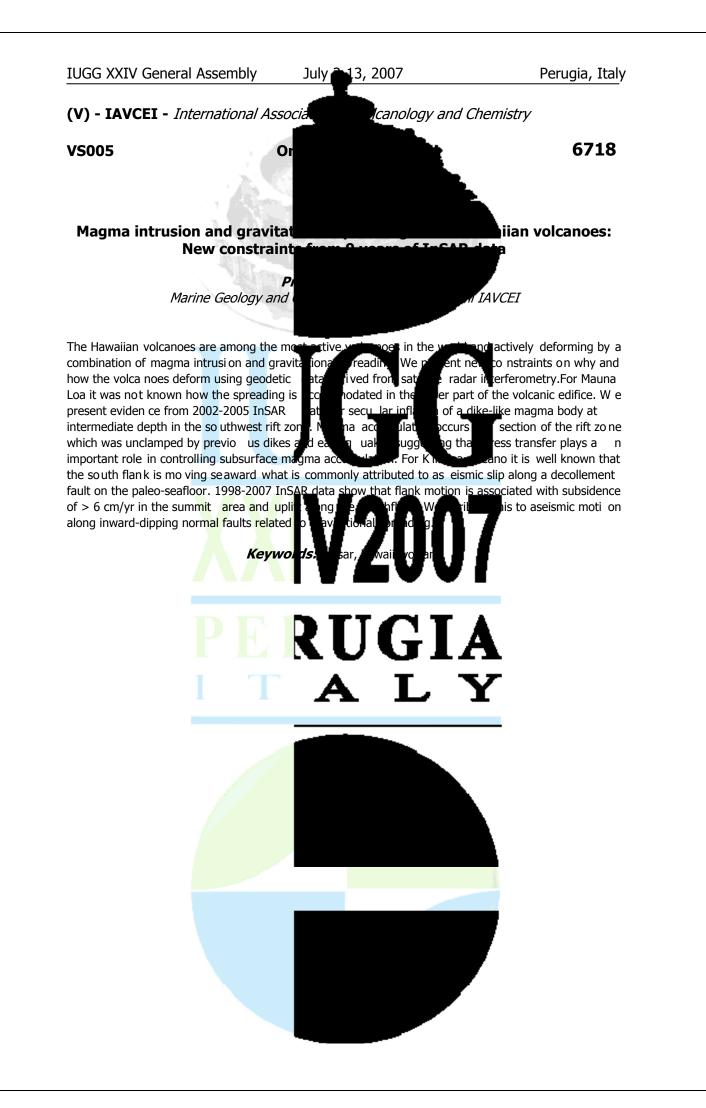


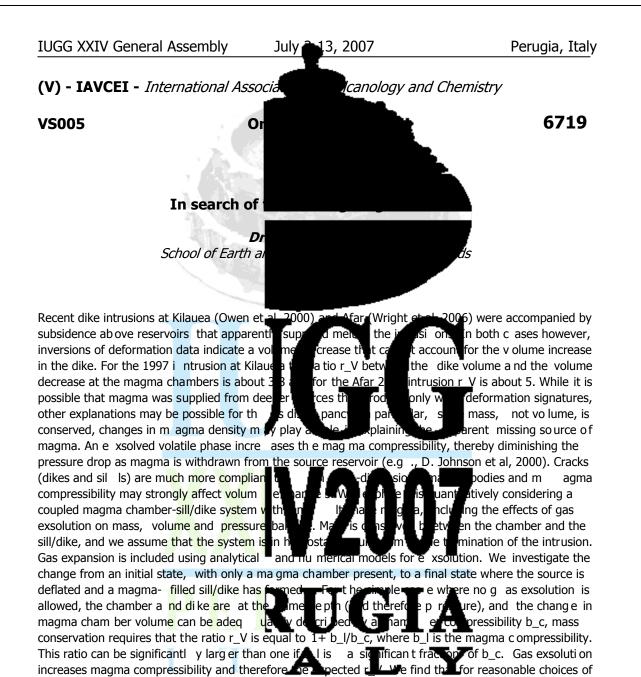
Etna by posing the question: are the voluties uslv a impli ed by SO2 flux ndog ed ma measurements consistent with the depth an eformation? The period between ude of rou m 1993- June 2001 was characterised by the of flank eru s, and is therefore chosen as the bS test period. In the following we present bot ets f<u>or that t</u> period, and then in our discussion examine if re asonable para meterisations o mpre nagma the integrated bul k litv modulus of the medium between the defo i) the geometry of the urface sour c d tl nat deformation source, can unify our observations in mine the implications of del. this analysis, with particular focus on the sources of deflation during eruptive episodes.



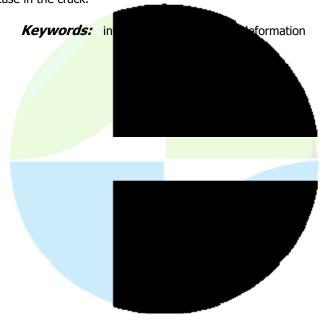


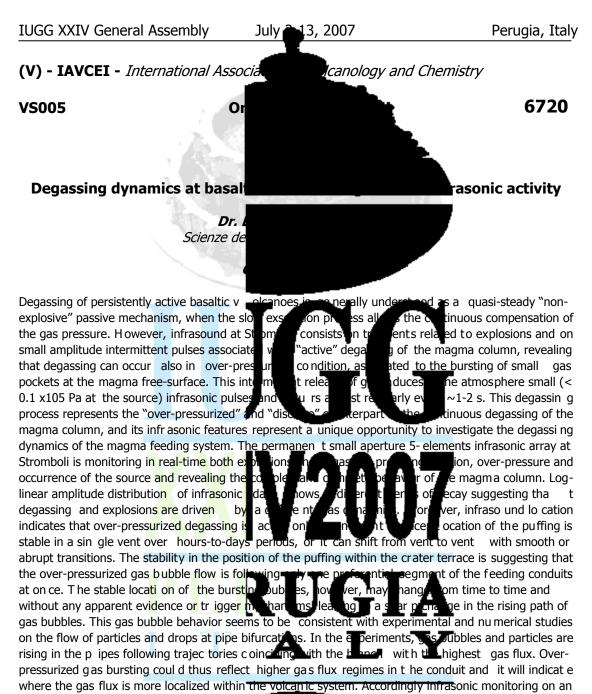






parameters volume decrease in the source reservoir can be up to one or more order of magnitude less than the volume increase in the crack.



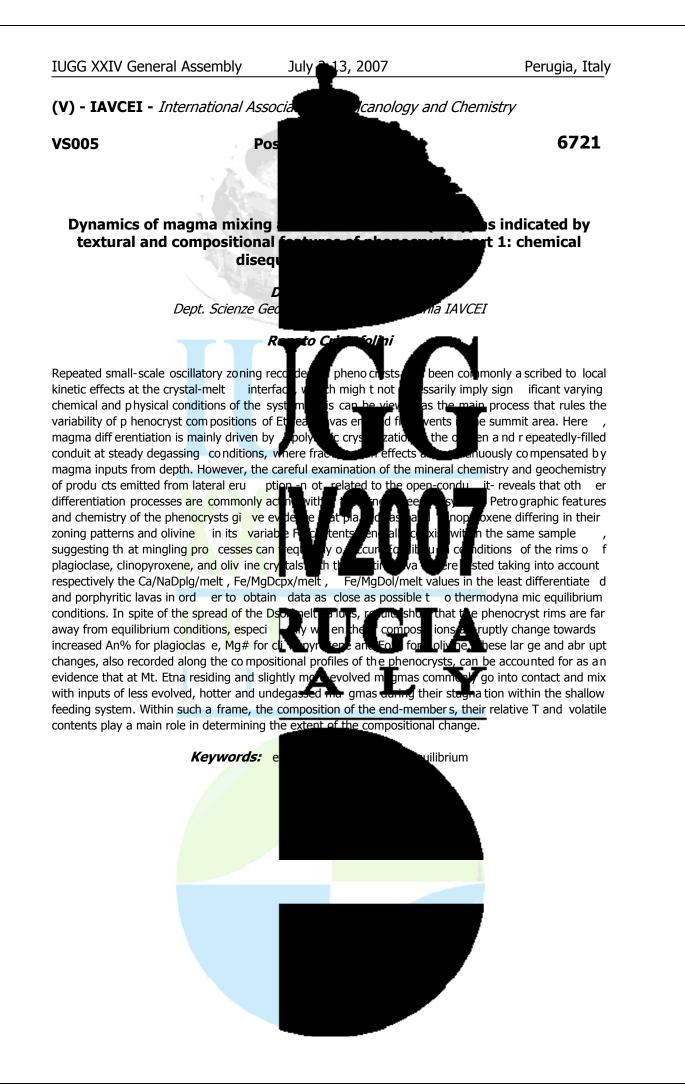


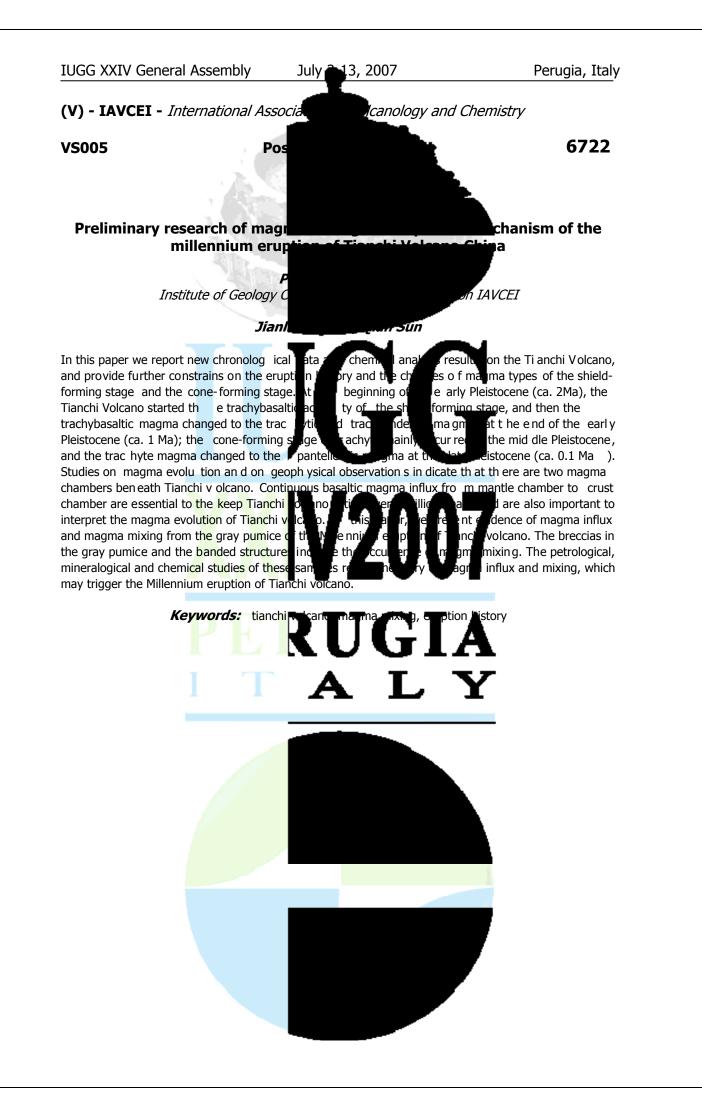
with but also to track changes in gas flux

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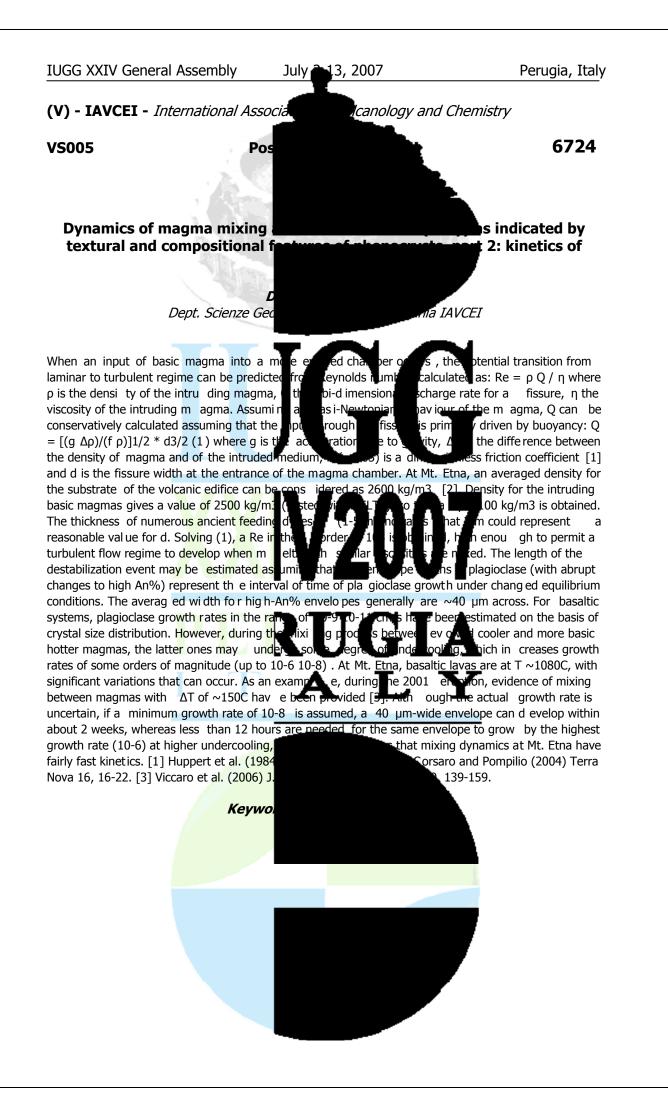
active volcanic systems would not only hel

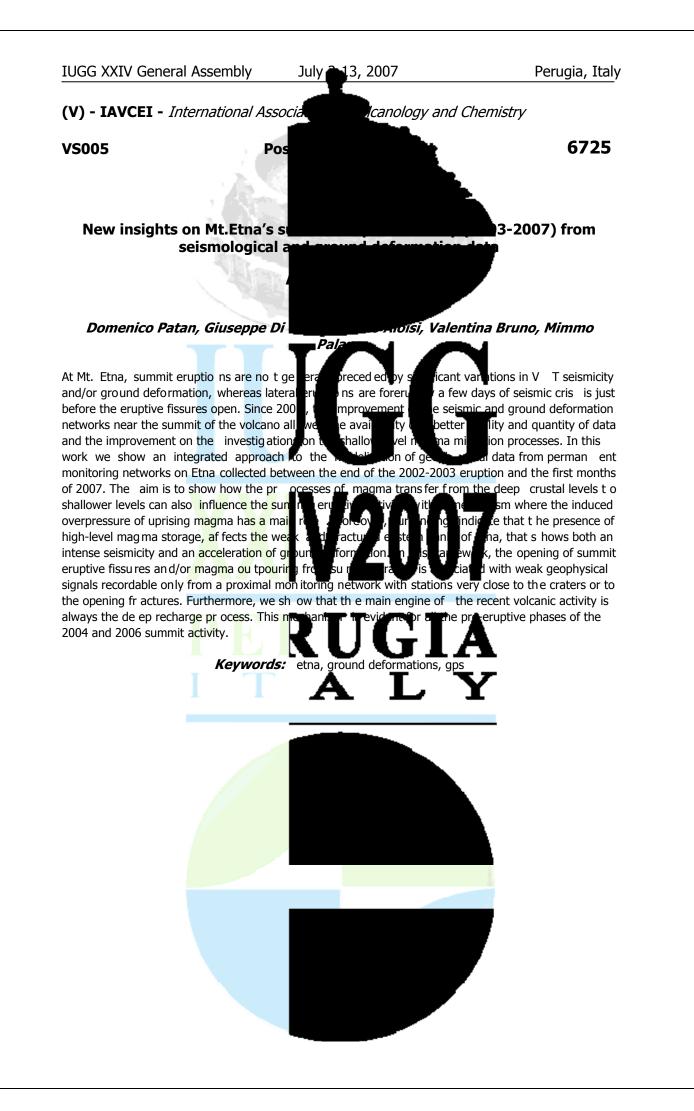
regime.

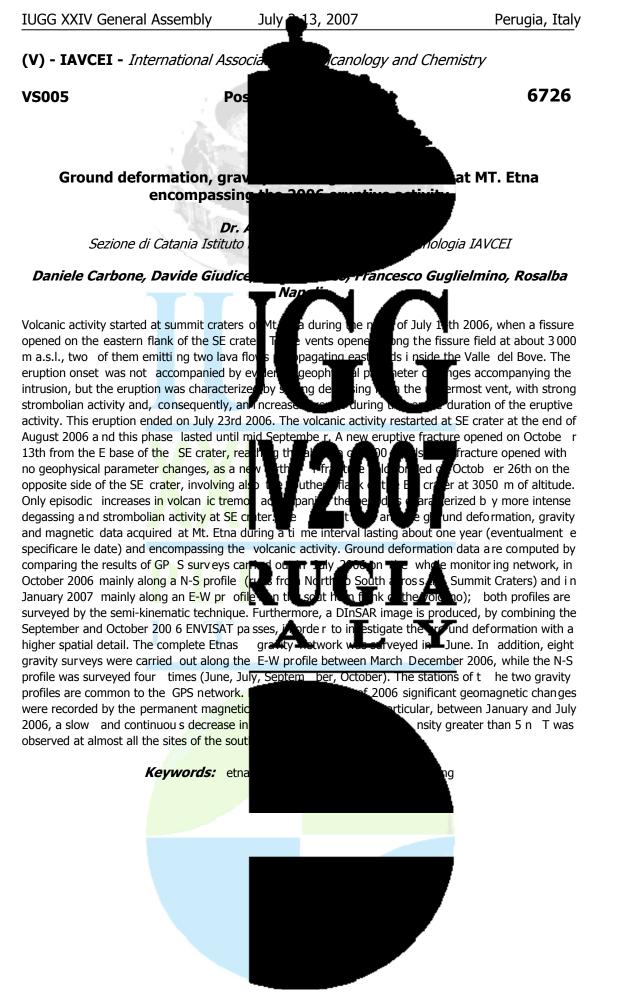


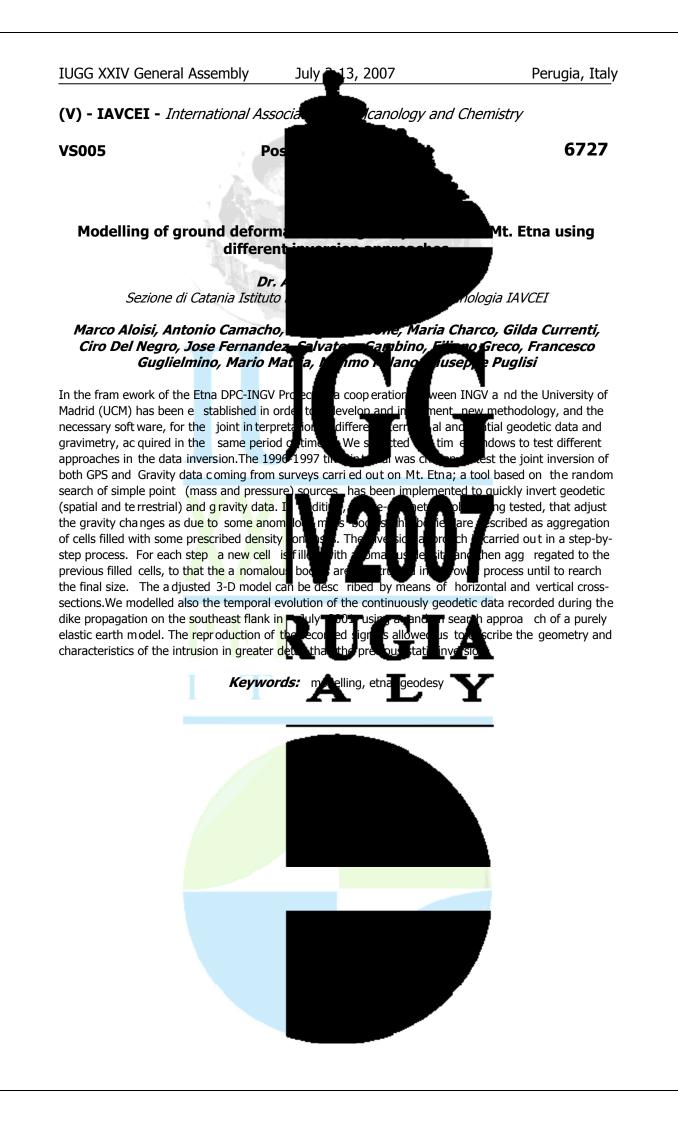


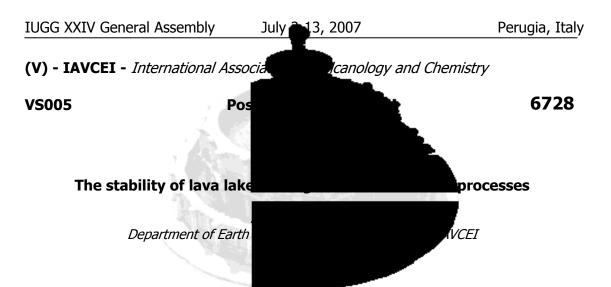












A physical model of a generic lava lake system is de existence of an equilibrium lava lake in which mag the pressure in the underlying magmatic r investigating the response of the system to system's press uredepth profile, to predict t simplicity of the modelled system, we find a through the magma is ignored. In this case are prone to s udden draining, are predicted. The

lake-conduit geometry, the s olubility and g as expansion I aws and the magma's volatile content. We show that an unstable lake must collapse <u>to</u> a new, stable equilibrium. Subsequent r echarge of the system by, for example, conduit overturn, to cyclic behaviour. Such a mech anism is o 'O'o eruption of Kilauea. When the rise of t that stable lakes must drain over time. W lakes, such as those observed at Mt. E rebu effective conduit convection mechanism or an ex

d. We deriv static ssure The stal lity ler ation. We dev ubse<u>quent b</u>e be hay ab le, the s

requisite conditions for the of the conduit balances his lava he system is tested by a graphical method, based on the iour <u>of the s</u> ystem. Despite the ally, the rise of bubbles and unstable lakes that bwn to be controlled by

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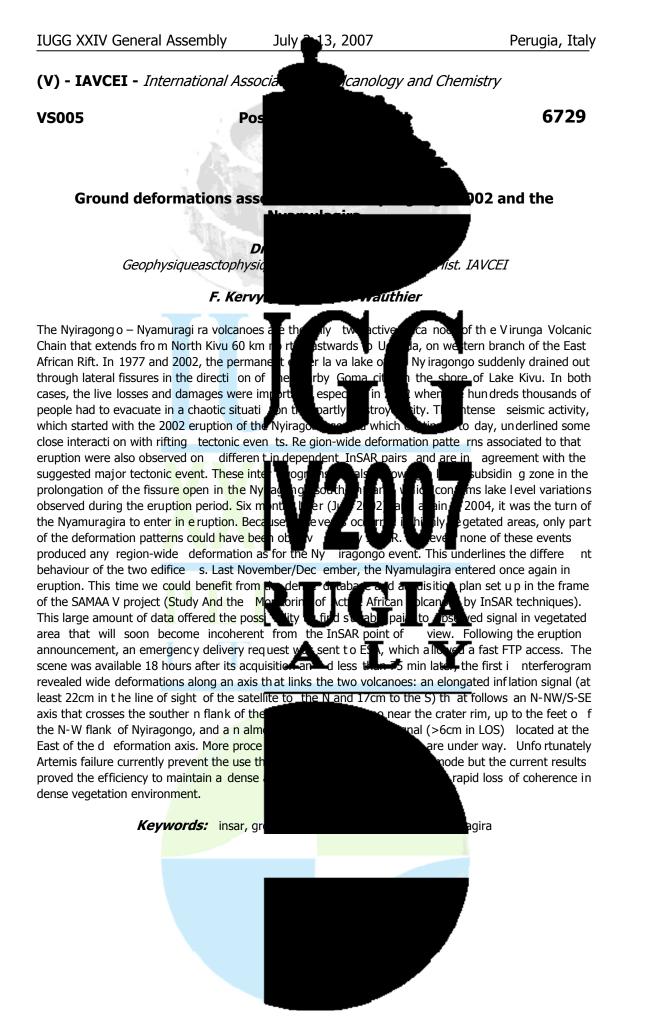
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of bubbles from depth.

equilibrium, giving rise ing the 19831984 Pu'u ered, our model predicts y degassing, stable lav a ea, Hawaii, must have an







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Symposium Calderas I - Calderas and resur

Convener : Prof. Giovanni Orsi, Dr. Sa

Caldera collapse and res urgence are comr interpreted in light of inflat ion and deflat significantly as well as their dynamics. Such volcano, including its magmatic feeding sy existing regional structures may influence stress propagation or may be reactivated during collapse and resurgence. The regional stre ss field may derectly shape of calderas and resurgent blocks. Th comparison of caldera colla pse and resur sedimentological, stru ctural, petrologi cal, g modelling, are solicited . Co ntributions on implications on volcanic hazar ds assessment for Field T rip F7 on the Neapolitan volca resurgent calderas will be illustrated.

Icanology and Chemistry

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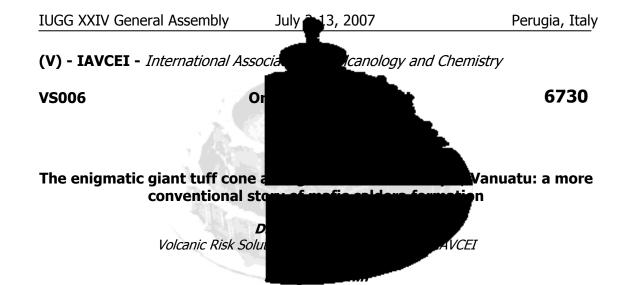
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crust, and regional stress regime. P renew structures, and the con the presentation, discussion and ontributions based on geological, as well as on v ariable types of Balderas, with obvious encourag ed to register pants bi Flegrei and Ischia tive (

oirs. Their size can vary

on both the characteristics of the



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Ambrym is one of the most voluminous activ km east-west elongated island, with a paral considered to be oldest and it is inferred t central part of Ambrym ~80 0 m above se volcano, M arum an d Ben bow. B oth volc. eruptions ranging fro m strombo lian event occasions (c. 30 yrs) larger sub-P linian and V ulc

volcanoes provide huge volumes of teph ra to the caldera, creating an exceptio nally large ash plain as well as sed iment-choked fluvial systems lea ding out from the caldera. Although it has been proposed that ~2200 y ears ago a cataclysmic phrea construction of a giant tuff cone. There is mapping. The giant tuff cone theory, he stratigraphy from a mosaic of sections loca mosaic of sections cannot possibly be related to Instead, we show that the island is a composite st ructure, formed of many generations of coalescing

monogenetic volcanic structu res. The type localities of the giant tuff co ne (dacitic) py croclastic flow deposits in the northern shoreline of the isaltered), phreatomagmatic fall and surge from the calde ra outflow and forming valley appears to ha ve existed for at least the la st c. 2200 years, and the se diment produced is gradually infilling a deep ly eroded older terrain. Appa rent naturity of the volcanic

formation of the present cal dera, and this structure through degassing and eruption from the central caldera vents. There is no evidence for a climactic (or for that matter dacitic) explosive caldera-forming event. Magma rise in the centre of the island from a long-term stable shallow magma storage degassed magma along lateral dykes is inf km wide caldera slowly for med. Hence w nessesary to explain the volcaniclastic facil Furthermore, the volcano should not be us explosive phreatomagmatism.

it consists of a 35 by 50 he northern part of the island i s coalescing stratovolcanoes. In the Idera occurs, containing 2 active non thly) small-volu me (wee matic lo sions. On mo re rare two frequently active

> era formation and event based on new

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1) mafic, (hydrothermally ther 🖌 fluvial se diment derived har te mor e prese nt-day Ambrym seem has r esulted in the til evolvine a response to mass loss

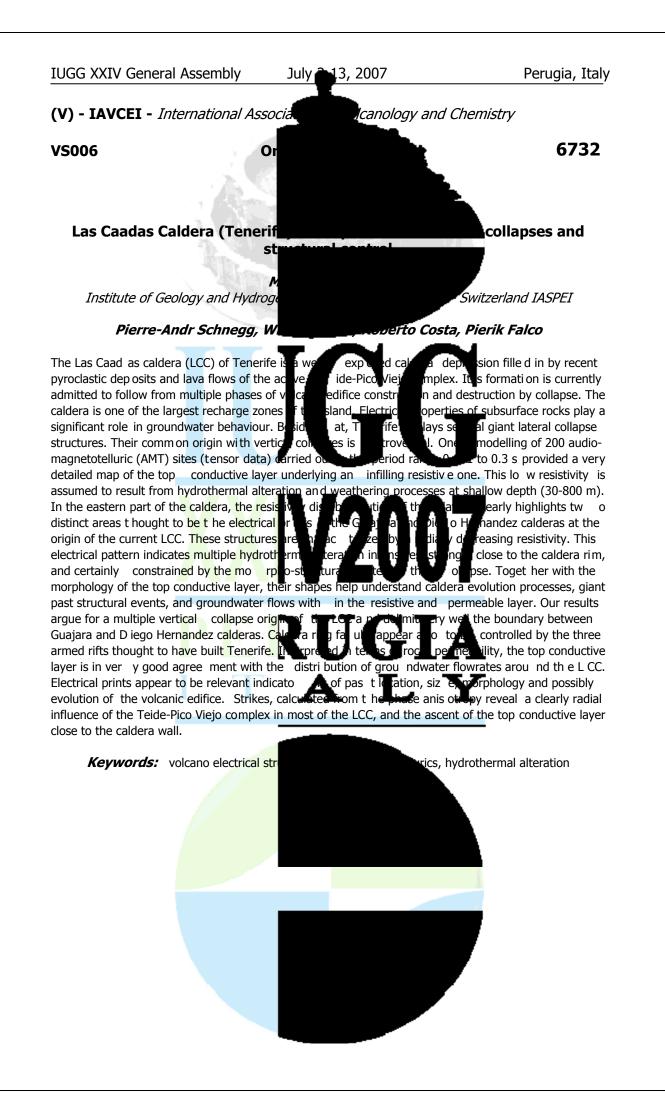
inte nse degassing and drainage of r m eans by which the central 12the giant tuff cone is not the formation of its caldera.

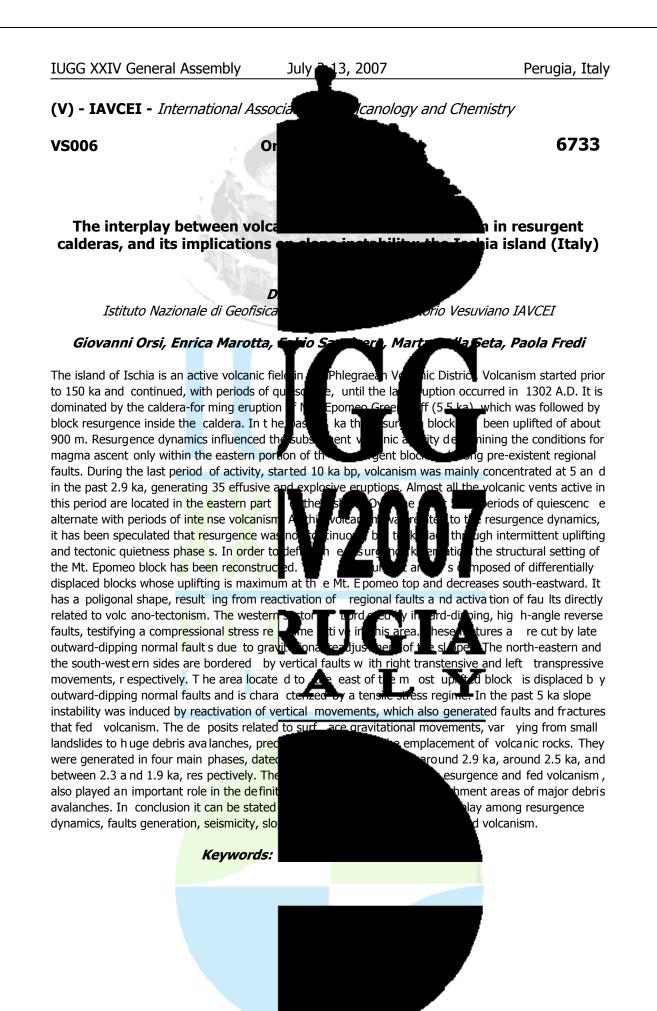
ve example for large-scale

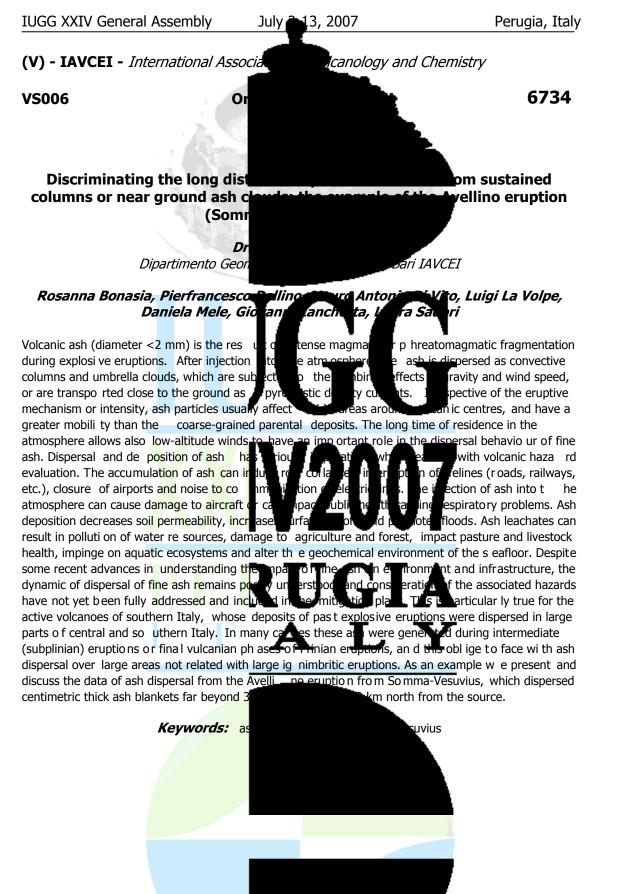
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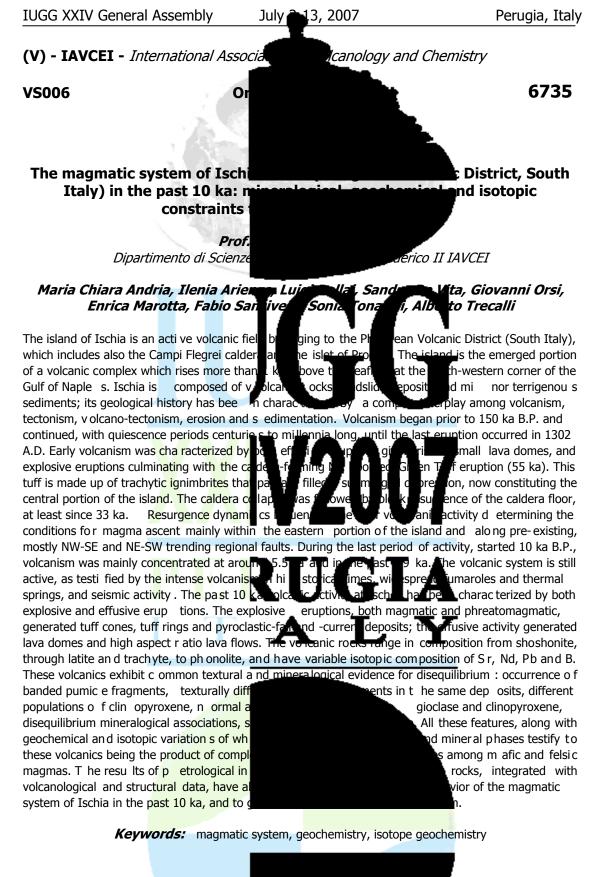


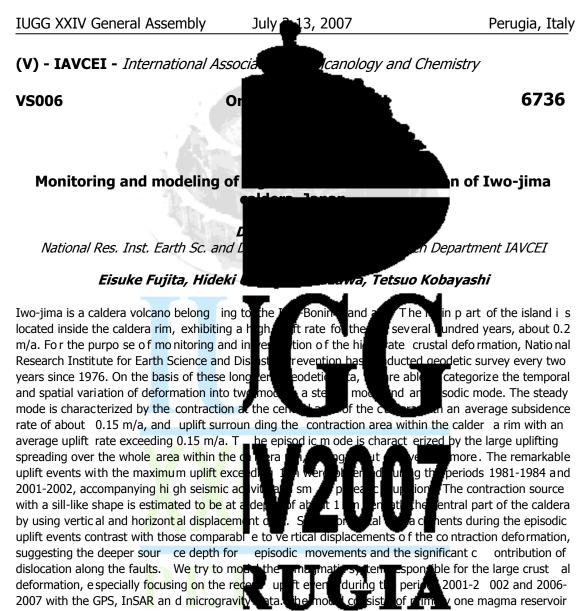




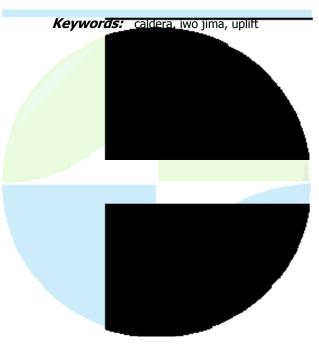


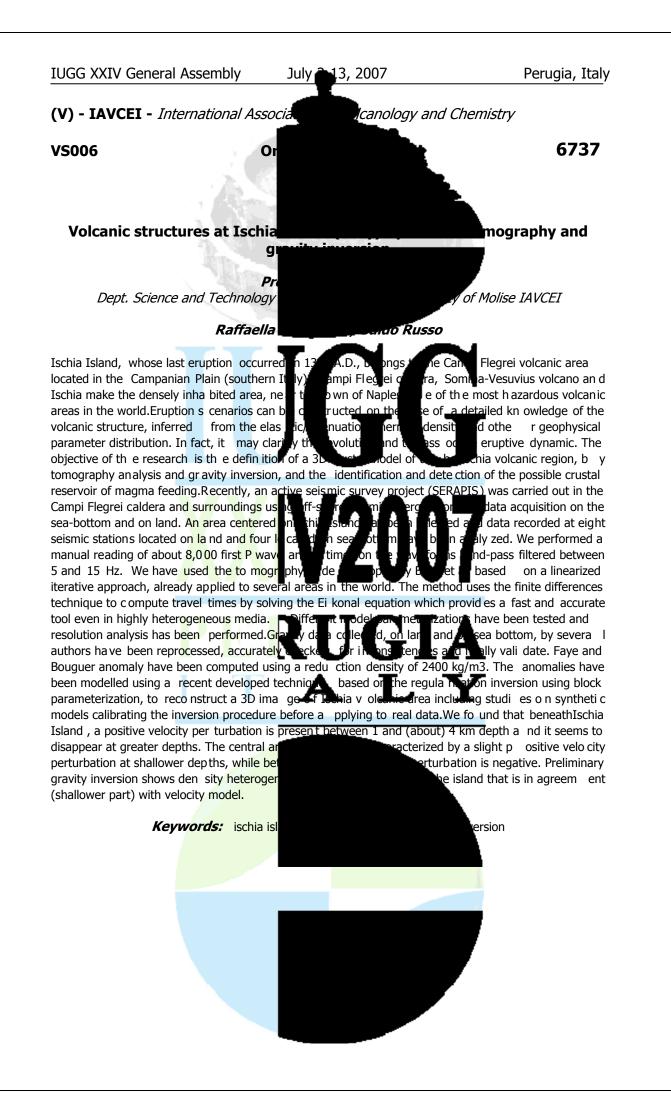


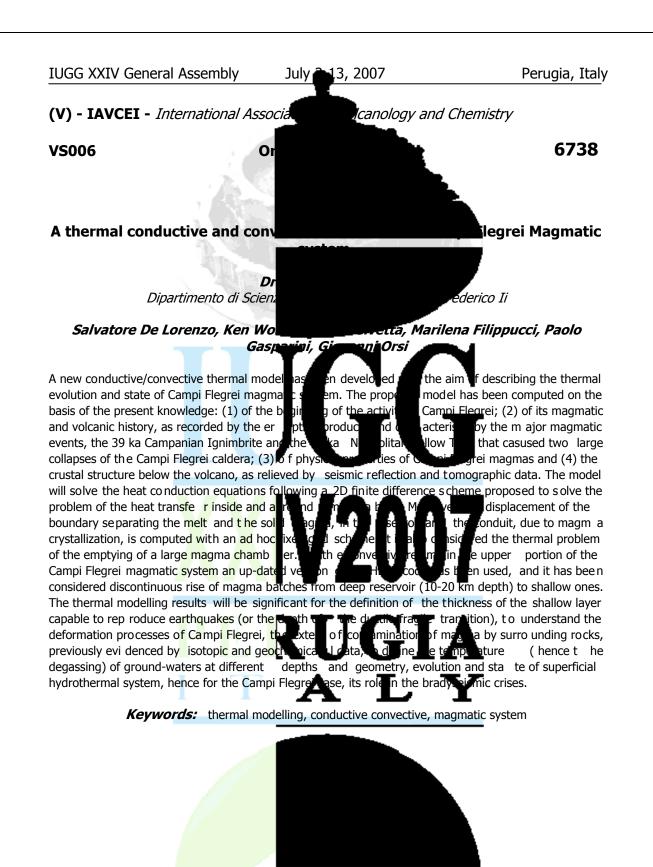


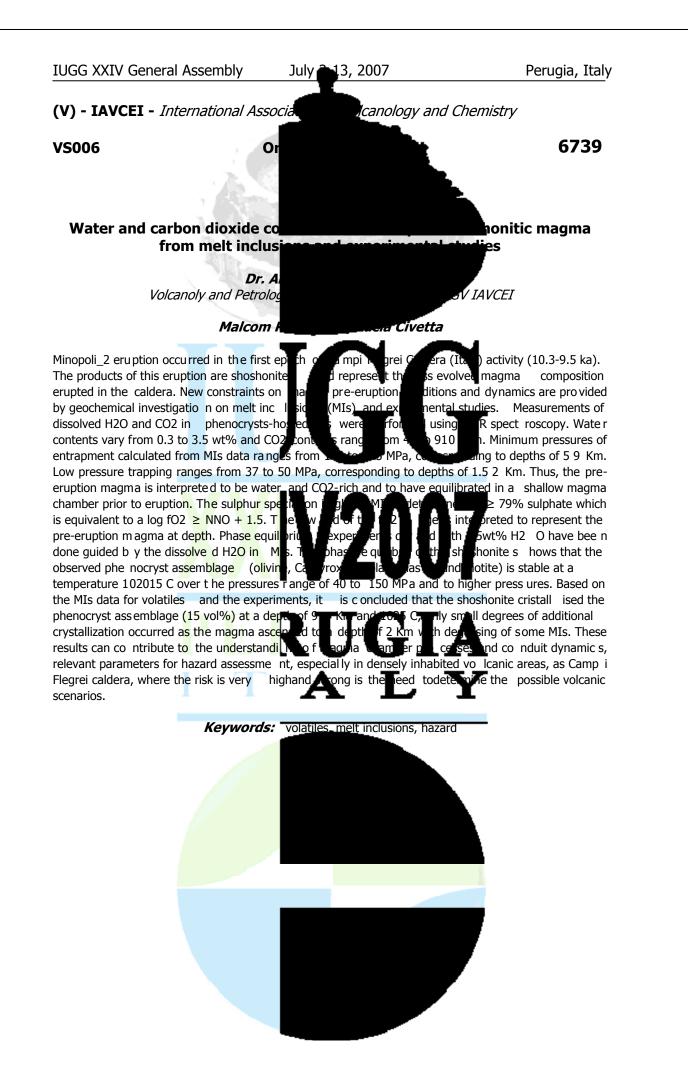


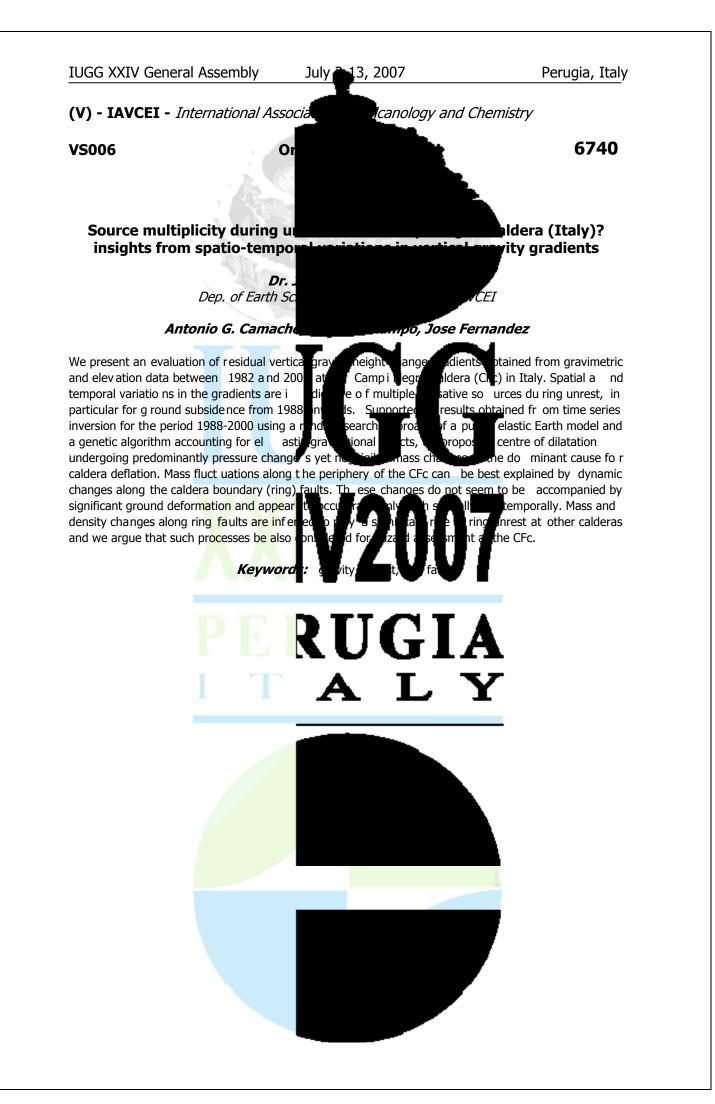
with an episodic magma supply at the deeper part. The top of the magma reservoir suffers contracting deformation due to cooling a nd/or degassing. We discuss evolution of the deformation modes and the possibility of eruption, taking account of available geological, geomorphological, and geophysical data.

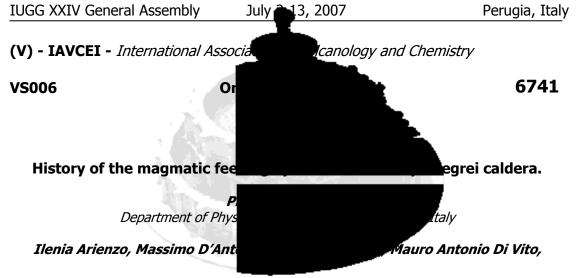












The definition of the magmatic feeding syste crystallization time-scale, re lationships betv position of the vents, and magma processe evaluation. Investigations aimed at defining mineralogical, geochemical and isot opic ana the Campi Flegrei caldera is characterized b reservoirs (20-10 km dept h) mantle-der ved ma

components, variably interact ing with the

through timeThe relationships between the

of activity, and the isotopic composition of

reconstruct the architecture of the magmat

composition of the magma feeding a future

continental cr ust. In the shallow

h of ie vo oes position of th of paramou Cam<u>pi Flege</u>r r, Nd, ep a nalle erentia

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chitecture, composition, fms č and structura I upted magmas importance for volcanic hazards gma<u>tic svste</u>m, include detailed natic feeding system of The r agma ervoirs. In the deep ontaminated by bre c

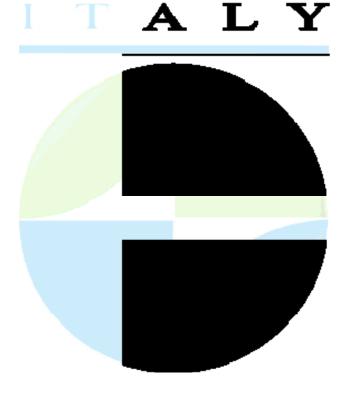
reservoirs is otopically distinct magmas, further

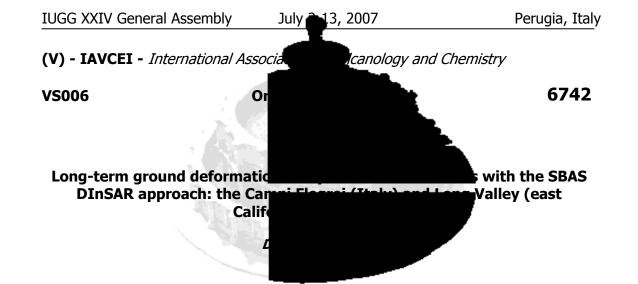


differentiated, contaminated, and mixed and mingled before eruptions. These processes generated isotopically distinct am pi Flegrei caldera hts, during the last 15 ka egrei caldera allow us to chemical and isotopic

on

Keywords: magmatic system, volcanic hazard, caldera





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We perform a ground deform ation analysis California) calderas that are localized in two Differential Synthetic Aperture Radar Interfe spatially dense deformation maps with centil the Small BAseline Subset (SBAS) algorithm and corresponding displacement time series algorithm permits to proces s data relevant to exte

volcanic source, and regional patterns associated to Flegrei caldera we used all the available El (track: 129 frame: 809) and descending (track: 129 frame: 809) and descending (track) time series extending from 2 002 to date; separate the vertical and horizontal comport revealed the start of a new uplift phase with the

renewed activity showing a maximum velocit y of about 2.8 mm/year. The shape of the deforming area seems to be modulate by main shallow stre the most recent unrest episode occurred in although with a significantly slower velocity been performed by using ER S-1/2 data acquired from a single (descending) orbit; the dataset includes 21 acquisitions relevant to the 1992-2000 period \_\_\_\_\_ e computed mean de deformation time series highlight three different de

of Hilt on Creek Fault (Sierra Nevada); a very larged deformation pattern affecting the overall caldera region. In this case the DInSAR results cle caldera have a maximum in corresponden sequence of three different effects: a 1992 a 19982000 subsidence phase. Moreover, us to map the extent of the zone with a t detected three-phases deformation pattern the mapped area clearly extends outside t has been dedicated to a comparative analy the two investigated calderas; this allowed us to point out some peculiar differences o n the detected

dynamic of the two volcanic centers. In particular, it clearly emerges for Long Valley caldera the primary role played by the magmatic source on the deformation pattern geometry. Indeed, in the Campi Flegre i case study our analysis reveals the signification floor that introduce a modulation on the de Rossi, M., Car mona, C., Ardagna, F., Peltze field of the Landers earthqua ke mapped b Fornaro, G., Lanari, R., and Sansosti, E. (

d the Long Valley (east texts. In particular, we exploit the ho logy that allo ws us to pro duce racy More specifically we applie d formation velocity maps meai nages e remark that the SBA S out 100 x100 km), thu s

te ctonic processes. For what concer ns the Campi rde nď 0

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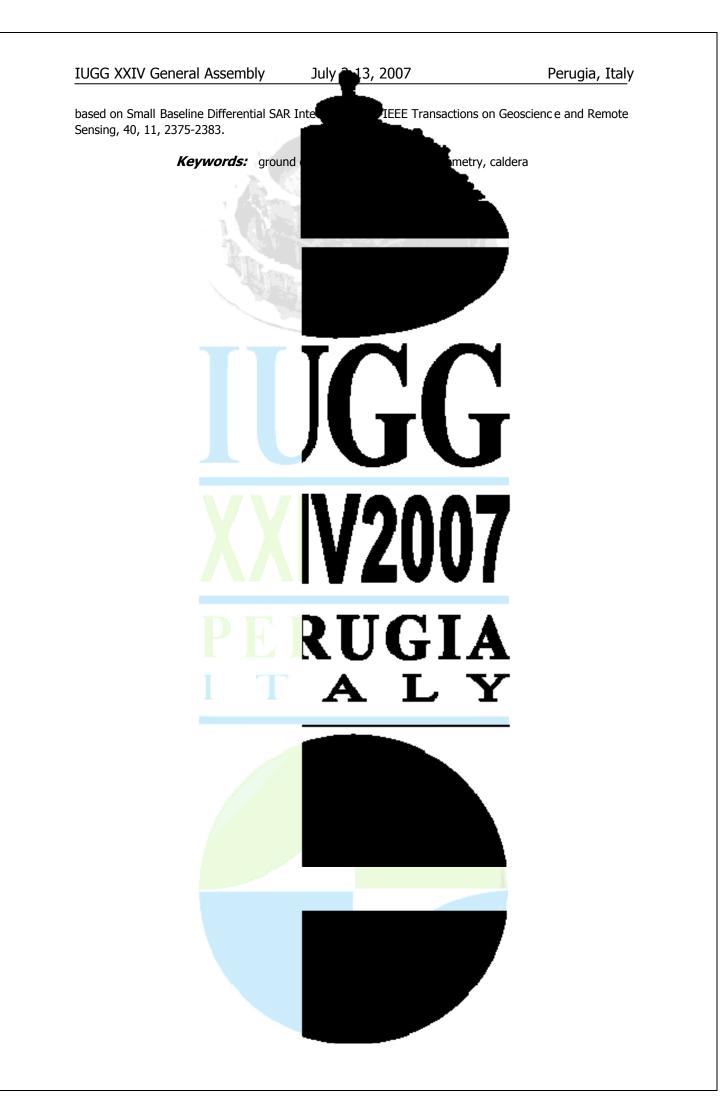
allowing to highlight the relationships betwe en local deformation signals, i.e., displace ments due to a ath I2 from ascending o generate deformation nd descending data to ty. The processed data

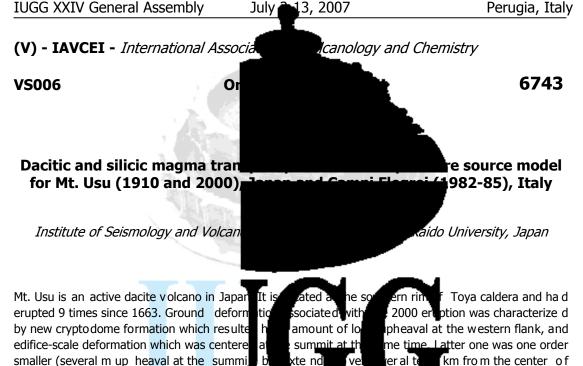
v clear starting from June 2005. The area of maximum deformation is localized in the centre of the town of P ozzuoli, with the also / how that with the respect asting for a longer time, enon (all / a h and su rroundings has mation velocity map and 🛃 a subsid g area localized in the volcanic Pahoa Island in Mono Lake; a deforming zone at southwest of the caldera, localized in the area splacement phenomena affecting the

e and are chara cterized by the 1998 unrest phenomenon and lacement time series allows ighly correlated with the e. We also showed that es. The final discussion n patterns detected for

res forming the caldera rences: Massonnet, D., 1993), The displacement 4, 138-142. Berardino, P.,

ace Deformation Monitoring





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smaller (several m up heaval at the summi deformation. Edifice-scale deformation excl flank showed sharp decrease of displacement with reversal (at around 6 km), a nd a small subsidence and contraction at further distances ( at 7-40 km). This concentric deformation pattern can be modeled by the simple com bined Mogis p ressure source

model with shallow inflation (2 km) and de be applicable for the past eruptions of Mt. can be explained also by the application o were estimated at both deep and shallow d magma transport from the depth to the sh deeper source corresponds to that of magina chambe r suggested by p etrological study. Twin source model with s hallow inflation and deep

deformations in the world. Deformation as subsidence over wide area. This was well e deformation data obtained in the 1982-85 Vertical displacement data rapidly decreases with distance at 5-6 km. This is quite similar deformation pattern to that of Mt. Usu. It was well ex

shallow inflation (3 km) and deep deflation (10 km) of dacitic or s ilicic post -caldera magma tr ansport exists between Mt. Usu and Cam pi Flegrei. Both deformation activities were characterized by the 108 m3 orders of ma gma intrusion from the dept h (around 10 km) and its signifi cant part rer Campi Flegrei (1982-85) was possibly induc (e.g. Dvorak and Mastrolorenzo, 1991; Bell 1914 eruption of Sakurajima volcano (Air important physical rules were obtained. Th transport from the depth, shal low doming activity. The application of Mogis tw in precould successfully evaluate these different

new btodome at the western eav (at ard m), uplift-to-subsidence

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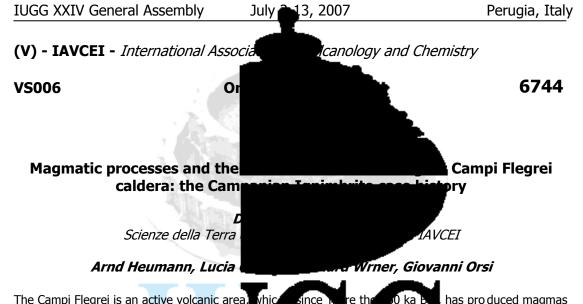
t intrusion process can -82. Each leveling data arly equivalent vo lumes 000. This model indicates da, 2007 ). The depth o f

deflation was applied to t he various types of volcanic on of ilauea showed all ground flatic ource. O n the contrary, Moais leg a g dera showed all uplift. mp plain by the simple Mogis two ressure source model; n volumes Common physical process (2-3 km). Resurgent dome growth at

onal new magma from the depth model was also applied to the eformation study, sever al cesses, which are magma without shallow doming tion and deep deflation

*Keywords:* usu, campi flegrei, ground deformation





The Campi Flegrei is an active volcanic area, which with variable chemical and Sr isotopic compo itio shoshonitic products. The most striking strue caldera resulting from collapses following the Yellow Tuff (N YT-14.90.4ka) eruptions. The largest of the Mediterranean area in the las 200 and Th isotope ratios and concentrations have be

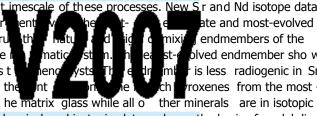
aimed to investigate the mag matic processes that occurred in the Phlegr aean magmatic system before the Campanian Ignimbrite eruption and the Limescale of these processes. New Sr and Nd isotope data on glasses, minerals and whole ro cks re erupted magmas, are combined to recons trachytic to phonolit ic Campanian Ignimbr equilibrium between host glass and the me and Nd than the most- evolved magma. O evolved samples are in equilibrium with t disequilibrium. On the basis of these geo

results, we hypothesed that the most-evolve less radiogenic pumices, represent isotopi mingled before and during the Campania relatively constant U/Th ratios and are c 238U/234U isotopic ratios measured in wh

isotopes are not affected by chemical alteration samples yield ages of 433 230Th/232Th initial ratios for these samples is similar within the errors; the weighted mean value is 0.88 0.02. If our isochrones reflect the c indicate that (1) the CI least-evolved mag the end of its crystallization, i.e. between the two Campanian Ignimbrite endmember information recorded in the Campanian Igr between magma batches could also happe

e th ince with the ligh ement of the ania<u>n Ianim</u>ł an ian ' mb dep os Measu ent ed on t

has produced magmas atios detected in the least-evolved pi Flegrei is represented by a wide (39.280.11ka) and the Neapolitan 200k m3 DRE) is the Şr and isotope ratios and of U hian Ignimbrite products



ate and most-evolved endmembers of the blved endmember sho w is less radiogenic in Sr roxenes from the most chemical and isotopic data and on the basis of model ling

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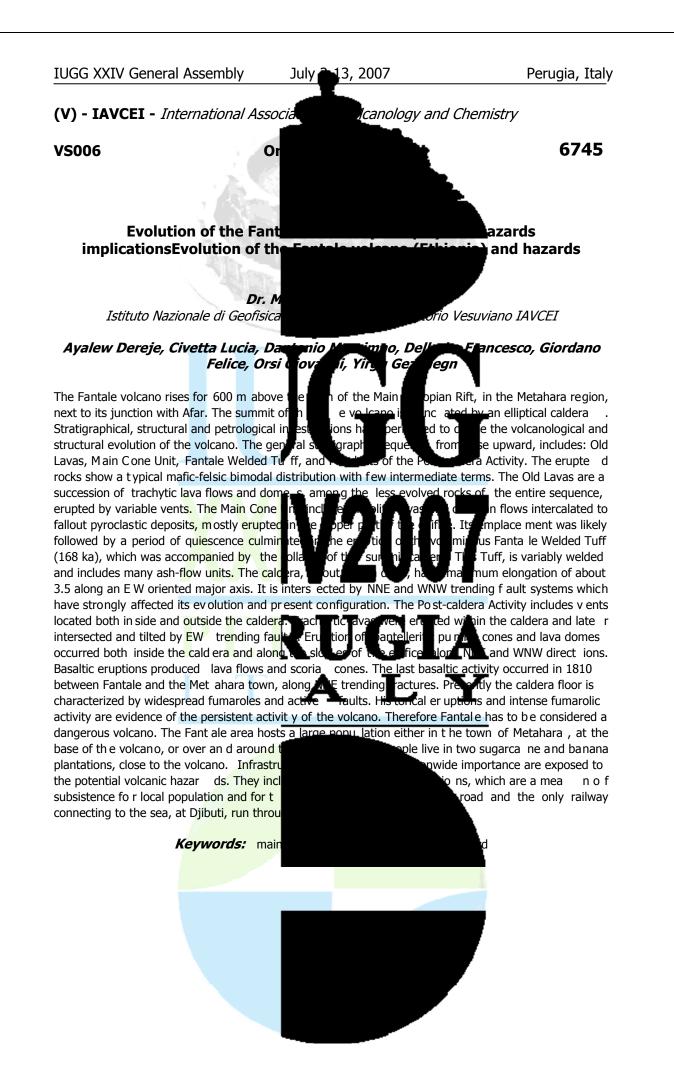
nices, and the least-evolved and ent n mas which mixed and diff on whole rocks show rier da haracterized by an excess 238U relative to 230Th. The and se cted minera Mow that the U-serie ternal m nd isochron for thre e least-evolved ka, 463ka a nd 487ka, the weighted mean value bei ng 452ka. The

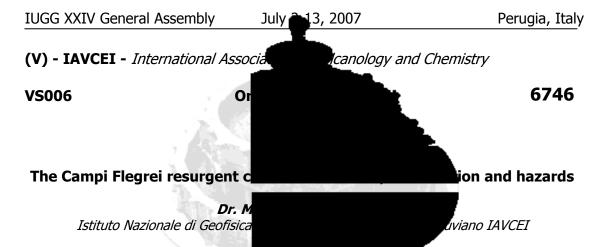
> the analy zed phenocrysts, then they tively short residence times since (1) the mixing process between lowing to preserve the time ggesting th at coalescence eruptions.

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Orsi Giovanni, Dellerba Francesco

The Campi Fle grei caldera is a nested and r whole structure is subsiding, while the cen occurring through a simple-shearing mechar vertical uplift of about 90 m. After the last c of activity, alternating to p eriods of quies intersection of two fault systems delimiting vents in each epoch is a go od tracer of the active

surface through time. The great majority alternating phreatomagmatic and magmatic explosions. They were low-t o medium-magnitude events, except two hig h-magnitude events, one in between the I I and III epochs, the stress activity of the system and its explosive ch such a high ha zard and inten se urbanisati affected by deformation over the past 2 k since 1969. These short-term deformation events

ductile and a brittle component, both generated by pressure and tempe rature increase in the magma reservoir due to arrival from depth of smal magma. The s triking similarity in shape bet resurgence dynamics is not changed sinc Furthermore the long-term results from the summation of many short-term deformation events. The collected data have permitted to perform a volcar and deformation history of the caldera to be tree hypothesised. Although it is impossible to d efine the time when an eruption will occur, according to the

reach the surface, have been defined. Accord dynamics of t he caldera, the areas within variable eruption scenarios have also been been represented in volcanic hazards maps

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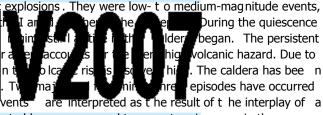
vo major collapses. The

Caura, Marzocchi Warner,

aldera is af fected by res urgence mation has generated a maximum was concentrated in three epochs d in 1538 AD at th e areal ibution of the eruption ured magma upraise to

of the re cognised eruptions were explos ive, almost all

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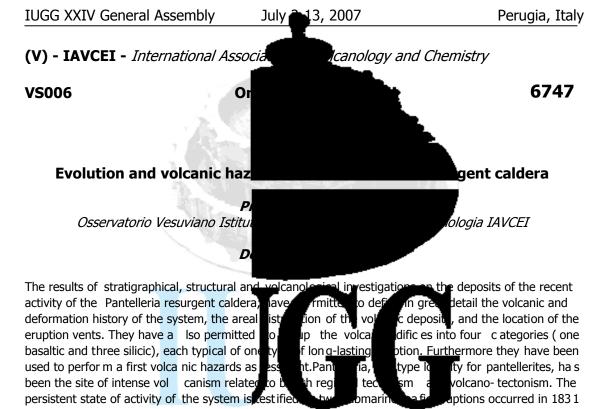
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During the guiescence began. The persistent volcanic hazard. Due to . The caldera has bee n episodes have occurred

ed and hotter than the resident d ef nation suggests that the e II and III epoch. etw/en

hazards ssessment. In portion of the volcanic deration for **T**uch a pur pose, has been nto con dynamics of the caldera deformation, the structural conditions which have to be acquired for magma to distribution in the past a nd the ongoing Ild open have been located. The ritory of such a scenarios have



and 1891 in the vicinity of the island, and by an in tense fumaroles and hot-water springs activity. The last caldera-forming eruption occurred about 50 ka bp and generated the Green Tuff. The subseque nt history has been subdivided into six silicic distribution of the ve nts ac tive during each

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structural setting of the island is defined b b faults system divides the island into two se t o major volcano-tectonic features include cal era Such a resurge nce, begun aft er 33 ka bp, nas genera teo uplifting and tilting of the Montagna Grande

erupted in the last 10 ka trough NE-SW align genetic relationships between resurgence na assumed as typical of I ong-lasting eruption re the pantelleritic Cuddia di Mida Te phra and pumice cone, Cuddia Randa zzo partially co llapsed edifice, and Cuddia Sciuvechi shield volcano. The results d attempt of volcanic hazards assessment in case of n it is possible to define the portion of the known volcanic history to be taken as reference for volcanic hazards assessment. Furthermore it is possible to put forward some hypotheses on eruption scenarios,

vent location, time, and characteristics of a of some of the past eruptions as reference evolution of some volcanoes as reference f

Keyword

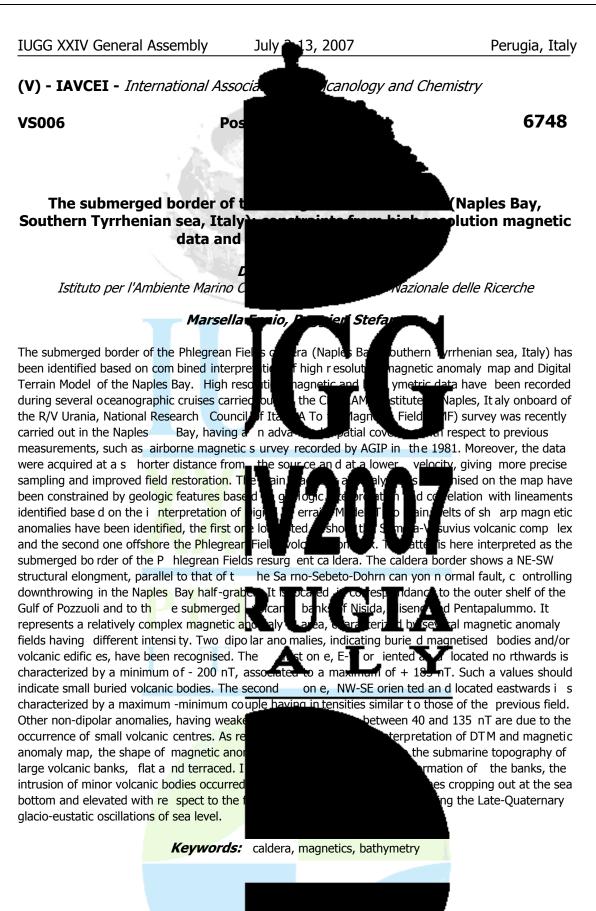
saltic eruptions. Areal ctural features.Th е ic lineaments. A NE-S W rustal discontinuity. The in the youngest caldera.

block, through a simple- shearing mechanism. The north-western side of the block is buried by rocks of these vents clearly indicate a ected volcanic edifice s,

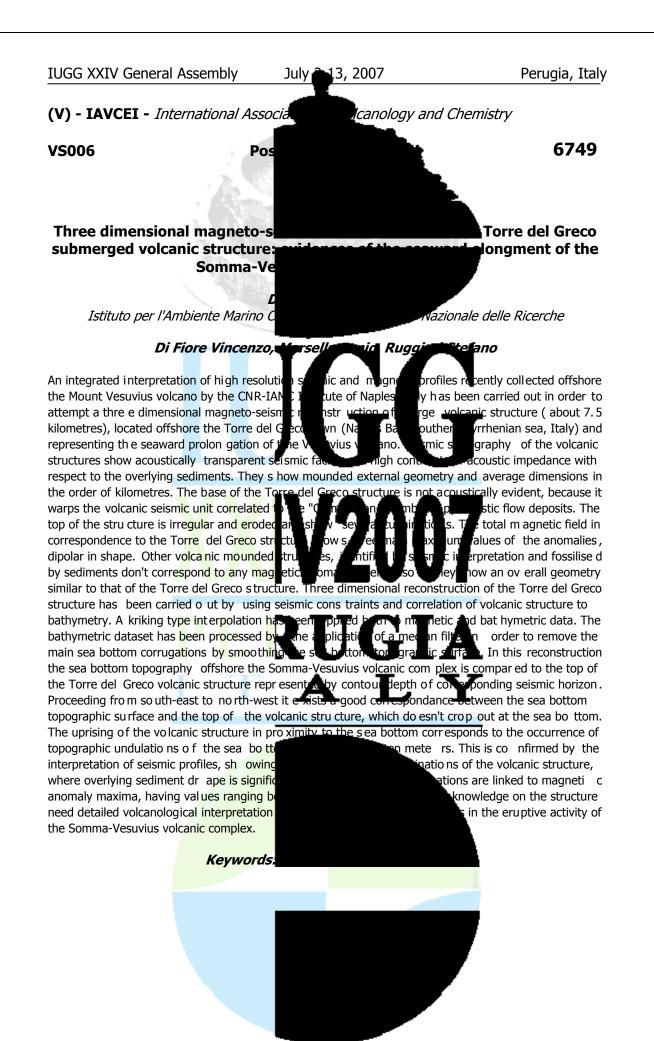
he. blcarlism nd , rsia, es and I ava flows, and rried out ha permitted to make a first t-mid terms. In particular adism in sb

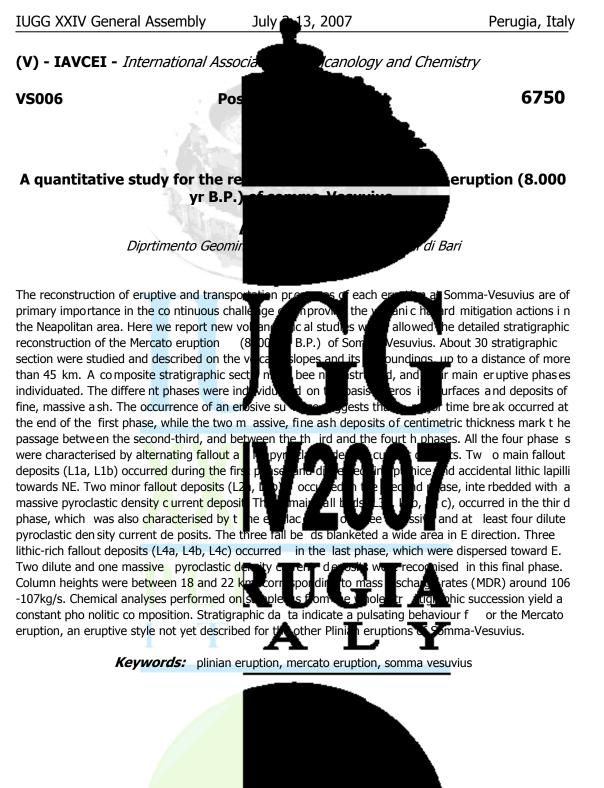
ssible also to assume the characteristics

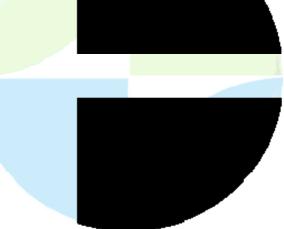
of variable magnitude, and the

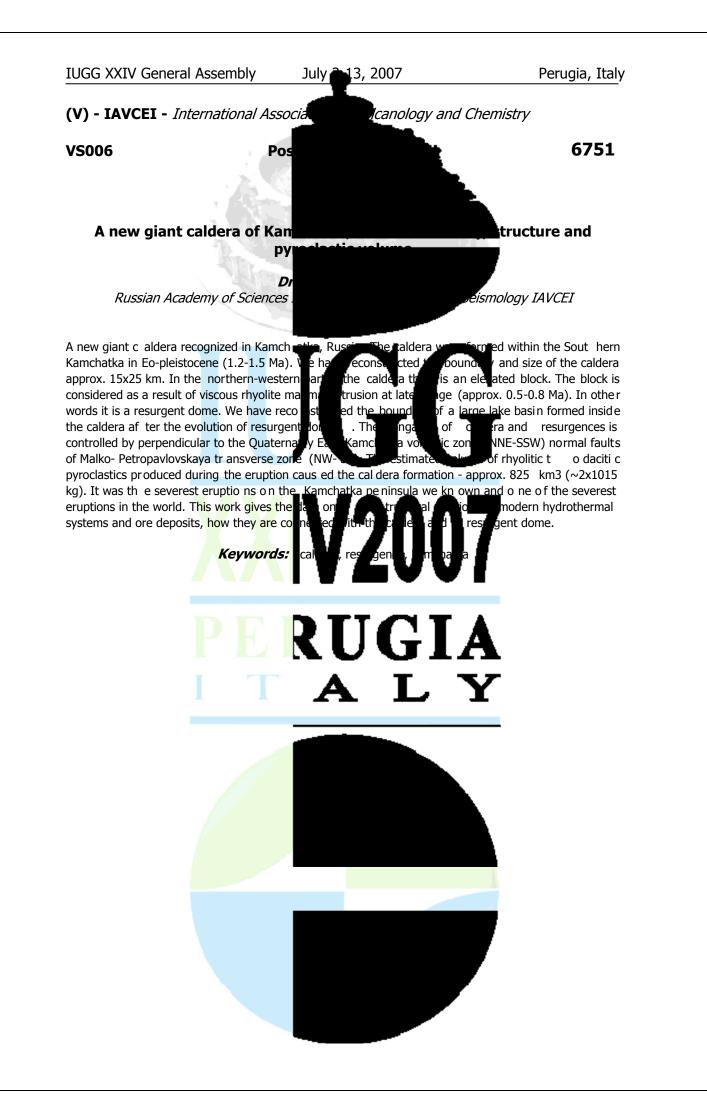


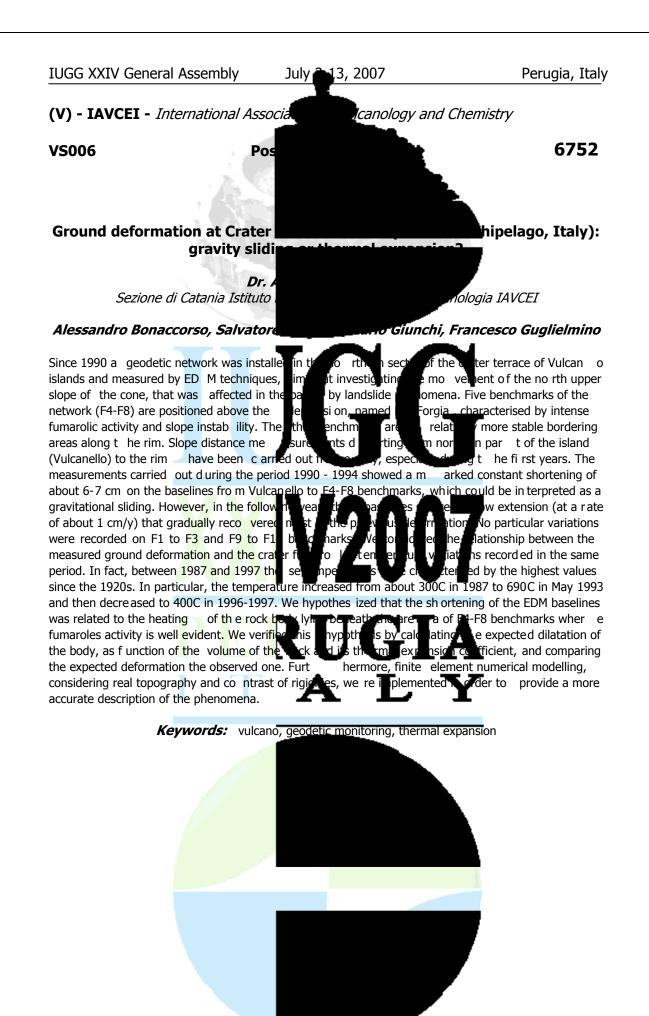
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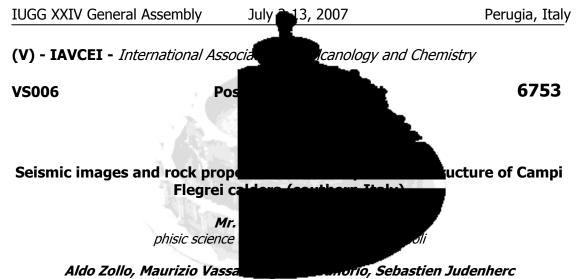












The volcanic area of Campi Flegrei has been activ eruption of Monte Nuovo in 1538, by the 1 activity, and b y the hot springs. In Septe carried out in the gulfs of Naples and Pozz shallow crustal structure of the Campi Fleg depth (the Serapis Project). The present st and tomographic images bas ed on the Selapis da

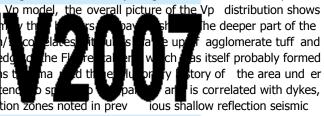
obtained by the 3D data gathering and thr Pozzuoli Gulf (z<1000 m).From the refined\_ the presence of a complex arc- shaped and anomaly (beneath 700 m, with Vp>3 500 interlayered lava, which form the southern following the two large igni mbritic eruptid study. The upper part of the anomaly that end volcanic moun ds and hydrothermal alteration zones noted in prev

the ent r an d1982-8 001, an exter ith th<u>e aim o</u>f era, a 05 des nt a is achi

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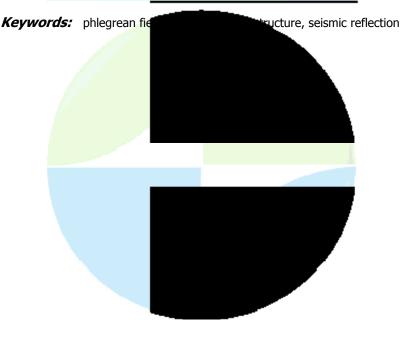
lustrated by the historic diseismic rises , by the fumarolic active-seismic in vestigation was estigating and reconstructing the ide nt g its feeding system at is of seismic reflection data flection s eismic sections

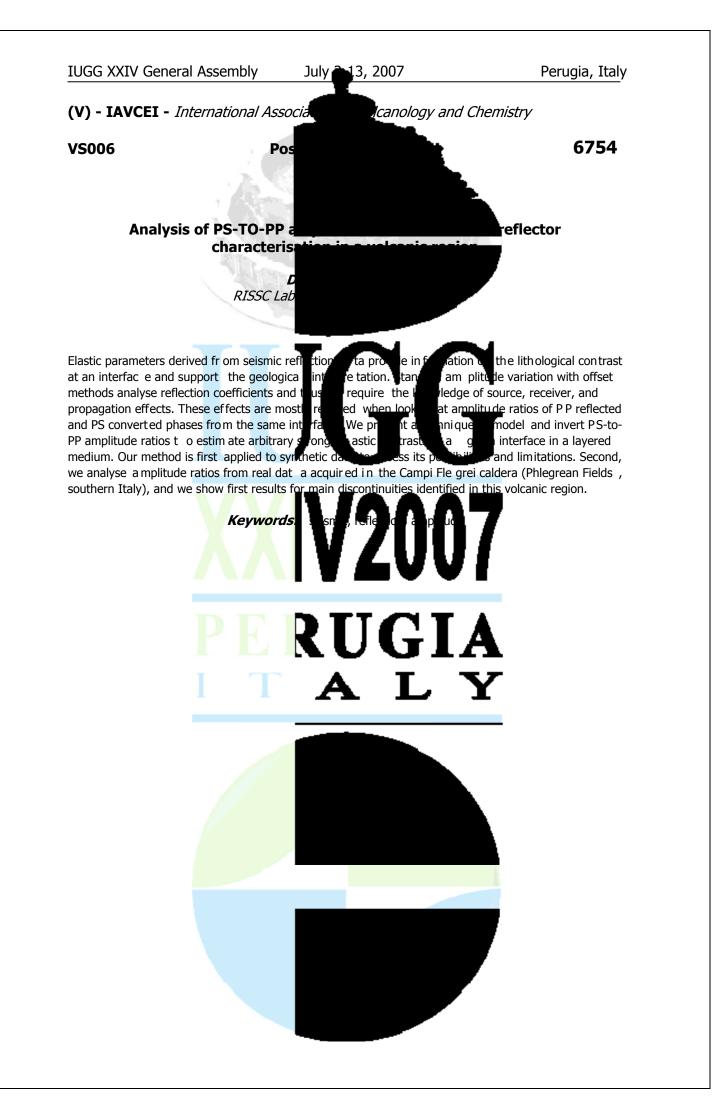
ough. refined Pvelocity imag es of the shallowest layer of



he deeper part of the agglomerate tuff and as itself probably formed tory of the area und er is correlated with dykes, ious shallow reflection seismic

analyses. The depth of the transition between the upper and lower parts of the anomaly is characterized by an abrupt Vp increase on the one-dim tt ed from the 3 D tomographic model and by the presence of a strong refle /0 .7 'ıt O vo Way T ime (TWT) on of the POP and P-S reflections at Common Mid Point gathers. The move-o ut bcit the layer bottom allowed to estimate relatively high Vp/Vs values (3.5 0.6). This hypothesis has been tested by a theoretical r ock physical modelling of the Vp/Vs ratio as a func of por osity suggesting that the shallow layer is likely formed by incohere water ted, volca ic and marine sediments that filled Pozzuoli Bay during the post-caldera activity.





## IUGG XXIV General Assembly July 13, 2007

Perugia, Italy

(V) - IAVCEI - International Associa

**VS007** 

## Symposium Calderas II: Calderas and calde

**Convener :** Dr. Gianfilippo De Astis, P

Calderas involve the rapid er uption of ma different tectonic environments. Some cald historical record and may have had large have bee n act ive fo r hundre d tho usands

volcanological research is aimed at under standing, the c onditions favo caldera-forming eruptions, how the magma related to the style of the causative eruption caldera-forming events, we are dependent we be able such data to forecast large calde the diverse methods, including field observa monitoring, and physical and numeric al mod caldera-forming eruption, its temporal evol associated subvolcanic magma reservoir.

ser in tin **Teve** g dired La bbse eol ogic reco hing eruptions tre nat cai lp d the tion an

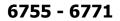
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contrast, other large volcano es ndera- forming eruptio ns. Much recent ring the o ccurrence of large the resulting caldera is and ion of uprest that preceded large precursoly eruptive activity. Wil I he future? This sess ion focuses on oc h gic studies, geoph ysical conditions leading to a rstand

**V**entura

e resulting caldera and



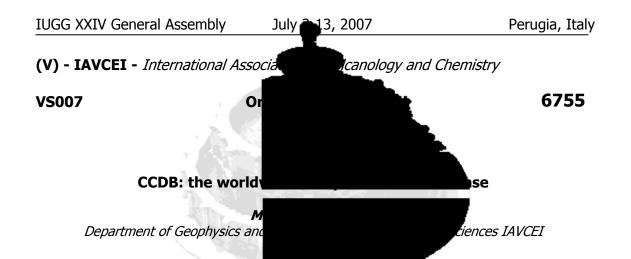


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Icanology and Chemistry

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Field studies constitute the most import nt wav processes. The reconstruction of past collap examples is a powerful tool t o understand t of this work is the elaboration of a worldwid useful and accessible tool for studying and records different types of information includi magma chamber and of the extruded depo associated local, regional an d plate tector cs.

published field studies of collapse calderas. More information has been summarized in a database link application. Thus, it is possible to visualiz according to different attributes (e.g. age, CCDB is to update the curr ent field base ld'

databases and complementing them with database does not include all the calderas further studies and analyses. Since communication and data exchange between scientific groups is still

an open and accessible tool for collapse @ elaboration of the CCDB web page. We are in order to streamline data and knowledge exc the database.

es igate an densta nd collapse cald era their paris anisms d coll se Caldera Da standing calde of the 50 post-d ſeforme

with present observe d calderation. The objective se (CCDB). The CCDB should be a ollapse processes. The datab ases dera, pr operties of the ra volcanic activity and ehensive compilation of

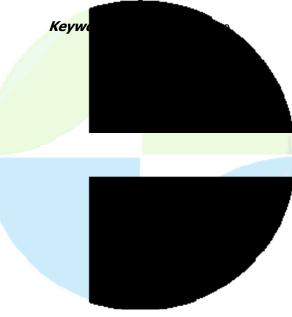
than 200 references have bee n revised, and t heir

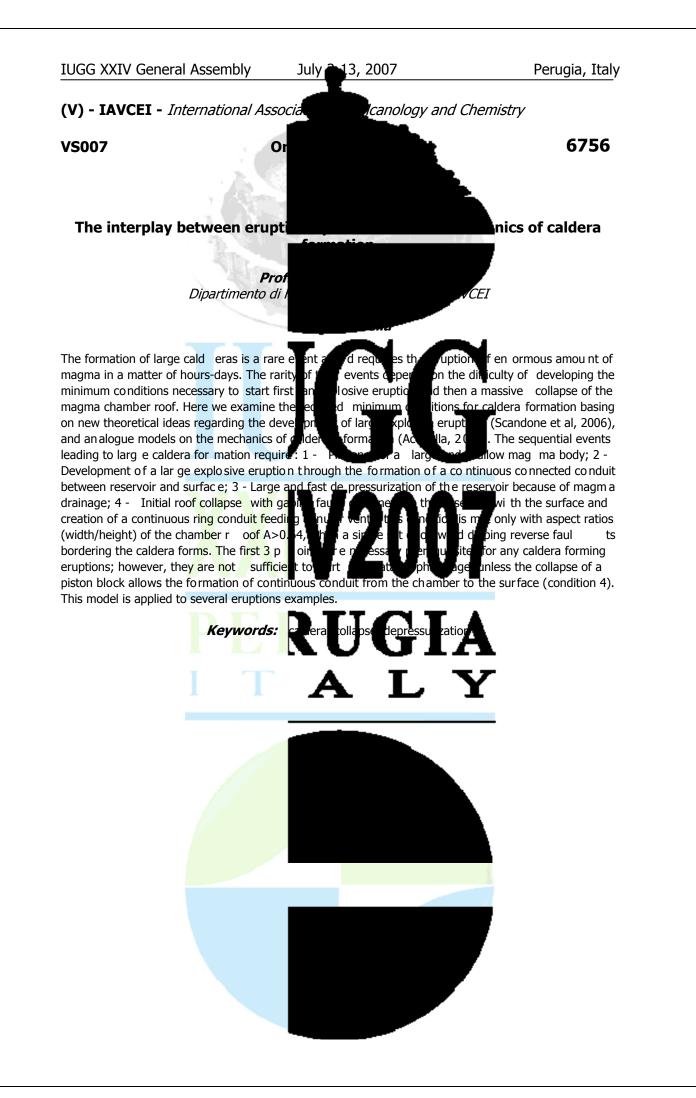


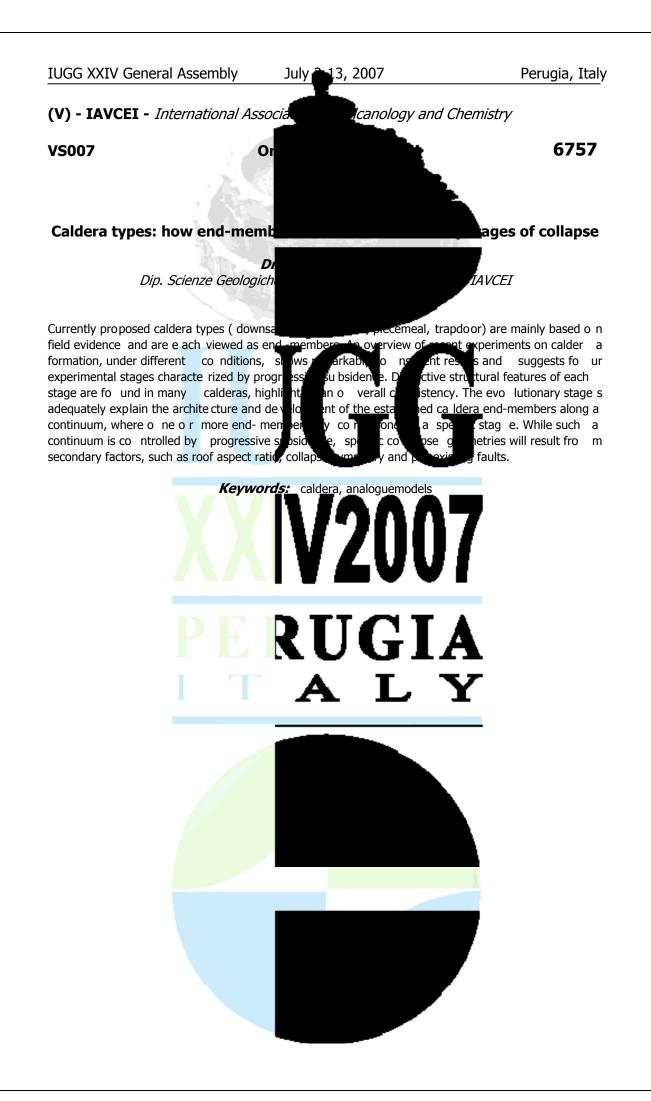
ed to a Ge ographical Information System (GIS) b and to filter them se. The final aim of the together the existing

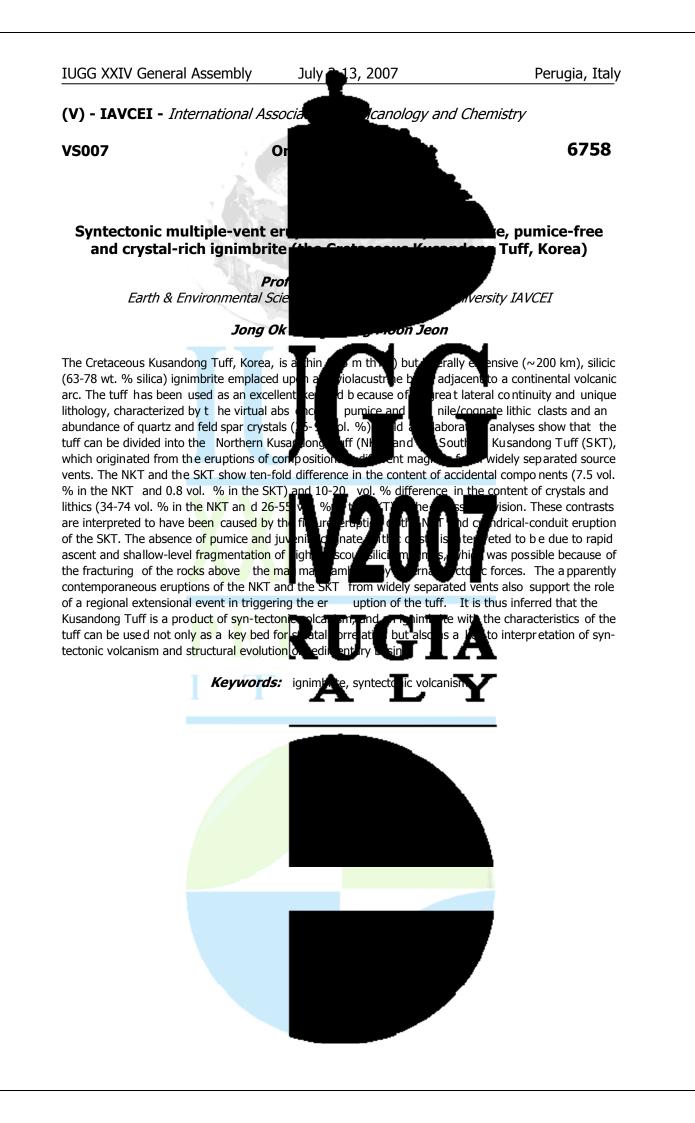
> bgraphy. Evidently, thi s epresentative enough for

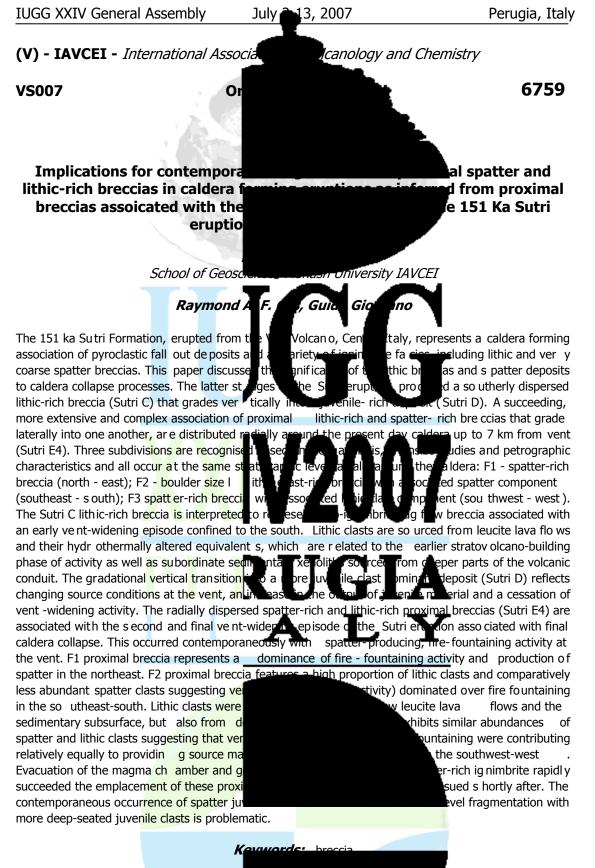
scarce and sometimes also unsatisfactory, the final and better use of this database is to convert it into pumose of thi s work is the commun ation ween research groups in D , ge users will be able to nt Fin ( sу nge Cor sed acquire (after registration and obtaining of a username and a password) the current database version, to propose cor rections or u pdates and to exchange informa involved in the study of caldera collapse processes. on with othe registered members also udes a formulary that will the CCDB in facilitate the incorporation of new calderas with their corresponding characteristics and attributes into

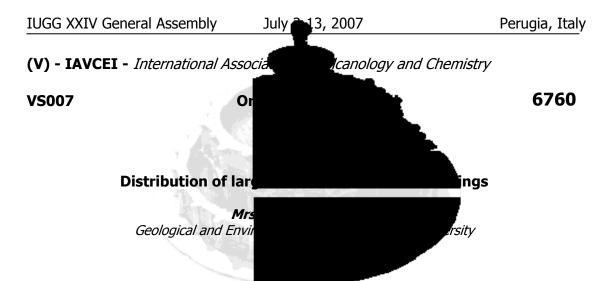






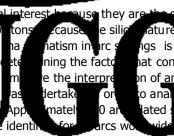






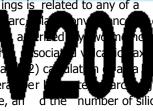
Silicic calderas are of fundamental geological interest magma chambers, the precursors of felsic pritons it resistant to subduction, the study of silicic continents form and persist. In addition, by silicic magma chambers in arcs, we hope to compilation of silicic calderas in arc settings scale silicic magmatism within volcanic arcs 2 Ma and larger than 5 km in diameter we

development of silicic volcani sm in arc sett ings is related to any of a thickness and age of the crust underlying t spatial distribution of silic ic calderas was distance between each silicic cald era a n stratovolcanoes of intermediate compositio arc that quantifies the number of silicic ca correlates with plate-normal convergence rate,



surface expressions of silicic continental crust makes hgs is key to understanding how hat control the occurrence of large n of <u>ancient</u> arc settings. A global the occurrence of largec calderas younger than e goal of this study was

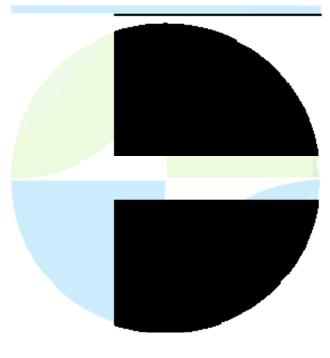
to analyze the intra- and inter- arc spatial distribution of silicic calderas, and determine whether the

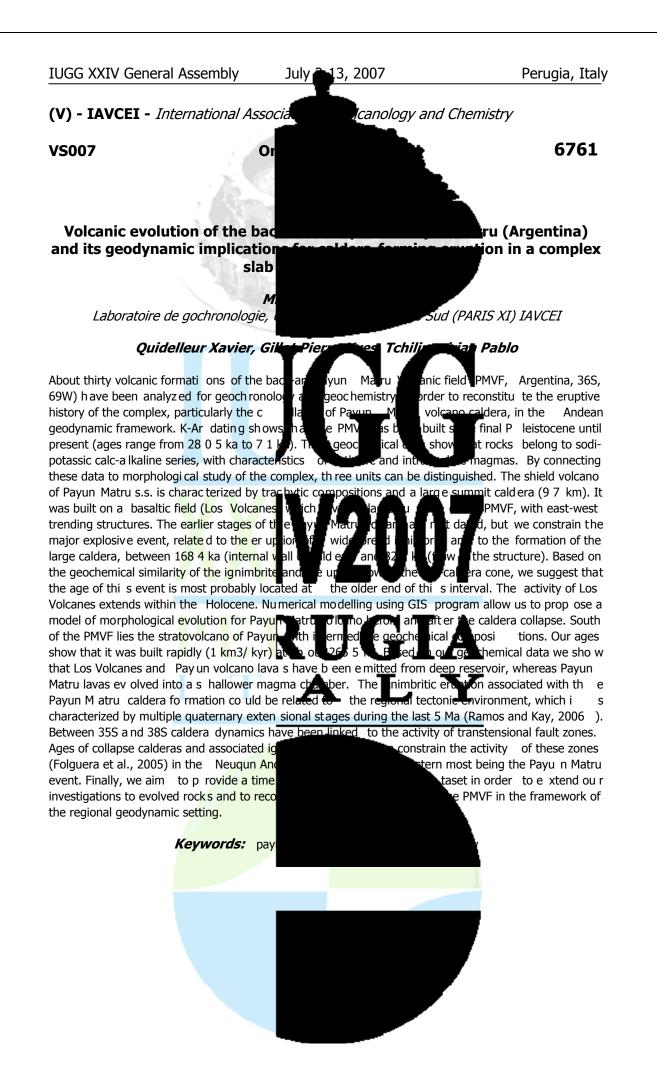


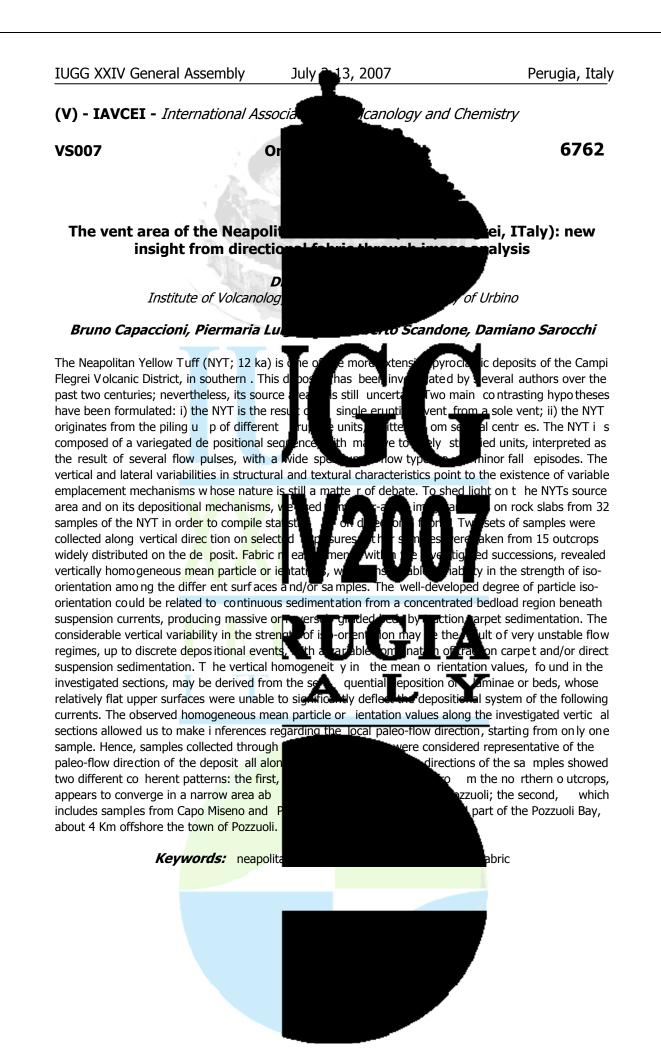
variety of fact ors including subd uction angle. The heasurement of the defined by the e arc, ensity for each individual aldera density positively number of silicic calderas decreases with

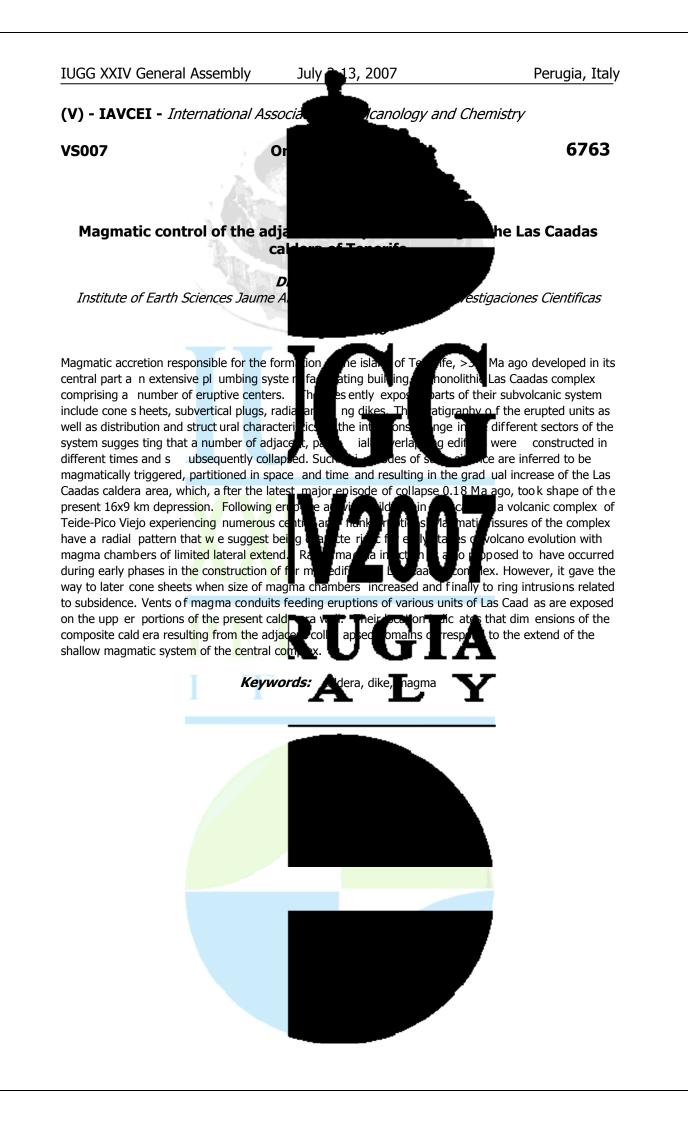
increasing distance from the volcanic axis. Both of these correlations are likely related to magmatic flux. Although calderas, in general, do not prefer tance behind the arc, in arcs on relatively mafic or young crus t, calderas ten xi s older continental crust, calderas tend to occur behind the axis. Cald nd the main axis of the arc are generally more silicic and larger than those closer to the axis

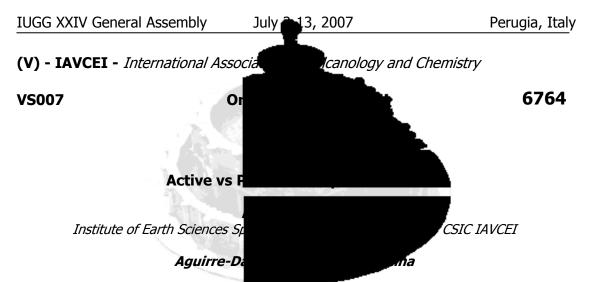
Kevwords











The formation of collapse calderas require the combin and tectonic condition, which are h ardly a very calderas are usually assumed to form by gra ta volcanic eruption. However, in some cases evolution of a volcan o-tectonic depression underlying magma reservoir. Such a caldera with large volume eruptions, (2) the rapid preventing ero sion and signif icant sediment ation

marginal zones of the basin; (4) the presence of co and, (5) the deposition of continuous successions of ignimbrites, sometimes several hundreds of meters thick, during single eruptions. Calderas of as the Sierra Madre occidental in Mxico, among other a reas. We propose to name contrast with the name "Active Calderas" t at passive calderas caldera-forming episodes thus indicating that the subsidence dynamics of the basin pre-and post-dates caldera collapse. This fact

response to the volcanic activity, as occur in

subsidence rate during the most important

of very spec evo la ring I collaps of t s may develo o th<u>e i nter</u>a ur is ca ce of

thermo-dynamic, mechanical Icanic system. Collapse oof of a hagma chamber during a ring an intermedia te stage in the of <u>the local</u> tectonics wi th an ssociation of the basi n nain eruptive episodes, ult-zone vents at the

-ignimbrite lag deposits in some m arginal zones;

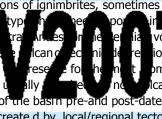
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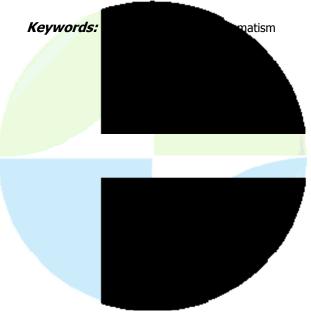


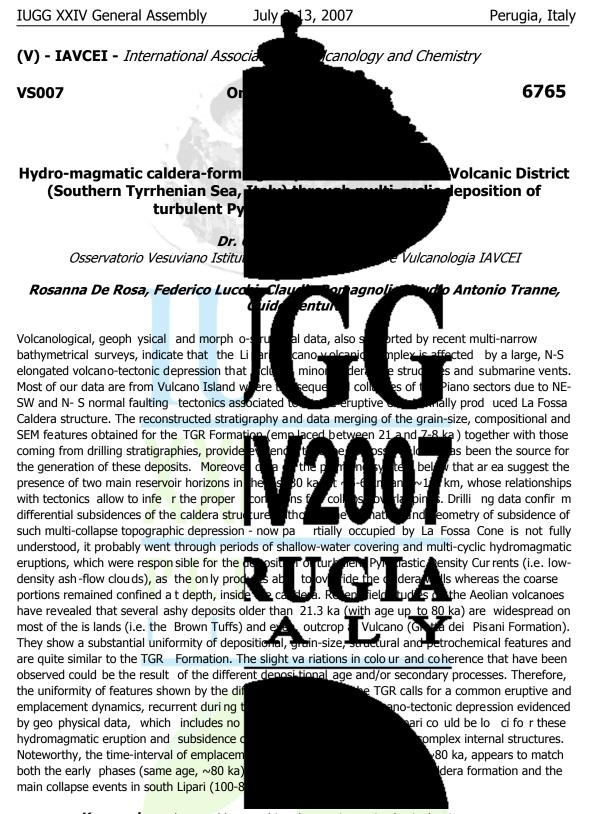
al volcanic areas such ism in the Pyrenee s as "P assive Calderas" in on caldera structures. In c or epicla stic sediments,

indicates that the subsidenc e structures is create d by local/regional tectonics and not as a direct caldenas volcanism cause a high similar to that of active this reason, it seems ггo

collapse calderas but after the subsidence ned reasonable to characterize th is process and the re sulting caldera-like str uctures as passive calderas, meaning that the caldera collapse is generated from supported by the fact that the volcanic acti vite de ubidence structure. This is a tectono edimentary ensively after the caldera episode and eases pro the basin returns to a norma I rate of subsidence at the same time than no n-volcanic sedimentation takes place again.

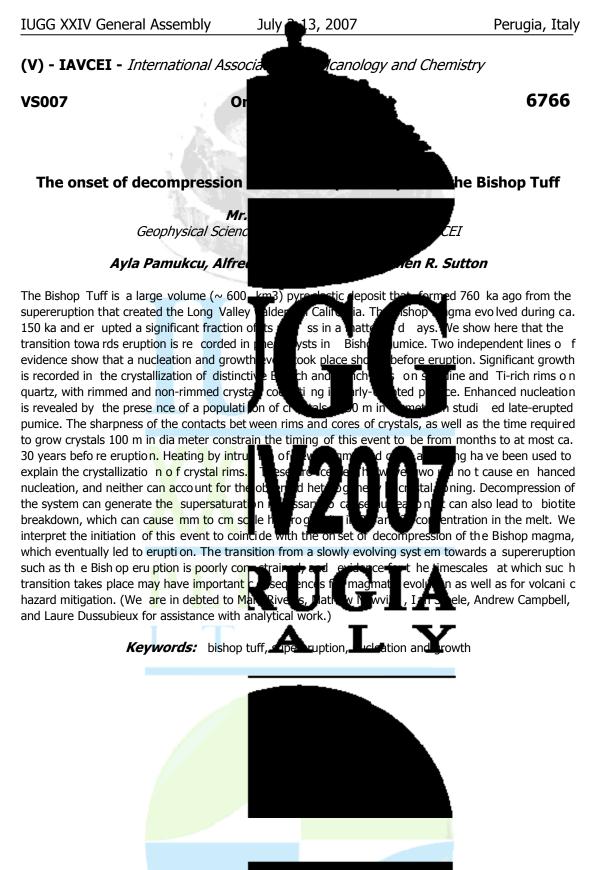
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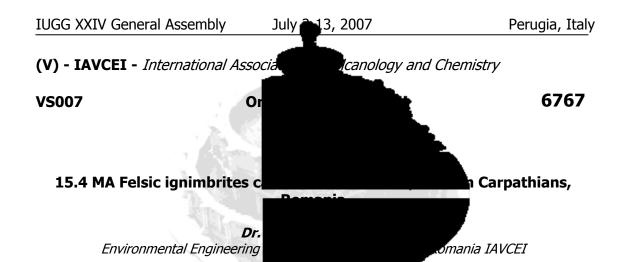




*Keywords:* vulcano caldera, multi cycle eruptions, piroclastic density currents







Complex v olcanic processes developed in Carpatho-Pannonian Region as a consequen Plate and the two microplates, Alcapa and area. Two types of volcanism have bee volcanism and an intermediate, mostly effus volcanism developed in the Pannonian Basir thick deposits of rhyolitic tuffs (eg. Lower, ] Tuff in Transylvanian Basin), mostly burie

reconstruct the so urces: some o f them have tentatively been o utlined, o thers are still unkno wn. A caldera has been outlined in Gutai Mts., Eastern Carnathians, in Middle Miocene, related to the onset of the volcanism at 15.4 Ma. A back-arc exter was responsible for the large sheets of rhy by the Paleogene flysch and mostly buried between 13.4 to 7.0 Ma. Log ging cores from reconstruction of the geometry of igni volcaniclastics, as well as their lithology

constituents and the chemistry of the juvenile pyrocl asts, pumice and crystals, emphasized the rhyolitic calc-alkaline character of the hetero geneo alterations (Flp , Kovacs, 2003). The fiame welding, as well as the explosive character of clasts suggests the transport on paleovalleys of the pyroclastic debris. Sedimentary structures reflect the rheology of the parental flows an d the evolution of the explosive so urce. Whiple units with massive structure, normal coarse-tail grading of lithic clasts compatible with successive mass flows emplaced by progressive aggradation from a steady, maintained

stratified suspension current, generated by texture, the cooling textures such as colum regime and/or low cooling r ates compatib emplaced as a single cooling unit, from th pyroclastic current, in subcri tical regime. compatible with boiling-over eruptive style southwestern part of the Gu tai Mts. The g the source location and the subsequent ev

towards east, from 350 m to 30 m. The ignimbrites extended to 10 km2 show den the proximal f acies, the intr acaldera igni succession of reworked pyro clastics inter l intracaldera, by mass flow reworking altern steep slope between the intr acaldera ignir

🗈 Gutz (Eastern ducti cene Dacia/ Tis fied : a felsic, c-type\_volca stern a Carpa d Upr Tuffs hv actua ut in pl

in the entire h**i**ans) as involving the European broce oping the actual In tracarpathian t ly explosive, e xtensional-type The Mi ddle Mioce ne extensional hs, where extended and ne Pannonian Basin, Dej e are few attempts to



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mostly terrestrial style, counterparts, underlain andesitic volcanics dated in outcrops enable the e co-genetic rewor ked

provintic and adularia- sericite gesti he dens e to moderate of the accidental lithi c ha\_ acte 🖬 se- tail 🤿 ding of pumice clasts are pyroclastic current. This is in a ccordance with the rheology of the basal layer belonging to a density-

> ns. The eutaxitic text ure or welding e pipes reflect a volatile retention The ignimbrites are therefor e paintained, stratified dilute and volatiles regime is ation is suggested in the ment i n reconstructing located on the south-

western part of the area have been used to outline the spatial distribution of the ignimbrites, as well as their geometry. The ignimbri tes have a we dge geometry with thickness decreasing from south-west constant 350 m thick and c oarse westernmost deposits of

ization. They represent 50 m thick complex ssion had been built up edimentation. There is a defined as caldera outflow.

## IUGG XXIV General Assembly

July 🚵 3, 2007

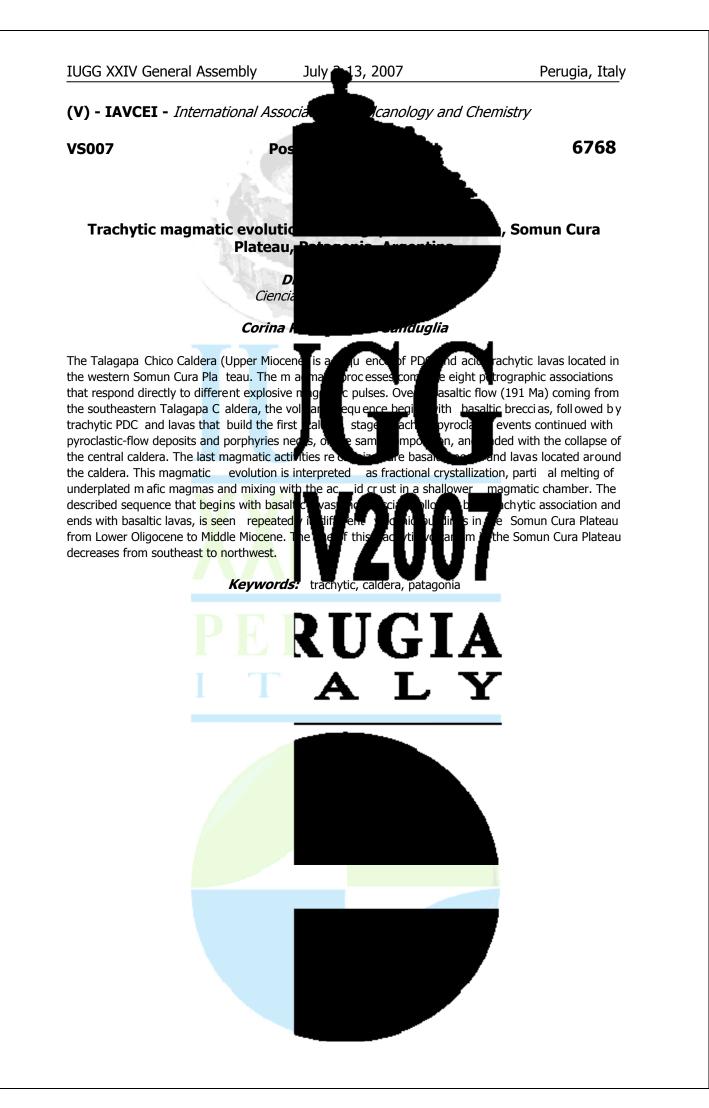
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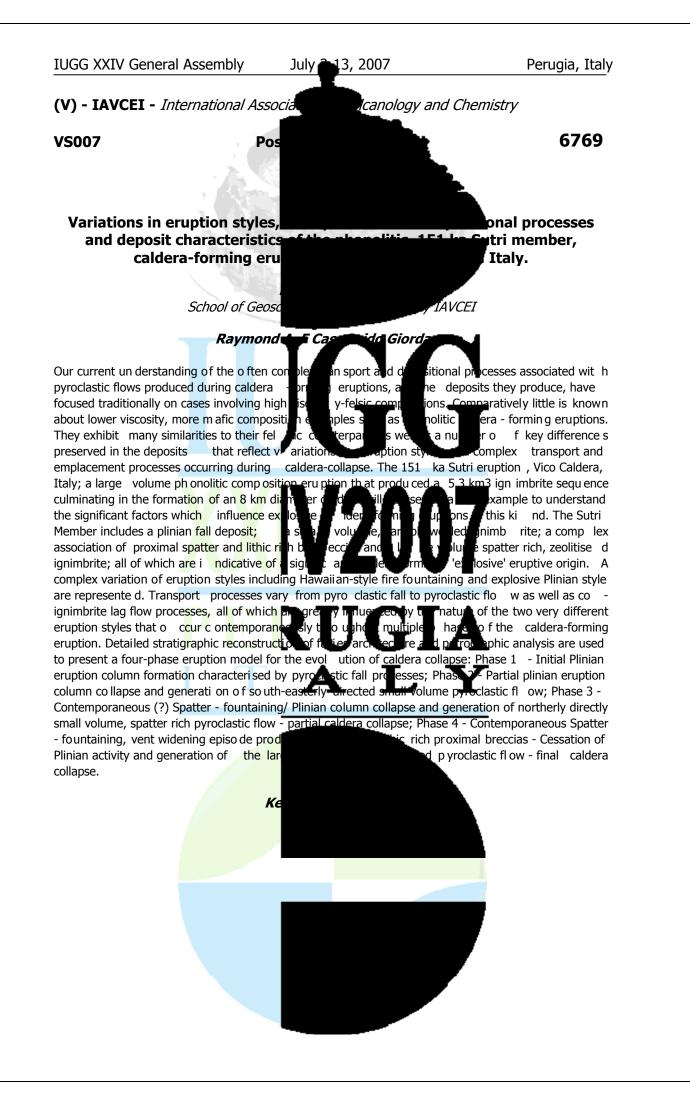
The outflow preserves the texture and the we deposition. The set of data was useful in dynamics of the eruption to the collapse of by subsidence. Cross se ctions through the which underwent subsequent tectonic proc 2003. Petrolog y of Badenian ignimbrites, Bolyai, Geologie, 17-28 haracter in accordance with the hot state intracal mod el of the ignimbrites, from the intracaldera sedimentation entrained in morphology of the ignimbrit es References Flp, A., Kovacs, M., Studia Univ ersitatis Babe?

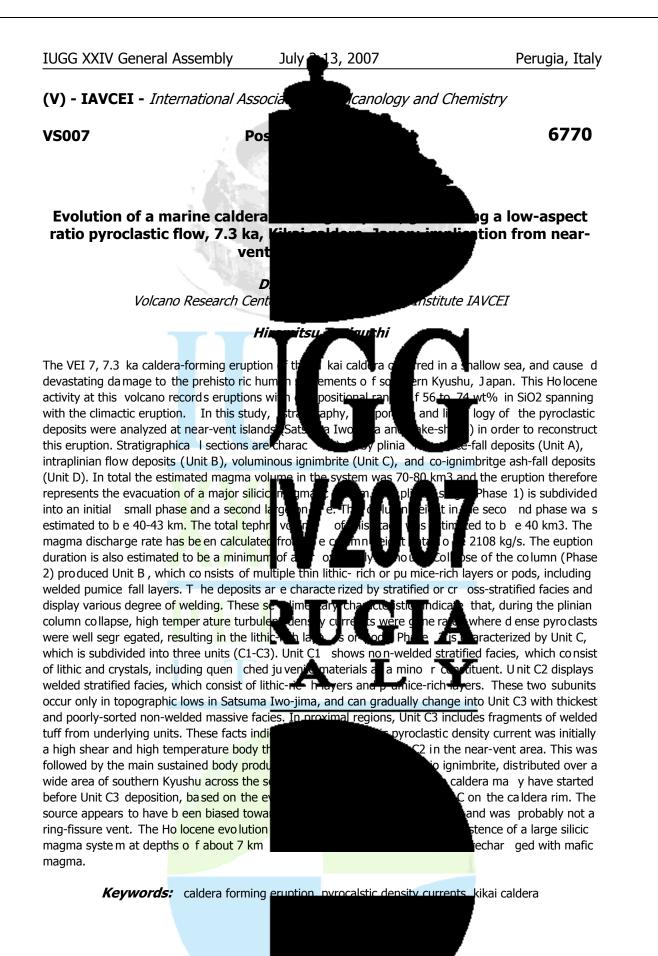
Keywords: ignimbrites, felsic, pyroclastic

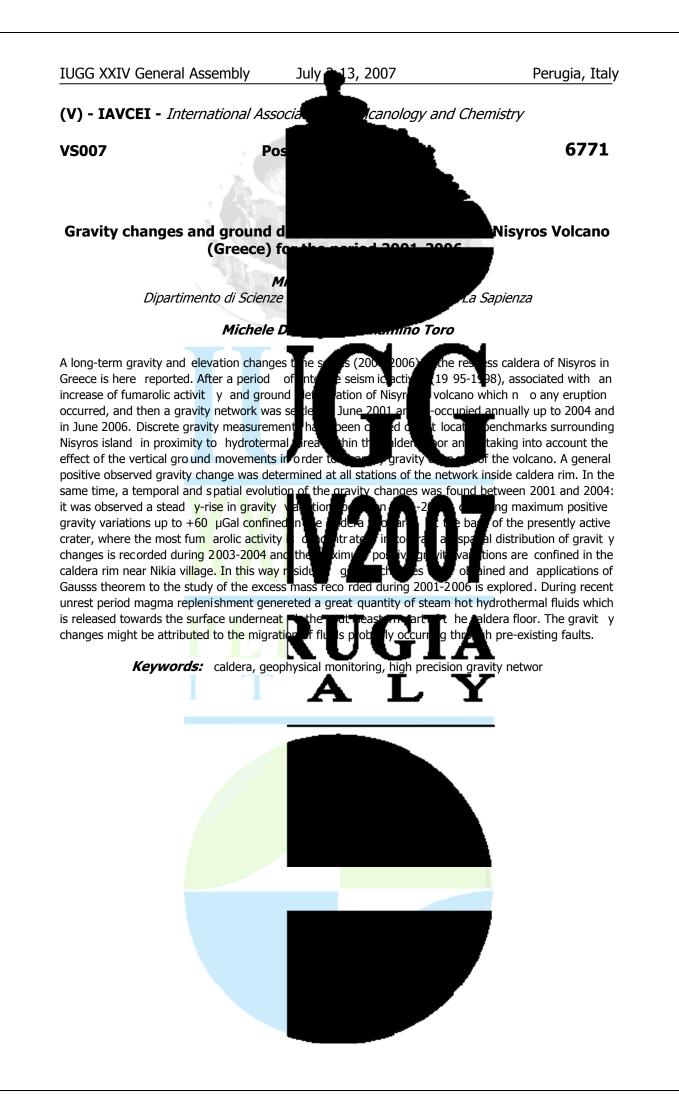
## I T A L Y

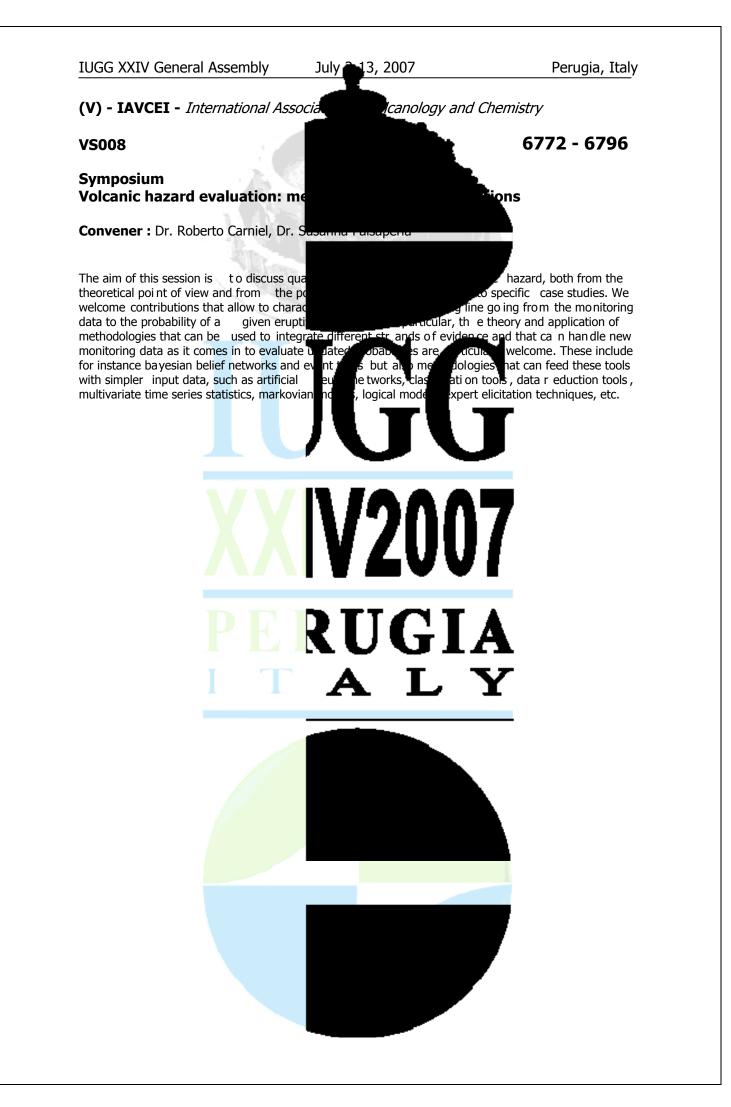


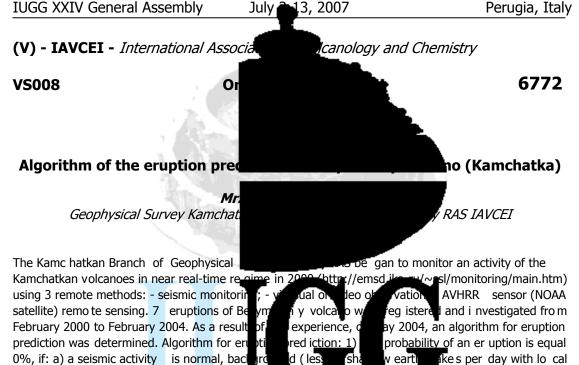






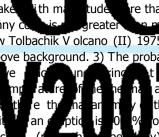






thermal anomaly temperature at the Bezynnanny d thermal an omaly of the lava flow at the New Tolbachik V olcano (II) 1975. 2) The probability of an eruption is equal 50%, if se ismic activity is <u>above background</u>. 3) The probability of an eruption is 90% for the next 30 days if seis mic activity is increases with the growth of the maximum cone relative t o the maximu m temperatu Tolbachik Volcano (II) 1975. 4) The proba bilit activity is abov e background, and rock ava anc seismic network. Inflation of the Bezymianny dome is

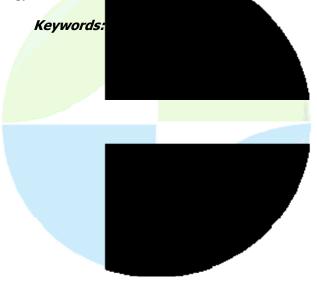
magnitude 0.75-1.25 and no shallow earthq

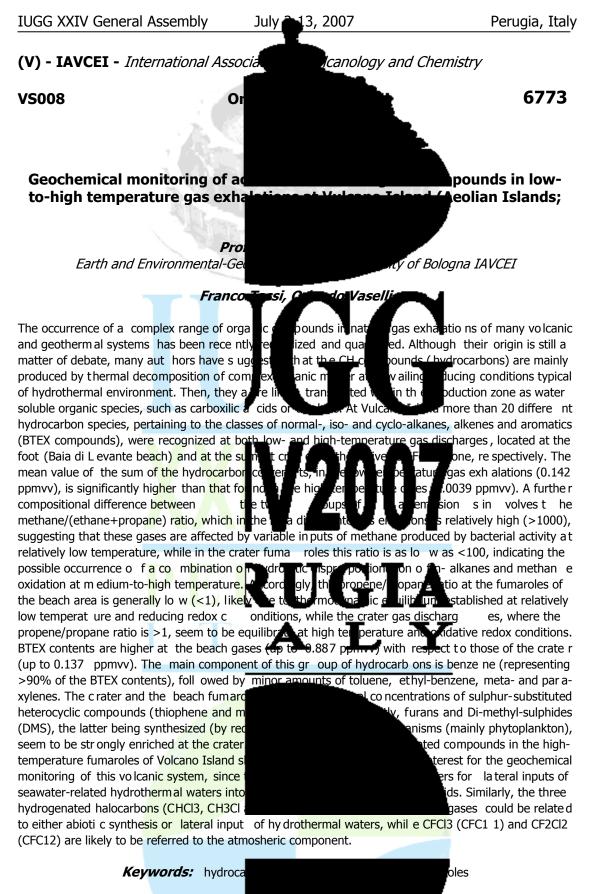


kes per day with lo cal 25); a nd the maximum hum temperature of the

a nd seismic activity aly at the Bezymia nny ava flow at the Ne w e next 7 days if seismic e detecte d by the KBGS the likely cause of an a ppearance of the roc k

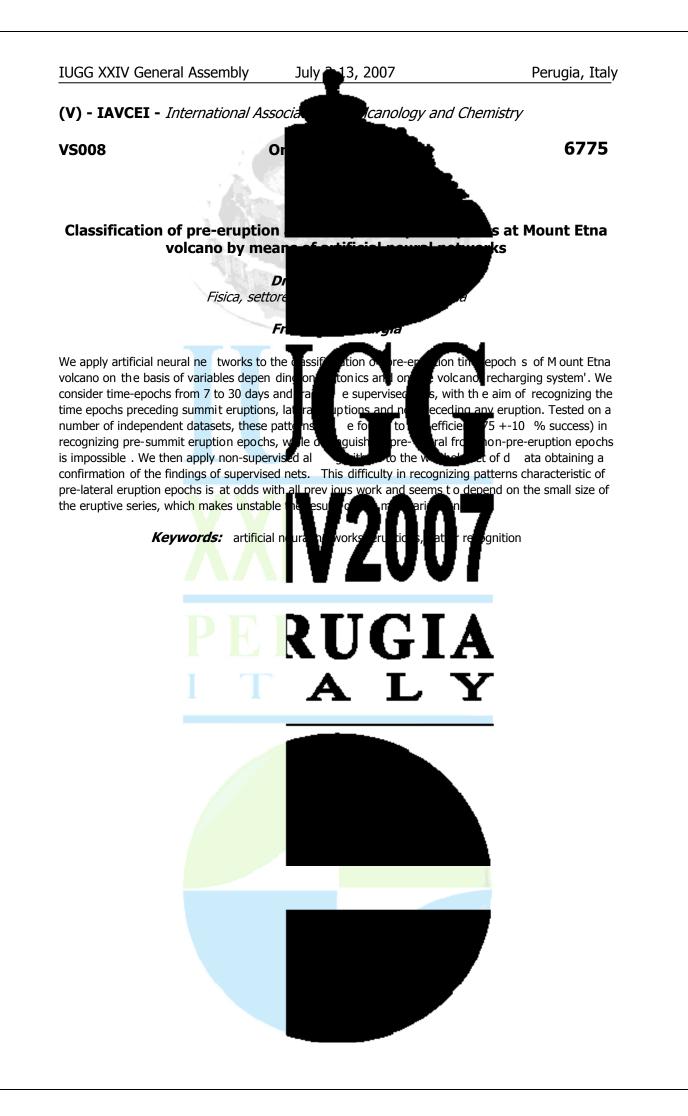
avalanches. 5) The type and size of the future eruption can be estimated based on the i ntensity of the preceding seismicity. The int ensity of the roximately pr oportional to the intensity of the eruption. 6) Recent seismid to ٥v s to c arry out thes e investigations only if the amplitude of volcal explosive erupti ons of Bezymianny volcano (June 2004, January 2005, tre luct vsk cano is less than 1 mpc at station CIR. All five last November 2005, May 2006 and December 2006 were prected using using orith m without fals e Toroject. For the 200 5-2006 events, alarm. All five predictions were passed to par cip ts of KV predictions were passed to Ka mchatkan Branch of Ru ssian Advisory Council. This algorit hm was tested by a joint, international, real-time experim ent in November 2005 and included participants from the Kamchatkan B ranch of Geophysica | Surve ano Obse rvatory, and the Institute of Volcanology and Seismology.

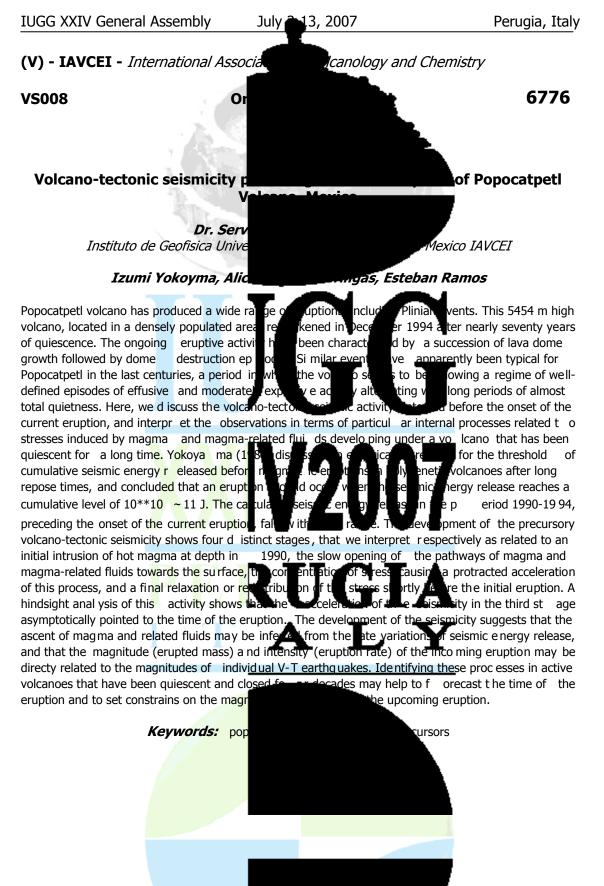


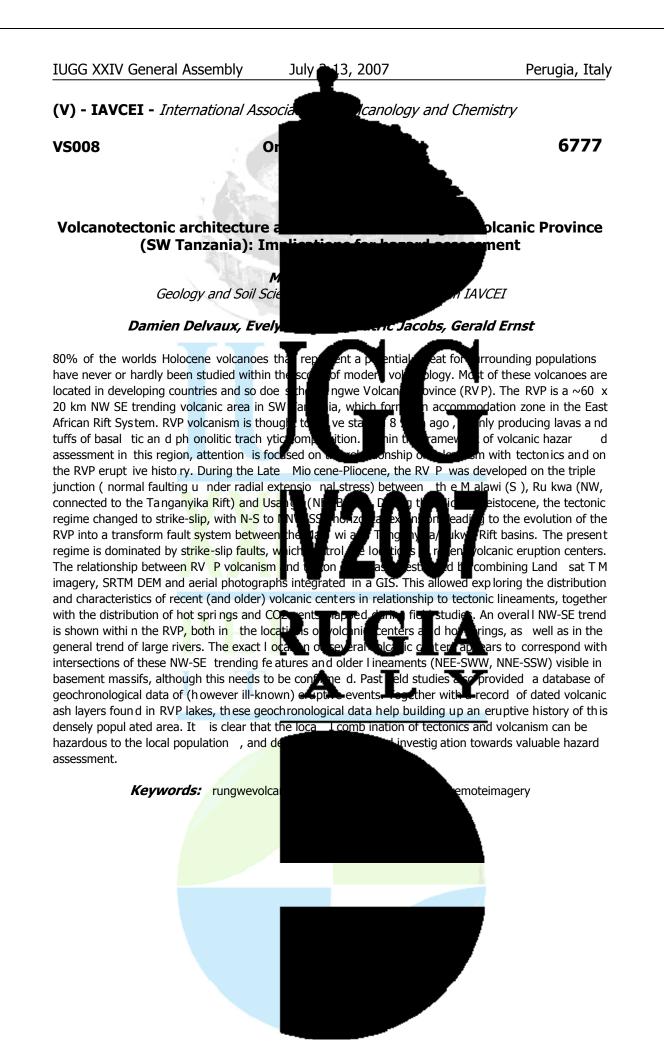


r**ds:** hydroca











dangerous areas of the caldera, but it is also quantitative volcanic hazard (VH) at Campi F meaningfull estimation of the probability at recently proposed by Marzocchi et al. (2007, 2006 Bayesian fashion.

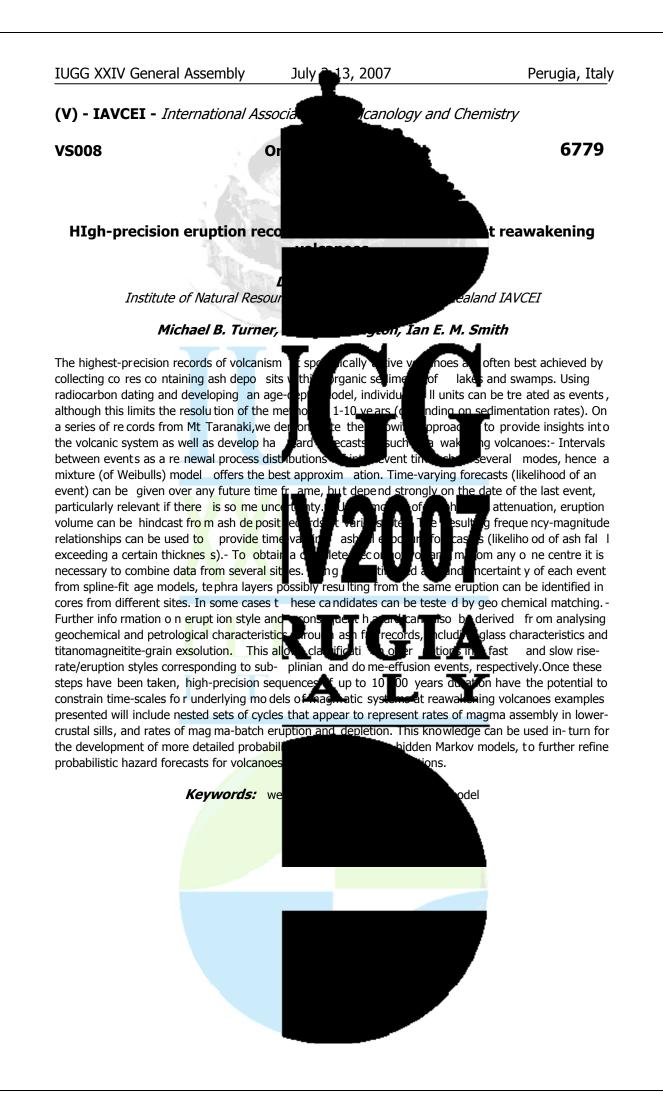
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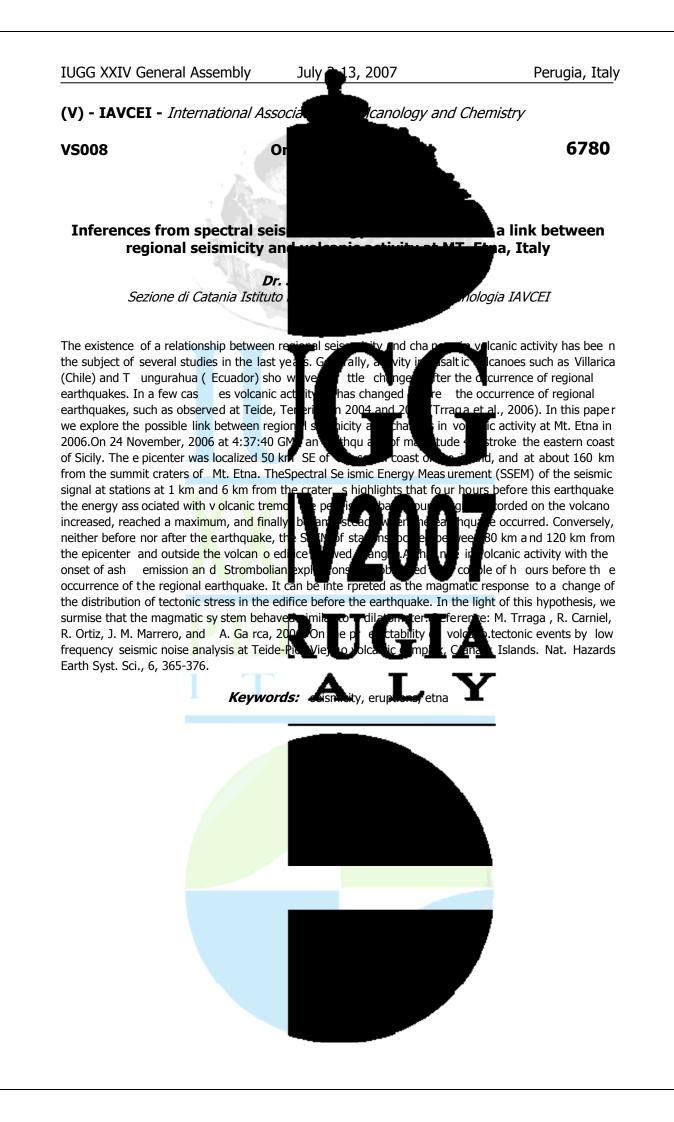
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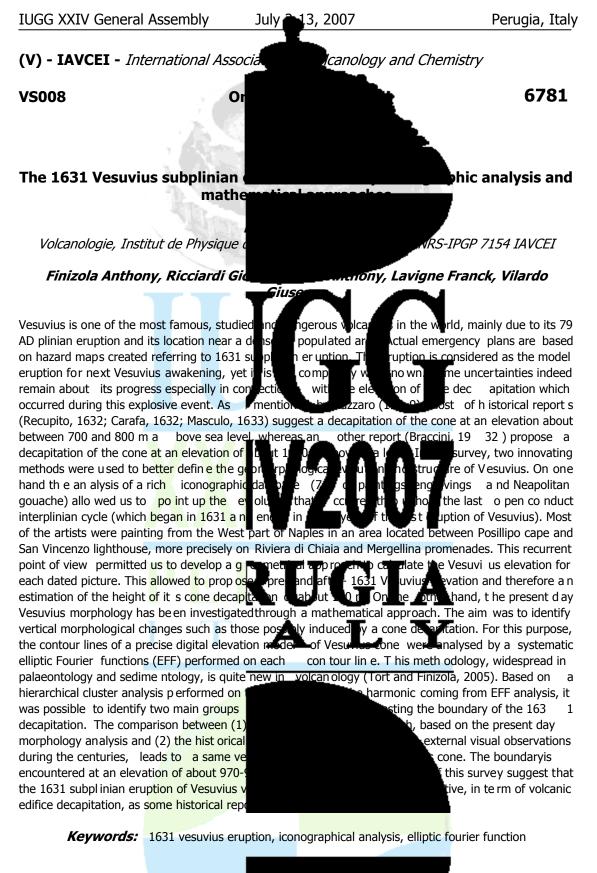
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ccess the more general problem of ed ma designed to represent a ayesi vent Tree (BET) model, VH in a fully structured,

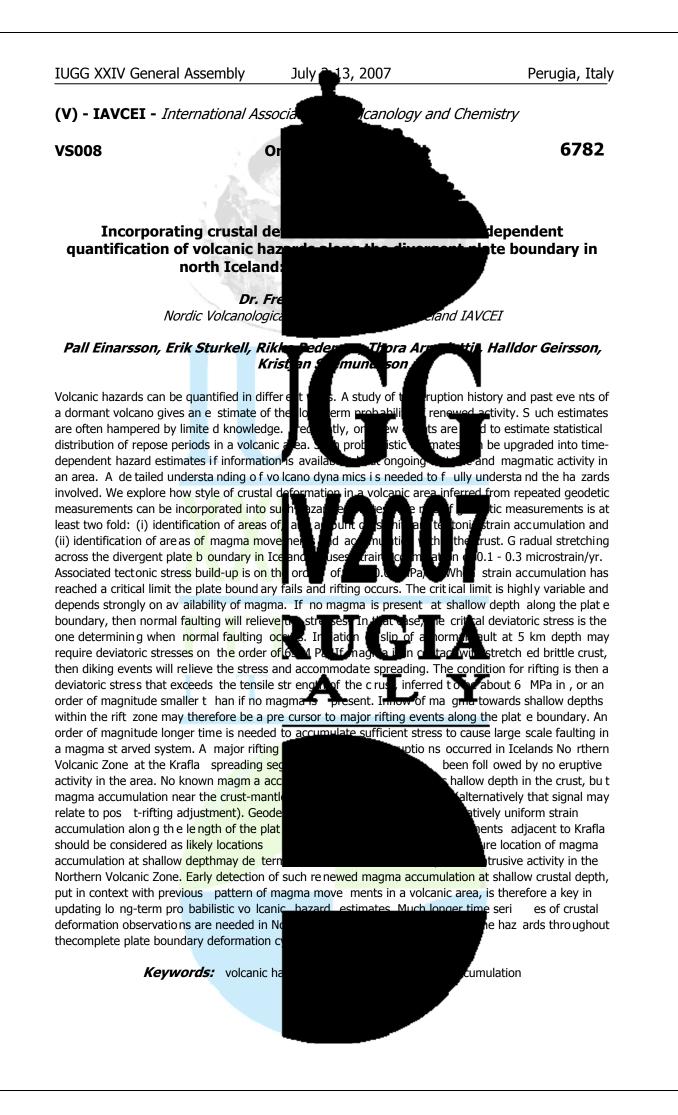


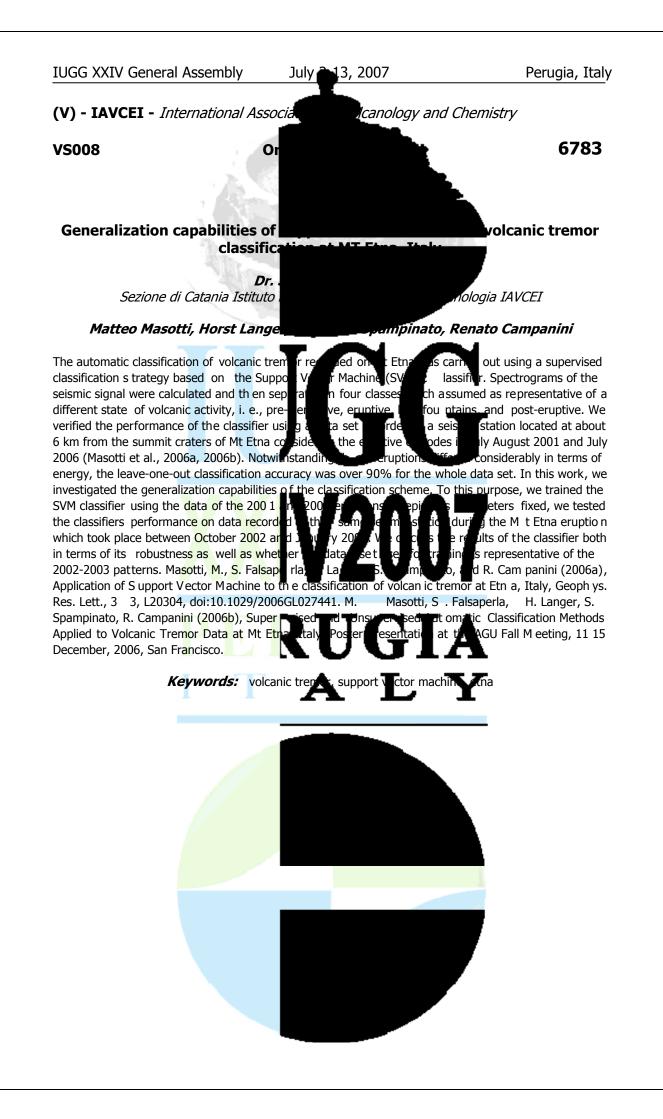


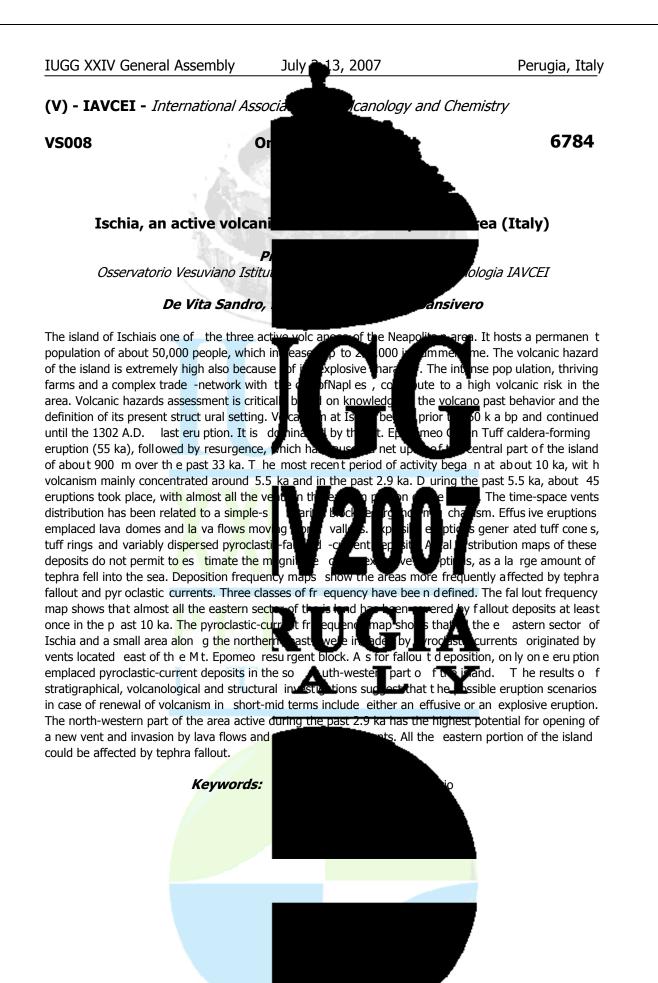


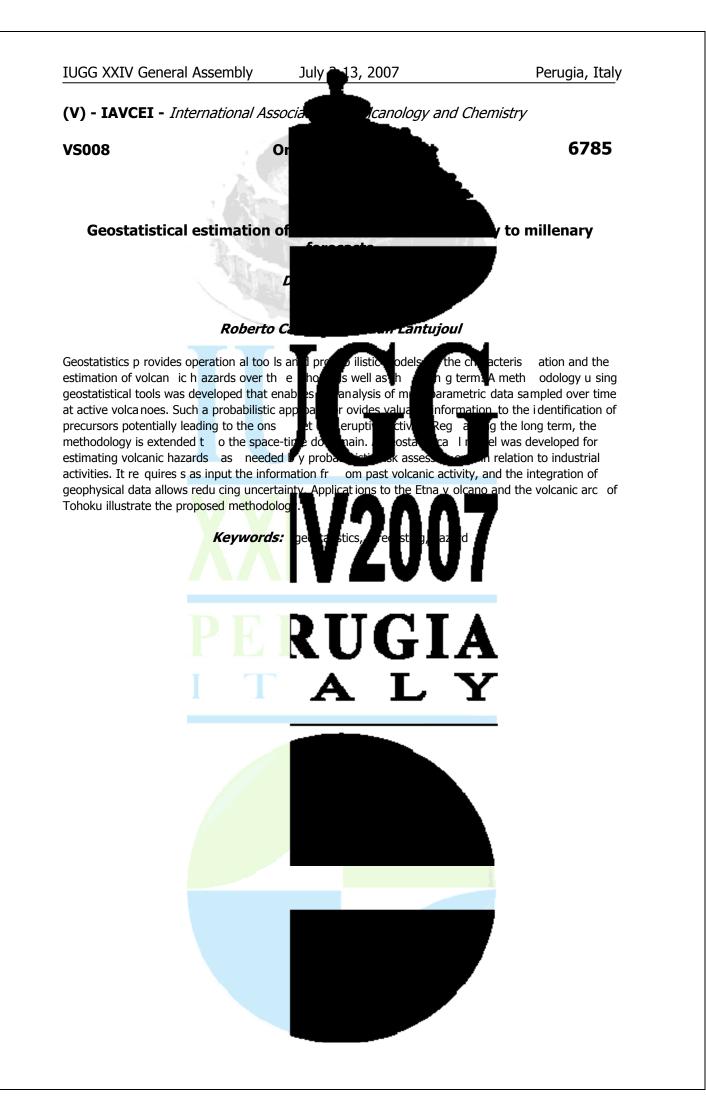


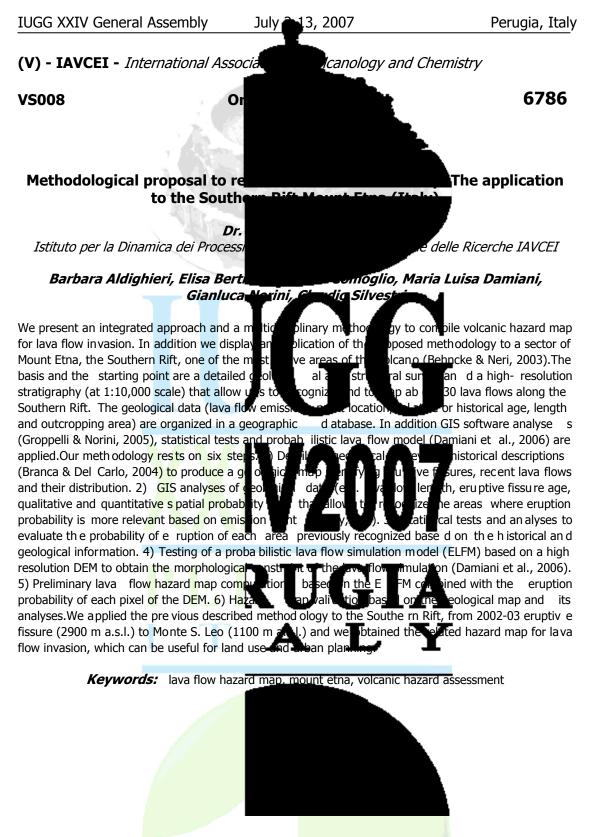


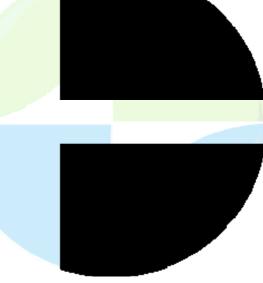




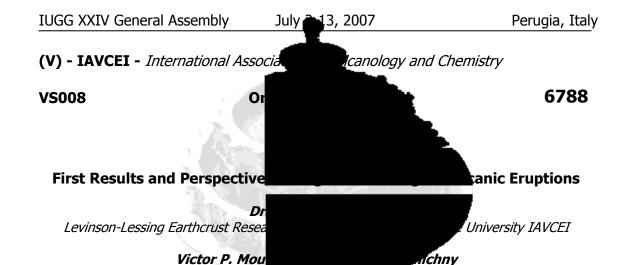












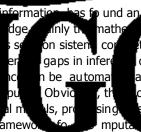
Logic, being a purely formal treatment of strict, and well-crystallized domains of know dge may show, inter alia, whether a theory obscured by natural language may be enlight to similar conclusion revealed. Logical infer involving many premises can be quickly con assessment, including the cre ation of physi Bayesian Belief Networks and other fr

expected results because of multi-discip linary and high ly intuitive character of knowled ge in this field, largely based on description of individual new kn owledge en gineering too ls in vol formalization an d organ ization of deive kr application of the logical n otation used for The architecture of event bush allows t

automatically, individual varia bles being provided by statements, predicate constants, by predica tes of all the statements, logical connectives, by edges of different types, existential quantifiers, by q quantifiers, by all the rest cas es. The event volcanoes was recorded in terms of PLL. The has vividly showed some general regularities in erup tive processes which are rarely p ut as such in the literature. Unfortunately, we could not avoid using system unlikely resolvable to the present d ay substitution of variables. We studied the deducibi apparently true, some, apparently false, from the natural-sequent calculus. The results must domain of knowledge describ ing the lava knowledge, and improve the quality of eru

Instituto Nazio nale di Geofisi ca i V ulcano volcanic hazard, a Human Capital Foundati NATO-Russia collaborative linkage grant.

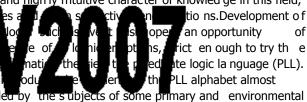
Keywords: logic



ation in the most perfect, ries. Application of logic tical t o ptior xpert

te and compact; contradictions det ected, and different pathways and <u>thus long</u> chains of inference uld be useful in hazard Igments, composition of zards, and decisio n-

making.Attempts of applicat ion of I ogic to the ge osciences including volcanology h ave not led t o



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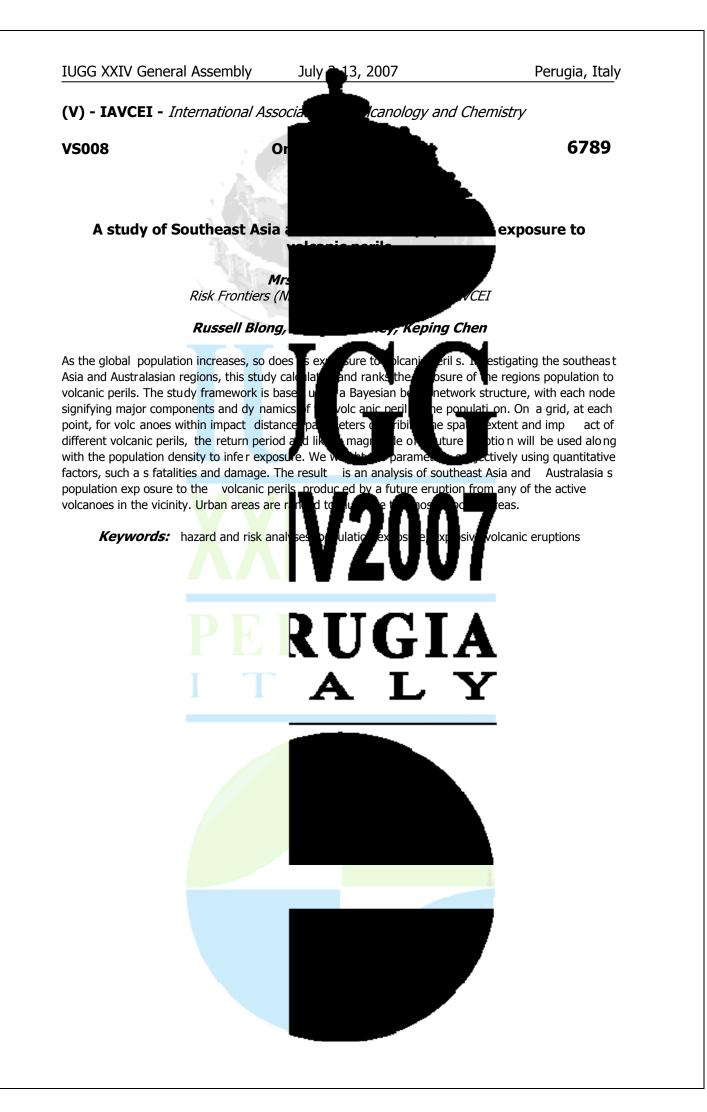
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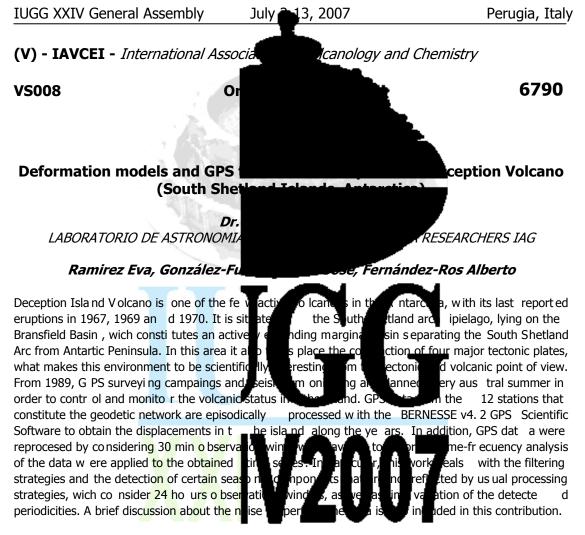
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tio ns.Development of an opportunity of fict en ough to try the te logic la nguage (PLL). PLL alphabet almost

e act of Soufrierre Hills-type rup ents several times. PLL of s le li lau hat makes the whole logical predicates perthis can b avoided by appropriate lity of various statements, some of them bein g knowledge stored in the event bus h, by means of about the degree of maturity of the the formulation of volcan ological is supported by Project V4 of antitative tools to evaluate cluster project, 2006, and

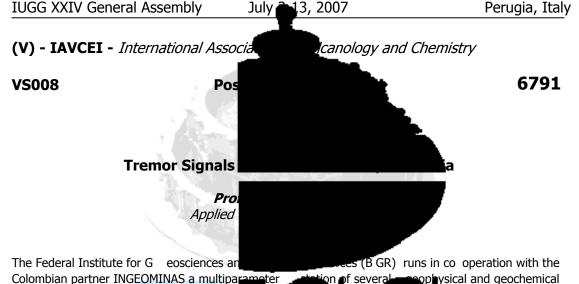
bush structure, and univers al





Keywords: volcano geodesy, time series, wavelets analysis





Colombian partner INGEOMINAS a multiparameter sensors at Galeras Volcano in southern Coloi bia monitoring of volcanic activity as well as for the activity. Three-component broadband seisme Two different groups of seismic signals are r Volcano-tectonic signals with their sources i and tremor signals emitted by the fluid syst broadest sense referring to all seismically

from non-stationary flow of the magma-gas fluid in the uppermost reservoirs of the volcanoes magmatic system. Tremor signals moved into the focus of scientific interest for their potential to give new insights into the flow regime. At Galeras we record singular signal types of the Tornill os, the l appearence in the vicinity of the 1993 erup for their possible value as precursors in eru been collected up to now since the insta especially during the periods of reactivation in writer 2000, autumn 2004 and 2005. In our work we do not focus on the analysis of the signature of only a few single events, but rather on the parametrization

on of severa This e 19 earch inte fun are installed with these the s e prod volca The ninus fluid

ter statio n aims for the tipara mental pocesses of this volcanic lose proximity of the crater area. adband seismometers at Galeras:

material of the volcano mor herein is used in its d noise, as it orig inates

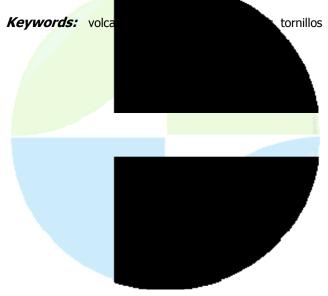
nals: the distinct and nic Tremor. Since their als are intensely studied set of tremor signals has tion at Ga leras in 199 7

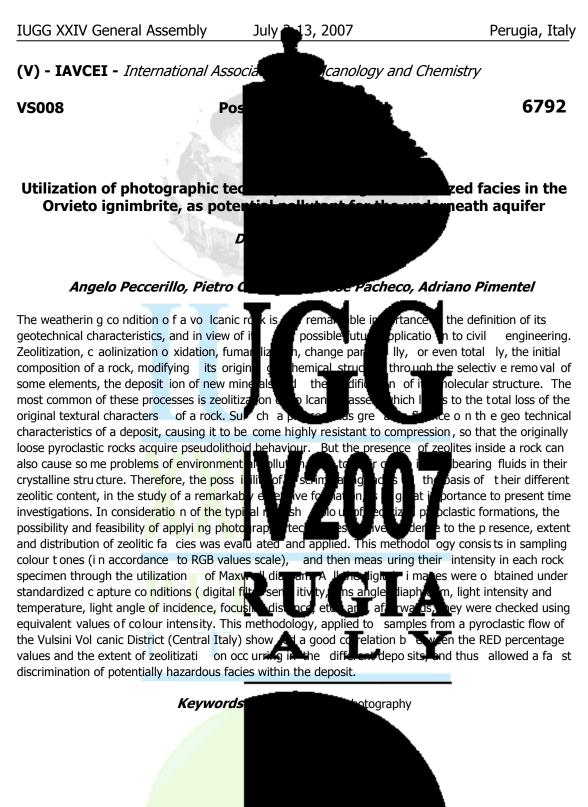
of a larger data set to use the distribution f eters as a base for physical modelling of possible oscillator and resonate llos precisely determined by 'Tor only a few par ameters, it has been sho wr οy ide rang e of kinematic kter at t S and spectral p arameters and parameter values. There are LP-Signals in habiting the ch aracteristics of Tornillos or Tr emors and thereby causing great liffi culty in the signal c **T**sification. One possible ignals is of a single source for the explanation for this heterogenity of the classi field LE e assumptio three signal cl asses, emitting in dependence of the stimulation mechanism the pure signal forms Tornillo/LP/Tremor or a continuum of transition forms. A general outline of the recorded tremor signals as well as first results of the parametrization

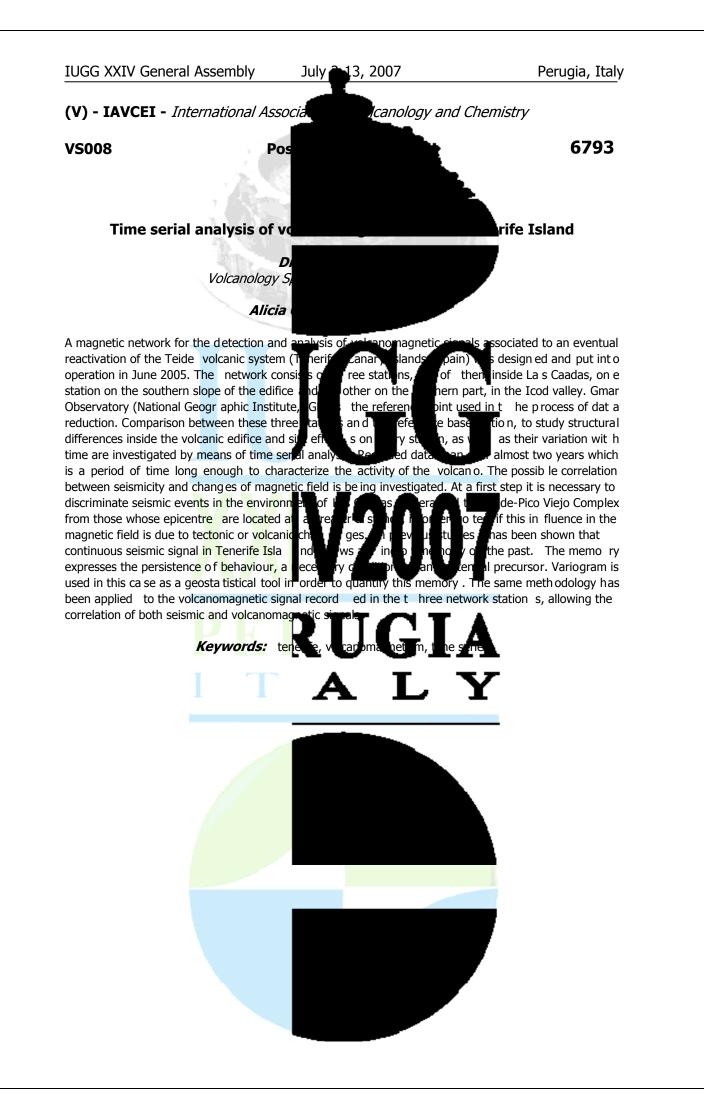
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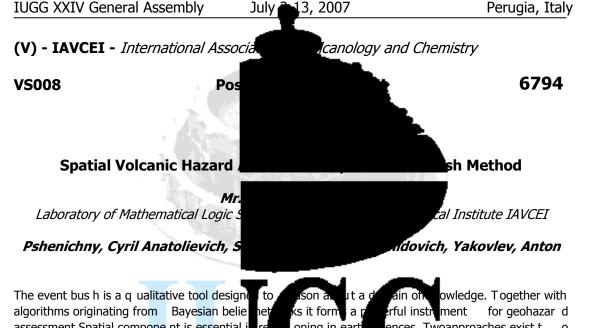
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assessment.Spatial compone nt is essential i handling geographic/geodetic data in the ev divide considered space domain(e.g., the ma of the basic event bush for each quantum, nn will describe this approach in detail. However, whi

not prone to t he hazard in q uestion e.g., summit zone cannot becovered by lahars, insusceptible area kilometers away from the foot of thevolcar time frames, and remote lo calities cannot inference is NP-hard, that is, there are no e cases.)Alternatively, one may note that each area(s). Thus every node can be regional assume that the contours ofcause and

otherwisetransitional no de(s) should be int roduced. In addition, each process can more or less selfpropagate in a linear or radial mode affecter lahar sweep acro ssthe ground, a lava do occurs in some localities, so does hydroth related to each node should be the sum of the area self-propagation. However, the definition of self-propagation a for each cause-effect pair, which can be based ent

theothers may create comp utational problems. Pre senting the two approach es to s patial hazard assessment by event bush, wewill demonst thecomputational complexity of thes e tech each of them, using Elbrus and Shiveluch examples. The study is being carried out ur Vulcanologia, Italy, Development of quanti Foundation grant, Information Technologies

Keywords: hazard ass

oning in eart hformalism.T rea) in 'nа crete i ig the ther bies ntervà

ences. Twoapproaches exist t 0 ost straightforward approach is to vals and produce a copy me unified manner. We ete approach may work

well, for large mapsthere will be way too much geographical quanta, a good deal of which mayappear

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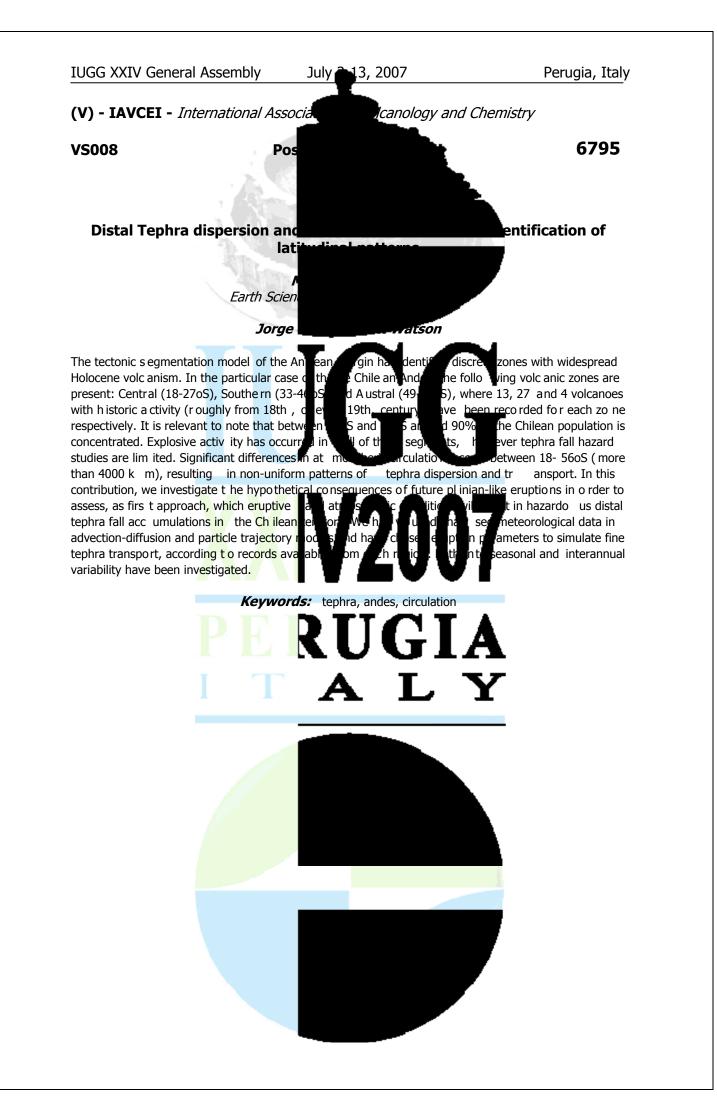


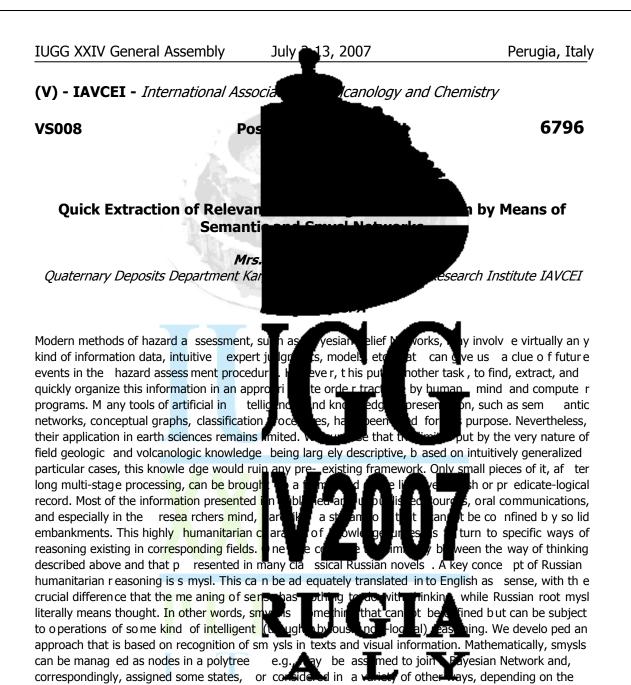
a dome in reasonable vs. (Notethat Bayesian cept for certain special ccursin s ome particular ontour on the map. We ush;

flow flows, pyroclastic flow or face smic disr uption ofrocks ce, the overall cont our Ėks, tc.) Πe of intersection with the cause(s) and the area of notae and intersection areas as for each on observations, or on ical model s expert judgments, is acomplicated task, and upda ting c omputation of areas of some no des given natical techniques within both, d iscuss to build anumerical mode I with amchatka, correspondingly) as titutoNazionale di Geofisica i ard, and a Human Capital

thods







formalism chosen. Each smysl can be attributed an unlimited number of names and properties. These may represent individual ter ms, or fragments of <u>particular</u> texts, drawings, <u>photographs</u>, etc. Names and properties may relate to more than on the second secon

those connecting smysls with names and

resulting framework allows to add and sea

also by visualized feeling of smysl and thu

hazard assessment. Presented will be frage

on published papers on the structure and

Kamchatka ). Financial support for this stu

Geofisica i V ulcanologia, , D evelopment o

Russia collaborative linkage grant.

Iges can be further specified. The v words or key parameters but that information for needs of n, semantic networks built and Shiv eluch Vo Icano ( of Instituto Nazionale di nic hazard, and NATO-

Keywords: semantic net, bavesian belief network, etna

## IUGG XXIV General Assembly

July 13, 2007

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Perugia, Italy

(V) - IAVCEI - International Associa

**VS009** 

## Symposium Models and products of mafic e

Convener : Dr. Jacopo Taddeucci, Dr.

This session will focus on the dynamics o multiparametric observations, and an alytic can erupt explosively after a broad range of include: surface bursting of b uoyant gas b water, overpressurization of cooled magma plugs,

saturated melt. However different these properties specific of mafic magmas, mainly to volcanol ogists is to relate , possibly in a highly-variable eruptive style s observed a no session is to b ring together scientists with explosive activity. In particular, three fields renewed effort on the study of mafic pyrog geochemical, and volcanological observation of

submission of presentations showing integr

canology and Chemistry

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s from produ ct an alysis, simulations. Mafic m agmas among those already identified, ous interaction of magma with external

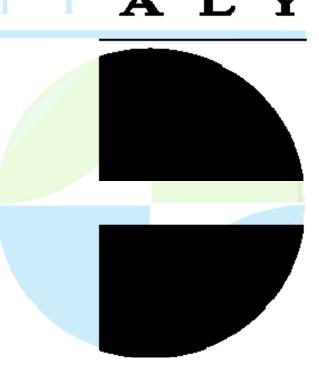
annular gas flo w loces bear, mav osity ar fast hм tive manner, te. oducts and h th nt ap**n** ch se to sts pyrd tic d ctivity;

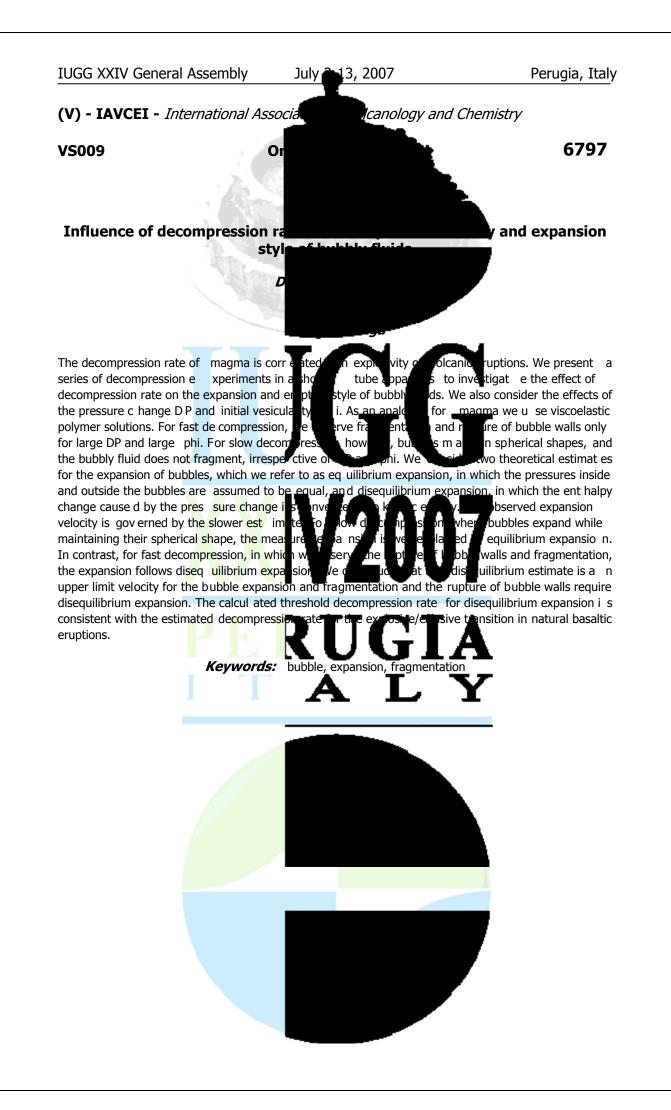
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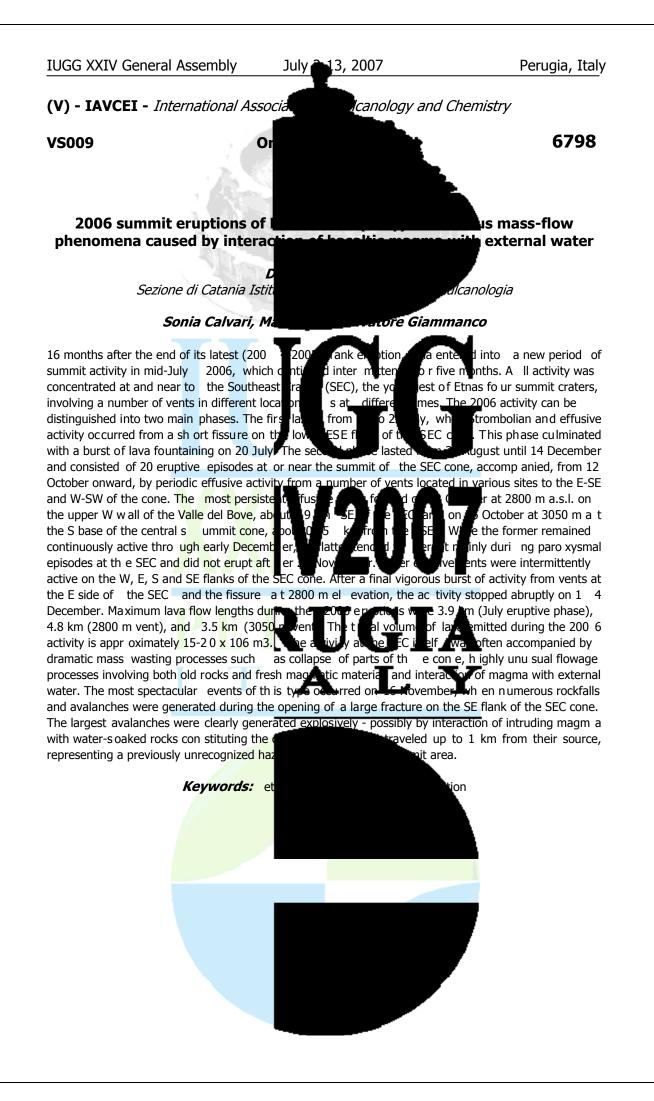
nd fast fo aming of vo latilect th e fundame ntal air tics. A cur rent challenge ernal kin bove physical pr ocesses to the ds t hey emanate. The aim of thi s that underlie mafic easi ng onverge on this topic: a parametric geophysical, sits; r lical, numerical, and

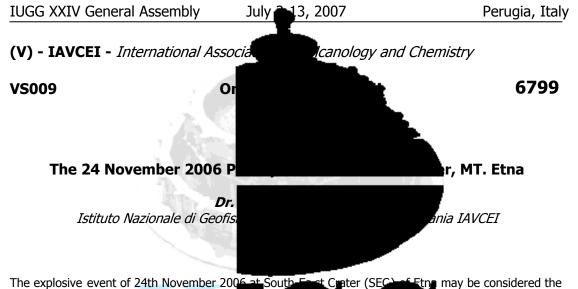
experimental simulation of eruptive processes. Al I the a bove (and other) specific aspects of mafic explosive activ ity will be wel come to this i nterdisciplinary session, and we specially e ncourage the

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best documented among small paroxysms in the 24th No vember, a weak eru ption co lumn eruptive activity, the favourable weather cor eruption cloud in real time, and to detail the ash samples were col lected over a relatively few hours aft er its end. Samples were sub chemical inves tigations, both by dr y sieving a n

procedure under Field Emission Sc anning Electron Microscope. The study of the fall out deposit was completed by drawing the isomass map and estimating the erupted vol ume. During the eruption, the wind direction changed from the east to eruption cloud and ash fallout. As a consed the the city of Catania after t he late morni g Airport. The observation of the eruption c bud one hand, and the results from analytical s reconstructing the dispersal process and interring the eruptive mechanisms. Furthermore, the grain-size

variation with time of both the column hei

rv of five sing ab ve : allowed us to t ed <u>tephra</u>f Cdu r ea a to o orph obse

From early morning of 5lov Du ring he almo st 10 ho urs of w the growth and evolution of the t in the distal areas. More than 30 the fallout and within a ical, co mponentry and d by a new a nalytical

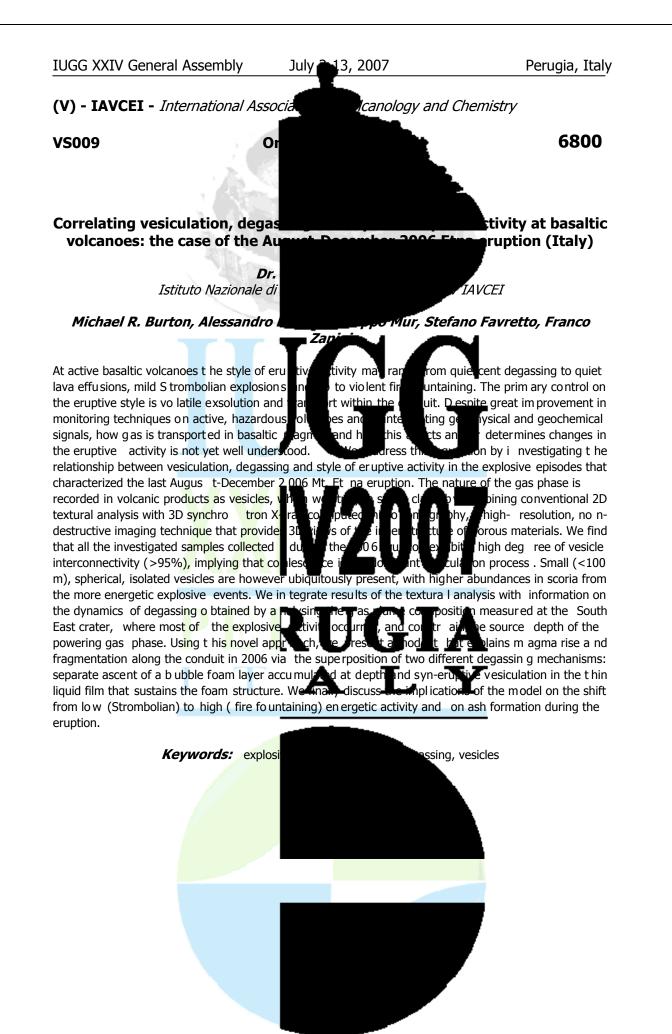
r otation of both the hd 30-60 a/m2 covered ernational Fontanarossa allout on the ground, on les on the other, allowed distributions of all the samples were used to evaluate the total grain-size of the deposit. By knowing the d direction, we simulated the pared it with field data.

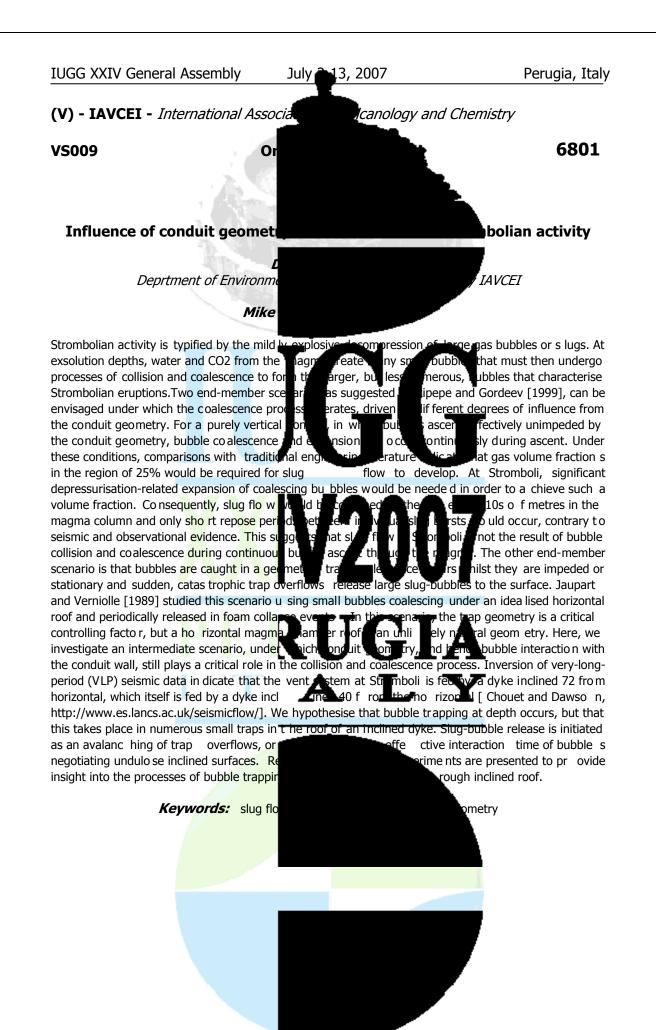
dispersal process and the de posit by using nd g e, simulating weak and The study of the 24th November episode at h th prolonged explosive events a t Etna require s detailed kno wledge of the wind dynamic and er uptive intensity with time with respect to shorter, high energetic 🛛 roxysms, w The eruptive condition s can be supposed almost constant.

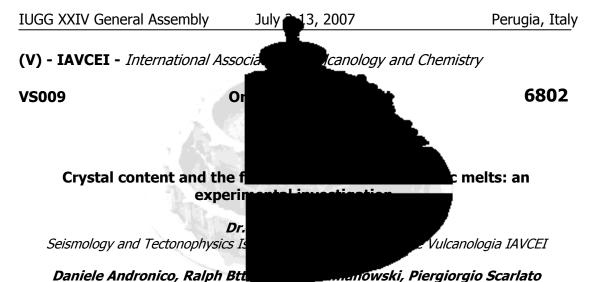
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Basaltic explosive activity varies largely in st violent S trombolian explosi ons as the m evidences, as well as theoretical studies, sug an important r ole in g overning the dominar the styles of basaltic explos ive activity, we crystallinity conditions at eruptive temperat e. similar to thos e of natural produc ts we melt na

pressure, electric, and seismic sensors, plus high speed camcorders, monitor the experiment. After each run we determine the text ure, degree of experimental products. Natural and experir and mineralogical assemblage. All other c experiments change dram atically as a f undior bulk of the melt reacts as a liquid when compremorphology of the melt left in the crucible and in a high proportion of fluidal morphologies in the clasts

le a oduc ype d nmon en**t** me at the crystall of explosions riment ra obta elts sample

spattering and ash-rich vith r ers. Texural a nd petrographic of the fragmenting melt may have investigate the factors g overning nt basa melts under differen t e of crystal-liquid rati os r ha along variable time-

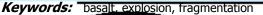
temperature p aths. Compressed gas releas ed below the melt provokes fragment ation, while force,



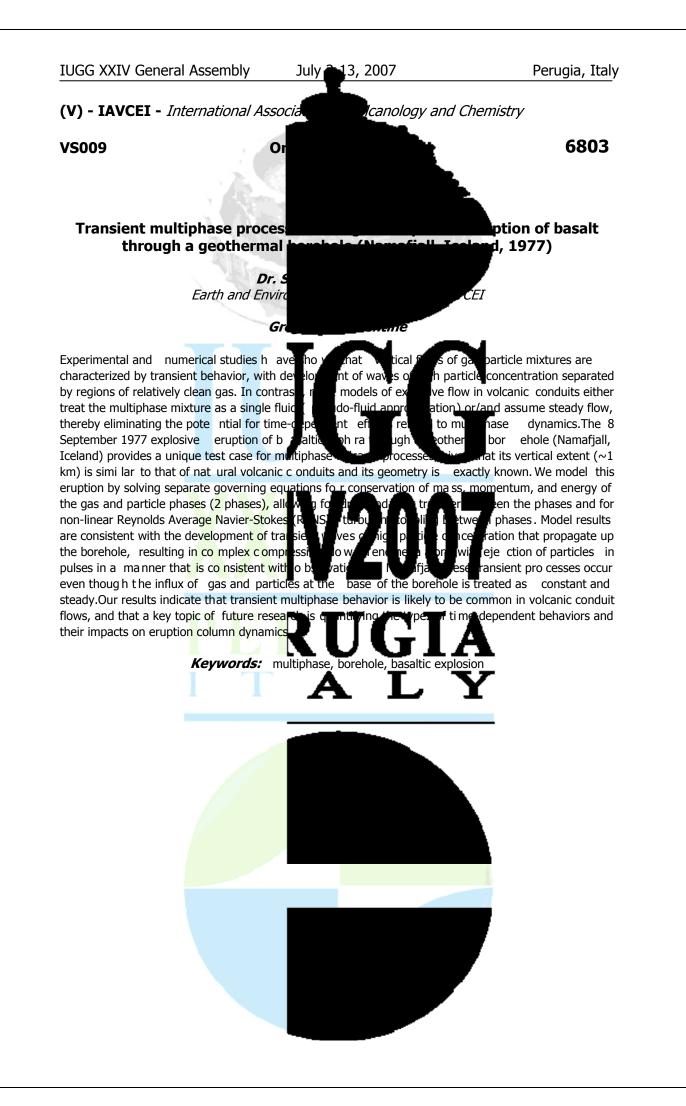
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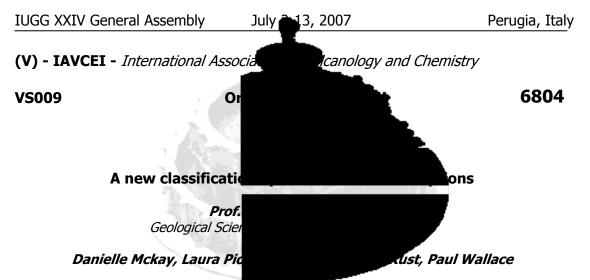
s crystallized in the e of shape, crystallinity, sults of fragmentation f poorly crystallized, the behaviour reflects in the

erupted. If ex tensively crystallized, britt le cracking domi nates during fragmentation of the melt , a s agments. We observe an testified by the appearance of large cracks off inverse correlation between t he crystallinit the time delay between cts a gas injection and melt fragmentation. This co stallization reduces the reased v elat lqq ability of the melt to dissipate stress by d uctile deformation, possibly by a combination of increased SiO2 content of the liquid (and th us a longer re erlocking of the crystal s ation time and ea rliè during viscous flow.









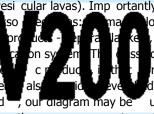
Existing confusion in the classification of basaltic eru traditional clas sfication schemes, which rely incorporate t he effusive activity that dor measurements have do cumented a wide va wt% in Hawaii an basalts to >7 wt% in ar exsolution depths and deg assing histories. F to rise and interact separately from the mell over several orders of magnitude. The compined r

(dominated by convective columns) to (dominated by passive eru ption of variably vesi cular lavas). Imp ortantly, these three end member eruptive styles represent very different morphologically and physically distinct erug that serve as the axes for our proposed cla advantages as schemes bas ed solely on r classified using solely field-based data. How First, as basal tic eruptions are often long-fived

is generate blel arac ón th s many ina in the H2Ō co ts) th<u>at</u> , in tu s the the re bh uptive

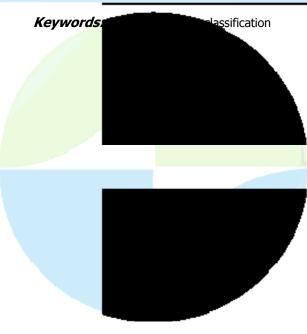
rge part by the inability of roclas tic deposits, to fics d ruptions. Moreover, recent ent of basaltic ma gma (from ?0.5 gi ves rise to a range of volatile y of b ic melt permits bubbles gma and gas may var y tes o

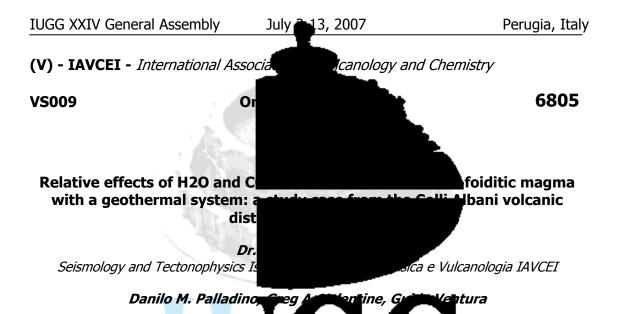
strombolian ( dominated by ballistic e



t ranges from plinian jecta) to effusive and produce three cones, and lava flows scheme has the sam e toric eruptions may be tages over past schemes. used to trace the temporal

evolution of an individual eruptio n. Second, as the axes represent eruption product s produced by ple wey to calculate the relative physically distinct processes, this method q importance of different eruptive mechanisms may be lated the relativ e velocities of the gas and liquid +/- crystal phases within eac tive pro duct po ses a CO πh en different potential hazard, the addition of such a sche me to probabilistic hazard assessment of basaltic gement. Finally our suggested ternary regions provides additional information for volcani diagram provides a simple deor of intermediate or risk mar ixed, eru ive styles the have defeated traditional classification schemes.





Understanding the factors controlling the recent critical point to volcanic risk a ssessment in multiple maar activity (70-36 ka) preserving and fluid-satur ated limestone country rocks Es integrated field and component analyses, in cate of ~0.3 km3 limestone reservoir representing ~159

CO2 pressure concomitant with strong C evacuation of population during the intens focus the inter ests on the Albano maar contact with a CO2 dominated geothermal K-foiditic dyke ascends thro ugh a CO2- H relative role of H2O and CO2 in interaction with

the P and T intervals of the CA geothermal system, inhibits the explosivity of magma-fluids interactions. Here we prop ose a two-sta ges quantitative dioxide interactions. During the first stage saturated sediments leading to an isochoric

magmatic events. In our model, depending during gas-pyroclasts mixture expansion to adiabatic case, in which gase s separate fro which pyroclasts and gases maintain the sa In both cases, water and carbon dioxide superheated s tate) or a re acting mixtur calculations show the dependence of efficient the mass ratio of magma to water and card

(CA) hy romagmatic activity is a ea. Here we focus on the Albano ractions between K-foiditic magma dit epi volume, by means of cava leaded to the disruption

seochemical data indicate volume that composition of geothermal fluids trap ped in the CA limestones reservoir is characterized by hig h si <u>na</u> processes. <u>Moreover</u> events such as the



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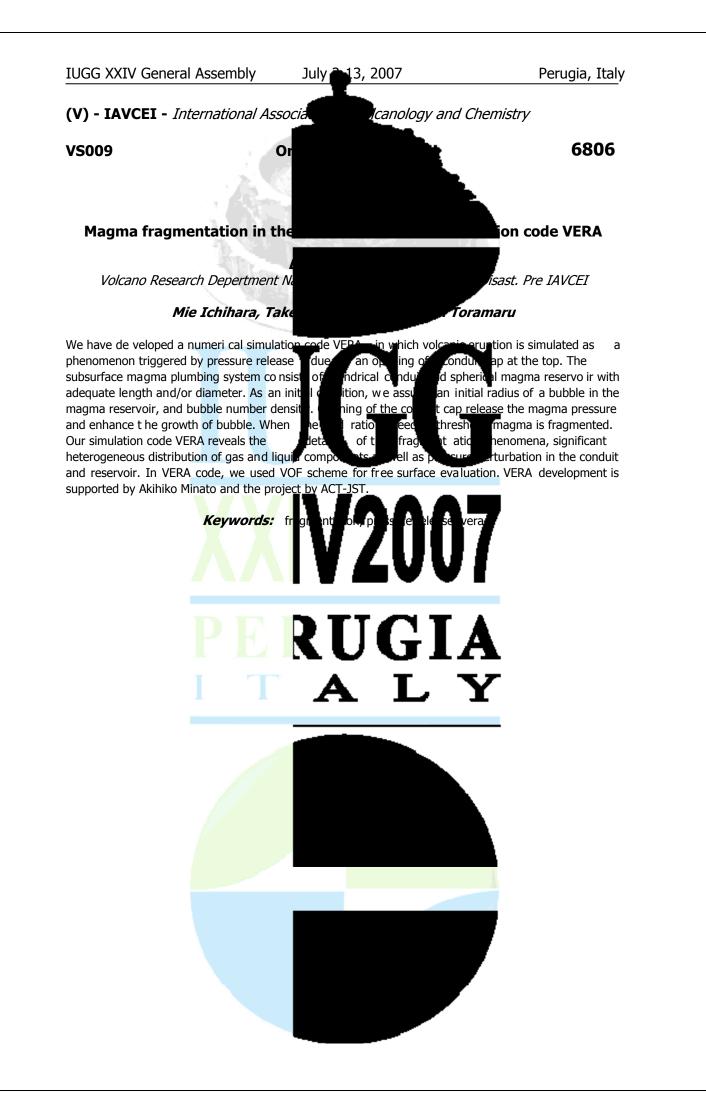
eismic swarm have to e e ffects of a magma ng a scenario in which a for understanding the

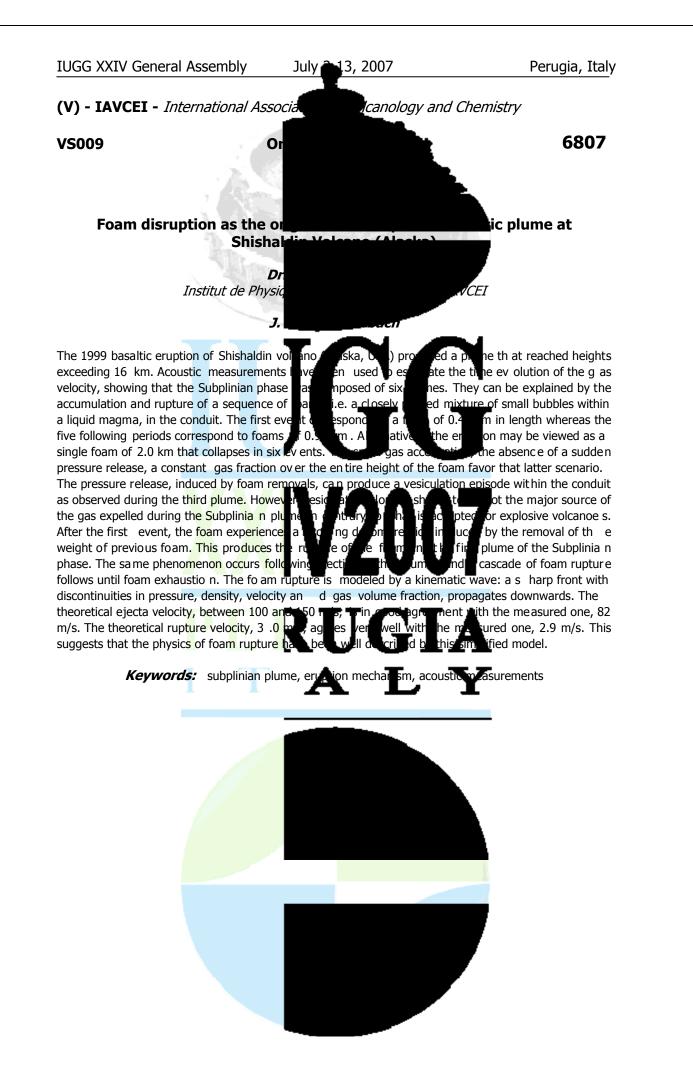
anic risk. Previous models (e.g., Wohletz 1983) suggest that the occurrence of a non-condensible phase, such as carbon dioxide in mamical of magma-water-carb on n uprising dyke to fluid-

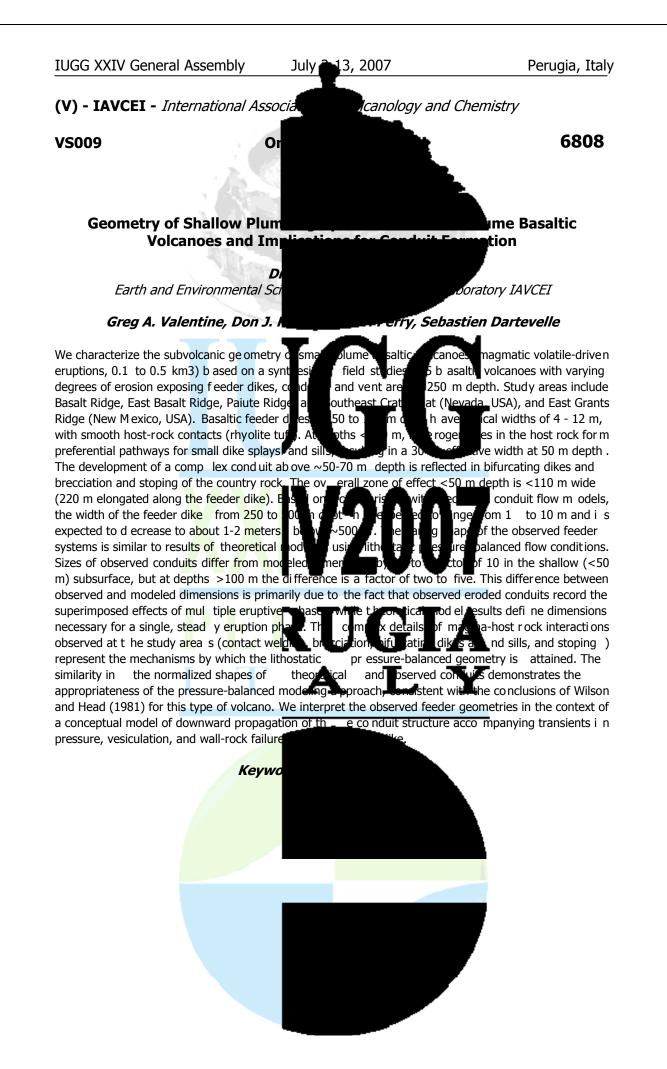
ntil ne seengt h of surrounding re rocks is excee ded. Then, we estimate the relative effects of CO2 vs. H2O mass ratio in explosive interaction with magma in terms of "efficien cy" (2010, expansion work vs. ma wha thermal energy ratio). Input data for efficiency estimations take interact tailed field nt t he eco nstruction o f Albano stratigraphy and component analyses s howing a wide ra nge of lithic/juvenile clas ts volume ratio (i.e., 0.05 to 0.95) recording diff erent eruptive mechan isms spanning from phreato magmatic to mostly

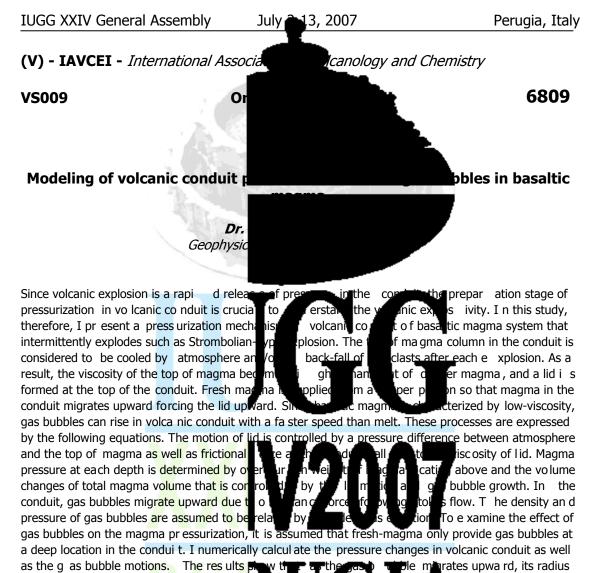
> heat t ransfer from magma to gas es consider two end members: i) an and ii) an is othermal case in heat flux during expansion. reacting mixture (i.e., the t hermodynamic superheated state upon



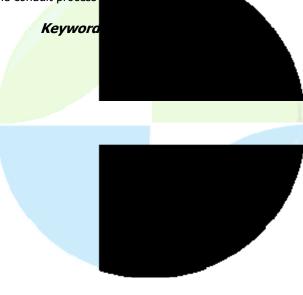


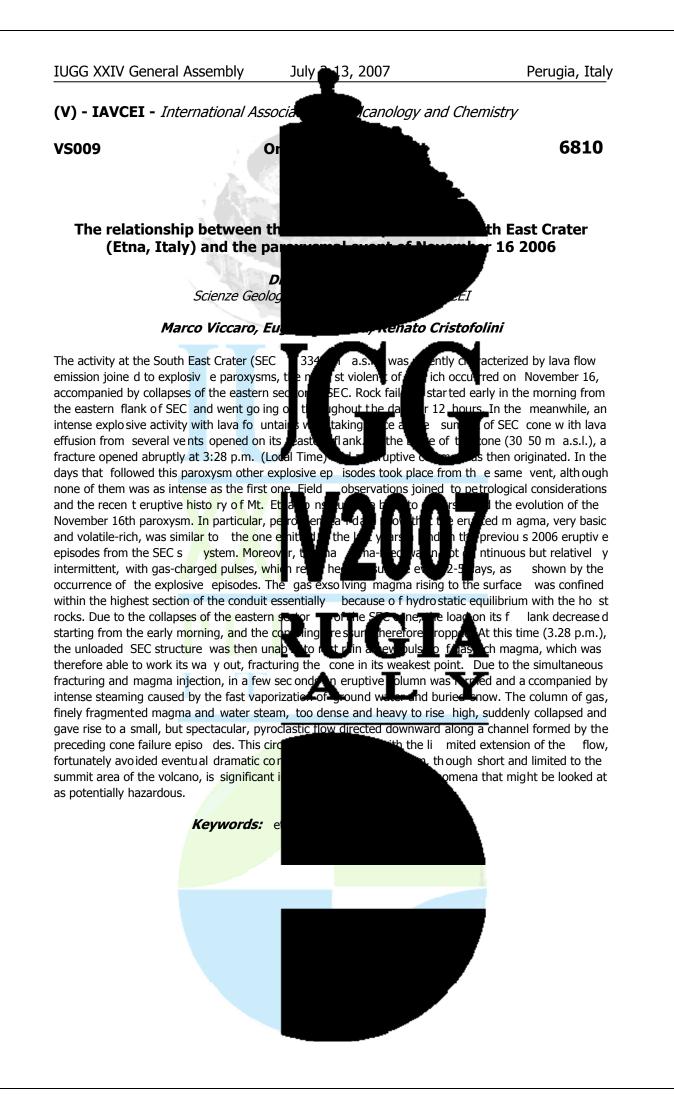


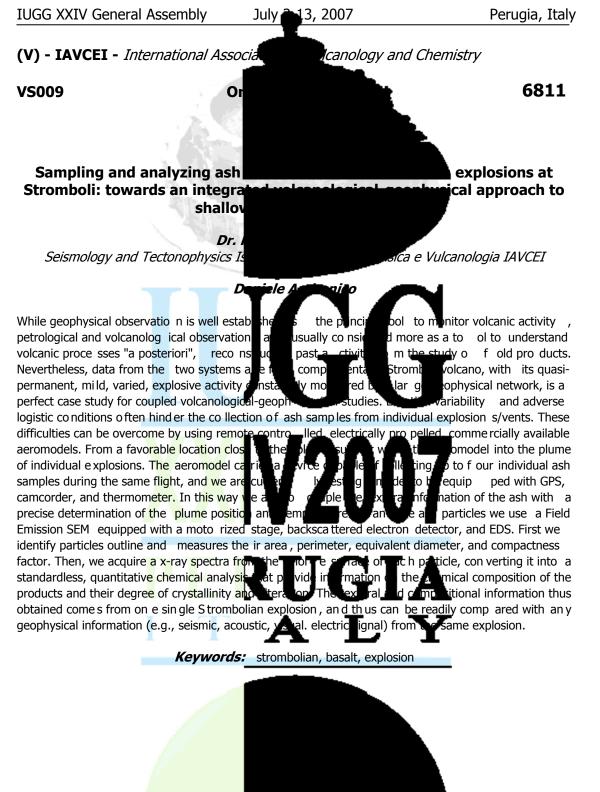




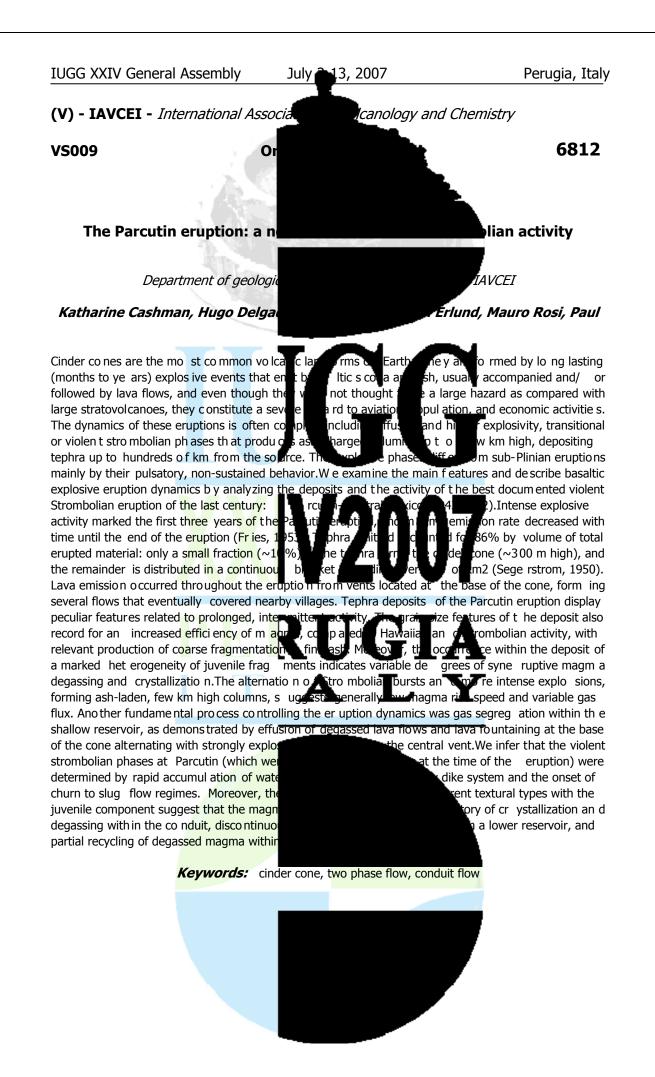
increases due to de-pressuriz ation to de c ge bi ncy force. Hence, gas bubbles increa se its volume rapidly and hey as cend. That is , he m e t a pressure of magma in the conduit rapidly increases with time. Numerical calculations in various settings show that magma pressure can increase more that a few per nt of lithospatic pressure due to the lid of conduit when sufficient amount of gas bubble rid up and/ viscosity large. In case of no gas bubble in the conduit system but a constant magma supply, over-burden pressure makes the pressure of magma in the conduit increase, but constantly. Lhese temporal changes in magma pressure can be round volcanoes, which enable us to detected by re cent geodetic measuremer provide constraints on the conduit process

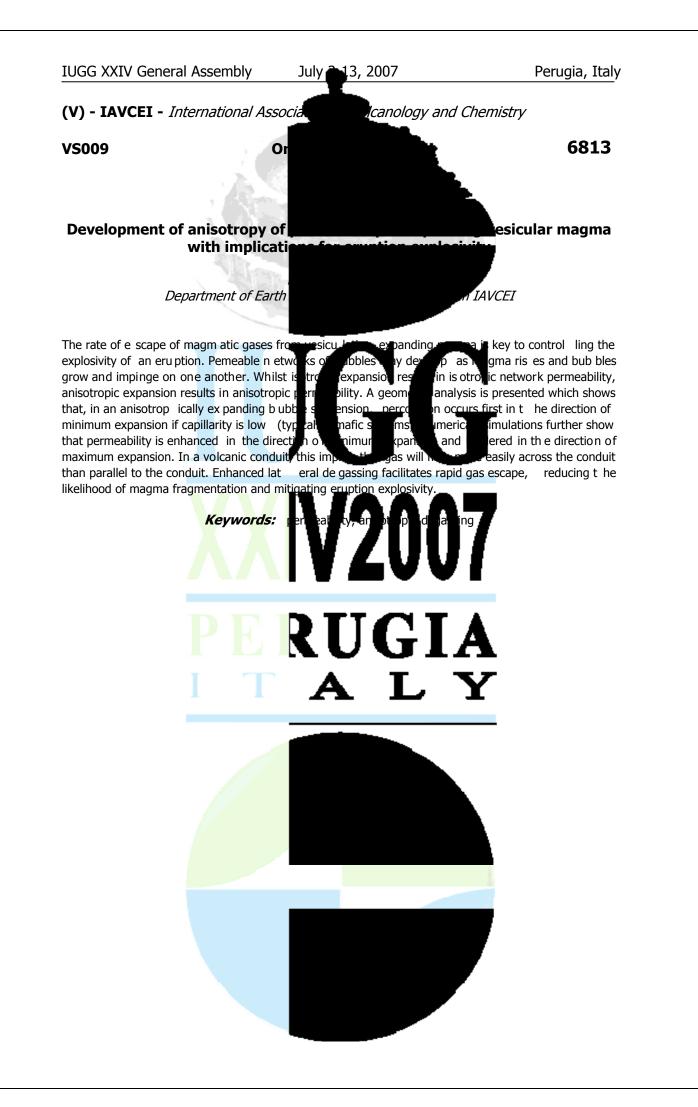


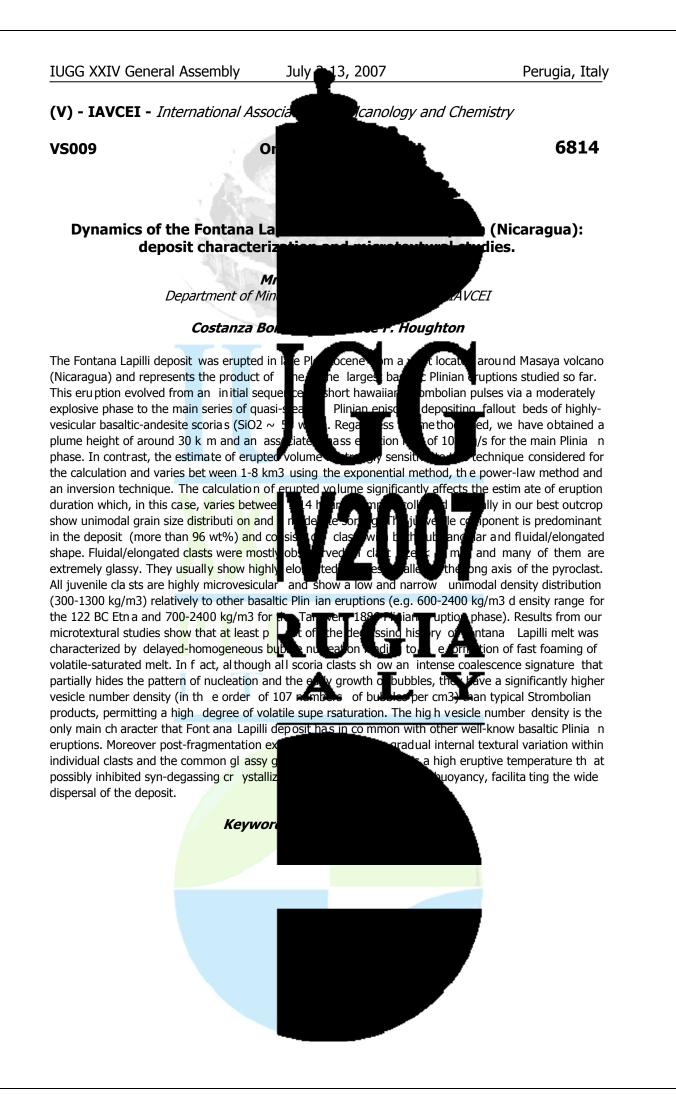


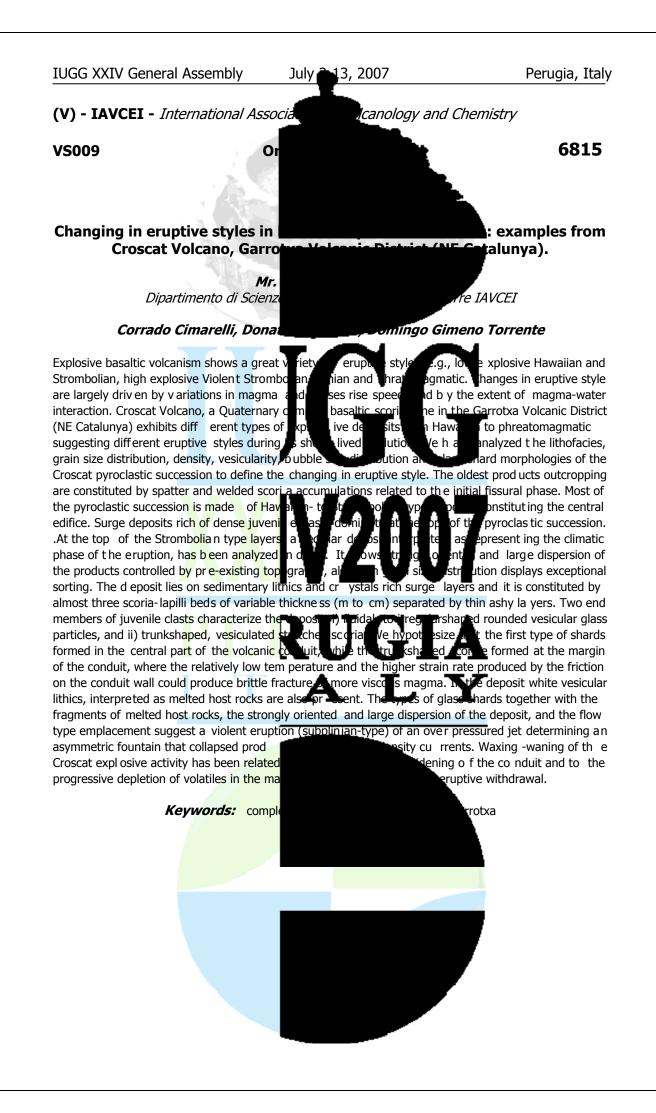




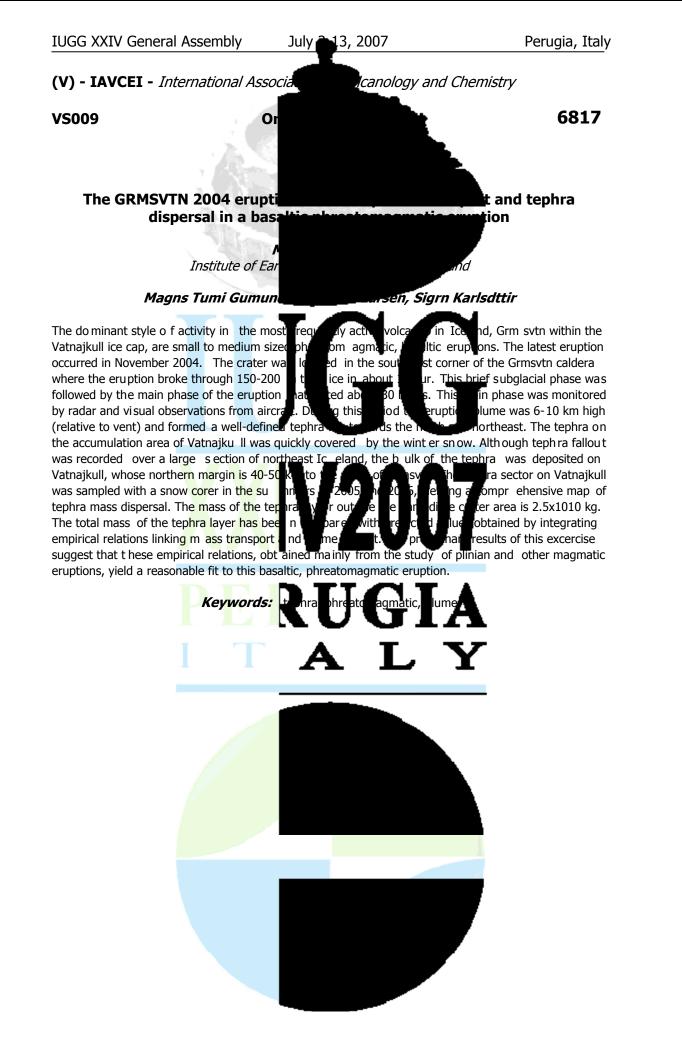




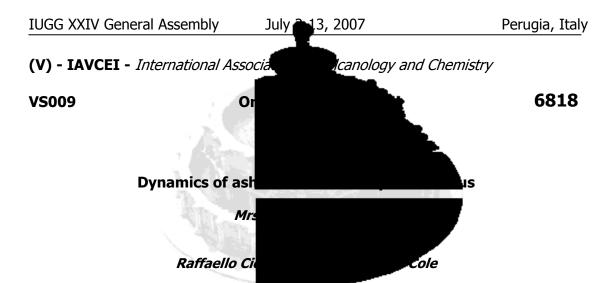








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Several erupti ons do minated by ash emiss eruption (3900 y BP). This type of activity have significant problems for emergency planning an on the depos its of s ome of these erupti environmental impact. The deposits of these laminated coarse and fine ash, interlayered to perio ds o f co ntinuous ash emissio n al samples are characterized by the contemporary pr

All these fragments show a large textural variability\_\_\_in terms of vesicles content, size, and shape a nd microlite content. The juvenile products ar assemblage formed by clinopy roxene pher leucite, pyroxe ne and plagioclase. Matrix between the composition of the residual categories of juvenile fragments were dis morphology a nd textural features, elaborated t analysis. Cryst al size distribution of the

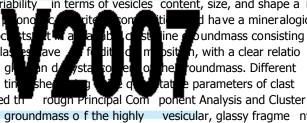
representing the most abundant juvenile f magma degassing and magma ascent. The primary role of magmatic volatiles in driving

ion have urred at Ves bee ed in over zard asse sm aimed at de bn s <u>are mai</u>n nor la 5e with ifferer

atter the A vellino plinia n ven though it can pose past, We present the results of a study their eruptive dynamics and rmed by thick sequences of thinly nis sug ts that they are related stromb olian activity.A ll vio uvenile fragments, from

light-brown, highly vesicular, glassy pumice, to black, moderately to not vesicular, microlite-rich scoria.

fion



d have a mineralogical oundmass consisting of h, with a clear relatio n oundmass. Different parameters of clast vesicular, glassy fragme nts,

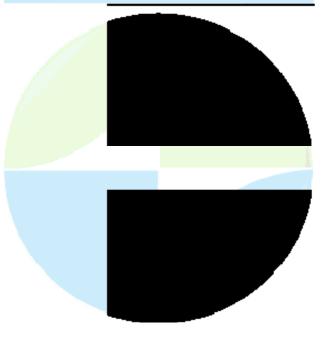
he timing and the dynamics of hly ν ular material suggests a poradic occurrence of ĥę ргп

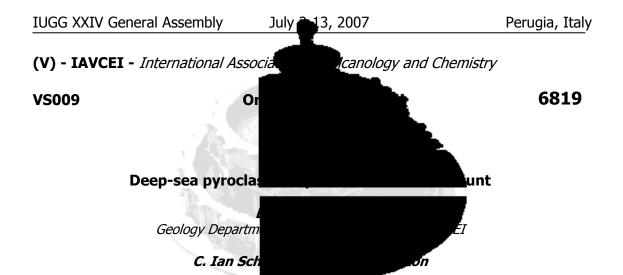
sedimentological evidence of phreatomagmatic activity in the deposits suggests a minor role of magmawater interaction in the dynamics of ash eruptions t Vesuvius

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Keywords: volcanicash, crystalsizedistribution, eruptivedynamic





Loihi seamo unt is the youngest and so uthernmost summit is ~1 km below sea level. Previous tudi >10 m thick, and in October 2006 such summit with Hawaii Undersea Research Lab interest. At the first, 20 m of coarse-grained basal few metres of the deposit are character internal layering, separated by discontinuou across. Upsection, deposits coarsen to lapil domin

containing co arse lapilli to small blocks dis persed throughout. Three populations of clasts typify the lapilli fraction of the deposits: ragged, high texture; and a variety of lithic clasts showing There are als o occasional mixed clasts, particles. We infer that all the particles are fall or currents from an eruption of a very minor comp onent of the de posits, sugges ting

transport or deposition. The fragments' vari able vesicularity (<50%) indicates a role for magmatic-gas exsolution in driving the eruption. Very glam water guenched the fragments prior to dep microlite populations, as well as particle s microlite populations, as well as particle s the population of the population of the population of the population. The second site of interest has a basal section dominated by thin lava sheets with cavities that contain sideromelane ash depos fragments termed "Limu o Pele". Several metre on lava sheets, comprising mixtures in var

absence of ap parent seafloor-weathering with more polymict grain populations. The considered to result from quench granulat seafloor lavas. Limu fr agments have beer eruption, or fr om seawater bubbles rising with sheet lavas favors the latter interpreta

Islands group, and its hiclastic deposits locally ad vo 5 dive onto Loihi's seamoun t es I V. Two sites ar e of particular ing and overlying pillow lavas. The parse ash with cm -scale beds nts, i seve ral centimetres lacking distinct internal

stratification. Beds are defined by la yers a few cl asts t hick of coarse r fragments, with some bed s

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with a crusted surface ent trans port histories. weathered lithic hses en deposited directly by

urce. Fine ash is a very rragmentatio n o r/and so me winno wing during

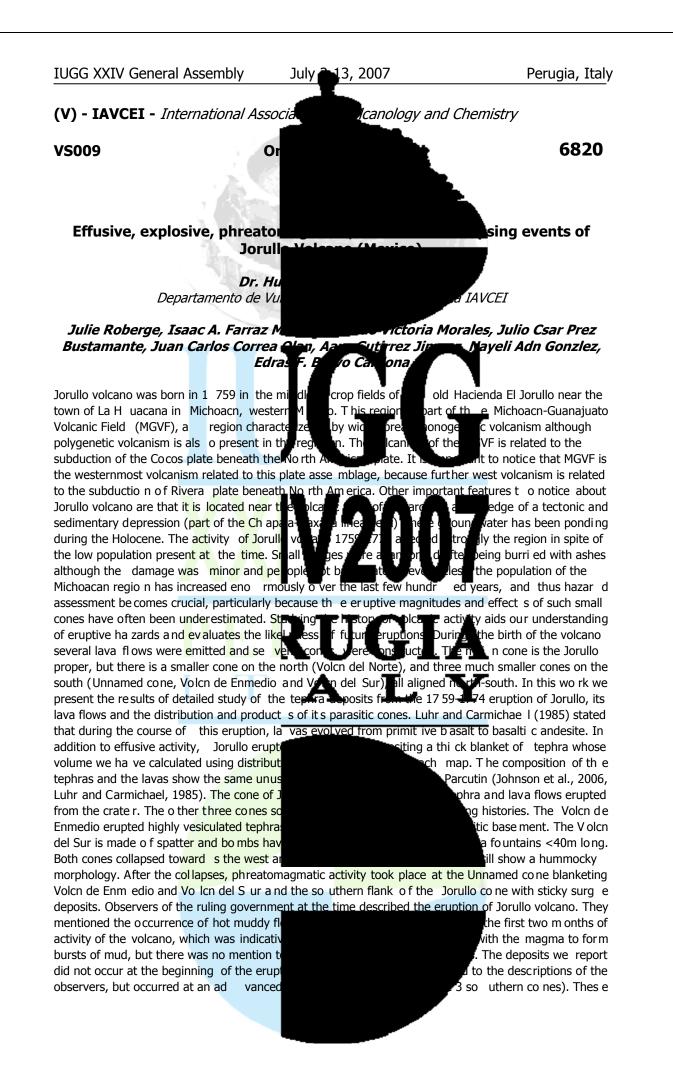
will assess vesicle and

ain int erpretations of

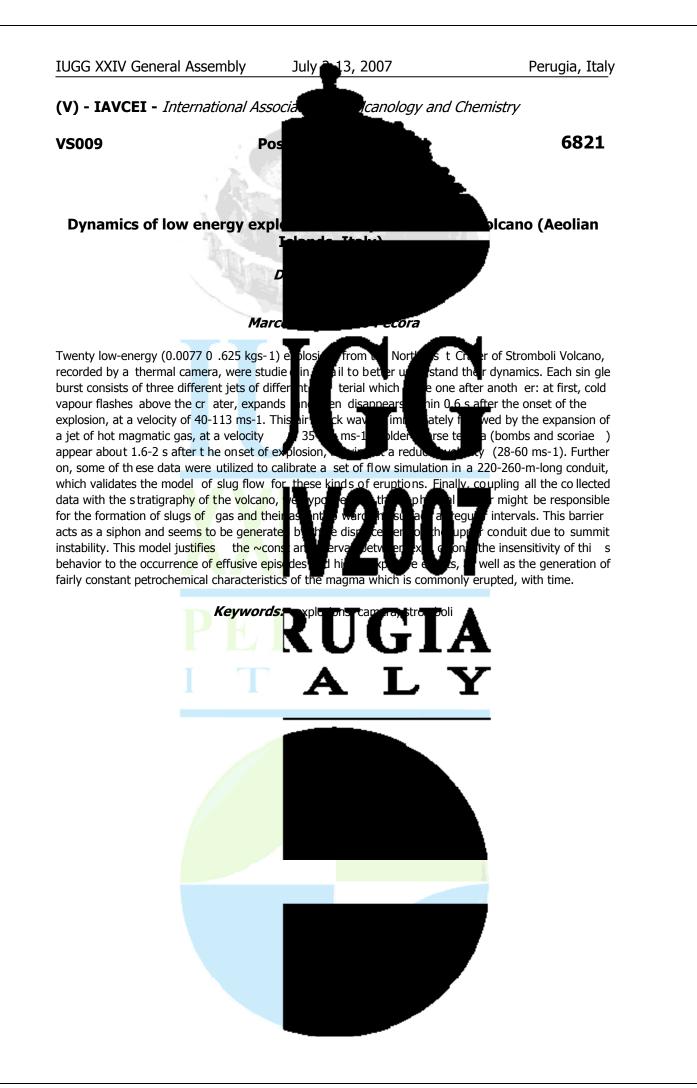
ding in the deposit suggest that

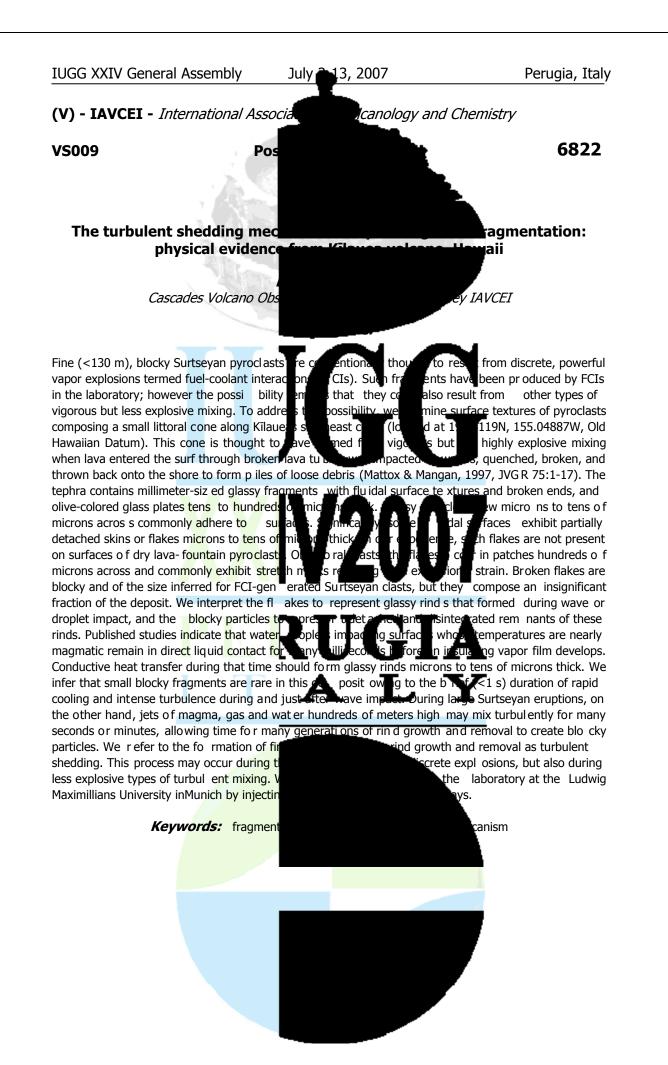
igh proportions of small thin sheet-form ed ash and billi-ash overlies the basal ying proportion of limu fragmen ts an d den se blocky sideromelane grains. We infer that these are primary deposits because of the consistency of clast type, ces, scouring, or inter vening deposits ed with the limu fragments ar e ociated with e mplacement of om a fount ain-like style of ere, the close association

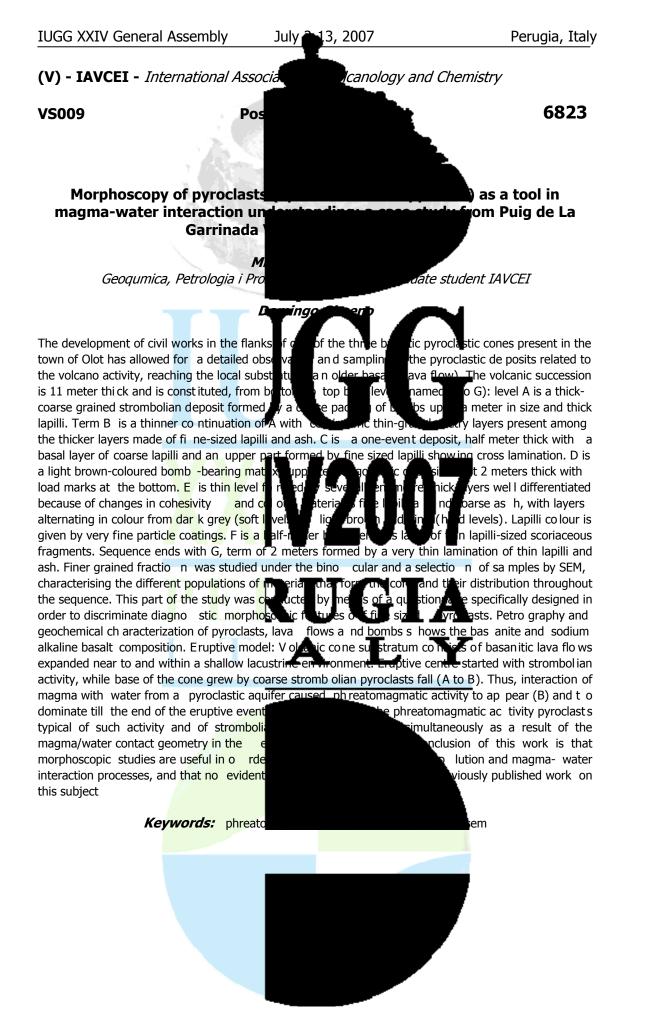
Keywords











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#### IUGG XXIV General Assembly

#### July 13, 2007

Perugia, Italy

(V) - IAVCEI - International Associa

## **VS010**

# Symposium Modeling the plumbing system petrological, geophysical and f

Convener : Prof. Benedetto De Vivo,

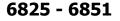
An understanding of the internal structure objective of v olcanology. The composition magmas are stored, the way they differen

levels have important effects on the style and magnitude to understand the behavior of active volca strategies and for forecasting volcanic erupt investigate the internal struct ure of volcano adequately investigated by employing a sing Investigation of major, trace element and is can furnish sig nificant information on the n inferences on chemicophysical conditions geochemical studies can provide basic information

on the structure of the base ment beneath the active volcances, on the depth and size of magma chambers and on melt migration within the of fluid inclusio n in pheno crysts and xeno pressure) of fluid inclusion entrapment, i.e. or that the maxi mum of infor mation on singl ad results of these different disciplines. Integr furnish t he m ost reliable models for volc

collecting scie ntists with di fferent expertise (P etrology, Geochemistry , Fl uid Inclusio n Petro logy, Seismology, etc.) to discuss the potentiality of integrated research in the modeling of volcanic plumbing systems and how this can help in es tablis more reliable forecasting of volcanic erupt and possibly other IUGG associations.

canology and Chemistry



### tegrated

or the ough often overlooked dep ths and conditions at which signation of melts at d ifferent crustal



ns These issues are crucial le basis for monit oring studies have been mainly used to kity of volcanic systems cannot b e res a multidisciplinary approach. lir const ituent phases

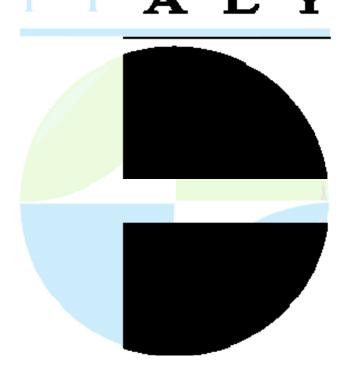
on processes, allowing ation. P etrological a nd e n working in the past

and, presumably, in the future. Seismic and gravimetric data have been widely used to have information

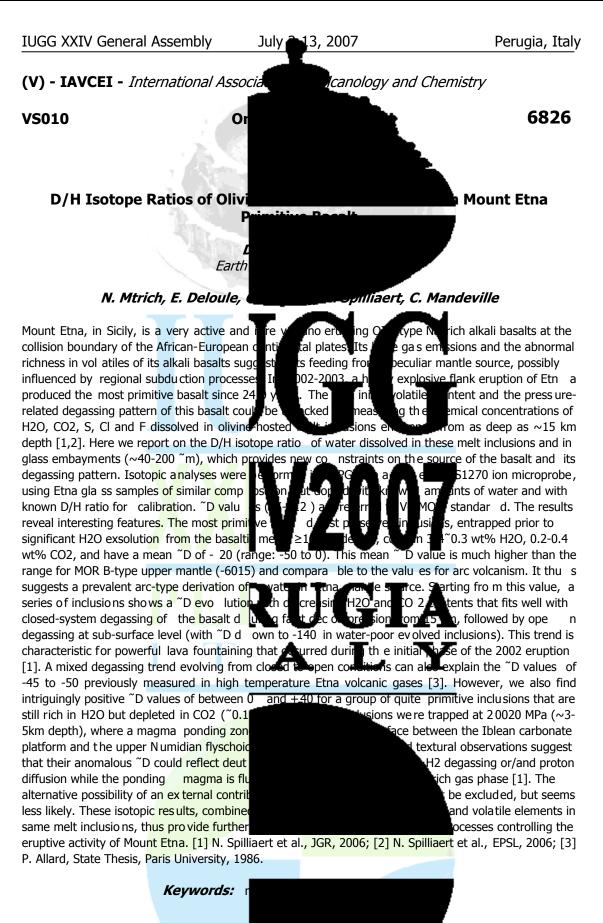


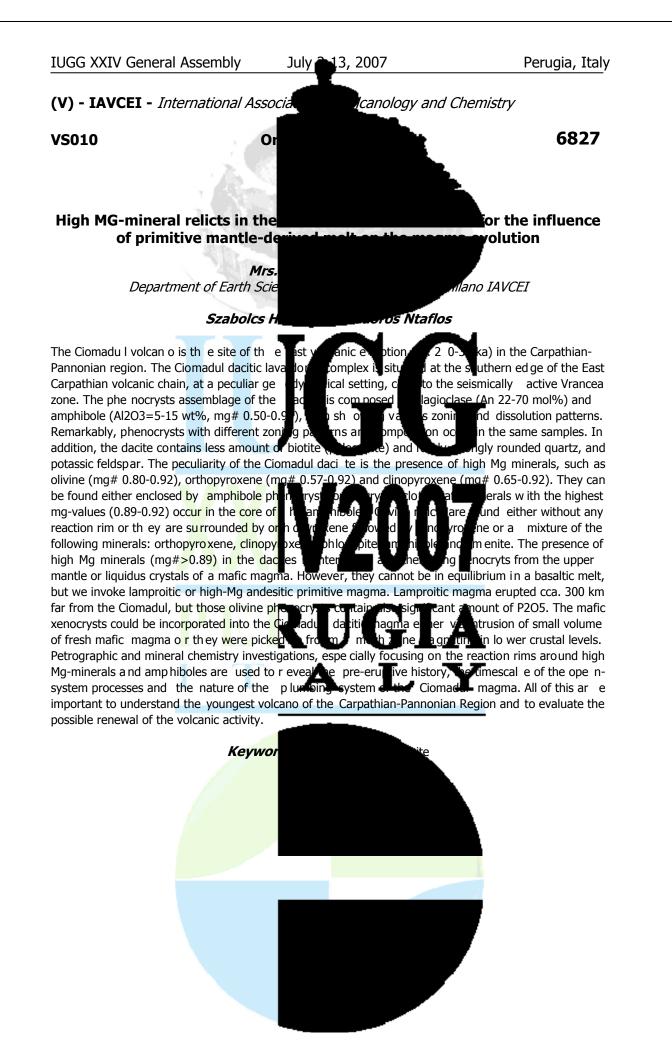
mposition and density straints on depth (i.e. . It goes without saying udies that integrate the on has the potential t o osium has the aim of

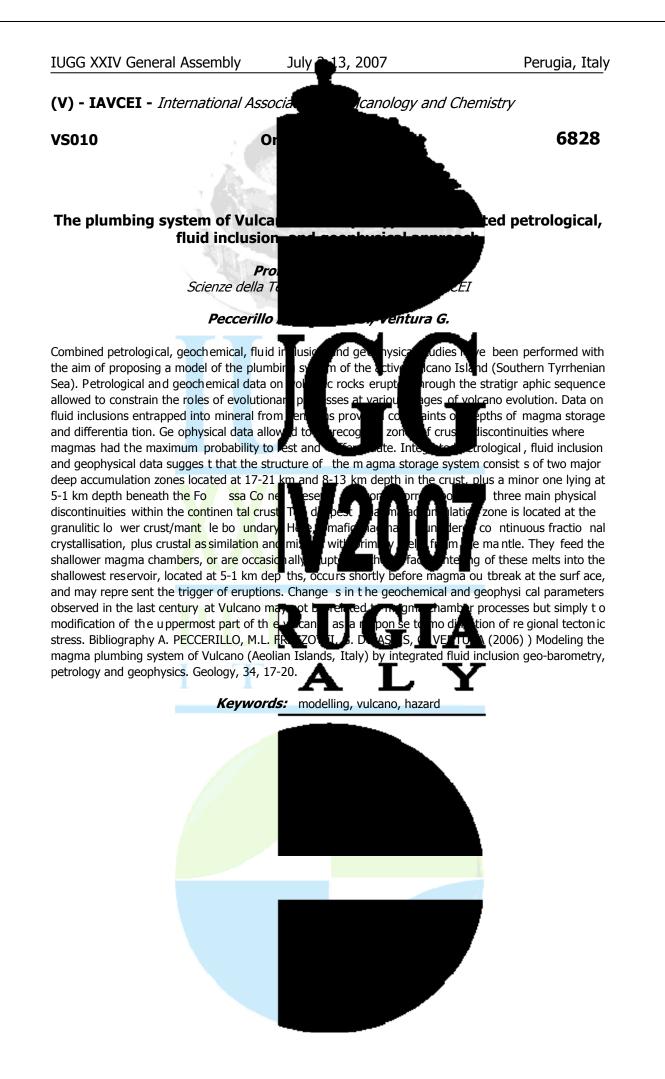
nitoring and in obtaining ano n erest to I AVCEI, IASPEI

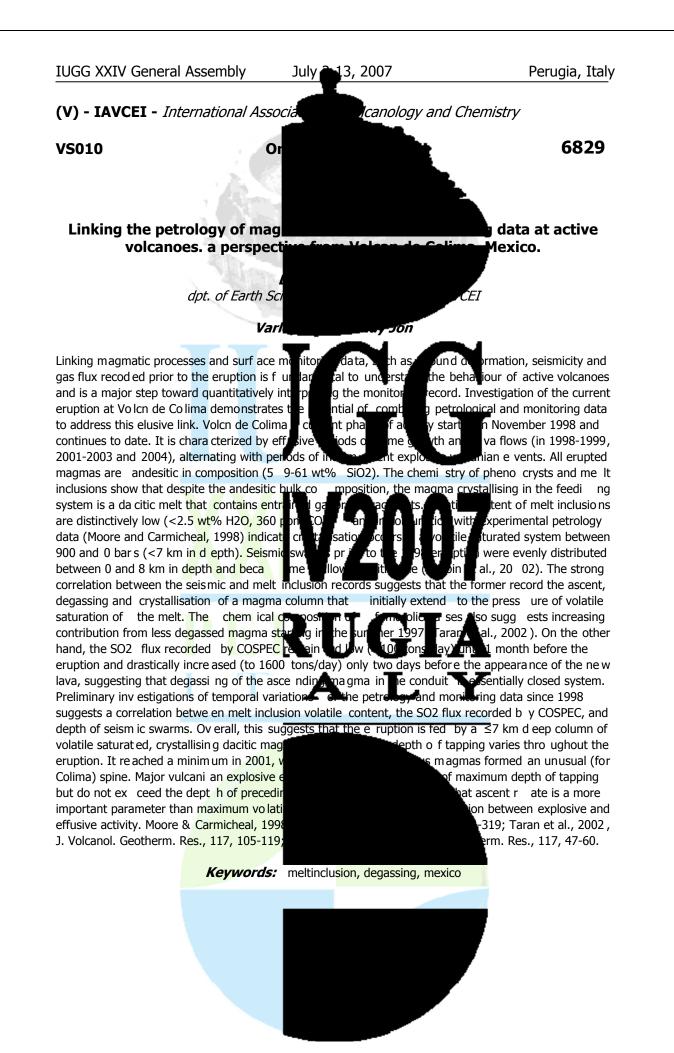


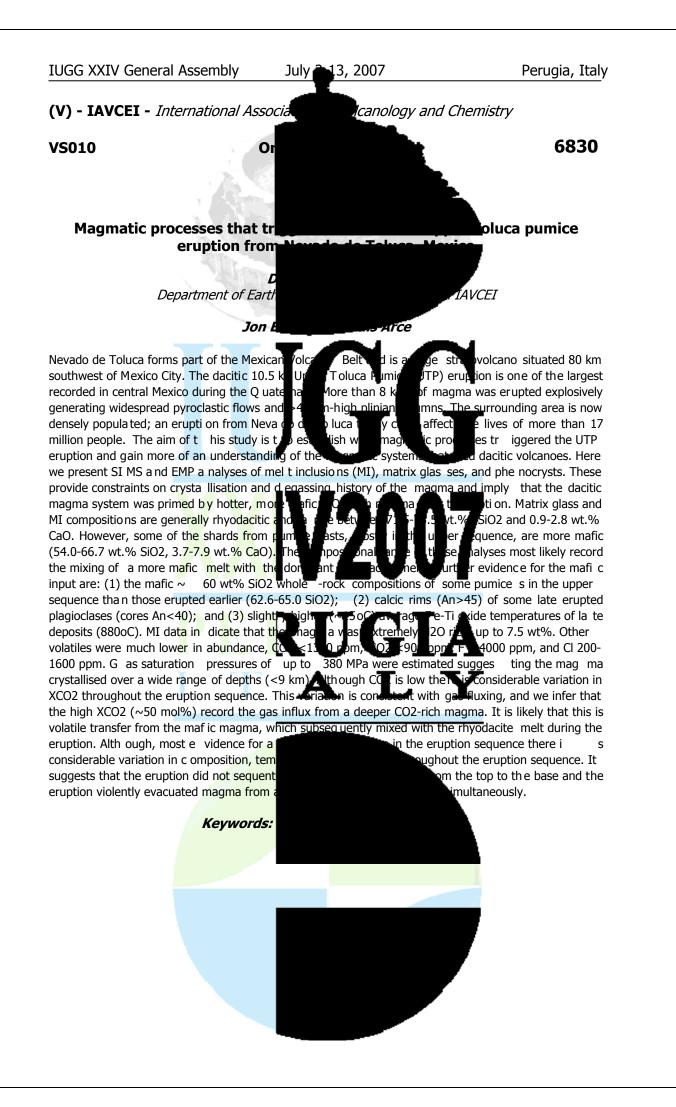


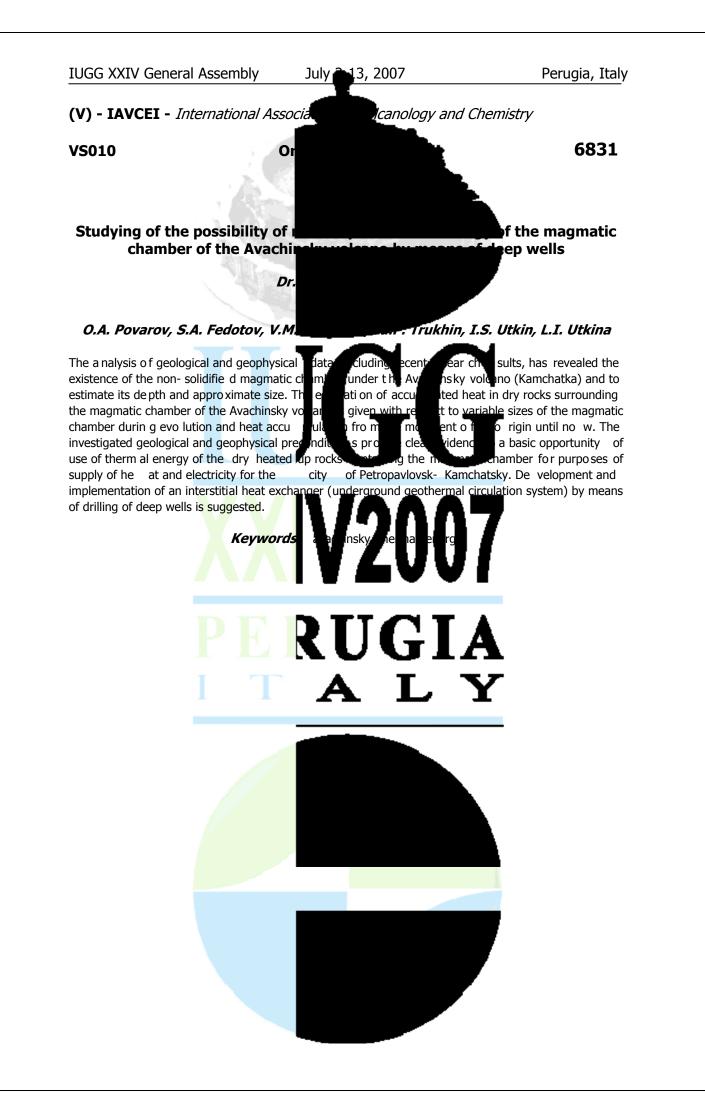


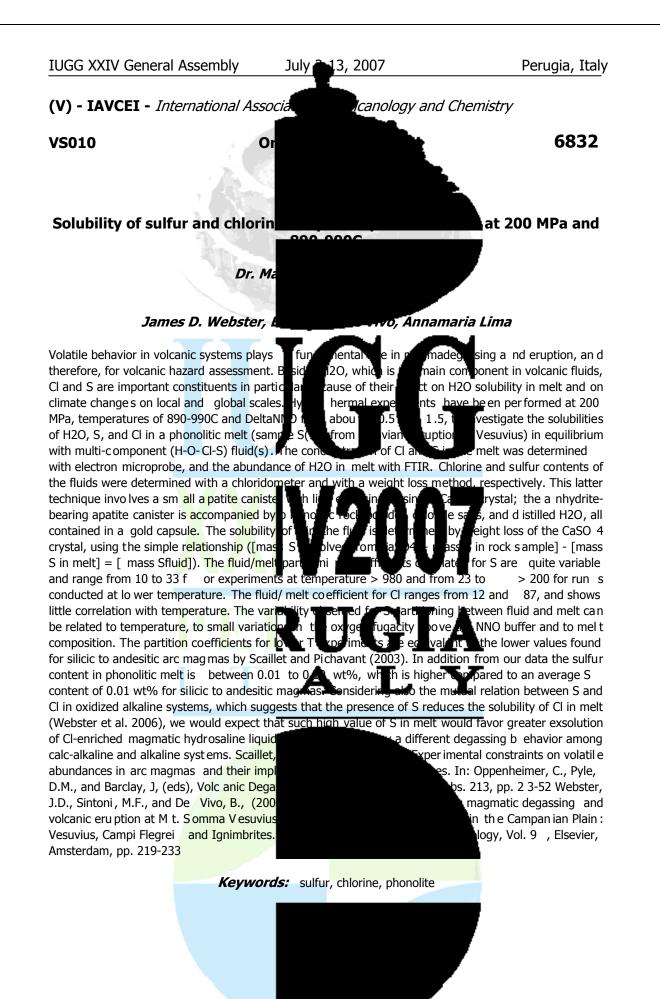




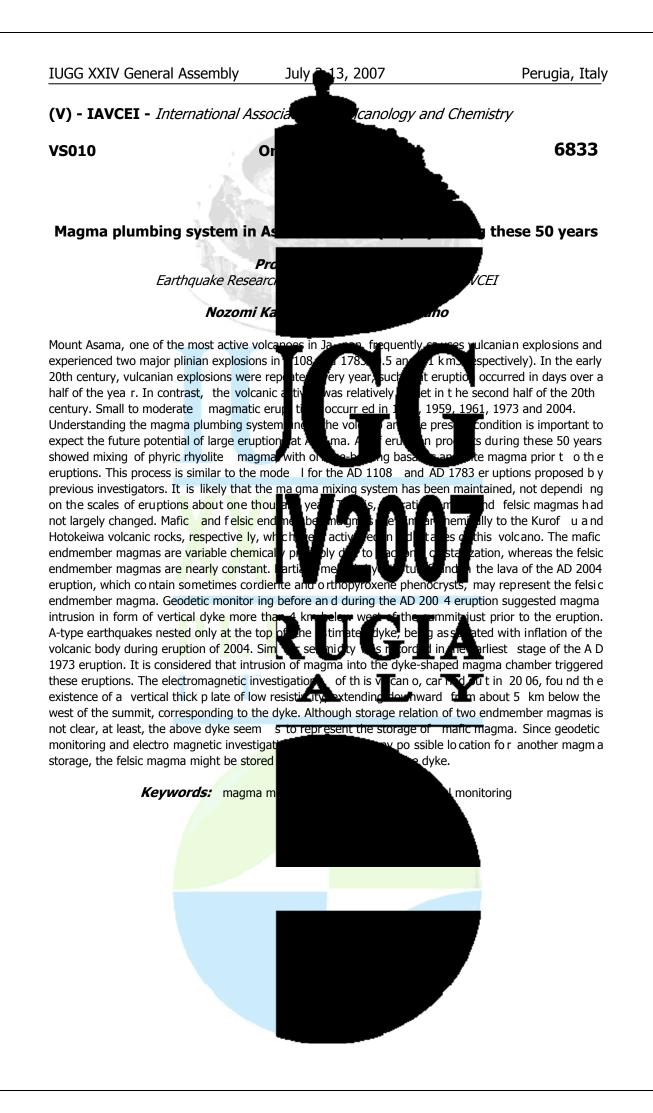


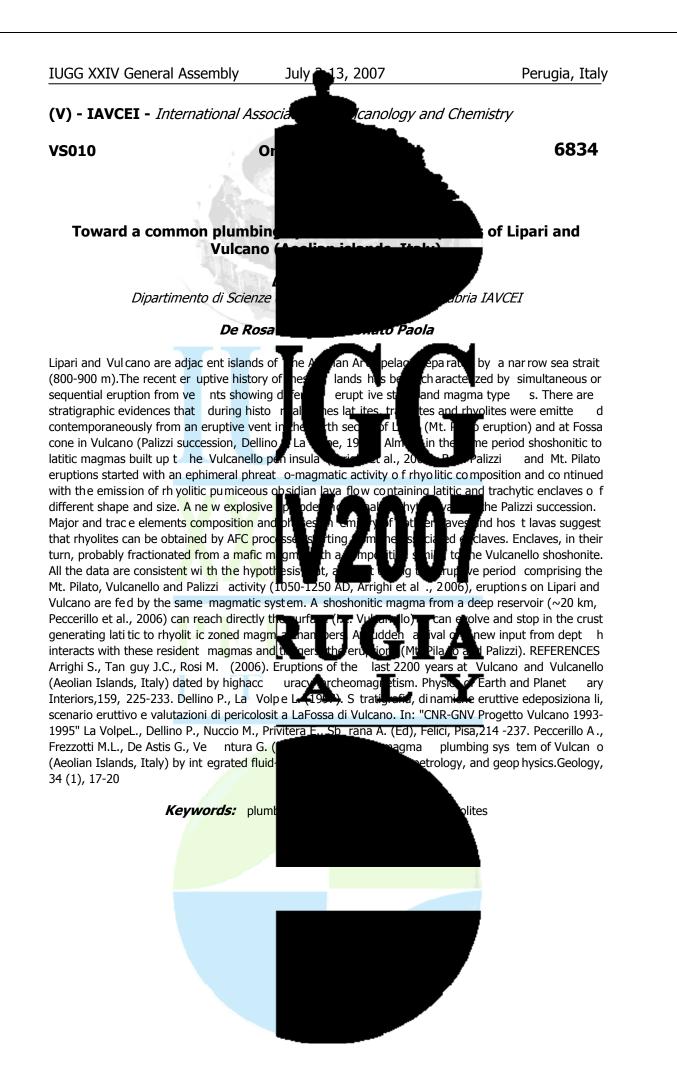


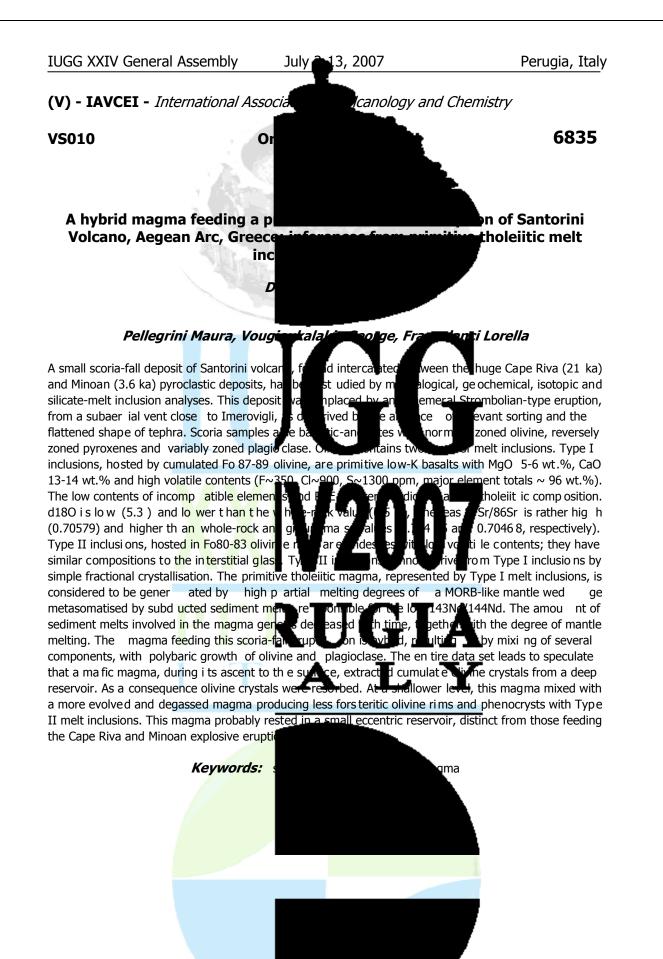


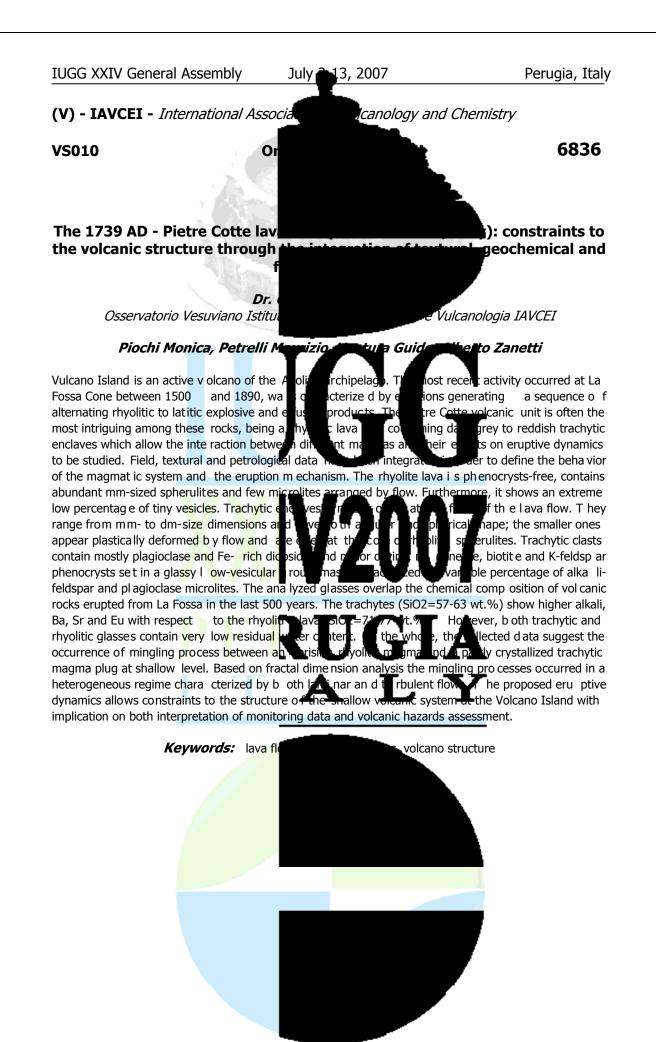


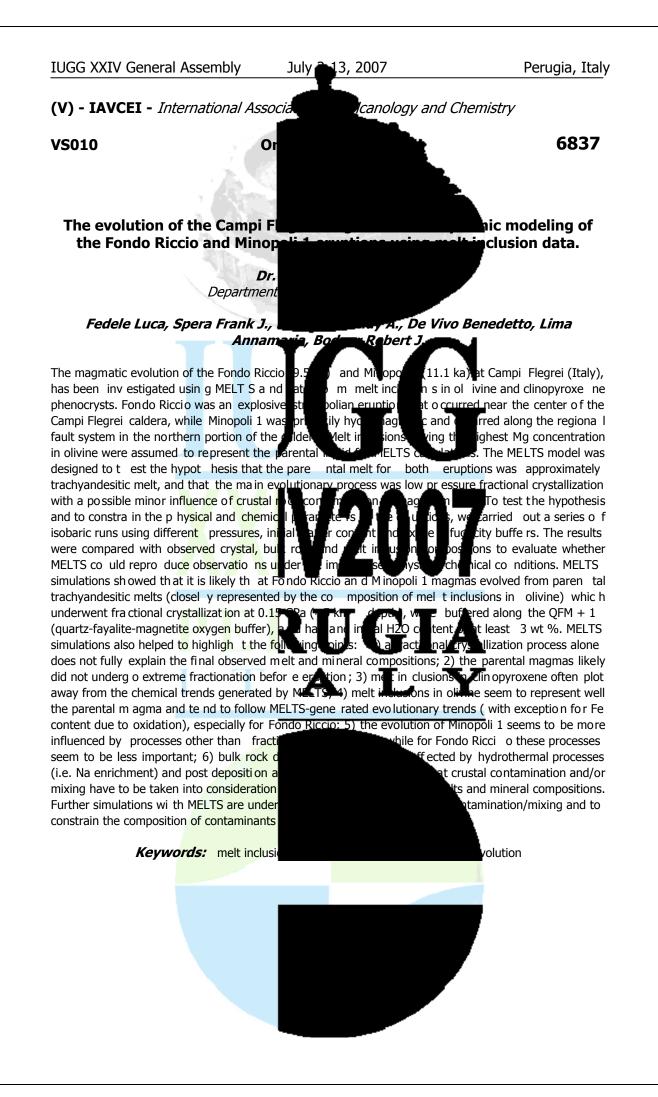
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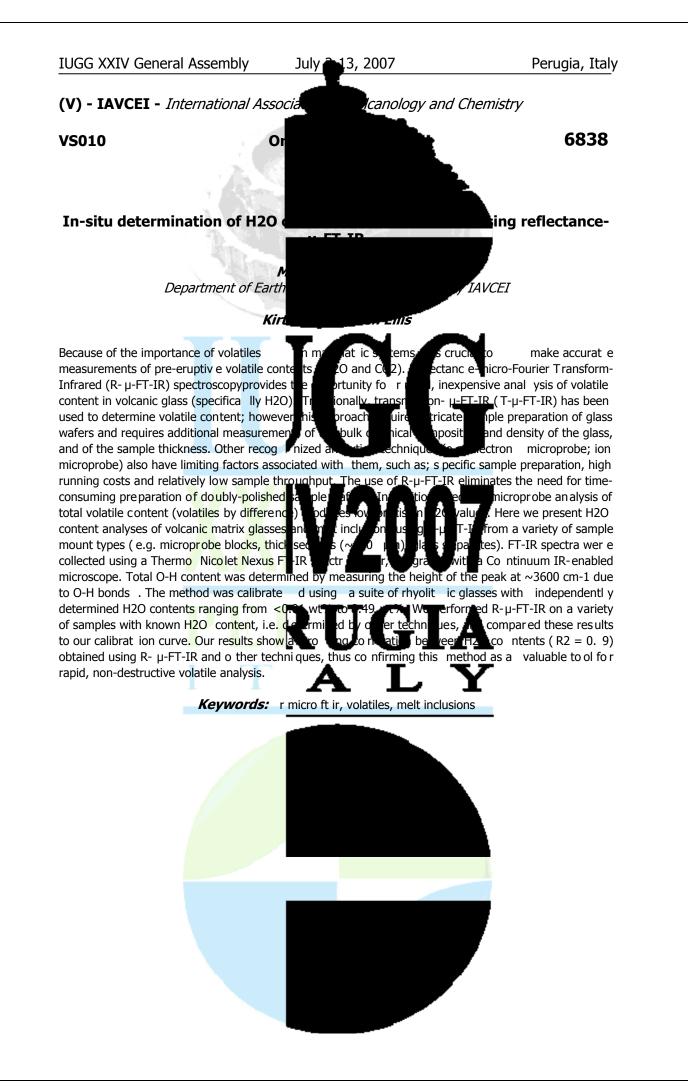


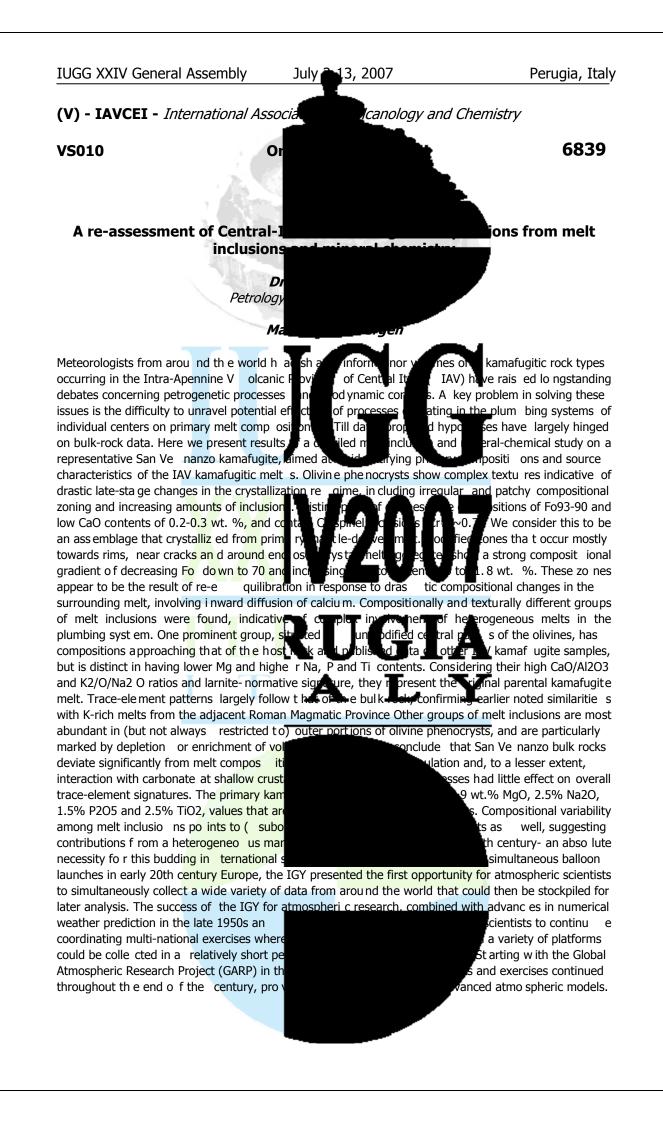




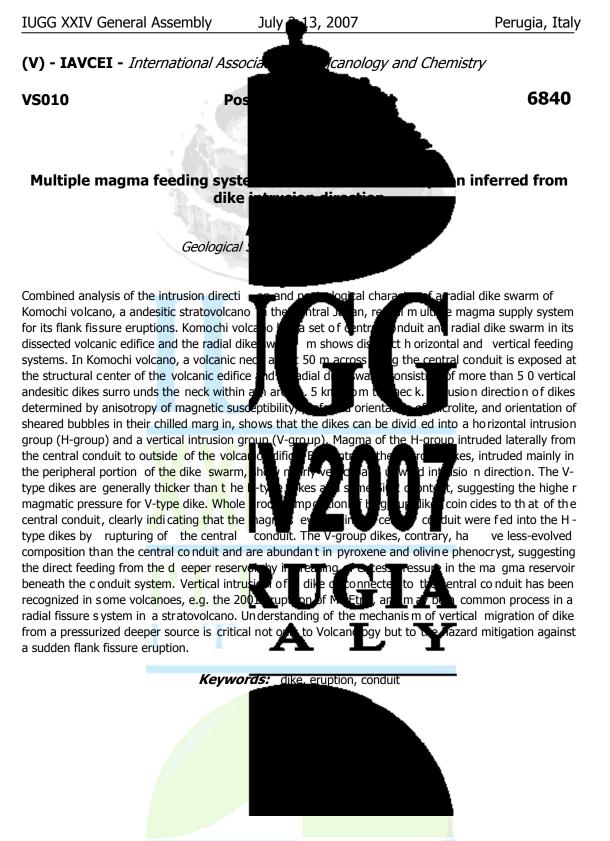




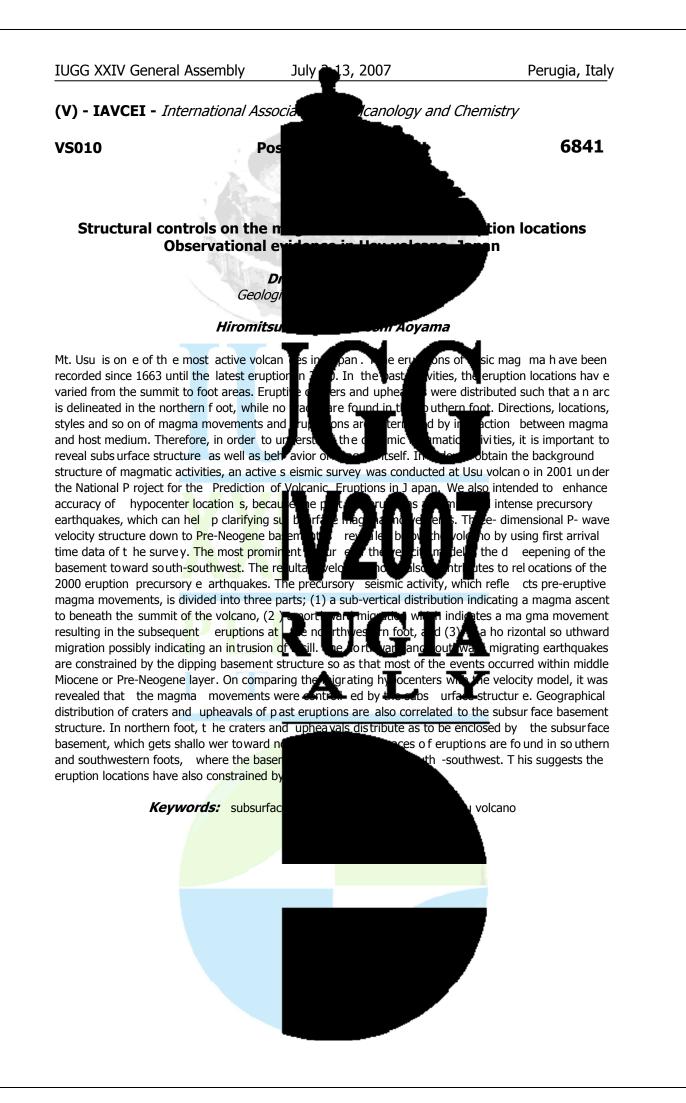


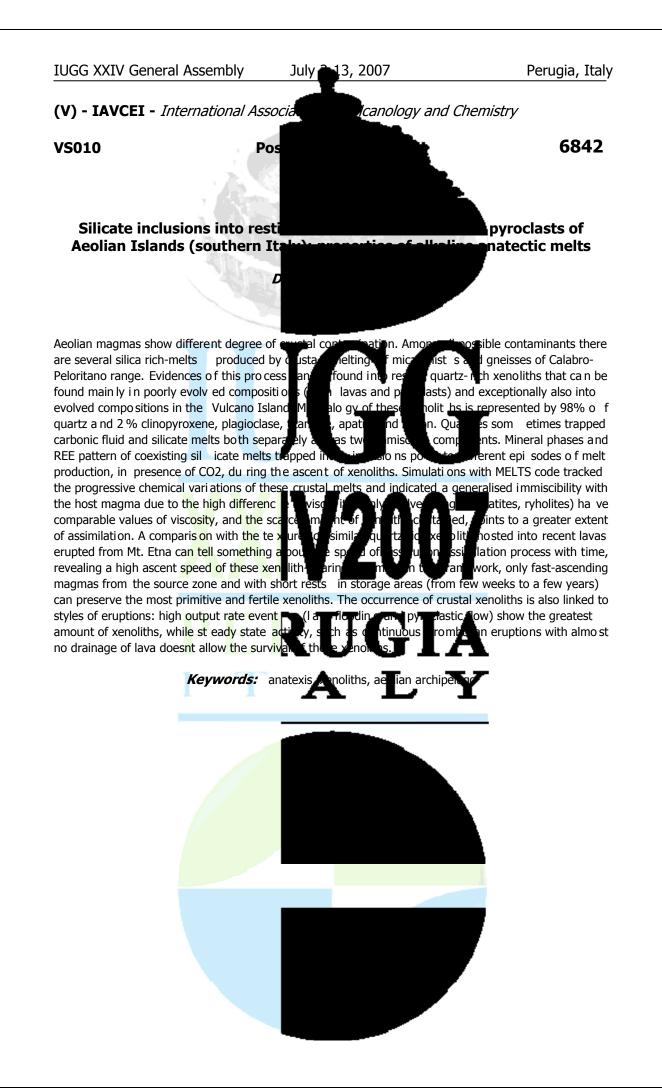


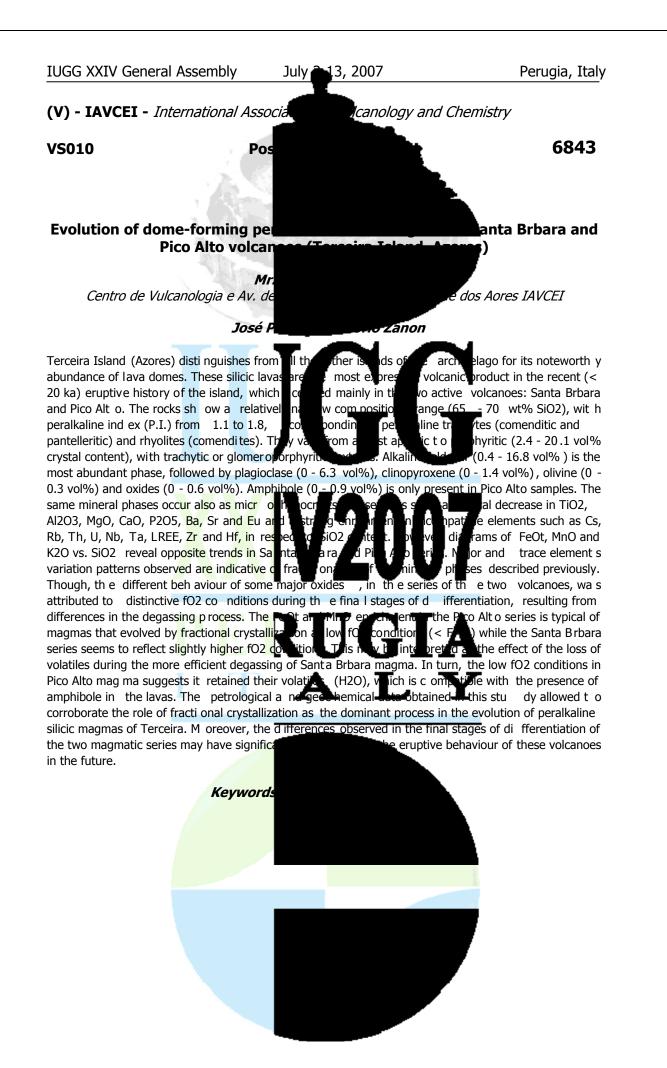


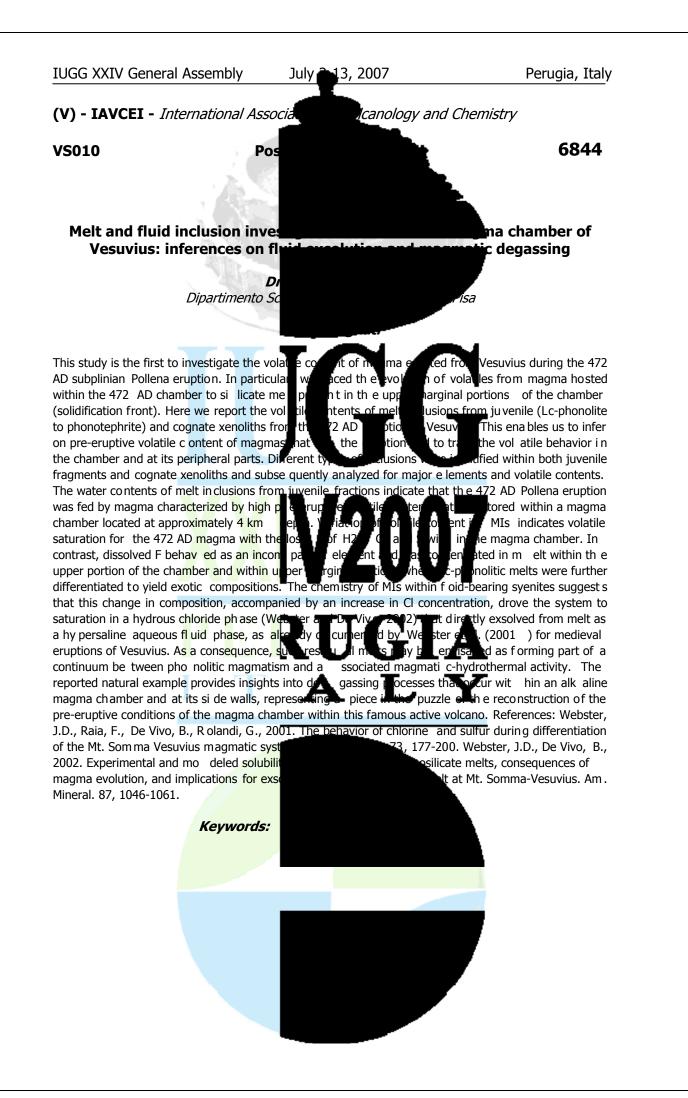


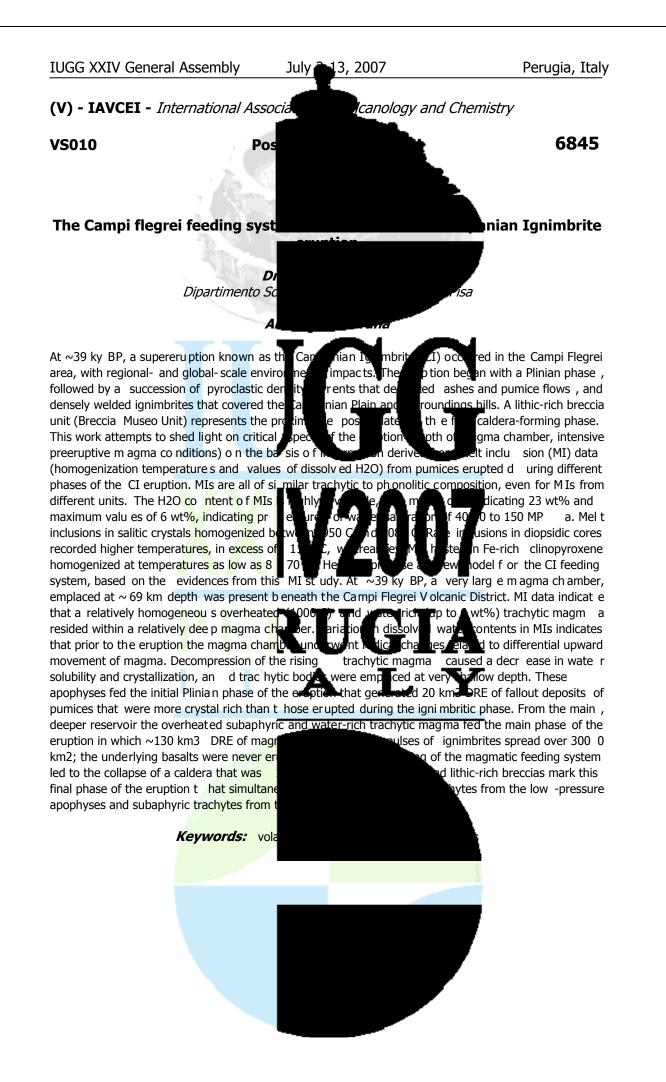


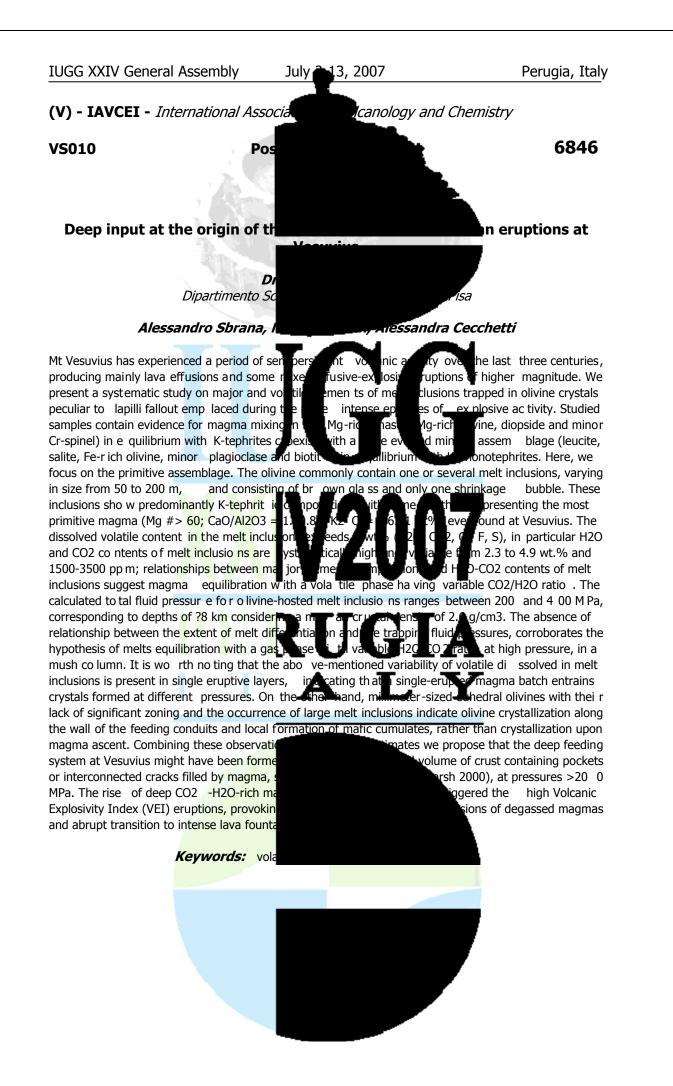


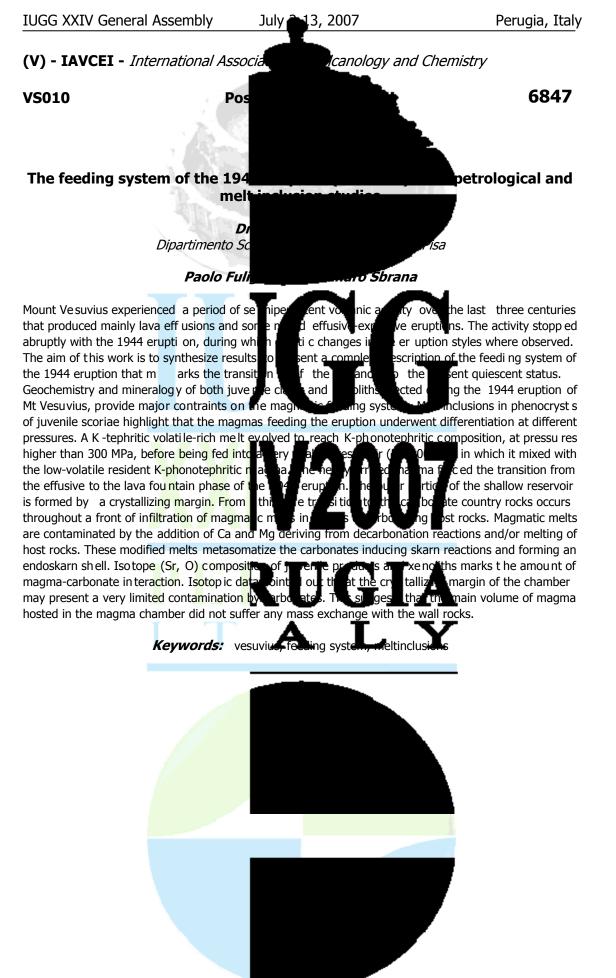




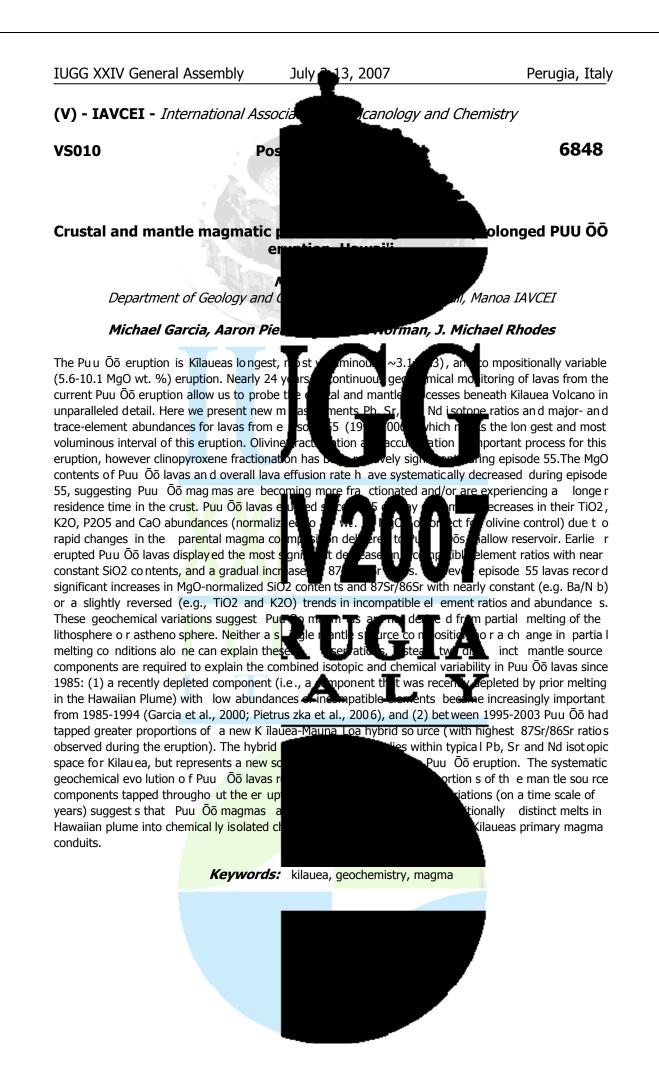


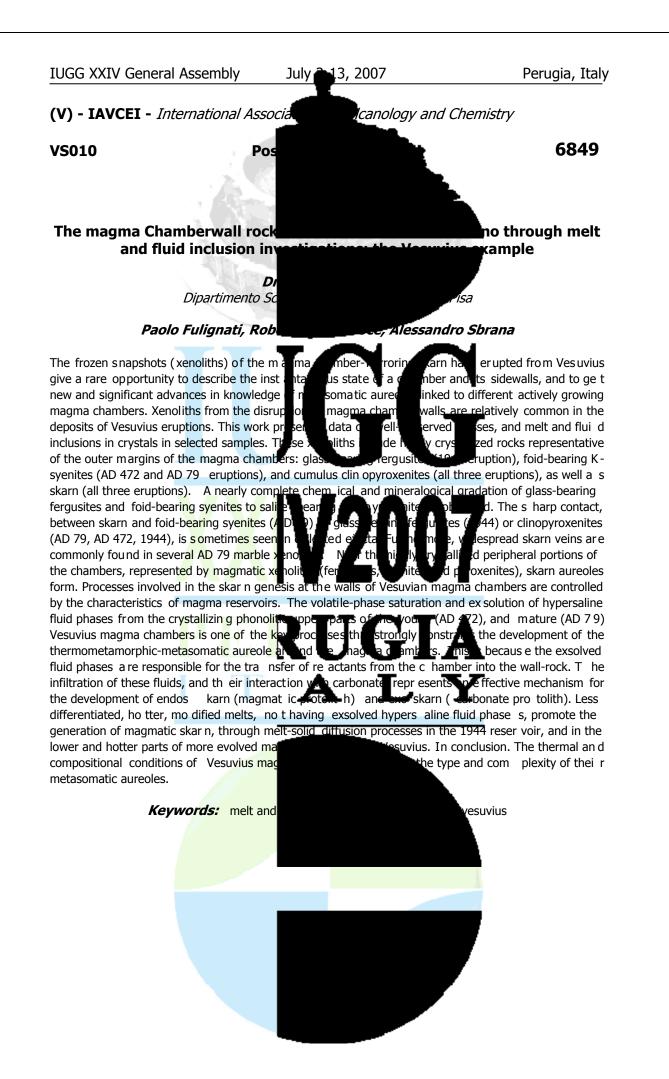


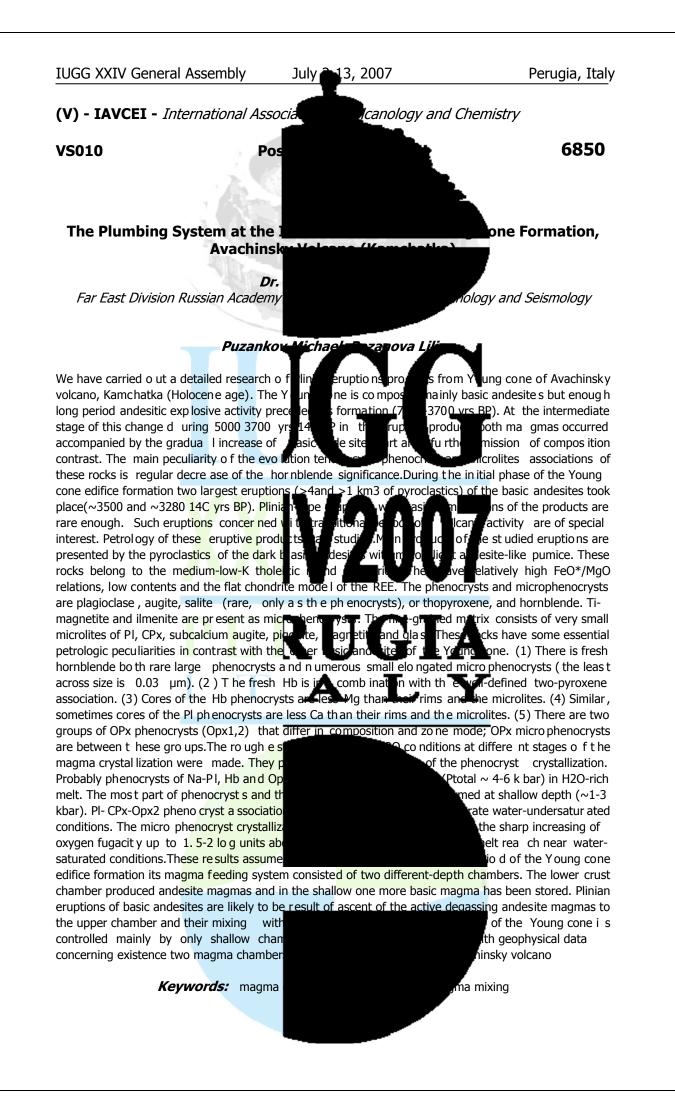


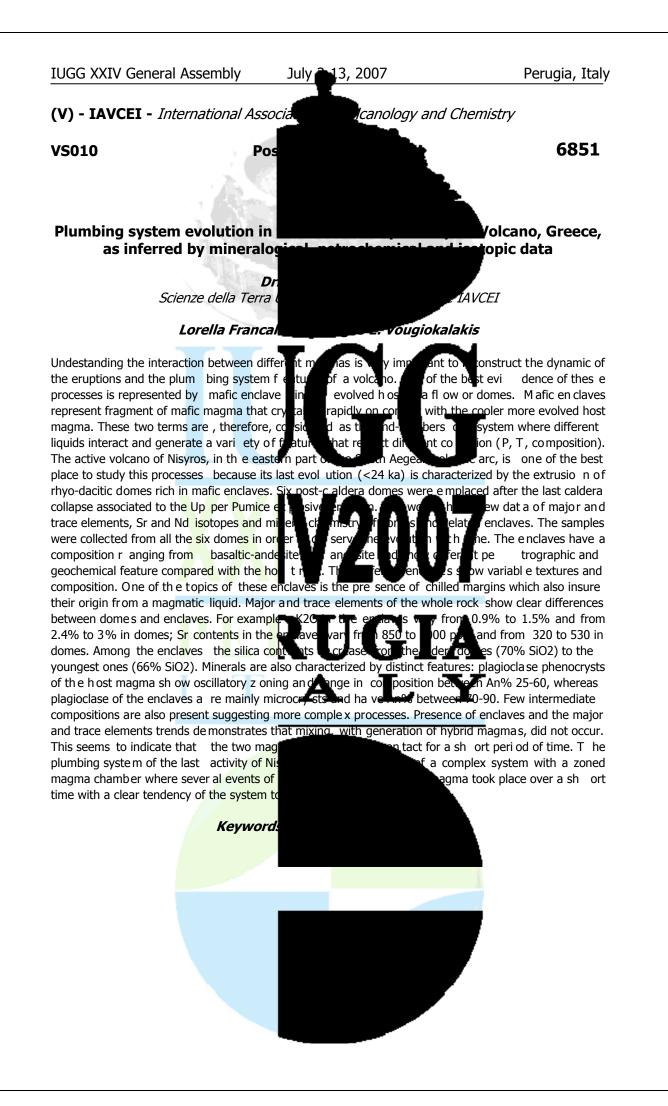


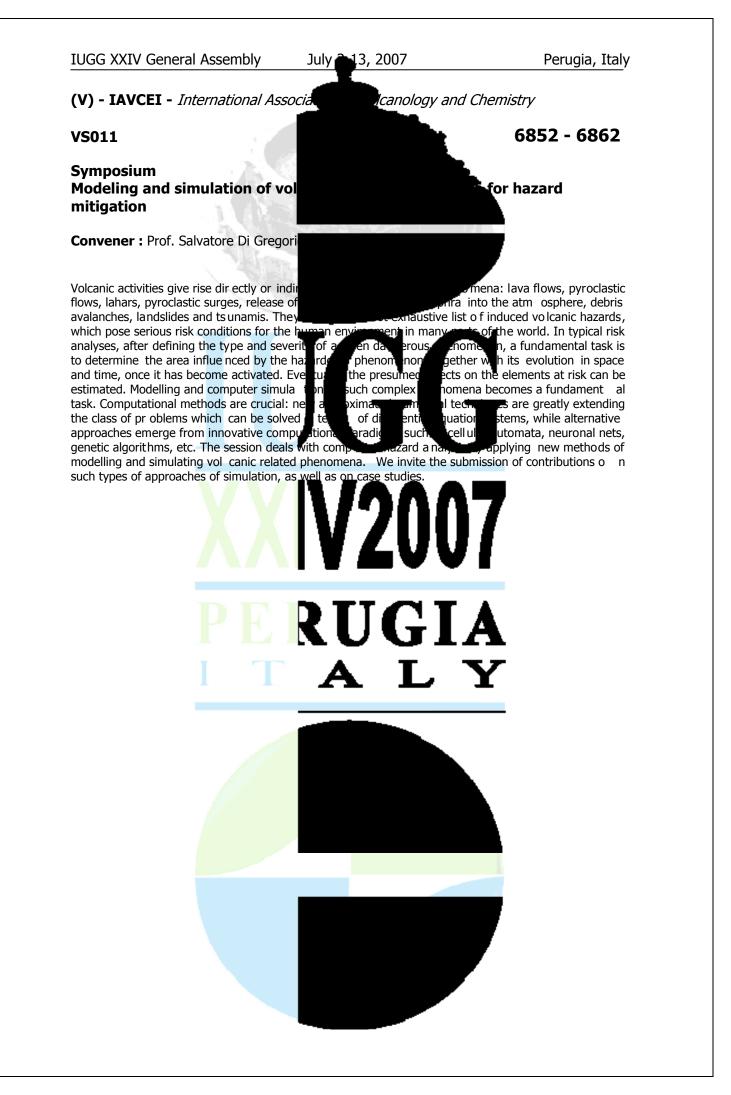
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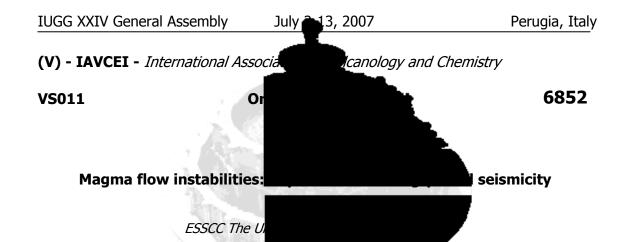












Volcanic eruptions are typically accompani from a small r egion of the upper conduit. These sign they often precede a change in the eruption the shear stresses are highest are thought t

have been unable to develop shear bands to strength mode ls. Presented here is a mod constitutive relationship dependent upon cr magma streng th analogous to planetary lith highly-crystalline regions a macroscopically discont

in deeper crystal-poor regions there will be mechanism. This will result in a depth where the brittle-ductile transition occurs, and here shear bands may develop. We utilize the Finite Element localization and the generation of shear ba Hills Volcano, Montserrat show the generat between the free-surface to a few 100 met where LP seismicity occurs corresponding t explain deep earthquakes in subduction zones as

period (LP) earthquakes which originate have the cap nds fa viao Shear pa ble of bro pths at which hich <u>the m</u>a his re y and for th fore le tvi

tepredict eruptions since the conduit wall where ng alc g the seis mic trigger. But m odels signals occur using simple magma str<u>ength</u> is deter mined from a s in a de pth dependent hosp here, in shallow ation will prevail, whilst

a macr oscopically continuous plast ic deformation

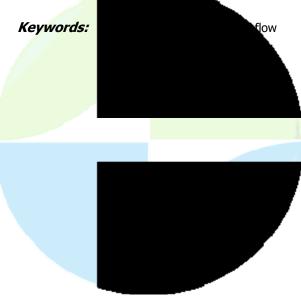
a result of

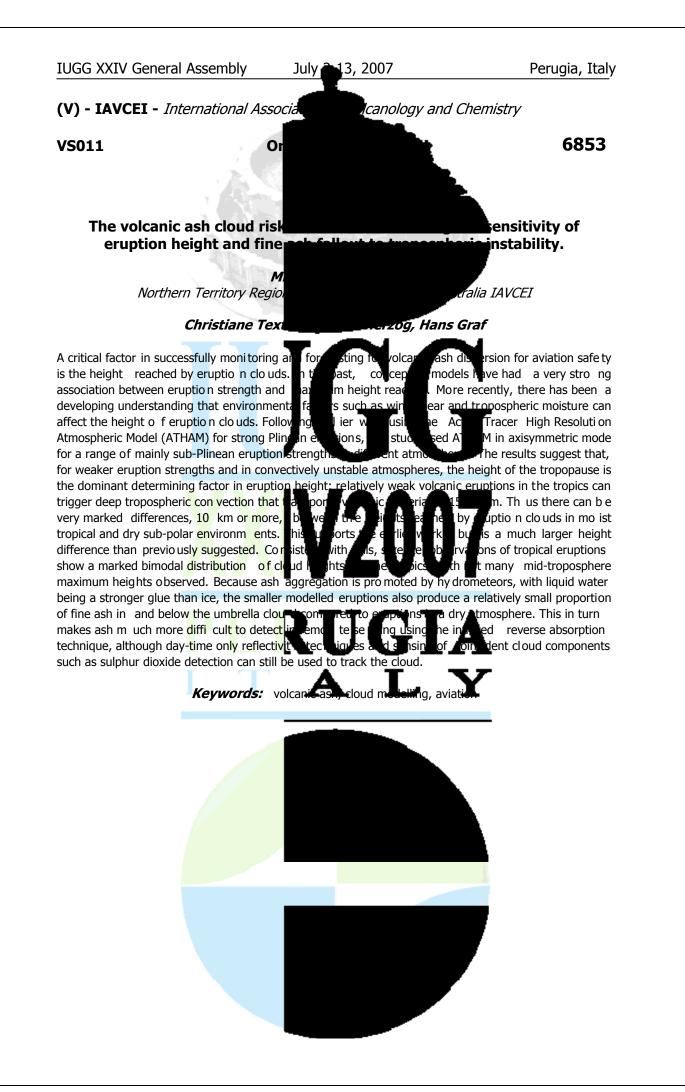
ates to simulate shear trained to the So ufrire er shear bands (forming ds forming at the depths ear has been proposed to instabilities in flow and this may also be the

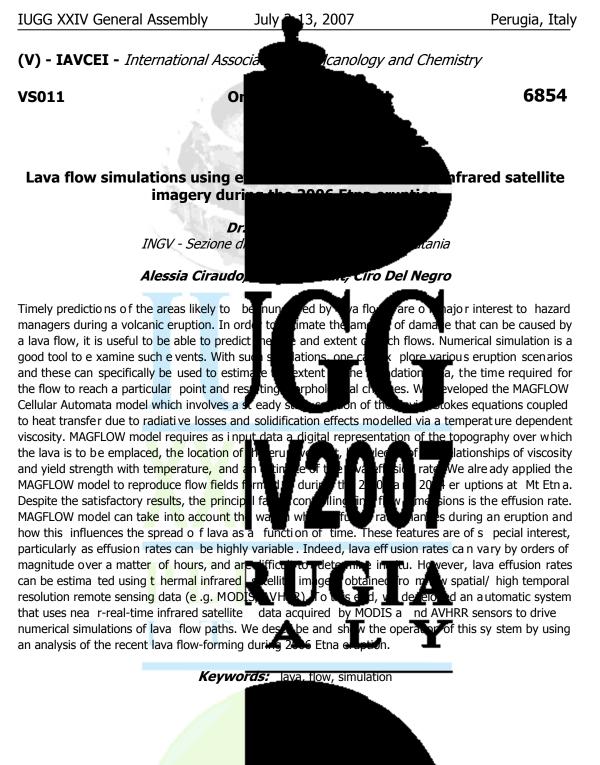
case for LP sei smicity. The viscosity within t hese discrete shear bands suggests a failur e and he aling cycle time that supports the observed LP s ddition, this she ar band model allows LP events to be present during the dif sion<sup>®</sup>reg les, e genous and exogenous, with exogenous growth occurring when upper nd endogenous growth preten con uit s ar occurring when upper conduit shear bands are not present. Shear bands are also found to have a large effect upon the ascending magma extrusion rates e cause she r-induced now alters the over-pressure for p resure increase and decrease predominantly in the upper conduit. This model proides a me in the upper-conduit possibly responsible for flank tilt due to the gene ration of shear bands. Pressure changes as la rge as several MPas develor p at depths of 100s m, values suitable to generate the observed tilt.

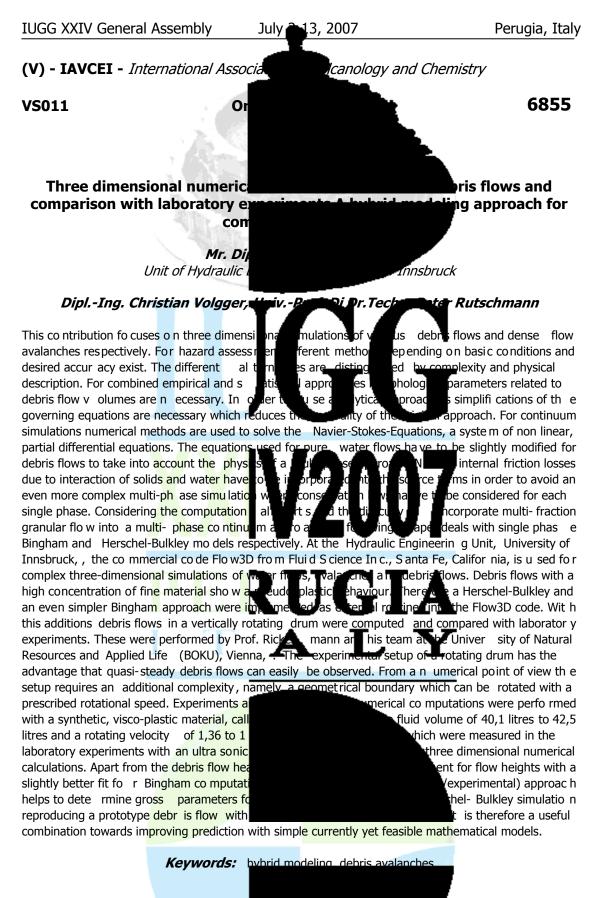
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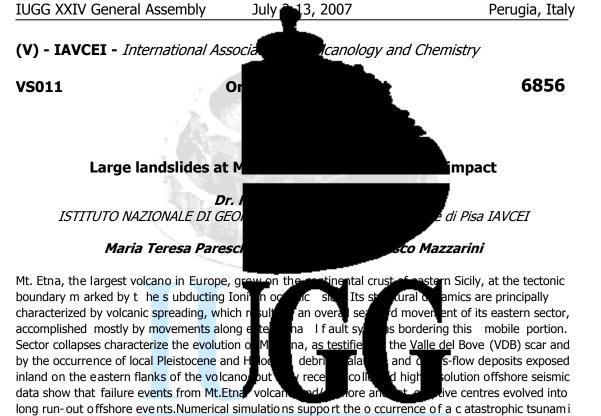
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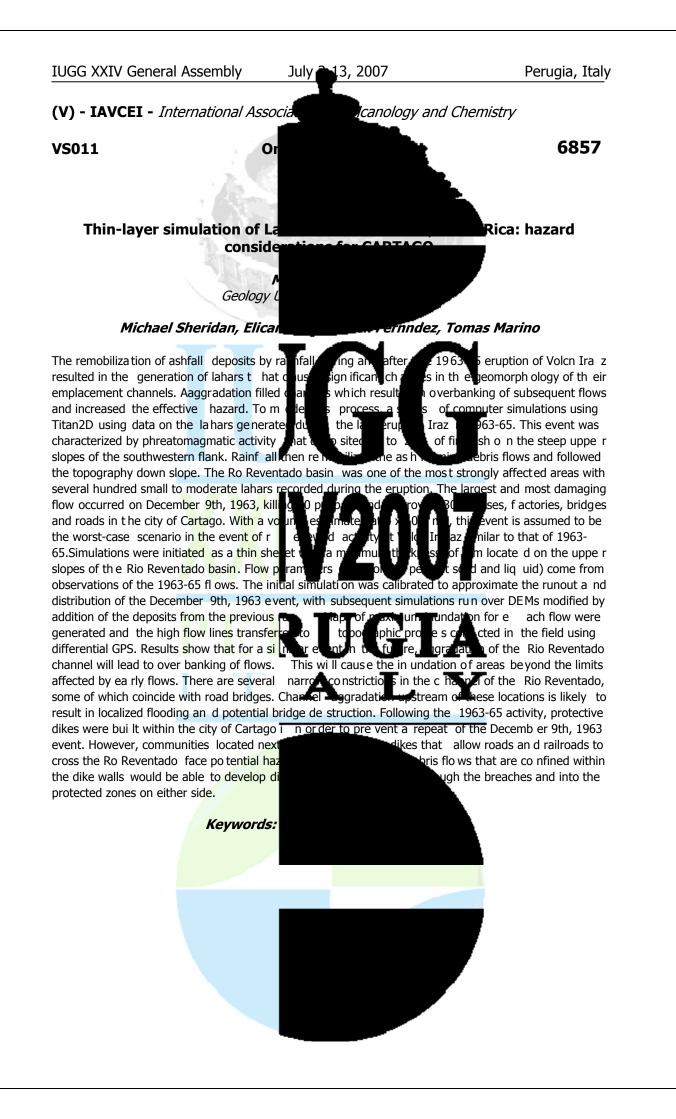
impacting all of the Eastern Mediterranean in early Holocene, related to the most recent Mt. Etna failure event (that related to the Valle del Bove s Mt. Etna (Sicily, Italy) which entered the I resulting tsunami waves were able to desta generated the well-known, sporadically lo widespread megaturbidite deposits of the I

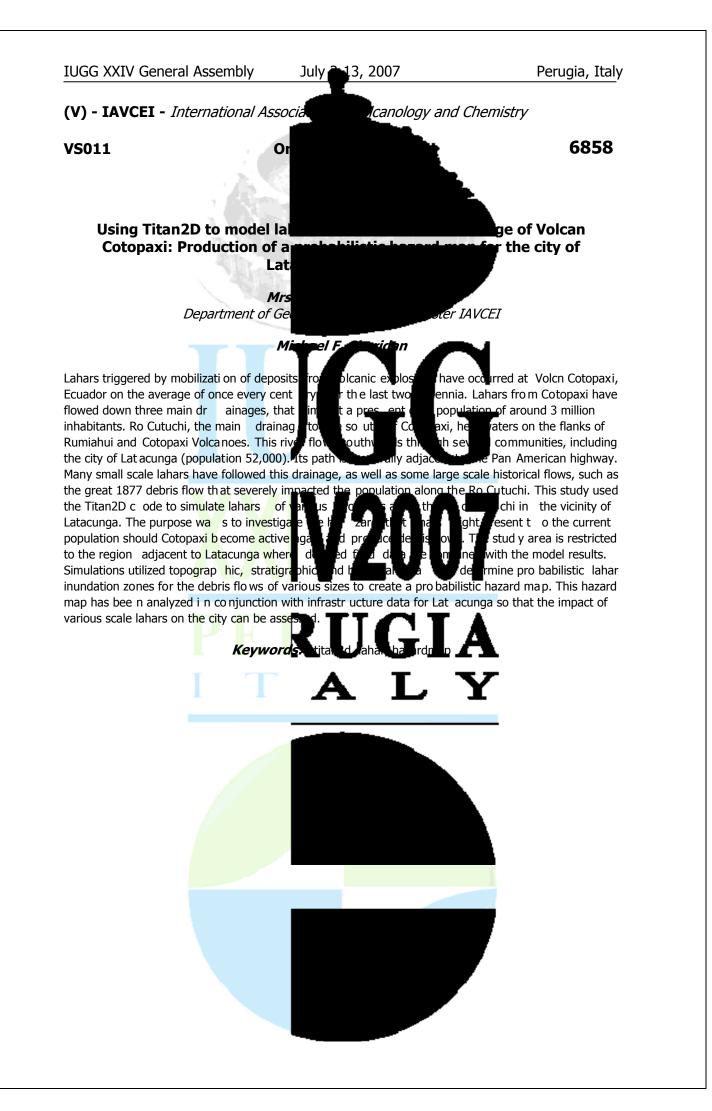
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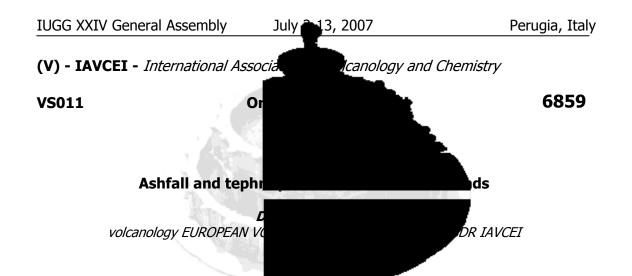
debris avalanche from hulations show that the he Ionian sea floor. This e I onian Sea, and the

Keywords: etna, landslide, tsunami









It is kn own that volcan ic eruptions can produce a destroy property. In histor ical times, thick and w eruptions, for example, thos e occurring on vo Krakatau (Indonesia) in 1883 and Hekla (Ice small volcanic islands is greater. Ashfalls are widely in their effects, depending mainly on or intensity of the eruption. Any one of th, after a volcanic eruption, but when eruptive activity

Therefore, it is necessary to develop prevention and public awareness on volcanic islands with new technological tools. In the past years, GIS and numerical modeling of pyroclastic dispersal have become useful tools for risk assessment in volcan offers the opp ortunity for many interdiscip hazards on islands with active volcanoes. T potential impact from general standard ged land cover and prevailing winds) and help will help to display possible transport and deposition or tephra using the transport equation (advection-

dispersion eq uation) to m odel the w ind profile , atmospheric tur bulent diffusion and particles sedimentation. The ap plication of mathem several fields together with G IS mapping population living on volcanic islands. With

general approach is a key point for risk mit warning, emergency management and land

ashfa bread ic islan de sug in 1947-48 a st co<u>mmon o</u> ume o om th te cts ca ties a a sma

arjety of ha

hich can kill people and associa ted with many fad D Ta mbara (Indonesia) in 1815, 970. The proportional impact on eruptive phenomena. Ashfalls vary uption and the duration hconvenience during or sland problems increase.

> tem (GIS) techno logy h GIS-type mapping of d mapping and display ogical data (topography, The numerical modeling

nin

available scien tific data from e ert ag e to civil authorities and g g fall cov erage, we can 6de a determine which areas will be affected by future volcanic events. By superi mposing volcanic hazar d areas on volcanic island ashf alls dispersion, and atural and structure elements, in a ne r critical beling calprovide an incomparable spatially correlated project, GIS in association with athematic dataset to ma ke decisions f or the develop ment, preparation, and emergency plannin g necessary for safe and sage cohabitation for active volcanoes. <u>An integrated programme that combines mathematical</u> modeling, GIS and scientific data will pro expert mode to civil authorities. This civil authorities to evaluate early

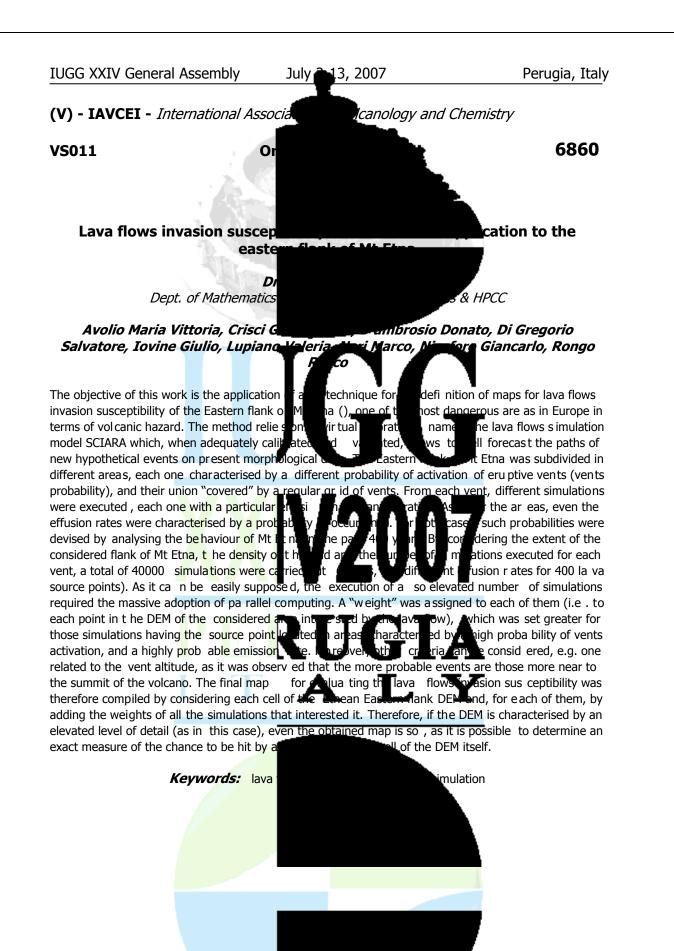
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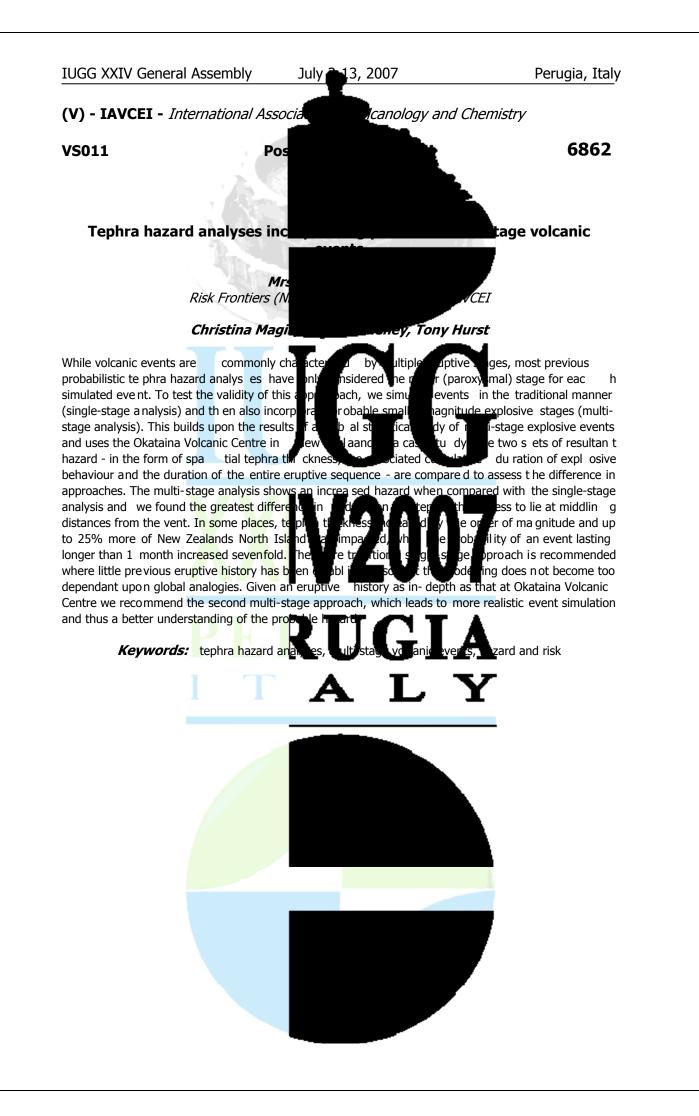
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#### **IUGG XXIV General Assembly** July

Perugia, Italy

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## **VS012**

# Symposium **Cities on Volcanoes: looking at** communities issues around vol

Convener : Prof. Giovanni Orsi, Dr. Ja Co-Convener : Dr. Claire Horwell, Dr.

Urban vulnerability to natural hazards is o By 2050, the world population is expected take place in developing countries, and par

the urban population, large numbers of people will be fragile landscape, with huge impacts on the atur almost 450 c ities world wide with a populati strains of rapid urbanization are nowhere me From today to 2025 about 80 % of the urba population densities place many more peopl The Cities and Volcanoes session will explor th managers and city officials. It will focus on collaboration between physical and so cial scientists

resilience to volcanic hazards.

canology and Chemistry

6863 - 6884

ology and

ssues in urb an development. eople. Almost all this growth w ill r cities and towns. By more than doubling

centrated in resot s sur ore than bne, h ( arent than in ation will be nd ir to a ks bet ר ז of mu disci he i ig volca

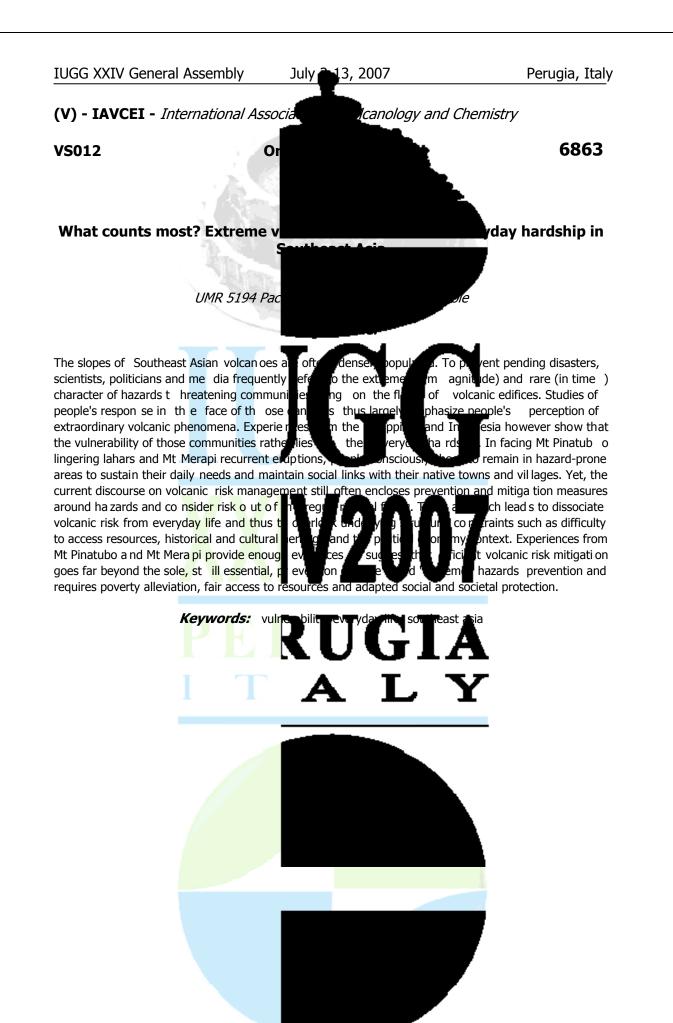
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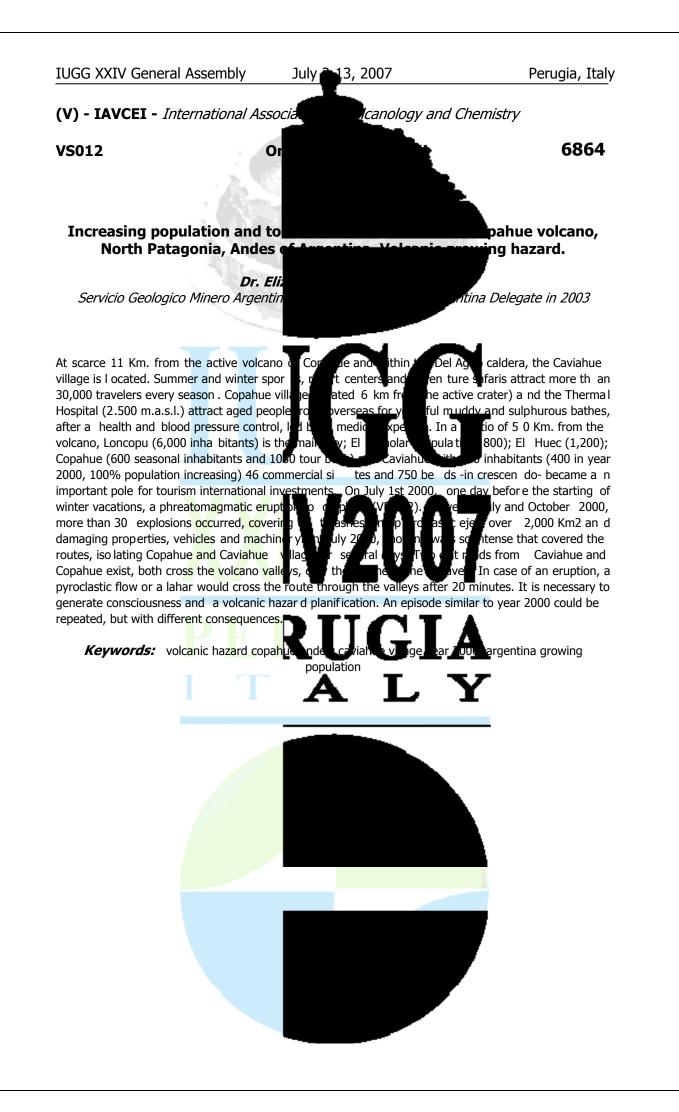
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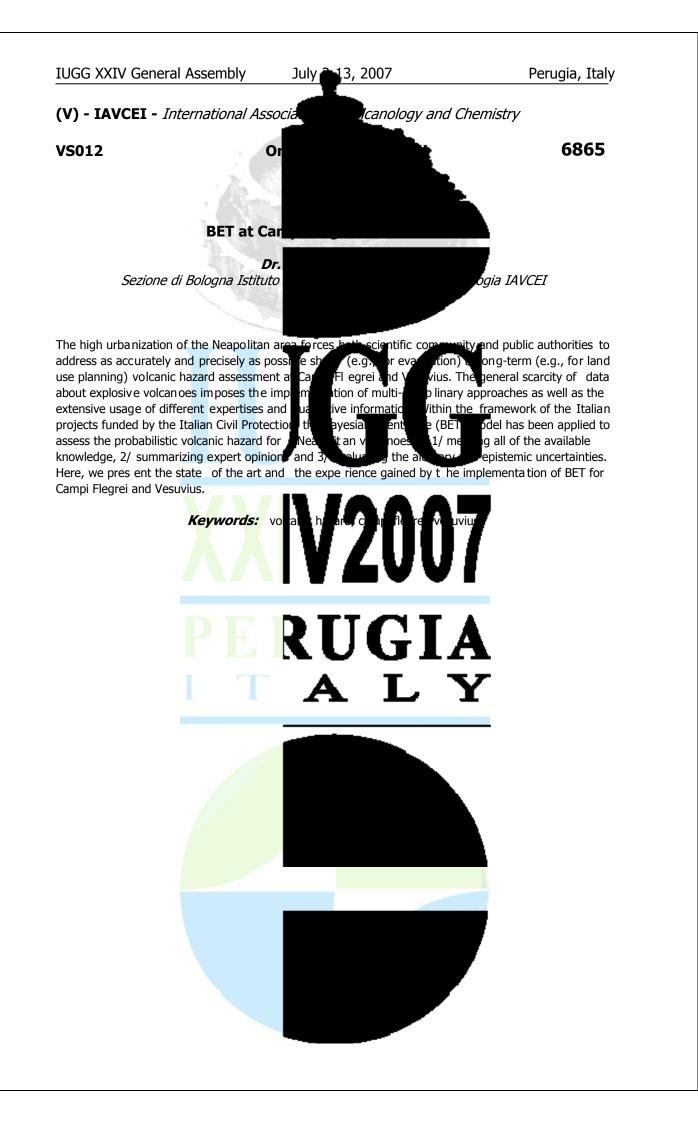
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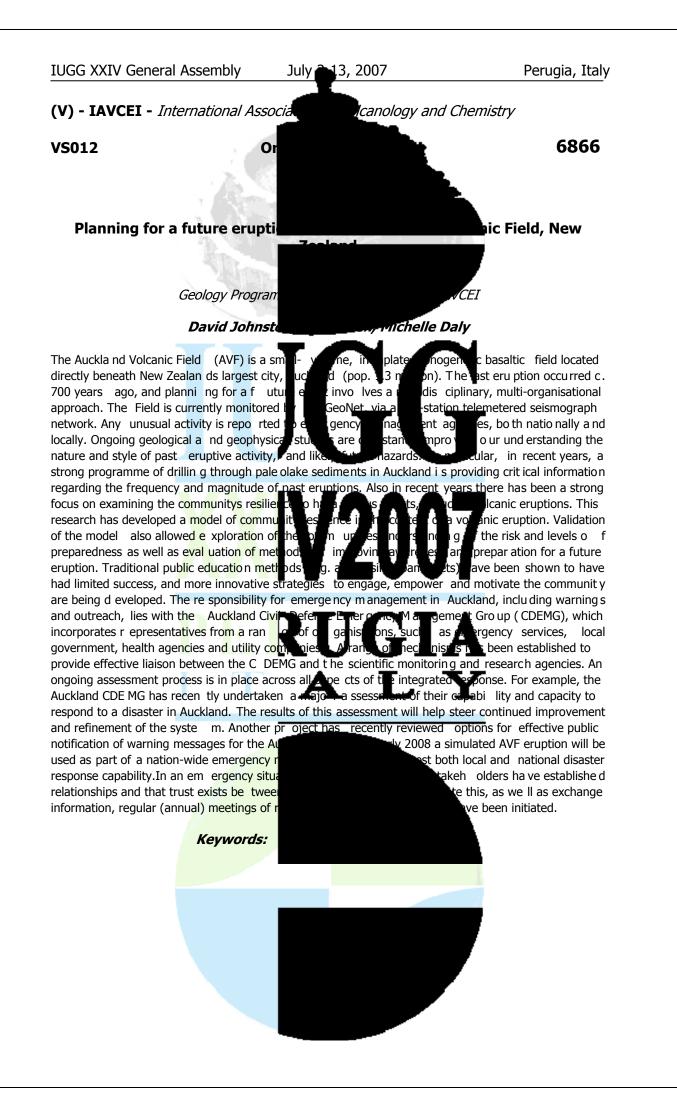
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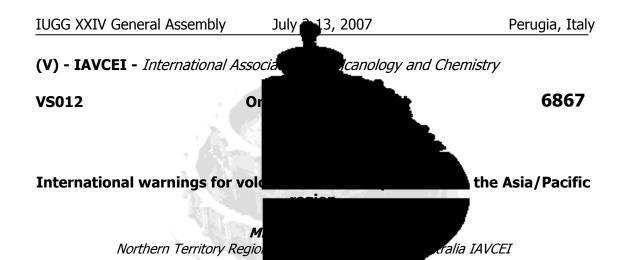
cities and their increasingly dina m. There are currently on inhal tants. The stresses and pcal areas in developing countries. e de veloping countries. Resulting Rularly volcanic hazards. community, emergency olcano d research and need for ry ap nd improving community











The risk posed to aviation by drifting volcaning ouds major city, because of the convergence of a craft the aircraft, and the chance that the volca warning system for volcanic ash clouds, the for aircraft at cruising levels, but sho uld als found many c hallenges in wo rking to ward including great resource disparities betwee issues, language issues, and even different safety

International Airways Vo Icano Watch, arrangements with local volcanological obs Centres more closely with the local meteo consistent suite of warnings to cover the There are also many scientific issues to cou clouds to aircr aft, enhancing our m odellin

atest where ards t citv may be a fect tional Airway on o<u>n the lo</u>o funct ŋq s, wid na ea

ce volca no is close to a port, nation ip m

e lower flying altitude of the airport i tself. The worldwide Icano Watch, is de signed primarily ale. In the Asia/Pacific, we have Airways Volcano Watch, ailure, communications sed on our experiences,

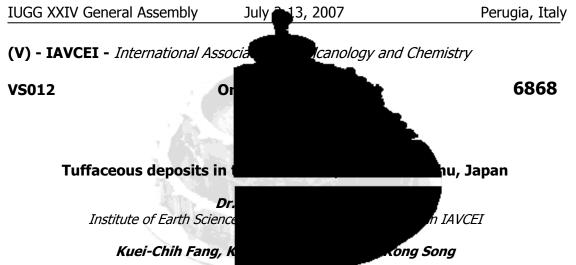
aircraft are no t yet safe from damage or worse from volcanic cloud encounters. To improve the

we can further develo\_

p aviatio <u>n-fund</u>ed enhanced servi ce olcanic Ash Advisories ensure that there is a bersion of ash clouds. sk of old versus new ash inificantly improving our

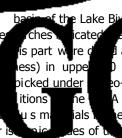


kid



The 1422 m I ength core was drilled in south central Honshu, Japan. The results of previous re core consisted of the lacustrin e clay and ag 1993). The m ajor 18 tephra I ayers (> 1 cm study. Pure and cleaned glass shards and m were analyzed in chemical and isotopic c respectively.Major dispersive mechanism of eastern ward wind. The K2O vs. SiO2 plot and Sr

the sources of tephra layers. There can be identi caldera of so uthwestern Honshu and thirte en layers from central and southern Kyushu. The 9.3 ka Ulreungdo eruptio n (U-Oki) in J apan Sea sources were identified from the Norikura wind direction.Based on the tephrostratigra 250 m of core were estimated from 0.16 m revealed a four times lower s edimentation should reflect the paleo-climate change.



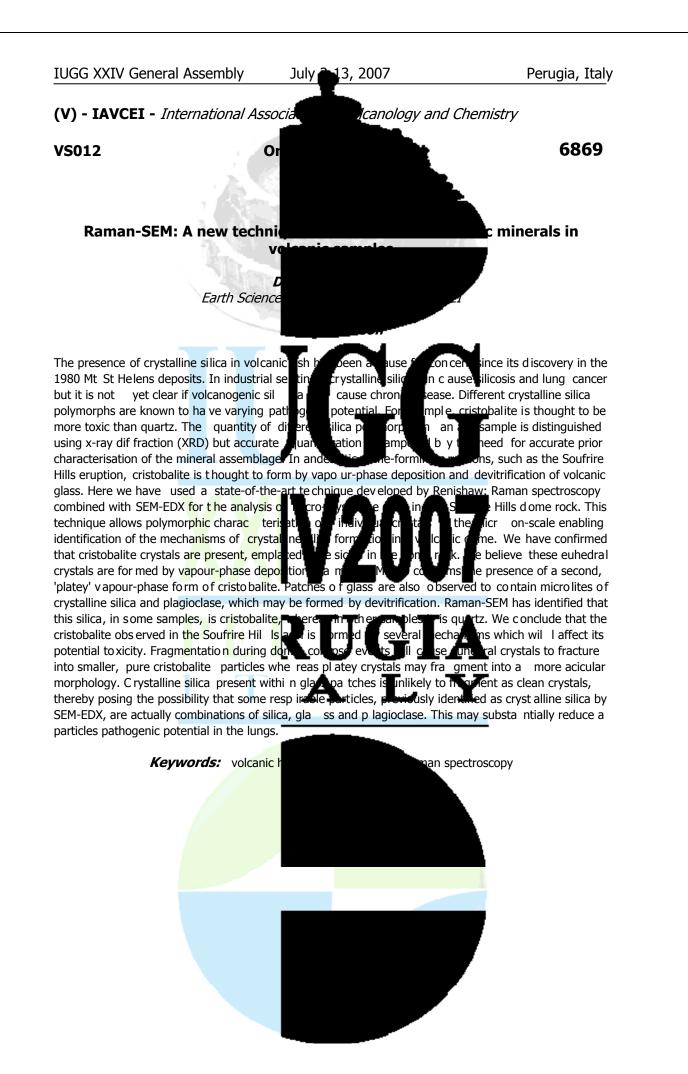
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(3512'06"N, 13600'49"E) in bst upper 250 m part of at the as with 430 ka (Meyers et al., m of core were chosen for this o-microscope in each tephra layer he TI MS in the IE S, ai e Lak wa was caused by the ards were used to verify

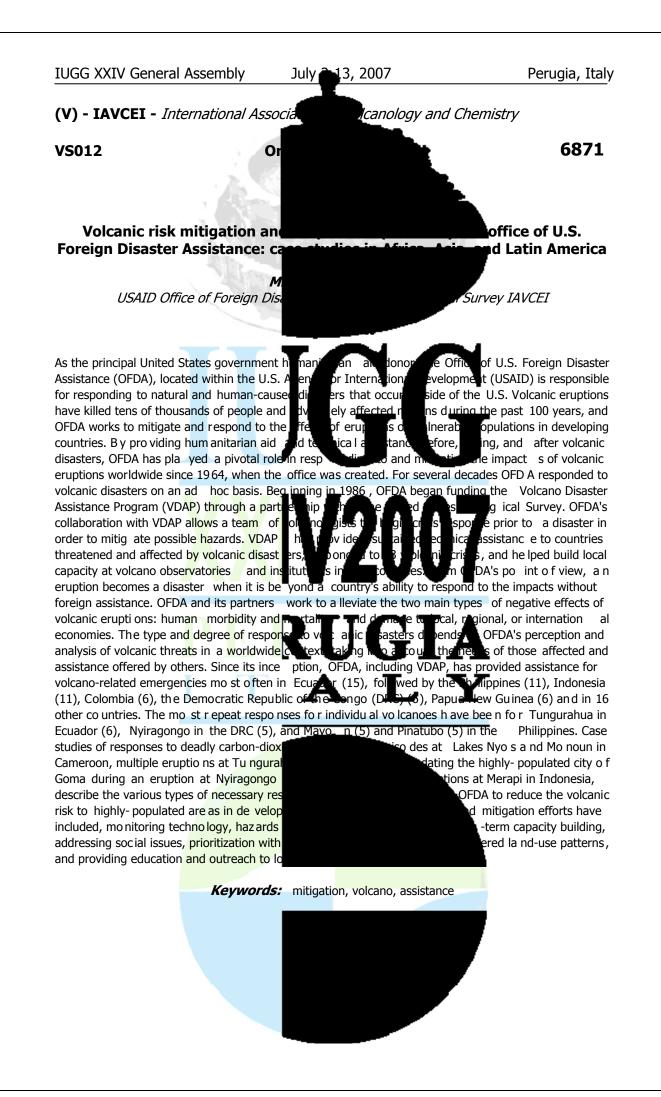
fied that three layers from the Daisen and Sanb e

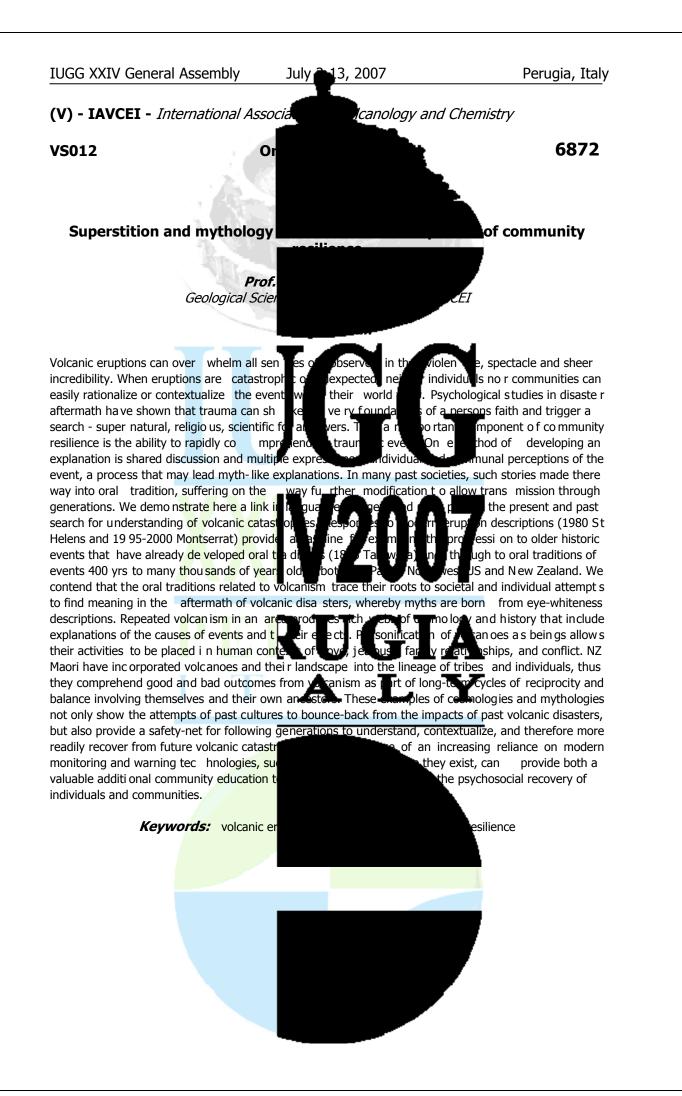
trast, three tephr а Honshu) in the upper ation rates of the upper e that there obviously 5 ka and 105 ka, which

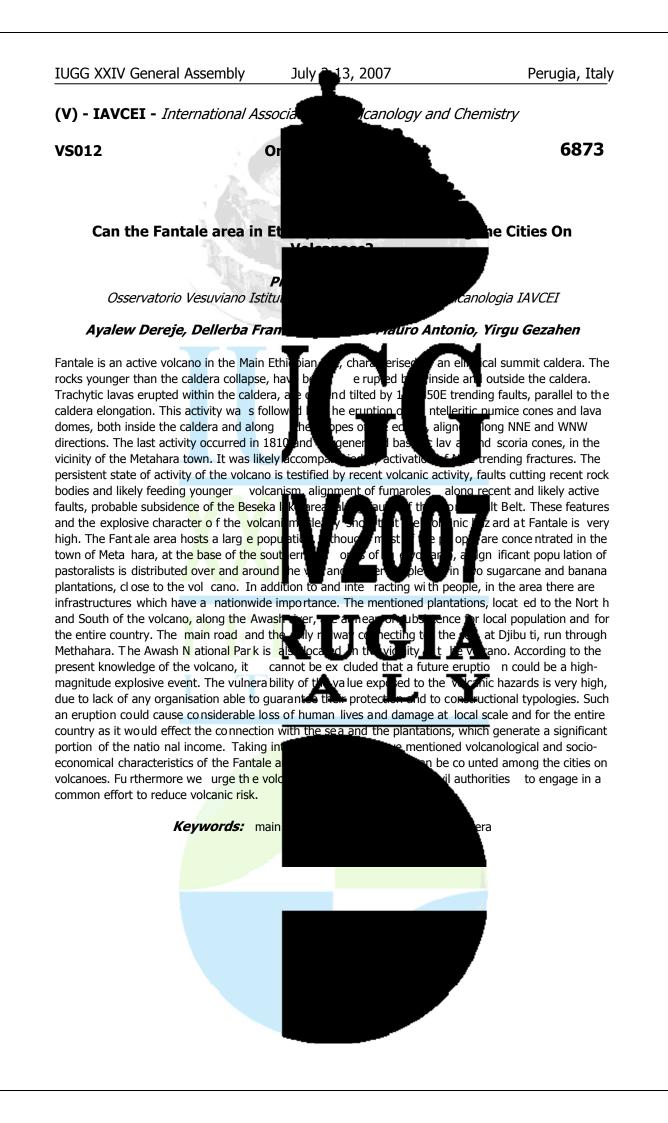
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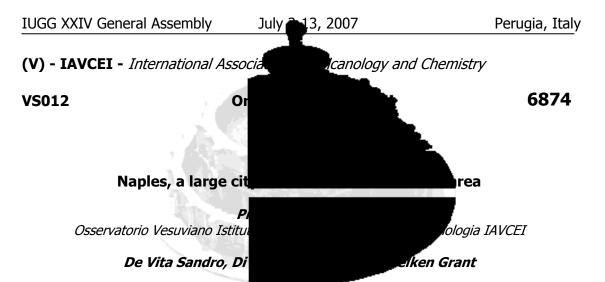












The metropolitan Neapolitan area forms an urbanised the city of Naples, founded by Greek settler in the with the surrounding towns within the Phle Vesuvius strato-cone, to form the Parthen be spread across two active volcanoes, each hid the best known examples of a high risk vold in 1538 and is the site of intense hydrother to o ngoing resurgence. The Somma-Vesuvius

interaction between active volcanoes and humanity for thousands of years. It has been growing through

time, because of soil fertility, a tempera Mediterranean basin, in spite of the volcar source of hazards. Volcanic a ctivity, accor emission, and their constructive (new rock actions have generated the conditions for

present geomorphic setting of the area, d

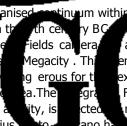
landslides. The use of the territory and volca

the conditions for further hazards. Fulfilmen

community, have realised t hat volcanic h

planned. Urban planning for the greater Na

landslides, and flooding. Historically, urban planing



active volcanic area. In fact, Parthenope, has merged d calle al ong the slopes of the Sommaens ely-populated urban area has explosive ch aracter, and is one of Fields dera, which last erupted recur unrest episodes related ntinuing seismic and

hydrothermal activity since its last eruption in 1944. The Parthenopean Megacity is a good example of



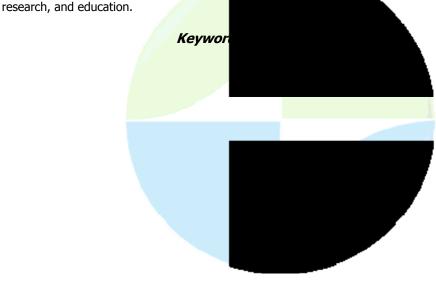
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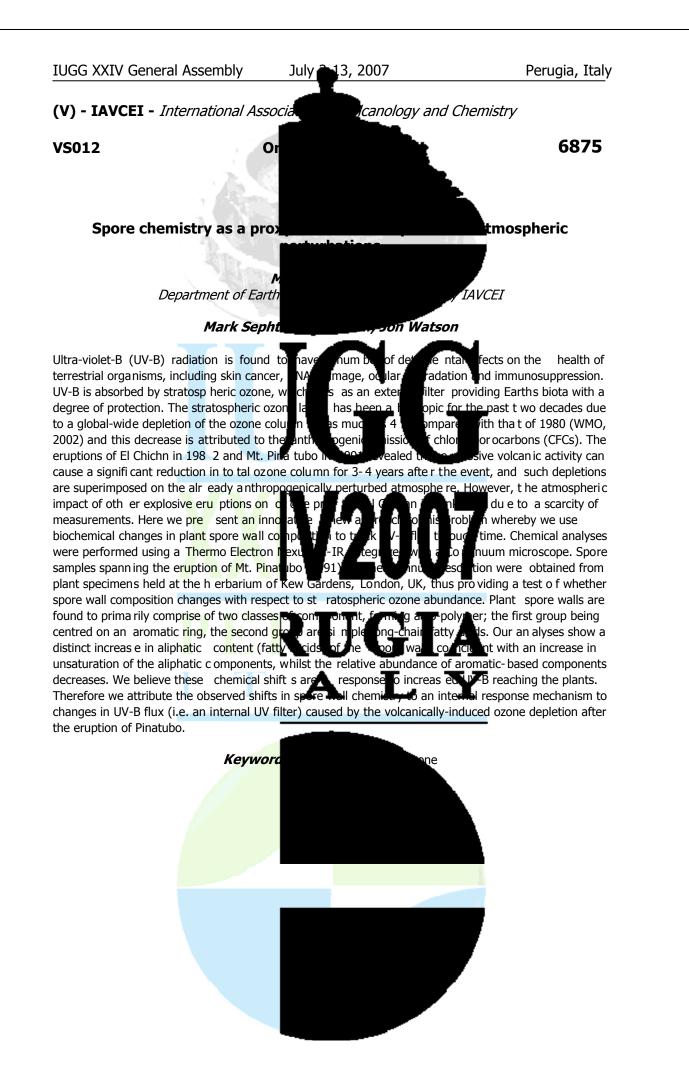
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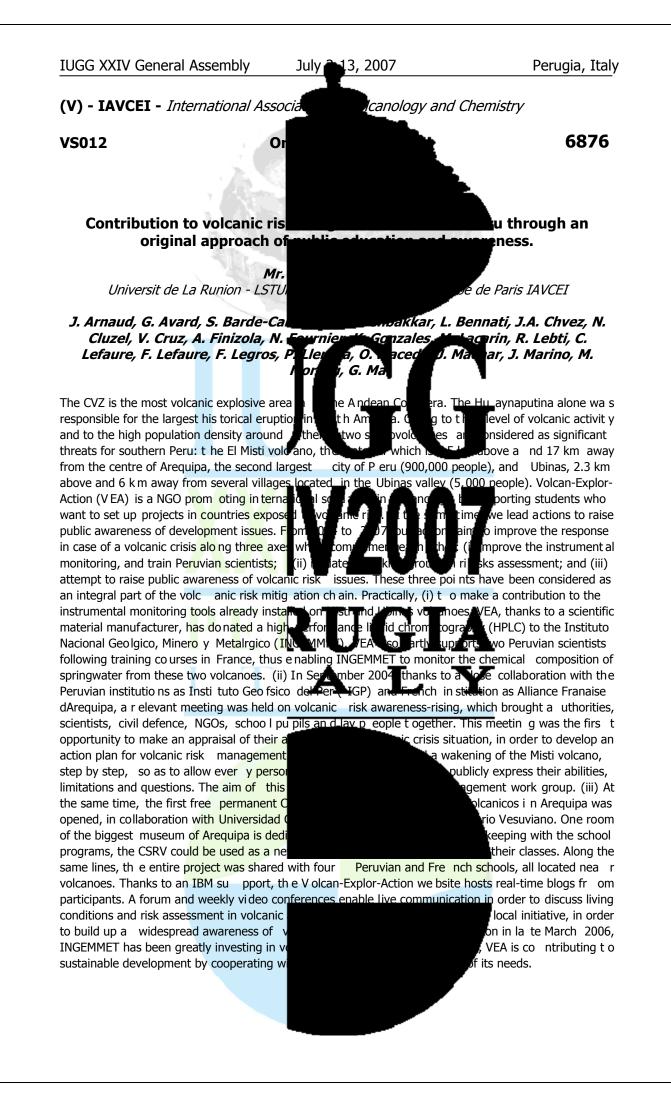
c positi on within the olcanoes is a continual micity an d diffuse gas o tect onic deformation) opulation and growing

economic activities have increased the risk. The Parthenopean Megacity is exposed to natural hazards in addition to vol canic activity, related to both ge ological setting and mill enary human habitation. The ctivity generates erosion an d nans for housa of years has generated las. usi ccess to water, building <u>'9, </u> and maintaining fortifications, required excavation of quar ries, wells, and cavities in lavas, tuffs and loose pyroclast ic rocks. The results of this activity are s out es of hydrogecogical hazards s uch as well as all over the world , litun area a the Nea

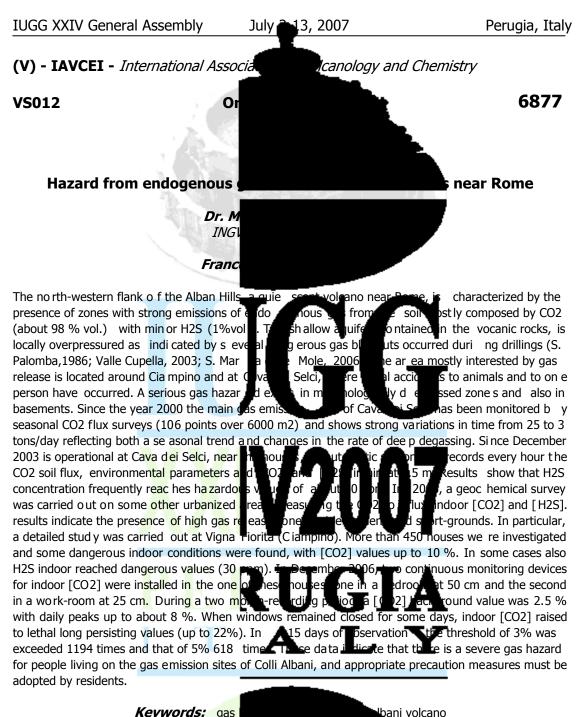
has not considered the catastrophic effects of a volcano, which has a longer recurrence time than that of a hum an l ifetime. During the last de cade, <u>Naples</u> civic authorities, solic ited by the scientific essed a nd risk mitigation actions b e for continuing volcano monitoring,

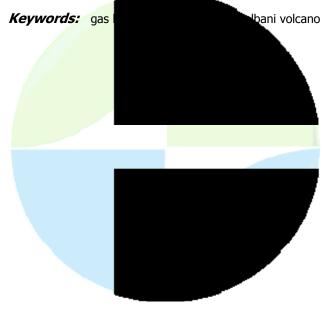


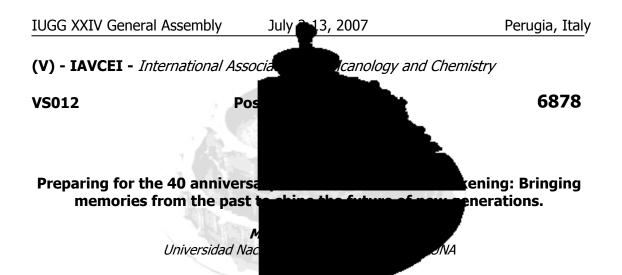












On july 29, 1968 Arenal volcano woke up taken along with important patches of gras swepted in a matter of few minutes. 40 testimonies from witnesses remain fresh in t vividly by those who suffered the loss of bel A national and international effort can brin entrepreneurs and friends of Arenal volcano that no one in the area forgets Arenal's funy and en

interested ones. Tourists and bystanders will be asked activities to take place in the most devel campaign should include am ong others; s meetings from different audi ences, etc. O businesses and locals will receive suggestid how to better recall a gray passage in the for the new generations.

ar 100 years its 🛽 icultu and land. em too fa r in emories. Day es. Su rvivors rces f we s will \ct e to ost 40 y

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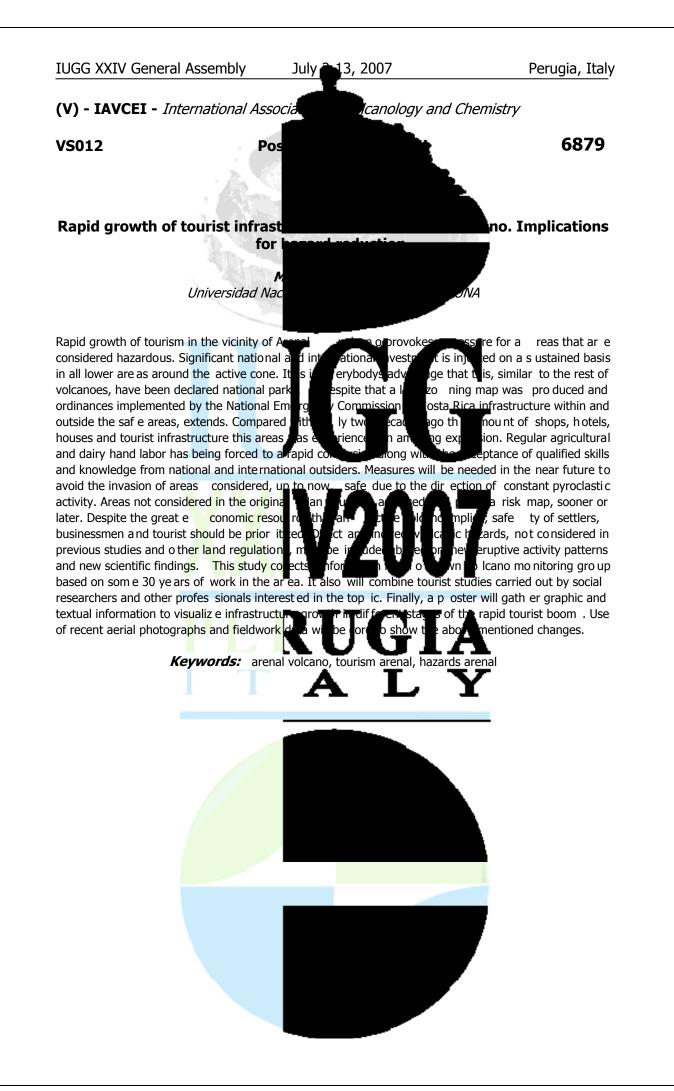
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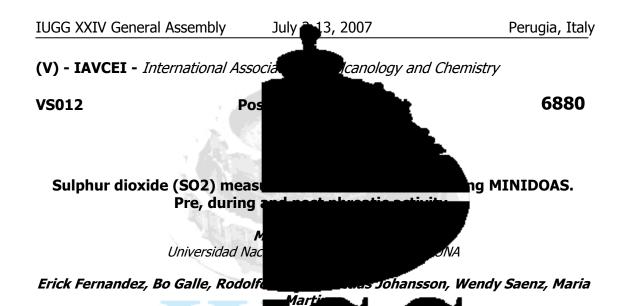
he 90 human lives were houses and roads were esto d past to remember. Nonetheless d nights of horror will be brought have a chance to tell their stories. embrance to visitors, served plann ell in advance to assure

n stop activity has given researcher a p ile of findings, do cuments and opportunities to share with the community and other to participate and become involved in all sort of una. A co mprehensive cts, scie ntific exhibits, ficia

ntal institutions, private les on volcanoes 2007 of orm it in a fruitful lessons







Pos volcano (10 12 00 N, 84 13 58W, eleval km N of San Jos, Capital city of Costa Rica. basin with a pyroclastic con e. There are magmatic period took place in the early 1 kept a acidification alley towards the SW of series of phreatic eruptions occurred prod Licing sig

expelling mate rial ranging from fine sedi compared in different eruptive conditions, hence purpose of data collection was twofold; to gather data that feeds a database and t o compare th conditions. A sulphur dioxide baseline is b control with other different methods routin place under what can be des cribed norma 61 ton/day. In April 2006 a second campa gn mentioned above took place. For similar traverses from 69 to 14 5 ton/day. The volume of SO2 estima tes will be later used to compare with impact

8m) is a' asa crater hosts al acc<u>ounts s</u>i s inter inc km lo de . 1 k omor

Indesitic trato volcano located 45 ot, hyperacidic lake that shares its 82<u>8 although</u> the last phre atoasification process ha s tain ed e end of march 2006 a anges in the crater and

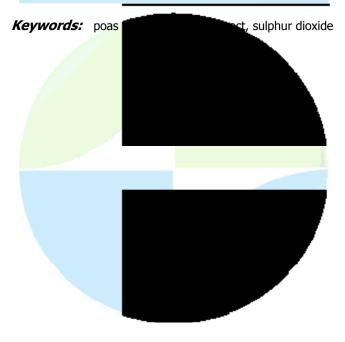
ments to met ric blo cks. Three Mini-doas campaigns are

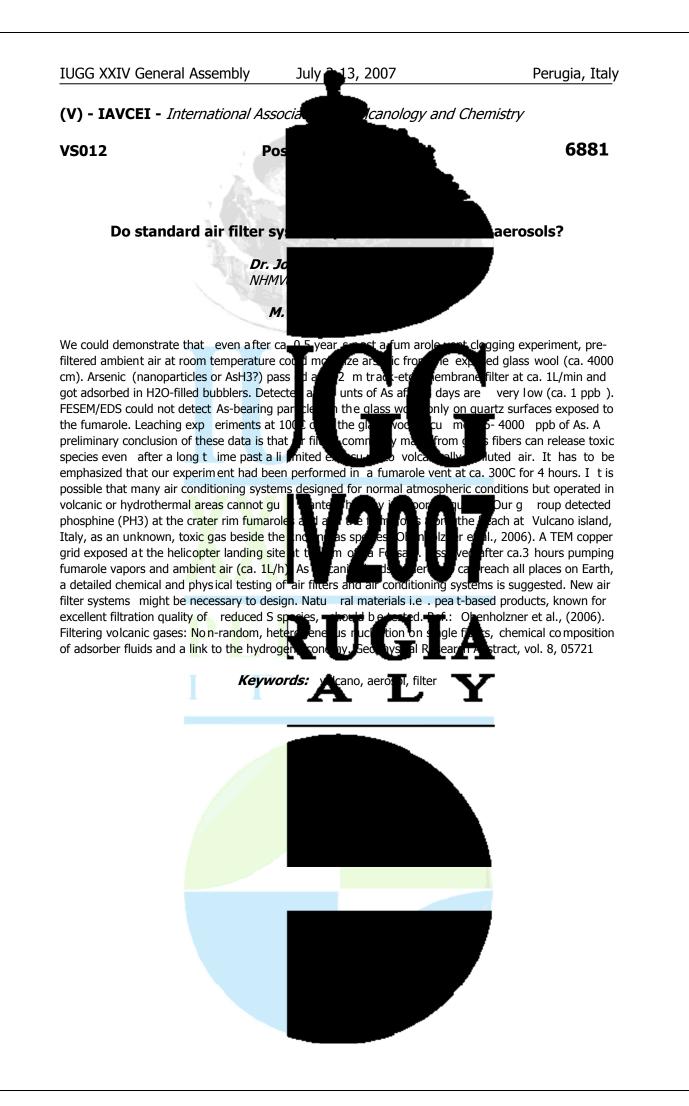


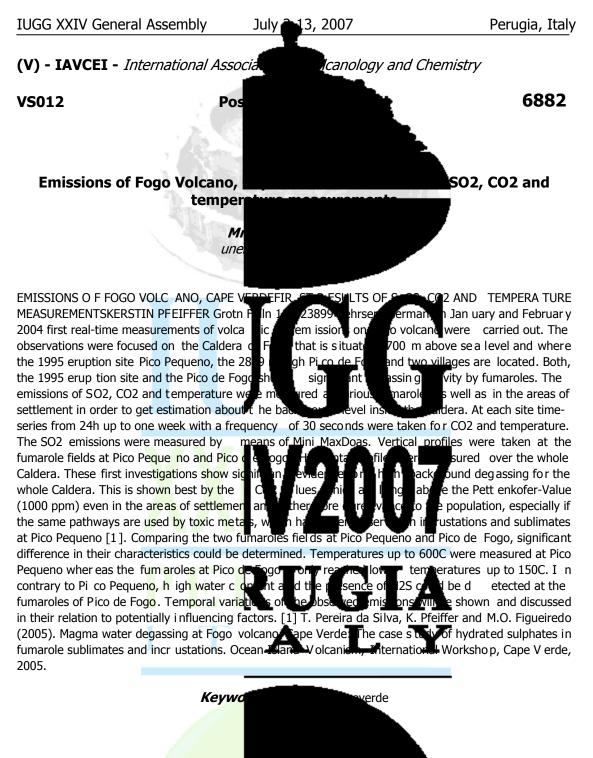
different degassing bring routines for cross the first campaign took

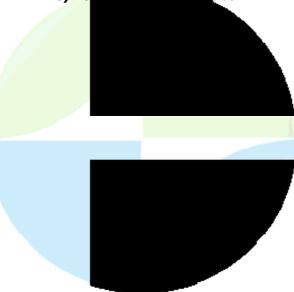
this cam paign is set to ter the phreatic episode than those from 2002 t he estimate av erage ranges

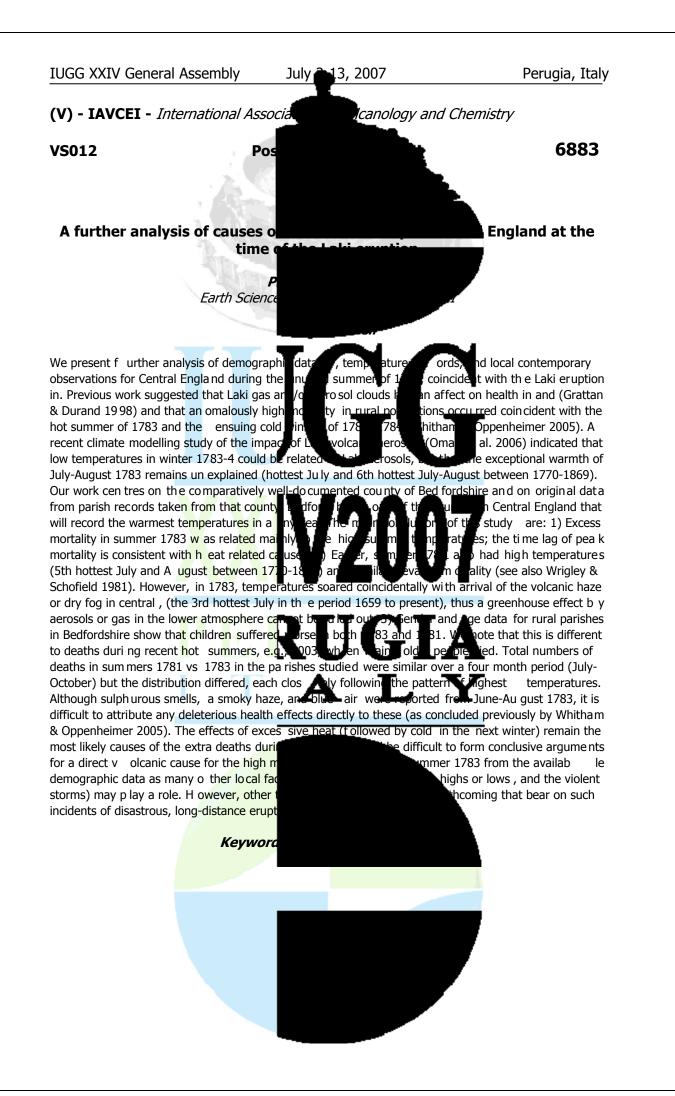
assessed in the vicinity of the emission sour volcano monitoring group have informed in the past about i mpact on veg hum life. The first and thir d campaigns were carried out on foot while the veh de baking traverses beyond sec nb ed he ' the summit area. Traverses on foot were made covering about 50% of the circumference of the crater rim, starting in the S side, walking clockwis e. An ster will show text and whic infor mation about these three Mini-doas campaigns. It will also show s me of th eacts on in astructure and vegetation documented for periods between visits.













#### IUGG XXIV General Assembly

#### July 13, 2007

Icanology and Chemistry

Perugia, Italy

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the Millennium

po verty and hunger by

ment. As well as being the

cities near volcanoes, especially

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## **VS013**

## Symposium Quantifying and expressing vo

**Convener :** Dr. Peter Baxter Co-Convener : Prof. Augusto Neri

The Millennium Development Goals of Un 2015, and natural disasters are recognise causes of huge economic losses. Volcanic where urbanis ation is unpla nned, and fo

and communications and, in this re gard, right red systems. Recent progress has been mad quantifying the consequences of eruptions this session is methodological advances for o and new approaches for reducing vulnerabil and exposure studies into formalised risk as a risk-informed approach as well as present

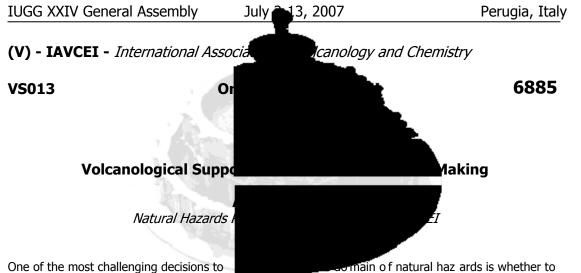
gies a on si loping truc n settlement ing and expre tri bu t, on v of eal c are

aevelo ping states. Eru ptions can ha ve regional and global climatic impacts which affect environmental sustainability, economic activity, travel ortant as early warning d multi-ascip linary methods for d human activities. The theme of y volcanic risk for decision-making, ntegr pf hazard, vulnerability h civil authorities within rfacin

velcome.

icular



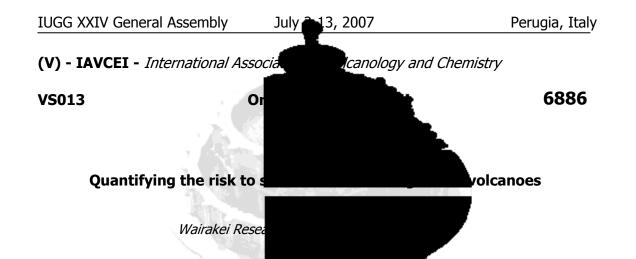


evacuate a densely populated region around a volcane The economic expense of mass evacuation is but could, potentially much greater if an evacuation is but could, communicated their expert s cientific views in vienno problem of decision -making to the em. Wi in the economic quantifying volcano hazard, and stochastic emission civil protection officials further by establishing probability criteria may be quantitatively expressed in terms of evacuation call.

a volcape that appears to is how, yet the cost of a rid, or is called in vision hazard the il it ecent development manion of sieve at procesilistic eria e rms of the procertion of

to be threatening a major eruption. possible human casualties is . Traditio ally, volcanologists have il protection officials, leaving the ent of probabilist ic methods for ation, and can ologists can support evacuum decision-making. Such warenes owing their lives to the





Volcano scientists are often perceived to be taking this has result ed in the tragic deaths of ca. dangerous working environments, but are th and potentially active volcan oes in North Is Science is tasked with monitoring these vol Earthquake Commission of New Zealand. In work in potentially hazardous environments high levels of toxic gases to sudden onset

Department of Conservation worker was sampling a crater lake on Raou I Island in the Kermadecs on advice from GNS Science when a sudden\_small phreatomagmatic eruption occurred. As a result, the worker is missing, presumed dead. As a co the risks to which GNS scien tists are exp tasks. This analysis will inform decisions at during and after a volcanic crisis. The first the return period for different magnitude e eruption, but the consequences for a person working close to the active vents were enormous. Most

studies of frequency-magnitude relationships of active volcanoes have been limited to large (V EI>4 or eruptions since the historical and/or ger of deposits. Most volcanoes have only had observations of small eruptions are few. Also vent area which are readily eroded. In Ne w Zealand, the frequently a ctive volcanoes (White Island, Ruapehu a nd Ngauruho e) have excellent observational reco records for a further 100 ye ars. Although the data average eruption rates for different magnitude eruptions can be calculated which can give likelihoods of fatalities during monitoring operations. Survival analysis of repose periods can also be achieved to give a better idea of return periods of activity the vent area. Results will be pr esented that

volcanoes are within tolerable levels of risk

isks in t n the ienti d sind s high as the New Z ealand larg<u>ely thro</u>u carry ring ad iay b a ran bjed tivitv

893. Volcanoes are m to be There are several active d in the Kermade c Islands. GNS the Geonet project funded by the ies, GNS scientists often f volcanic hazards from ely in March 2006, a

erests of the ir science, and

ded to better quantif y

he volcano monitor ing d be undertaken before, volcano is to investigate uption was a very smal 1

observations and preservation or decades, so detailed av 🗧 mi r depos its close to the

ls fo r ca. 5 Fars and less co mplete tect, particularly for e arlier eruptions, e is not us to a monitoring scientist close to the pring operations on New Zealan d a fatality per year).

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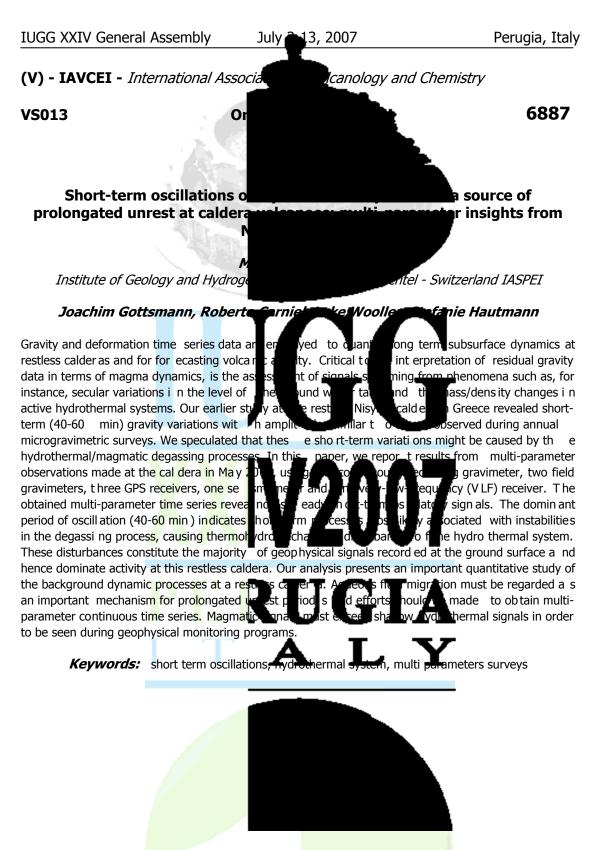
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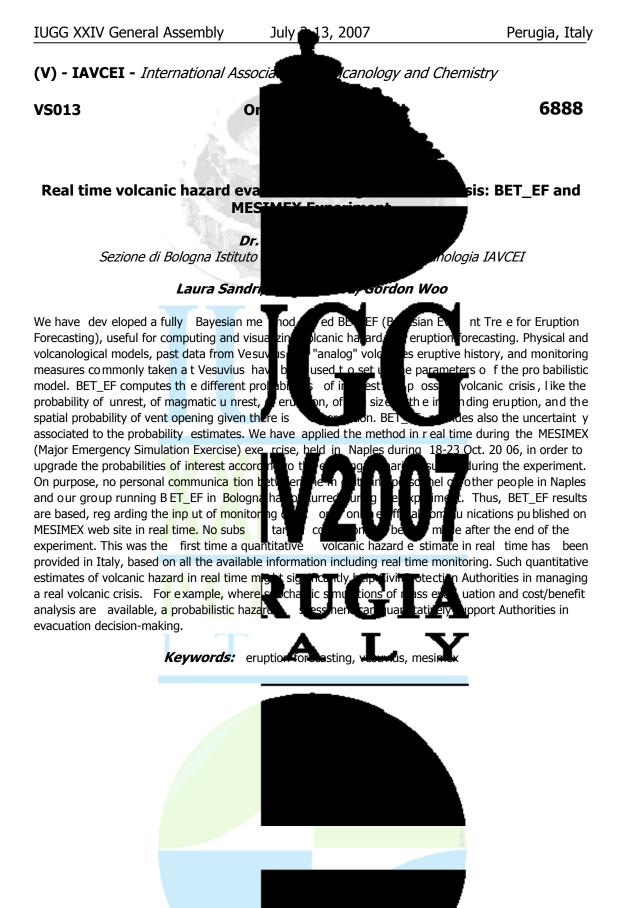
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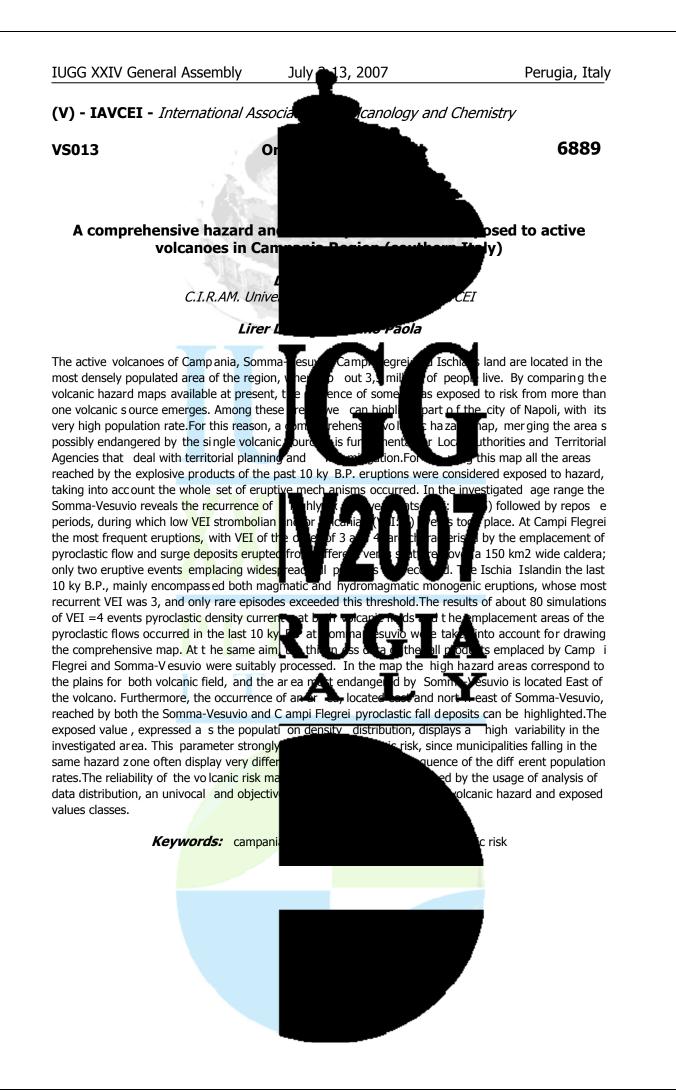
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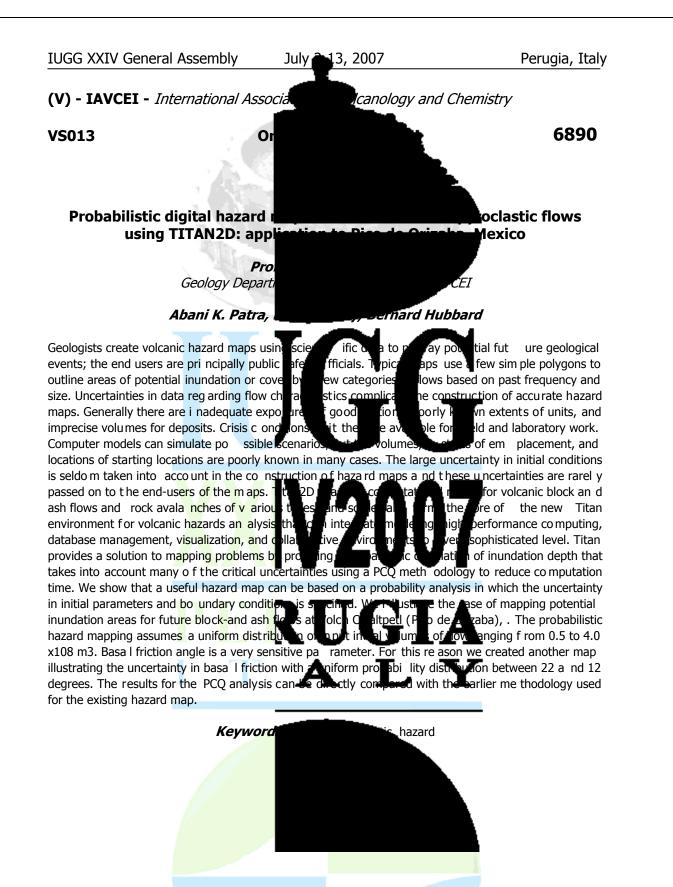


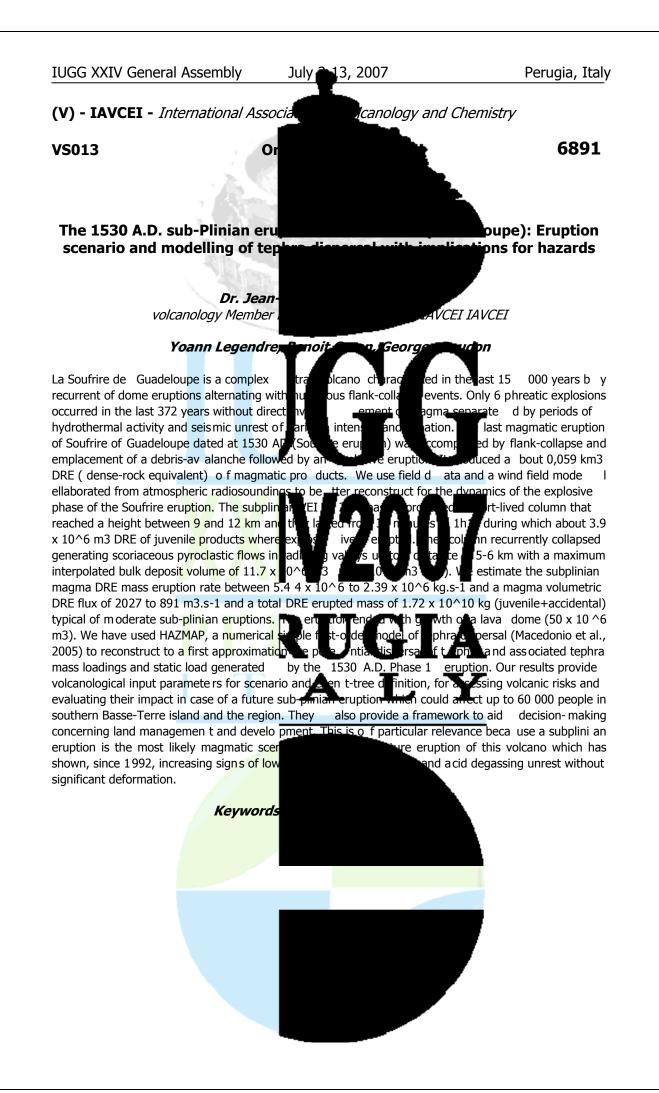


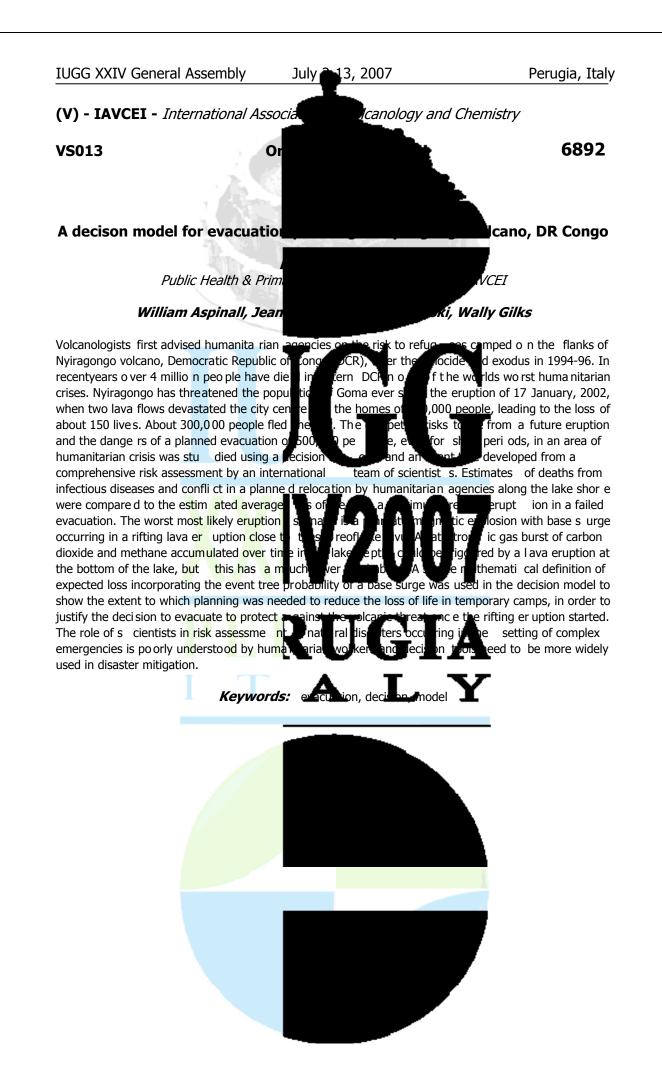


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#### IUGG XXIV General Assembly

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### **VS015**

## Symposium **New Techniques using Remote** Analysis: Observations, Integra

### Convener : Prof. Ken Dean Co-Convener : Dr. David Rothery

The use of remote sensing data and model has significantly improved ov er the past f eruption processes, distribution of eruption

launched that provide various facets of resolution not hourly to daily coverage has increased, and ome many as 36 sp ectral bands. Moderate spati and IR wavele ngths with pointing capabilitie spatial resolution up to 60 cm are also availa and volcanic processes can now be detected and details in the structure and composition an important tool for prediction of activity predict the movement of volcanic clouds. New field-

satellite data. Parallel to sensor developme nts, improvements in connec tivity, infrastructure and web browsers have resulted in better and more data are now received by local stations to Italy, Central America, Alaska, Hawaii, Kam new remo te sensing system s and techniq assessments are encouraged for this session ASTER and Landsat), we encourage preser this void.

canology and Chemistry

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# Monitoring and and Modeling

anoes, and to assess hazards, nde a better understanding o and assessments. New sensors have been



he number of satellites with as 250 m pixels and as as sr de multispectral bands in visible erage every few d ays. Data with ements, subtle increases in activity creases in thermal flux posits deling has also become iraph hange and wind-fields to ameras, radiometers, and

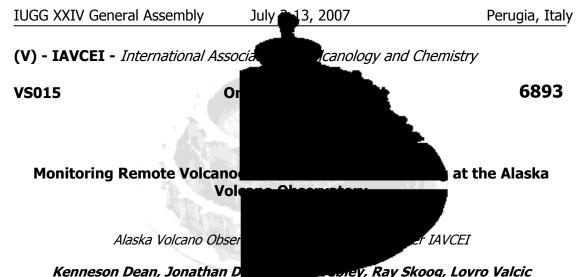
webcams detect volcanic pr ocesses not see n from satellites and provide calibration an d validation of

sue



distribution. Satellite ind assess volcanoes in ons on the use of these eruptions and hazar d resolution sensors (e.g. ent of new sensors to fill

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The Alaska and Aleutian Arc area of the Nor regions in the world. It is home to over 10 observed eruptions, which represents almos these volcanoes are located far from Alaska threat to human life and activities due to the of dollars of f reight and over 20,000 pas Casadevall, 2000) on air rout es between Asia and

Europe. The pr oblem with monitoring these volcanoes is that they are spread across ~3 000 km, with many lying in remote areas t hat often experience noor weather conditions. Currently less than half o f the regions volcanoes are monitored by se making expansion a difficult pro position. and unproductive method of monitoring. satellite coverage as the tracks of polar ort a robust means to regularly observe activ

Observatory Remote S ensing Group (AV ORS) primarily operates u sing data from three grou ps of satellites sensors: Geostat ionary Operational Enviro nmental Satellites (G OES) imager, Advanced Very High Resoluti on Radiometer (AVHRR) on Spectroradiometer (MODIS) on NASAs Terr acquired through onsite rec eiving stations the Fairbanks (UAF). Additional AVH RR coverage is provided by NOAAs Gil more Creek Tr acking Station. GOES data is supplied by the Naval Research Landatory at Minterey Bay, Salfornia, USA. For analysis purposes automated processing systems divide the regions. All images for each region are then checked by AVORS analysts to identify any new or ongoing volcanic activity. These checks are m ade twice daily under normal conditions, more frequently d uring

times of elevated alert status or eruption. algorithms designed to ident ify possible alerting duty personnel. During each moni previous session, using a proprietary image anomalies and volcanic clouds/plumes, red web-based interface. This interface is one support viewing and interpretation of the Further interpretation is also aided by mod visualization tools (e.g., Google Earth).

ey, Ray Skoog, Lovro Valcic

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ost volcanically active er 40 of which have experienced tive volcan oes. Alth ough many of canic activity still poses a serious over the region. Millions c pass tia n ever y d ay (Miller and g circum-polar routes to

> nd financial restraints s would be a spo radic ation, al lows frequent

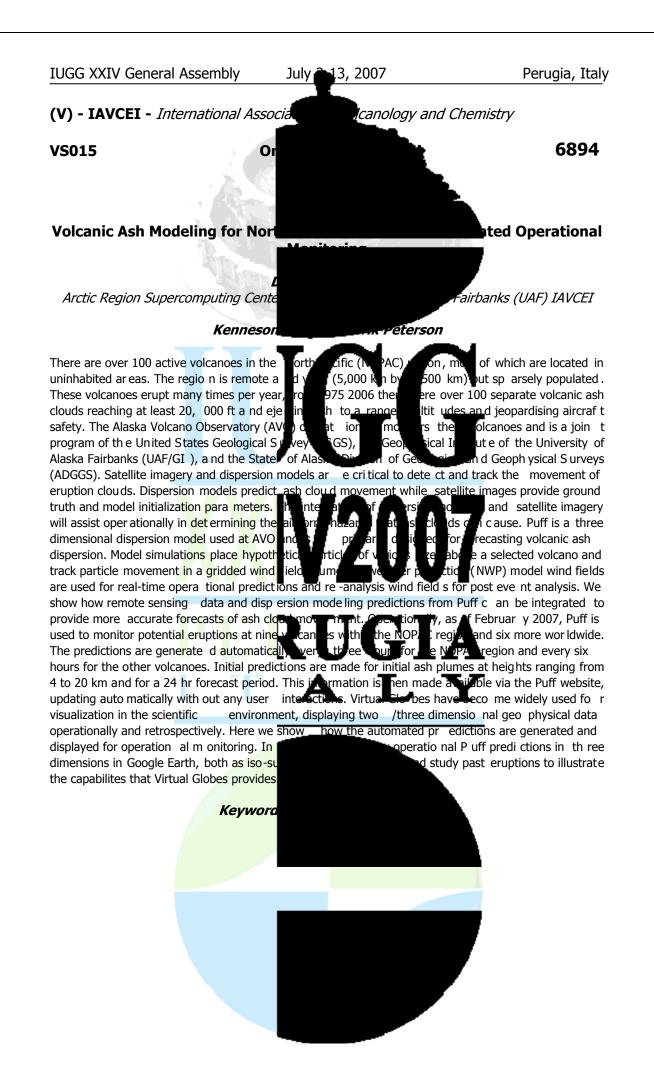
near the pole, providing oes. The Alaska Volcano

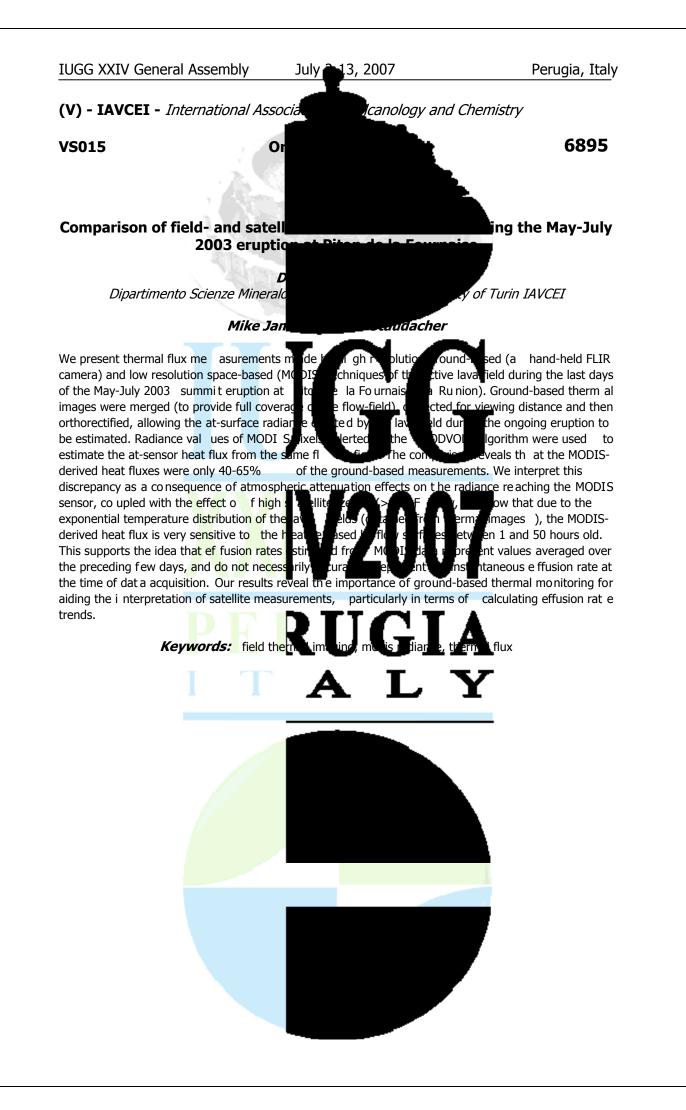
Modente Resolution Imagin g he 🖌 RR and MODIS data are ute GL, Univer sity of Alaska

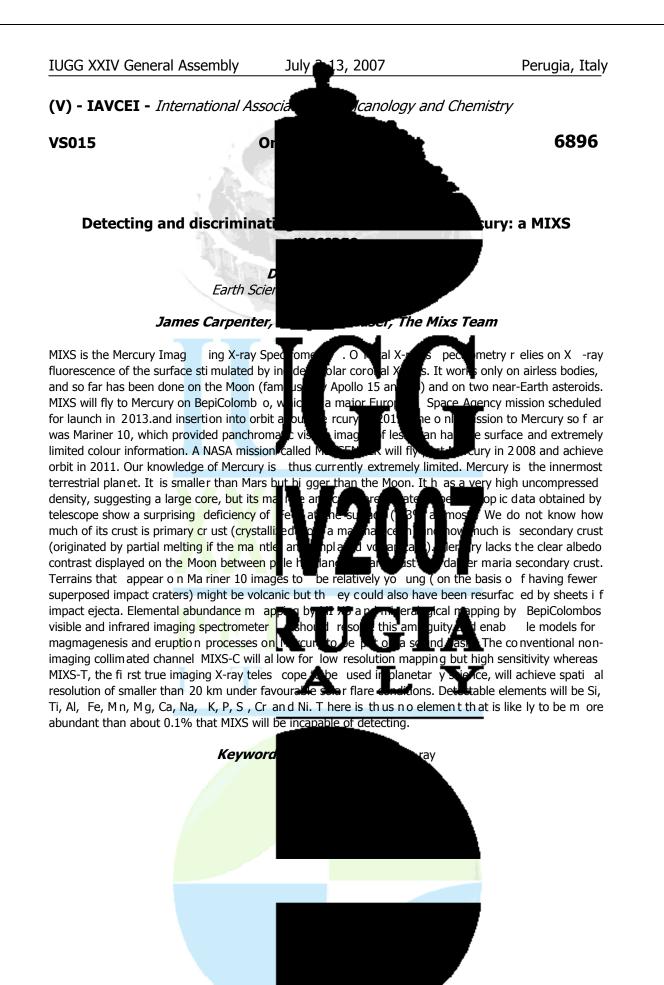
that c over the volcani c

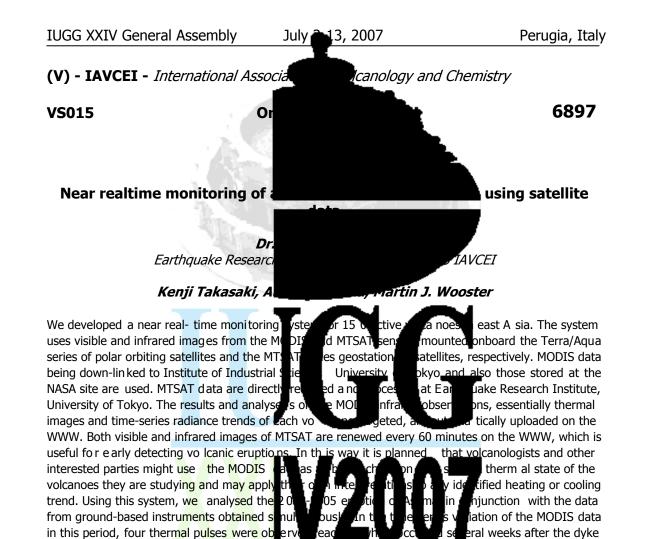
also performed by automat ed erate emails and t ext messages I new data acquired since the e analyst ide ntifies thermal base acce ssed through a developed by AVORS to sers Interfaces (GUIs). sh tracking model) and











intrusion at a deep level, as suggested by the GPS ground deformation measurements. The first and the second pulses involved Vulcanian or Strombolian eruptions, whereas the third and the fourth ones were

This indicates that magma had reached a shallow level then. Such a pre-eruption thermal anomalies is

ast Asia

evel of the conduit (1km

deep level. On the other

inning of the eruption.

ased monitoring system. The MODIS and MTSAT-

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non-eruptive activity. This means that m

bellow sea level) even if eruption did not o

hand, the first thermal pulse started rising a

This will be particularly useful in the remote areas

important for detecting a sign of co ming volcaning ruption, wi

based system is accessible at the URL: http://vrsserv.eri.u-tokyo.ac.jp/REALVOLC/.



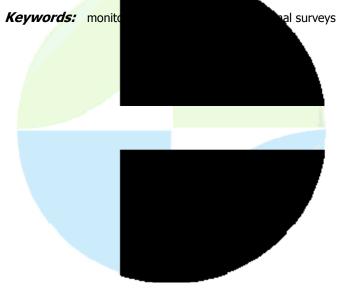
30 December 2002. The ING V Sezione di hand-held thermal camera. This instrumen anomaly recorded on the su rface of activ eruptions and eruptive processes. After the tw cameras have been installed on Strombol 1, Etna and Vulcano, allowing us to keep under control the eruptive activity, flank stability and as h emission. On Etna, we have monitored the 2002-03, 2004-05,

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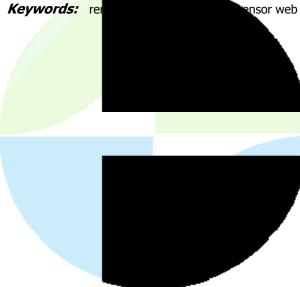
tive volcanoes using a er to detect any thermal applied to a number of Stromboli, fixed thermal

ar

s from helicopter allowed July 2006 and August-December 2006 erup surve us to follow the propagation of ephemeral nidden va tubes, as well as the ta <u>mera n</u>ave also been used to stages of inflation and deflation of the uppe ava fie ma calculate the effusion rate, the most import ant parameter to estimate maximum lava flow length, and also to detect ash plumes on Etna in good whether conditions. However, the three most recent 96, did not show eviden t eruptions on Etna, occurred on 2004-05, July 2006 nd Augu December 2 thermal anom alies on the s ummit craters before the opening of erupt ive fissures. Thus, the role of thermal anomalies and their meaning should be <u>compared</u> to and discussed with other geophysical data in order to understand when and if these d ecast eruptions.

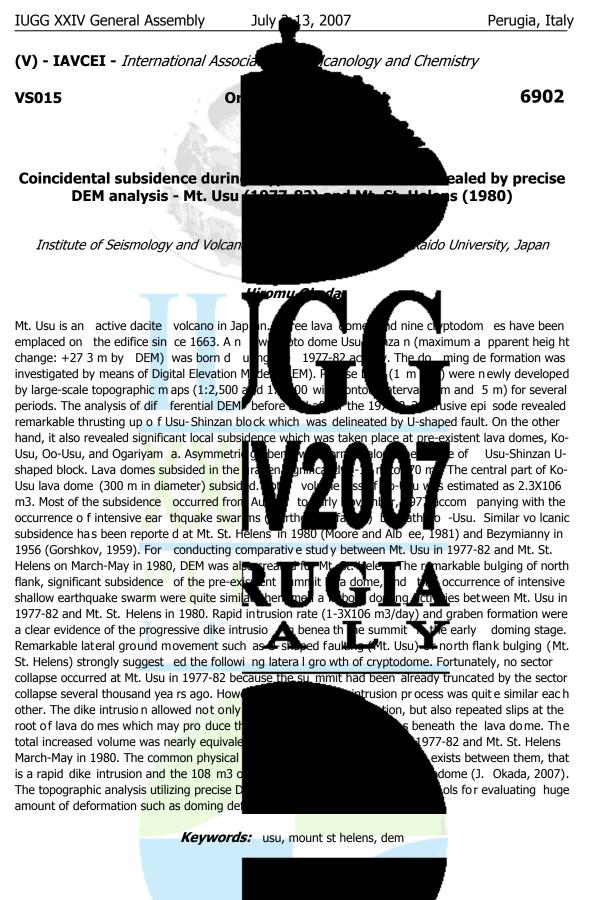


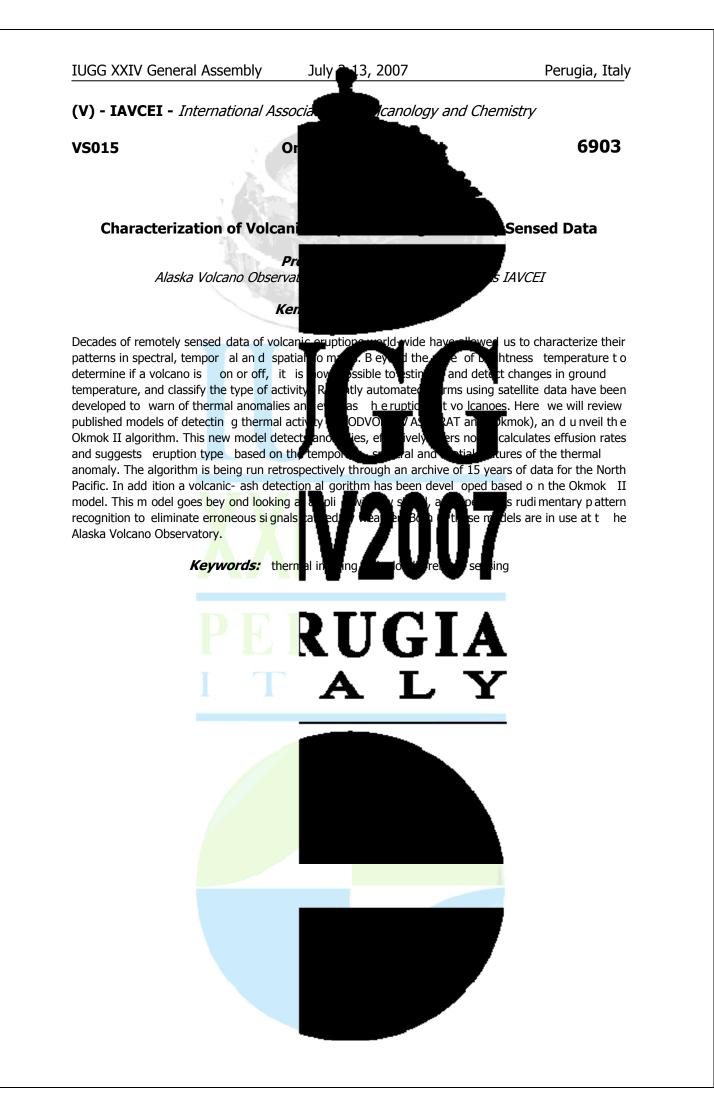




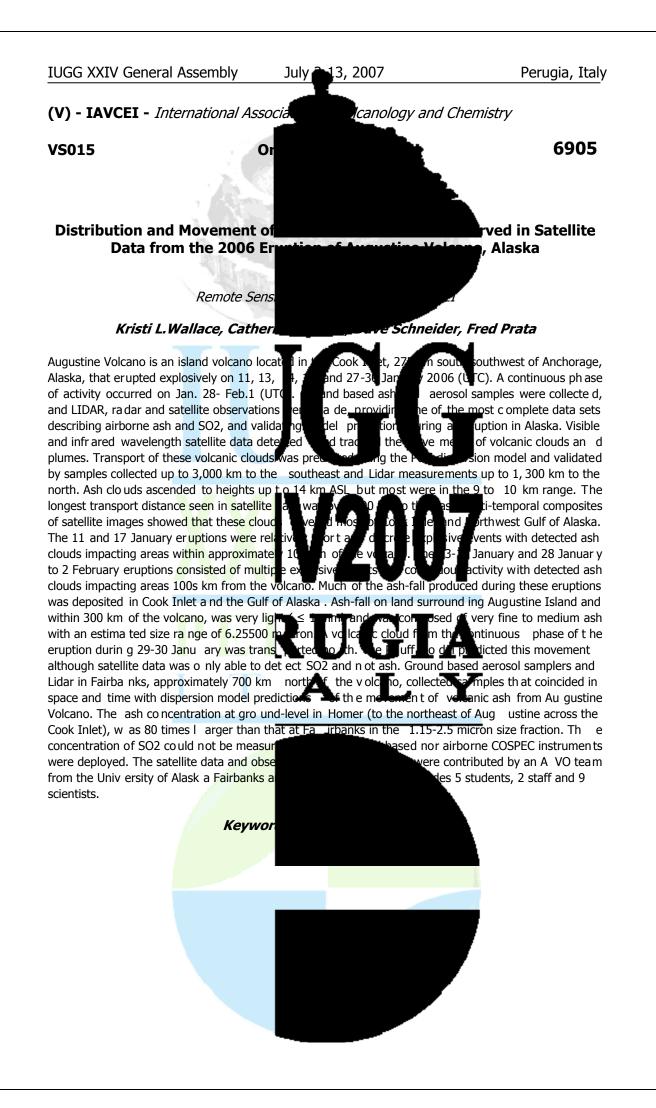


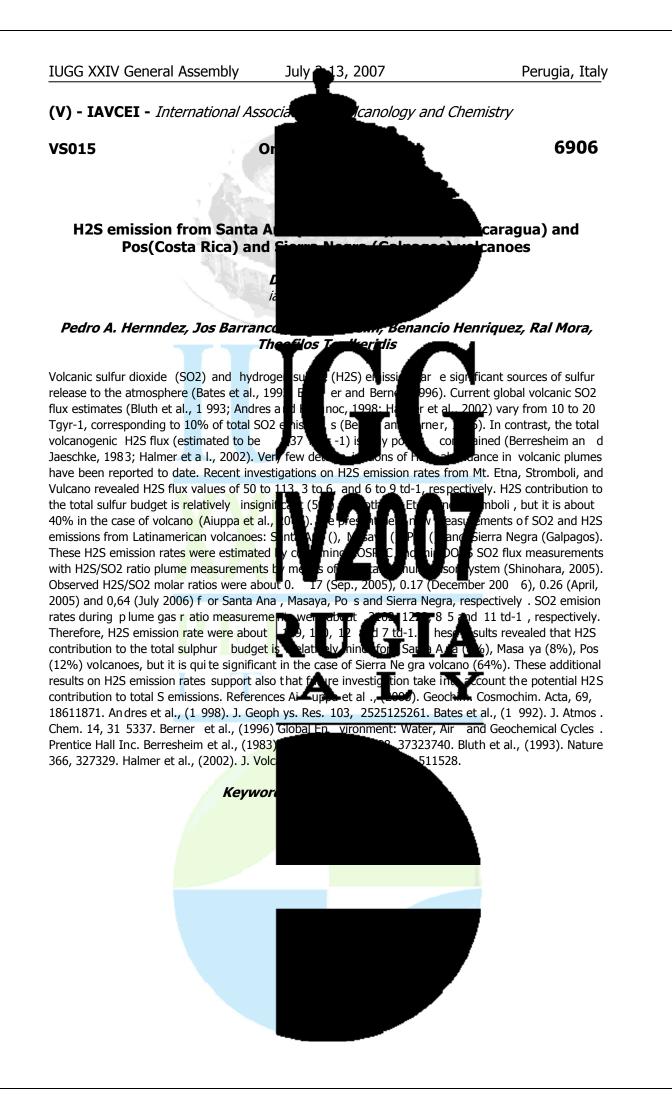


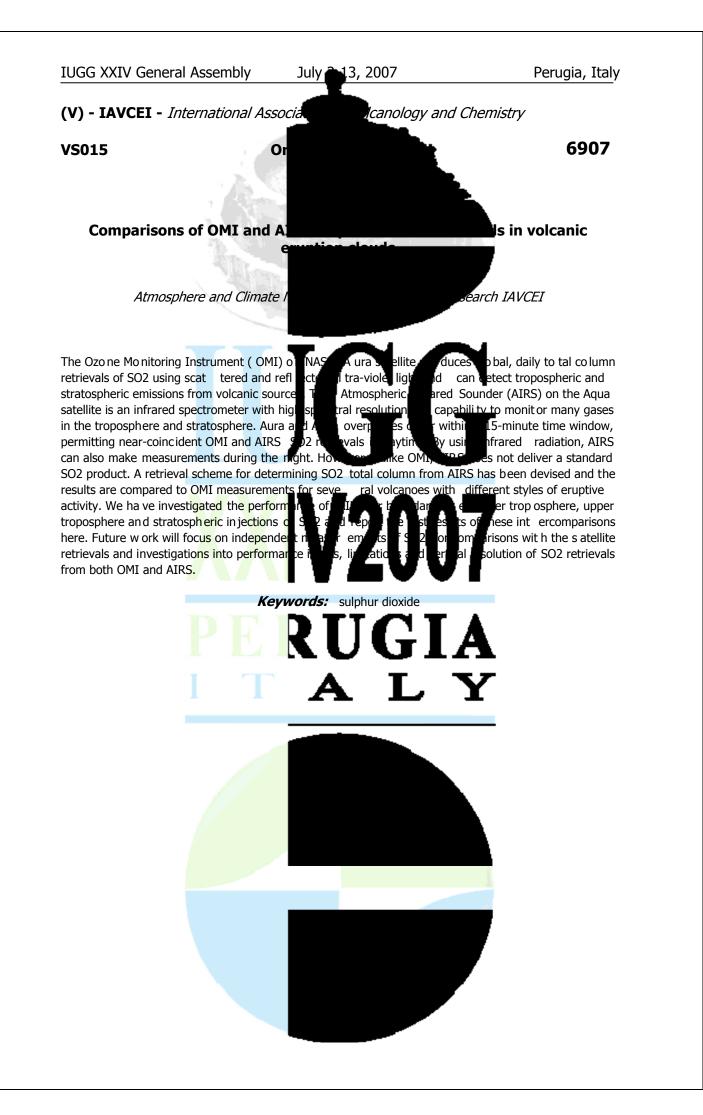


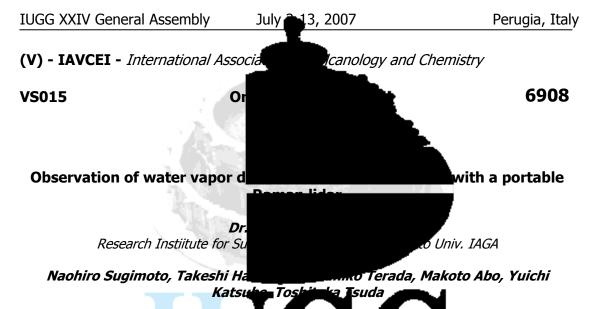












Observation of water vapor density and flux by water vapor is the majorsource of energy sensing measurement technique of waver va of the difficulty to distinguish water vapor in sensings such as DOAS and FTIR. A lidar (I a function ofline- of-site distance (range), and

distribution ins ide and outside the plume.A transp (SHG(Second harmonic generation):532 nm) and a telescope with a 35.5 cm diameter has been built in order to monit or water vapor profiles in th forSustainable Humanosphere), Kyoto Univ distributionin the fumaroles. The first expe m Japan in November 2005. Alternating observer /ati out in the directions towards the plume and the

ambientatmosphere. The water vapor flux velocity of 4 m/s measured with a video different from that estimated by the plume higher portability without deg rading the sensitivith The se cor carried out in January2007 at the same location observed in the plume in this experiment under will also be reported in the paper and the

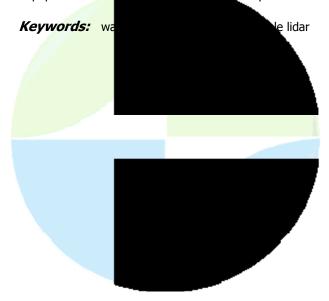
fumarolesis v rom volcanos th e <u>fumarol</u>e d outs is ca le o ould b

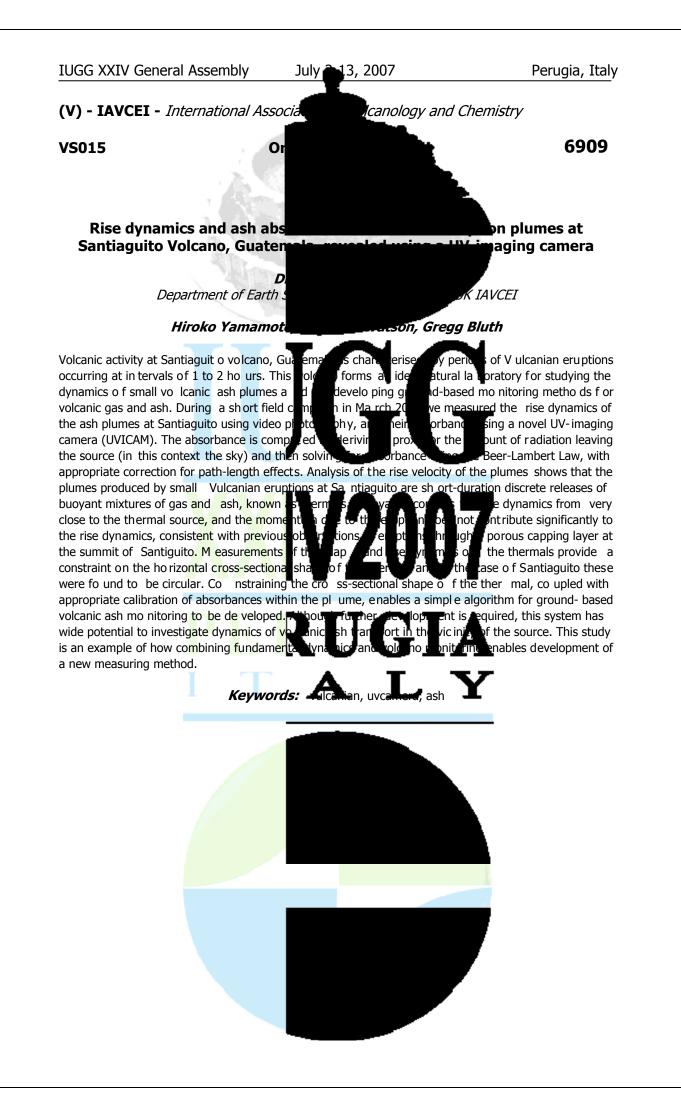
mportant pecause the latent heat the atmosphere. However, remote not <u>vet been</u> de veloped, because eans of p assive remote ume 🖞 asurir mospheric quantities as easurewater vapor

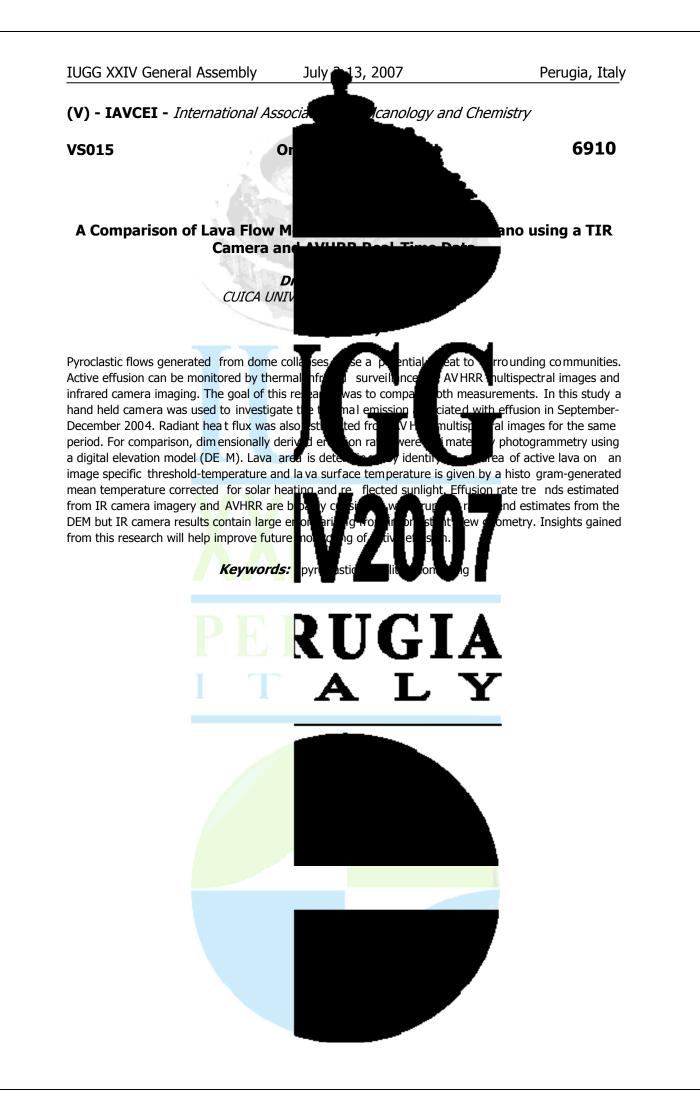
ortable Raman lidar with a pulsed Nd:YAG laser

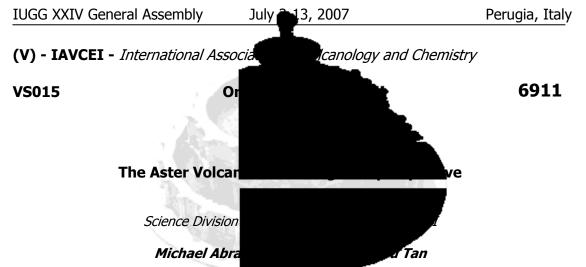
H (Research Institute leasuring water vapor of Mt. Aso, in Kyushu, water va por was carr ied marole was located at a

distance of 500 m from the lidar with a half maximum width of 200 m. The difference of the two beams indicated that the water vapor mixing ratio was ab out 1g/kg larger in t he plume com pared with the aboute100 kg/s using an upward ume his flux value is not too the further eveloped a smaller lidar .me (20 system by using moresensitive PMTs (pho tomultiplier tu bes) and a s maller telescope, whichshows experiment ing the new system was Mt. Aso, ncreasethe water vaporwas not the condition of relative humidity of ~100 % in the ambient atmosphere in the midwinter. More recent results from this portable system plannedin summer ped system will be discussed.









The systematic study of the world's most frequent vo its sisters MISR and MOD IS, the importance of w assessments. A significant e merging challer e burgeoning da ta archive in a way that allow information in a timely way. This issue is pa multi-spectral, high spatial resolution, featu granules; and specifically for our volcano volcanoes worldwide. To promptly and efficiently a

large ASTER i mage library, and to house

activity is a s froi h ext isi c II EOS il stru th survey, ext y ac<u>ute in ge</u>r ic tar data hich ally c foi ste nanac

elling arena for ASTER and anology to societal r isk ts is how to effectively access a on, and di stribution of important for ASTER, which has produced a e of over 1 million data ving over 1500 activ e us volcano data within a

other anci llary correlative vo Icanological da ta from MISR,

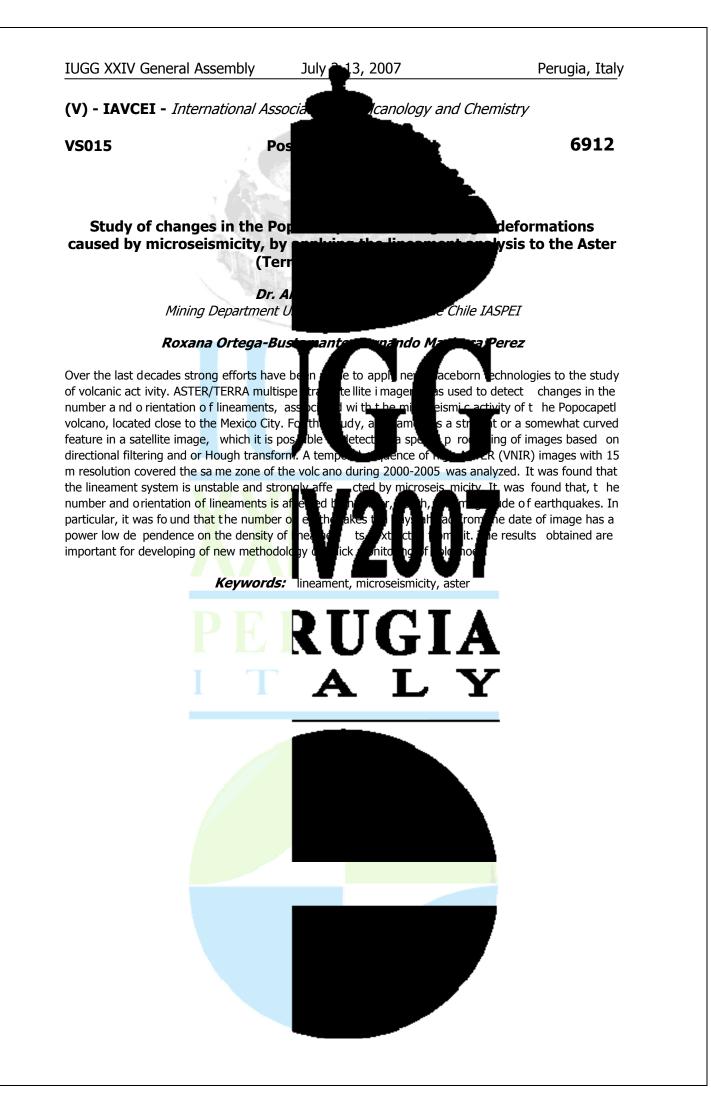
or

JPL ASTER Vol cano Archive (http://ava.jpl System (VDAA S). VDAAS pr ovides a fast intranet and internet access to these data. this unprecede ntedly large and accessible gld terms of data mining, data analysis, and da la d responders, the general public, and to educators

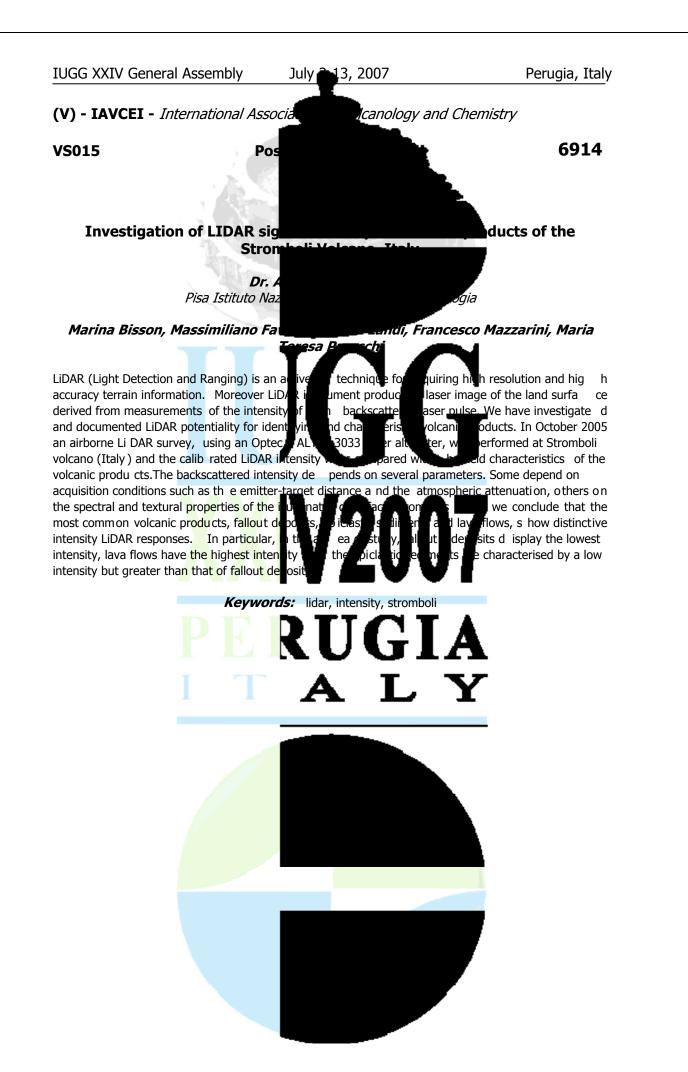
MODIS, EO-1 data sets, SRTM, and related in situ data, we have created a specialty domain called the olc

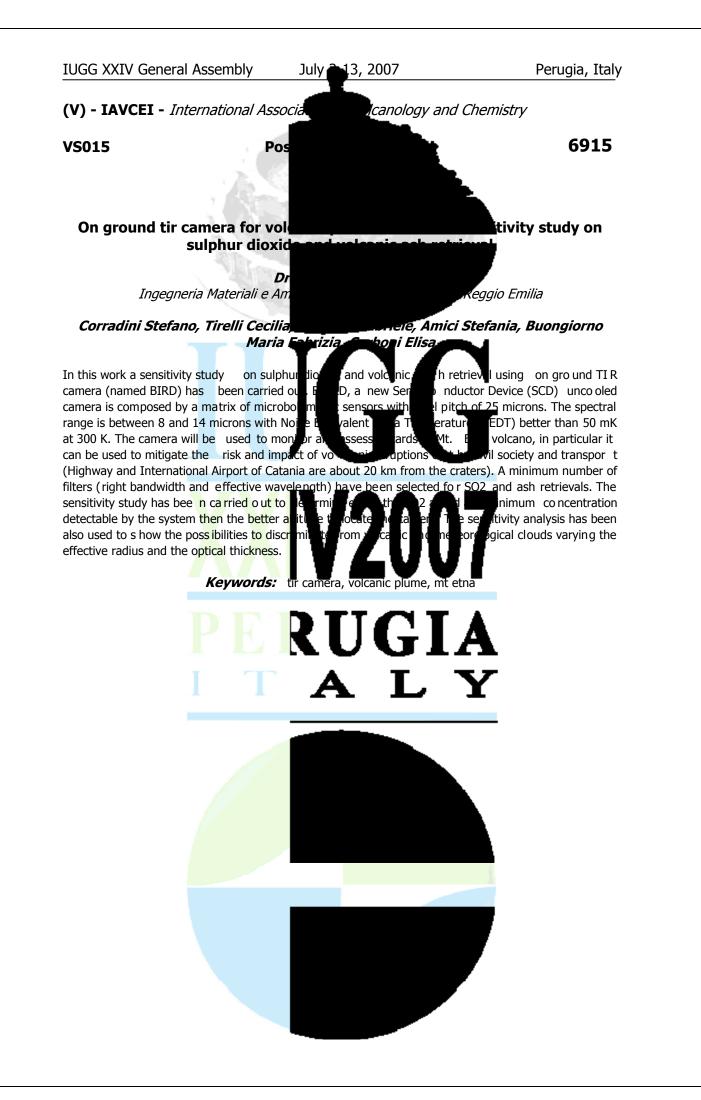
isition and Analysis nize a nd provide JP L s and opportunities that g data s et represent in commu nity, to d isaster

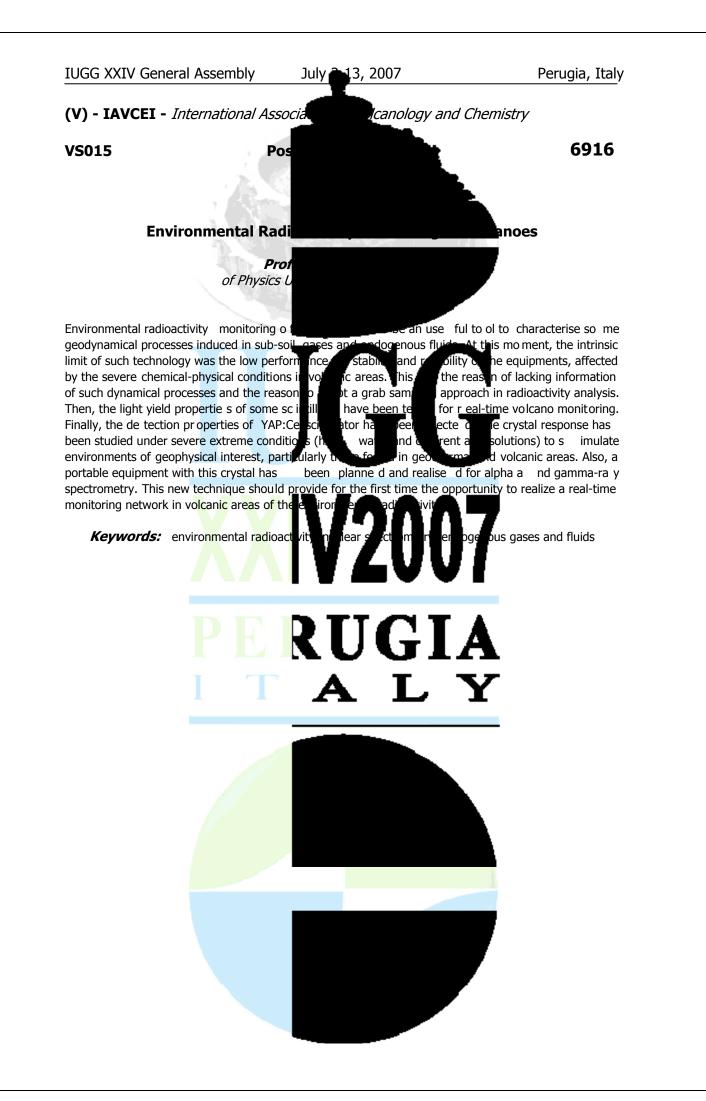


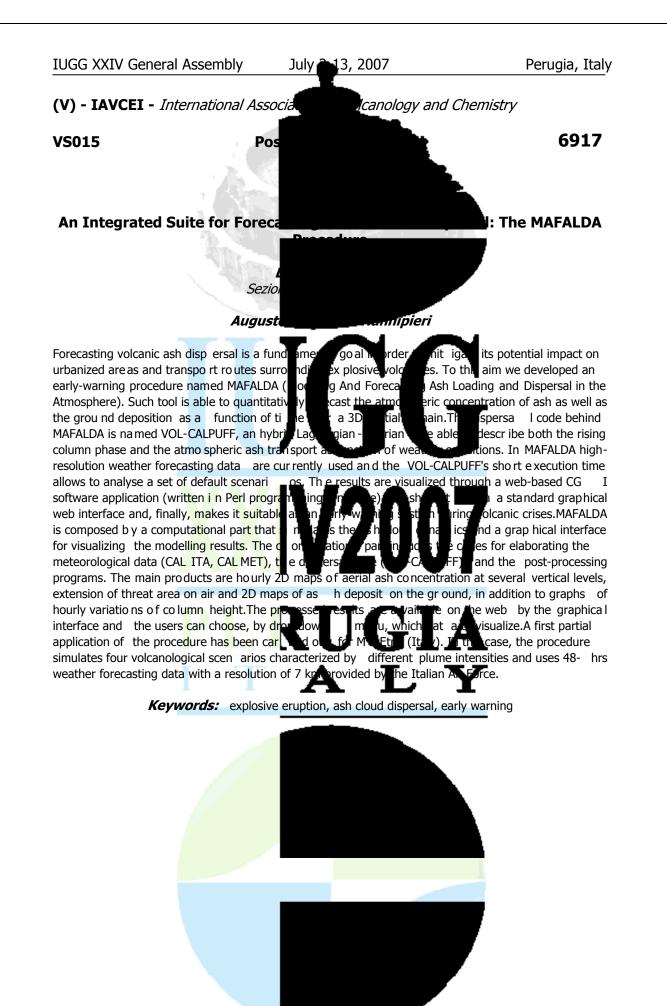


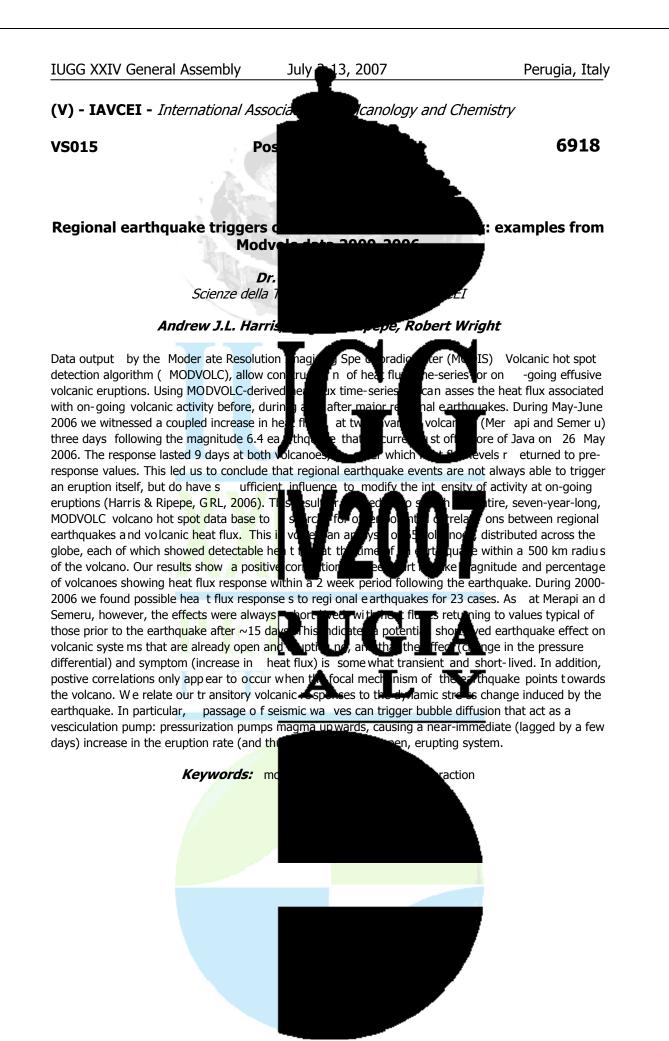




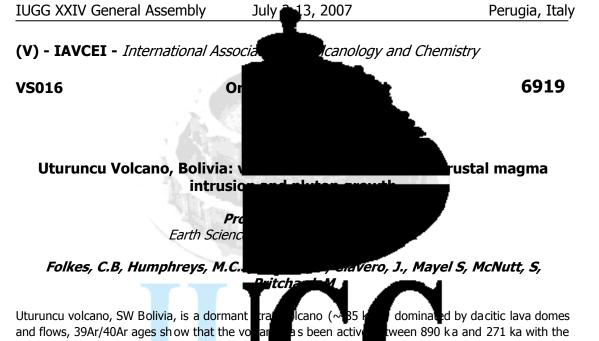












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lavas becoming younger and less extensive Geodetic satellite measurements record and with a central uplift rate of 1-2 cm/year. De rm beneath current local relief. D eformation indicates

average of ~ 1 m3/s (10-2 km3/yr). Th Altiplano-Puna regional crustal magma body. In a reat about 4 km depth below the centre of the has normal b values and is attributed to b magma intrusion. The porphyritic dacite lay magnetite-ilmenite assemblage and comme nodules and crustal xenoliths. Temperature about 980oC for the silicic andesites. C

the silicic and esite. Reversel y zoned orth explained by ch anging oxidation states du crystals and Fe3+ in plagio clase provide ascended, degassed and became more oxidised. The geophysical and petrological observations suggest that magma is being intruded into the Altip lanor na regional crustal mag abody at 17 km or mor e depth. In the Late Pleistocene dacitic and a newsiti

character of a ndesite magmas may reflect volatile \_-rich melt that deg assed, leaving the melt mor e oxidised before mixing with a more volumi anomalies and 270 ka of dormancy indicate intrusion. Uturuncu is thus inferred to be might eventually lead to a large eruption o caldera formation.

e deformat ion is attributed to magma intrusion into the eni te

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here are current signs of unrest. on field m 1994 to the present a sour t depths of 17 to 29 km 8 m3 over 12 years, a n

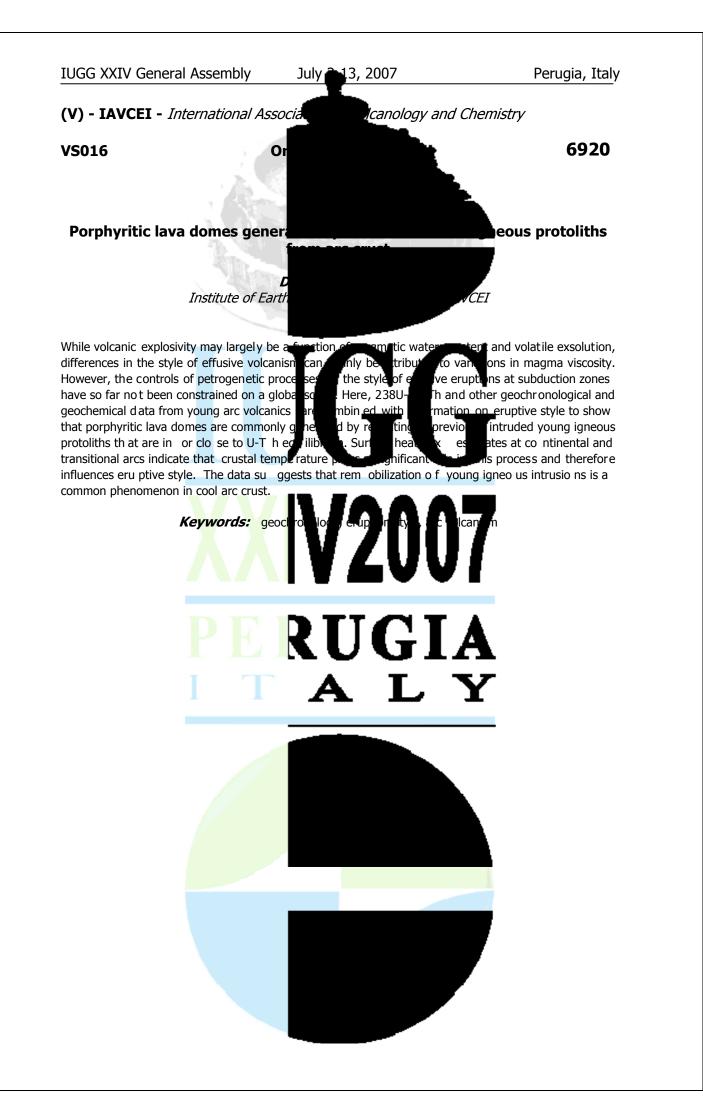
connaissance survey, seismic activity was recorded hmit. The seismic data ab ove the active dee p e-orthopyroxene-biotitehclusions, cognate norite 2oC for the dacites and ompositions and zoning patterns of orthopyroxene an d

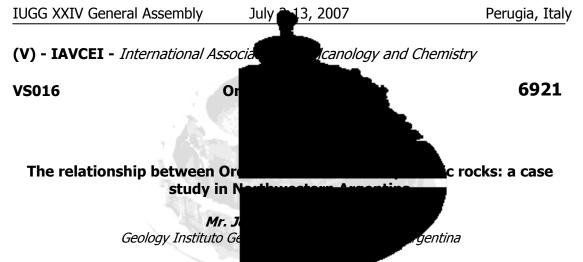
plagioclase phenocrysts indicate compositional variation in the dacites is caused by mag ma mixing with he and esitic end-member are -2+ os from orth opyroxene ceo helt that subsequently

÷. from 🖬 regional c rustal magma body to a shallow magma sys tem where they crystalli ze and mingle together. The late-stage oxidised current unrest together, geophysical

tem is in a prolonged period of growth. Such circumstances magma with potential for

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Pimentel, Marcio M., Viran

The geodynamic evolution of the Proto-Andrean subject of controversy. On one hand, it was that associated with the docking of several terra Coira et al., 1999). On the other hand, Dam Zimmermann and Bahlburg 2 003 and Franz evolution dominated by intracrustal recyclin prd The Puna region of NW-Argentina, records ediment

and Mon 1999; Becchio et al., 1999; Lucassen et al. discriminate different tectono-thermal cycle al., 2006) and sedimentological (Zimmerm Lower Ordo vician perio d the central-east

setting. This idea supports pr evious works isotopic data from metamorphic basement 1999; Lucassen et al., 2000; Lucassen and Franz 2005; Franz et al., 2006; Viramonte et al, accepted).

and acidic plutonic rocks as well as volcan two N -S trending belts about 600 km long Ordovician magmatic rocks of the south eas voluminous and widespread plutonic unit is exposed. It is composed of three silica-rich facies with U-Pb zircon and monazite ages of 475 to 463 Ma (Viraconte et al. accepted). S

plutonic unit, a thick volcanosedimentary sequence felsic metarhyolites and metadacites, intercalated with phyllites and metagreywackes is also exposed. A The bimodal nature of the m agmatism in

rocks have basalt and rhyolite composition isotope data. Amphibolites yield positive ratios of 0.7067. These values and the obs derived from a mantle source in a subduct 1.43 and LaN/YbN = 1.39 to 1.48) on a c ridge basalts (TMORB). On the other han subalkaline and display peraluminous chara

Mesoproterozoic model ages (T DM values rocks yield negative εNd(T) values betwee TDM model ag es between 1.5 and 1.7 Ga. magmas represent the product of meltin metapelites / metagreywackes of the base

nan, Becchio, Raul Alberto

ndwa jin of e rized as the mos et al., 19 , 19<u>90: Bec</u>c 2006 a ģ es w nino

ower Paleozoic is still a alt of replated sub duction events Ramos 1988; Rapel a et al., 1998; t al.<u>1999:</u> Lucassen et al., 2000; suggest a geodynamic au tl tribu of juvenile magmatism. orphism and magmatism

events during the period between ca. 510 to 440 Ma (Bahlburg and Furlong 1996; Moya 1999; Hongn

2000; Hongn et al., 2005) in which it is difficult to lies bad Łа.

of

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, 2005; Kirschbaum et aggest t hat during t he nsio nal intraco ntinental ogical, geochemical and Miocene) (Becchio et al.,

The evolution of magmatism started during the Early Tremadocian and it comprises mainly intermediate gengal, this magmatism forms

> hemi and isotopic data for una s pasented. In this area a adally associated with this ef bi mo metavolcan rocks, metabasites and

metarhyolite forming the b ase of the se quence vields an age of 485 Ma (Viramonte et al., accepted). ry unit is evident as the metavolcanic by trace elements and Sr and Nd d +2.5 and initial 87Sr/86Sr anomaly suggest they were ern (LaN/SmN = 1.23 to)

ple transitional midocean cs rocks are mai nlv They are enriched in

LREE relative to HREE (LaN/S mN = 2.17 to 3.26 GdN/ YbN = 0.94 to 1.47) with negative Eu anomalies and show neg ative Nb and Ta anomalies on mantle normalized multielement diagram. Plutonic r ocks yield negative ɛNd(T) values between -5.1 to -6.3, initial Sr ratios between 0.7216 to 0.7372 and

> Ga), whereas volcani c n 0.7089 to 0.7259 and ndicate that these felsi c e orth ogneisses an d / o r Our data combined with the

## **IUGG XXIV General Assembly**

literature about sedimentation, deformation a the most likely tectonic setti ng for the or According to this, the felsic magmatism astenosphere upwelling after tectonic swite long lived Hot Orogen (Collins, 2002; Hon Lithospheric modeling of the Ordovician for influence of a rc loading on foreland basi Lucassen, F., Franz, G., Viramonte, J. G. Viramonte, J.G. (Eds.) Relatorio XIV C ond orogens, tectonic switching, and creation Pichowiak, S., Harmon, R.S., Todt, W., evolution of the Central Andes: the baseme 101-126. Lucassen and Franz, 2005. The e (Eds.), Terrane Processes at t he Margins of Trumbull, R. B.; Romer, R. L.; Wi Ike, H.-G evolution at the Central Ande an continental and destruction - In: Oncken, O.; Ch ong, G M. R.; Wigger, P. (Eds.), The Andes Activ 1999. La deformacin ordovcica en el borde Argentino. Vol. 1, 212-216 (Gonzlez-B ong

ROM. La Plata, Argentina. Kirs chbaum, A Complex, eastern Pu na (NW Argentina): magmatism. Journal of Sout h American E H., Franz, G., Thirwall, M., Viramo nte, J.G the basement of the central andes (18-26 America Earth Science, 13, 697;715. Mend

de la Puna Oriental: Actas 5 Vol. 1, 134-15 2 (Gonzlez-Bonorino, G.; Om Parica, P. y R amos, V., 1986. El granito Catamarca. Revista de la Asociacin Ge T.E., Allmendinger, R.W., Mpodozis, M. Kay, S.M. of the central Argentine Chilean Andes. Tectonic Andes evidence for Cambrian continental c\_\_ C.W. Rapela (Eds): The Proto-Andean Marg of the Geologi cal Society. London. V iran Martino, R.,. O rdovician igneous and meta data and implications for the evolution o Sciences, accepted. Zimmermann a nd Bah Ordovician clastic deposits in the southern

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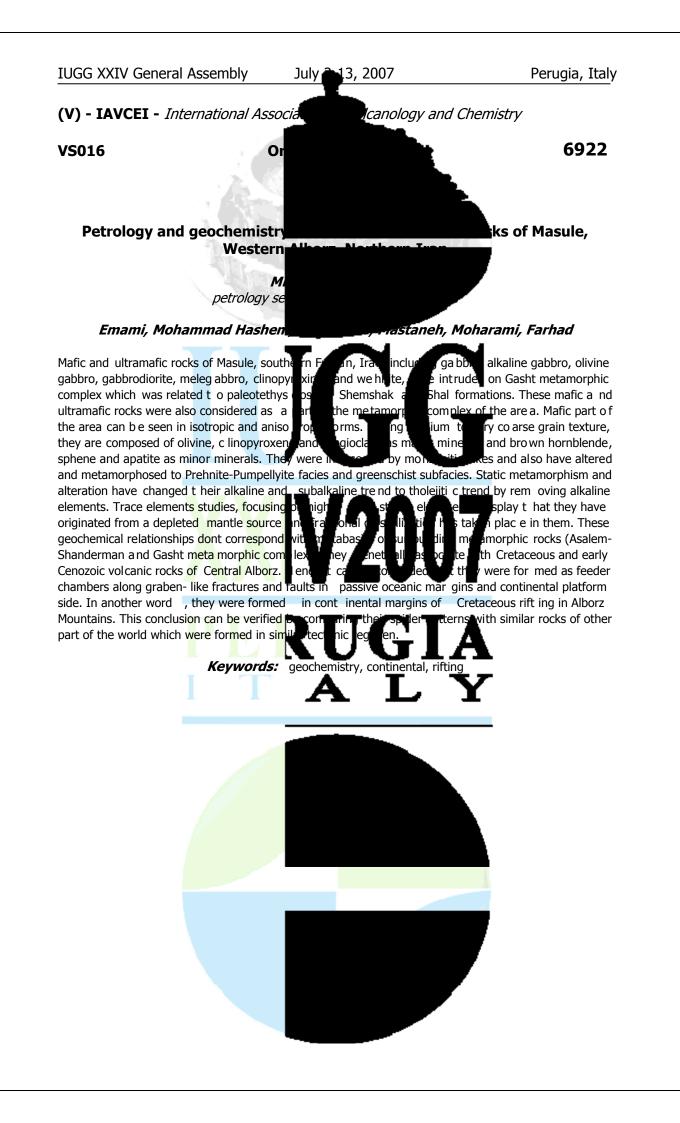
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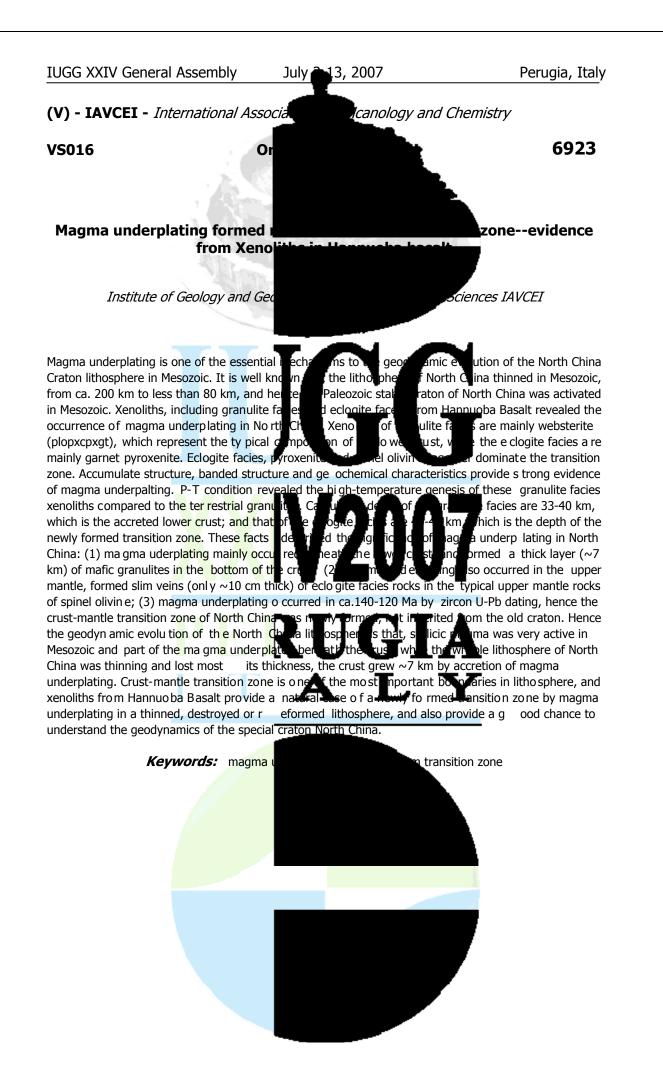
Perugia, Italy

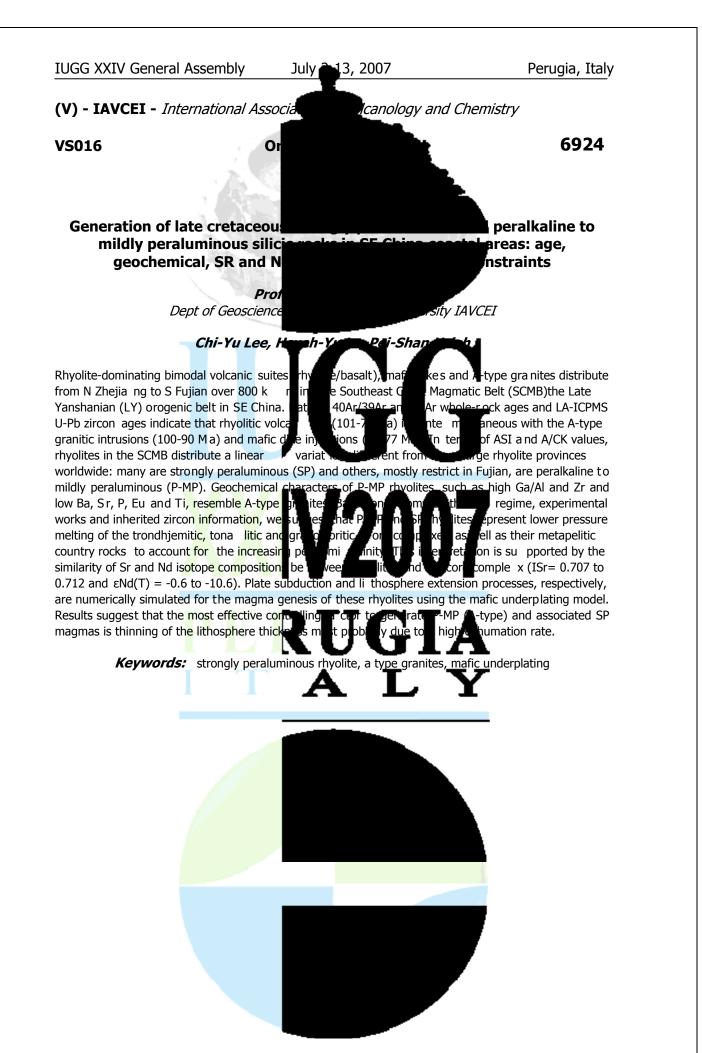
phism suggest that a retroarc basin represents agmatic ro cks in so utheastern P una. ive crustal melting associated t o subduction zone in the frame of a Bahlburg and Furlong 1996. western Argentina: on the 245-258. Becchio, R. Paleozoico inferior del Noroeste de Argentina (23-27S) metamorfismo y geocronologa. En: Gonzlez Bonorino, G., Omarini, R. y alta. Collins, 2002. Hot , 535-538. Damm, K., , H., 199 0. Pre-M esozoic of America Special Paper 241, the Central Andes: a non-collisional orogen comparable to the Cenozoic high plateau: - m: Vaughan, A. P. M.; Leat, P. T.; Pankhurst, R. J. cassen, F.; Kramer, W.; 273.Fr iebel, W. 2006. Crustal 5nte, J hio, R.; a geochemi ecor d of crustal gr owth, recycling G.; G iese, F ze, H.-J.; Ramos, V. A.; Strecker, ctio pring 64. Hongn and Mo n, de la F In latorio XIV Congreso Geolgico mari .; Vil ed.) S alta, Argentina. nt Hongn, F., Mon, R., Acua, P., Kirschbaum, A. Y Menegacu, N. 2005. Derormacin dctil intraordovcica en la Sierra de Cobres, (Puna Oriental-Noroeste Argentino). XVI Congreso Geolgico Argentino, Actas CD-The C obres Plutonic for Lower Paleo zoic F. Becchio, R., Wilke, aleozoic development of craton. Journal of South le , O., 1973. Faja Eruptiva a, [ Vid Congreso Ge olgico Argentin o, 4: 89;100. Crdoba. Moya, C., 1999. El Ordovcico en los Andes del n oroeste Argentino. In Relatorio del Congreso Geolgico Argentino, No. 14, .) Sa a, Argen tina. Palma, M., ectnico, Provincia d ific e Ramos V.A., Jordan Cortes, J.M. **19**86. Paleozoic terranes Palma, M 855;880 Ramos, V., 138. Late Proterozoic-Early Paleozoic of South America- a collisional history. Episodes II, p. 168;174. Rapela, C.W., Pankhurst, R.J. Casquet, C. Baldo, Saavedra, E. J. y Galindo, F., 1998. The Pampe an orogeny of the southern proto-Sierras de Cor doba. In: Pankhurst, R.J. and 142, p. 182-217. Special Publication amonte, J.G. Pimen tel, M.M. y Puna: New U-Pb and S m-Nd of South American Earth nd tecton ic setting of the

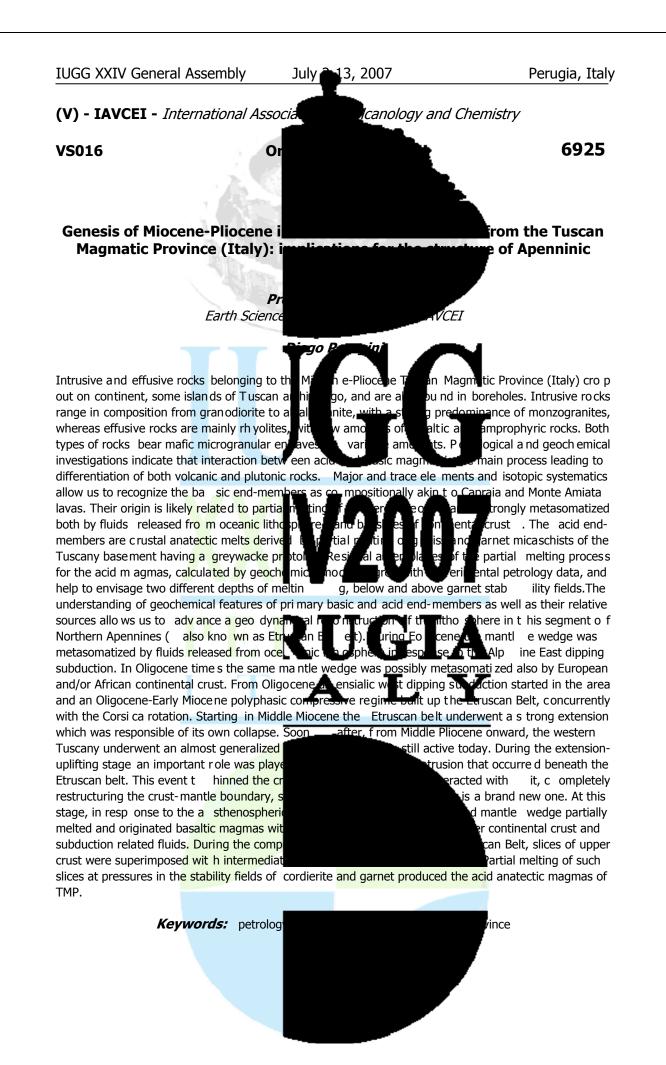
plogy, 50, 1079;1104.

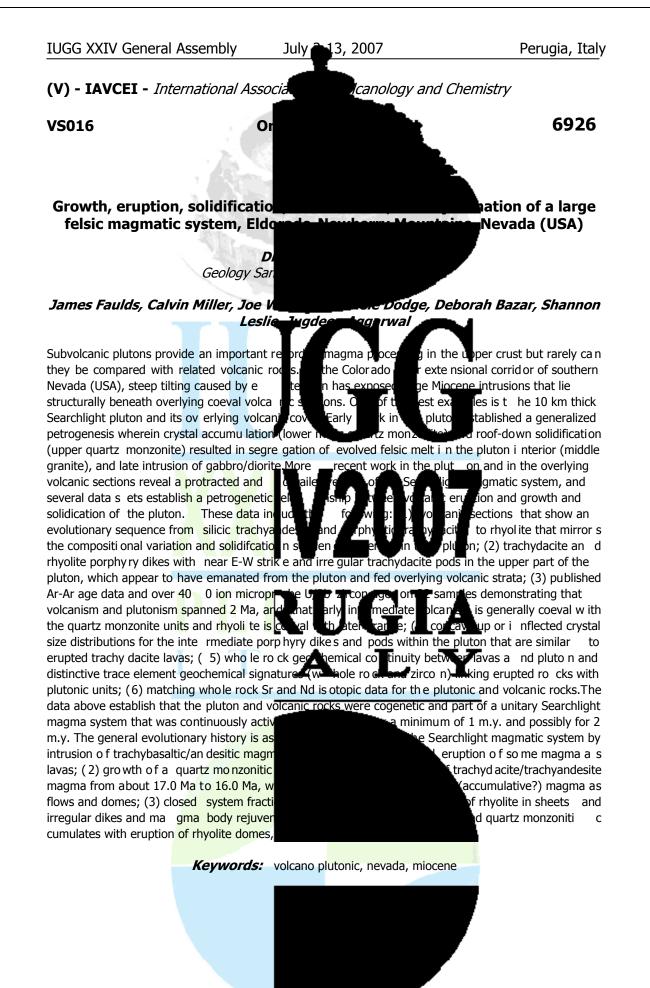
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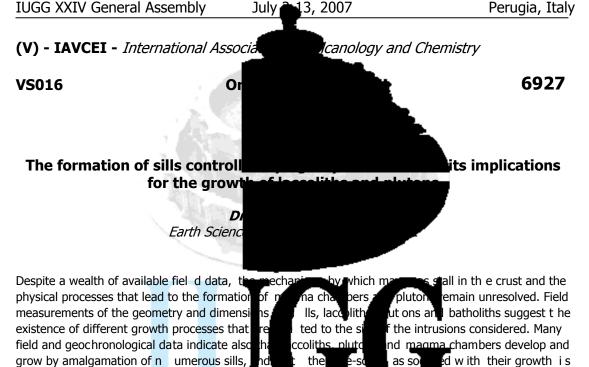












hypotheses for the formation of sills have been proposed decades ago but it is only recently that they

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have been tested seriously. These hypotheses have been tested by me ans of a nalogue experiments involving the injection of flui d into a solid conditions the formation of sills requires the only when their feeder dyke intersects an lower less rigi d, weaker layer. That litholo ica formation prov ides a mechanism for the g bwt

somehow related to their size. In many cas

plutons can grow by over-accretion, under-accretion or even mid-accretion of successive sills. In accord with field data , this model p redicts that a representing the cumulative t hickness of the lateral extend. The model also predicts that consequences for their size and sh ape. Viscoury-c lithological d iscontinuity they origin ally for

discontinuities and thus woul d provide a saucer-like shapes that are commonly obse

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and plutons. Yet, the mechanics and dynamics of s ment i

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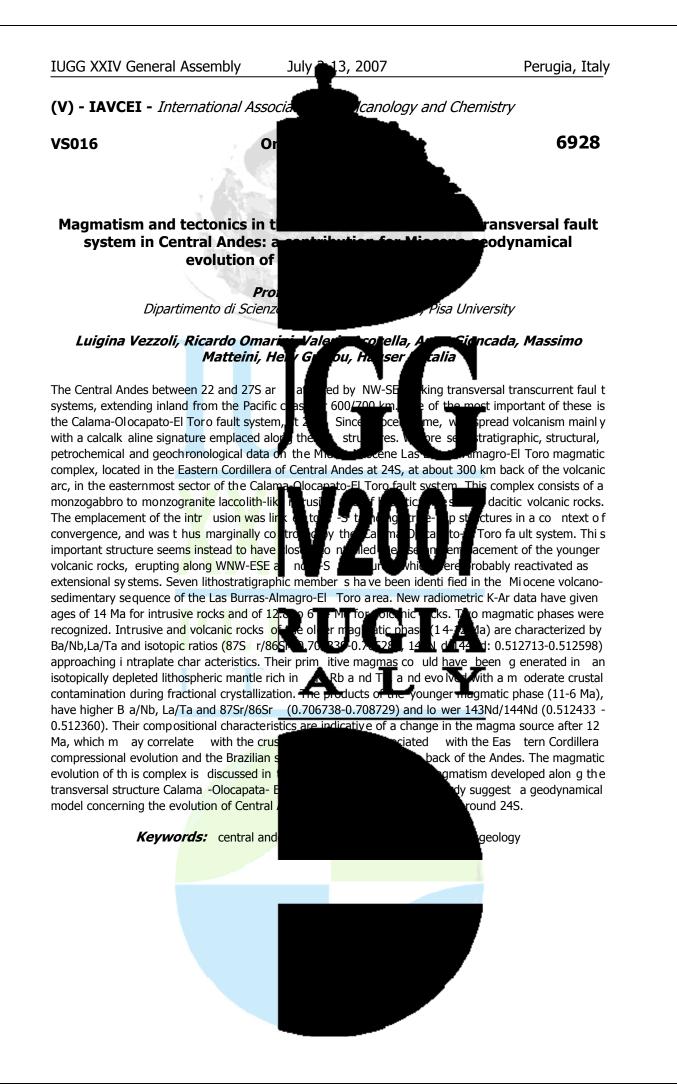
d with their growth is as so uildin cks of larger laccoliths ly understood. Different

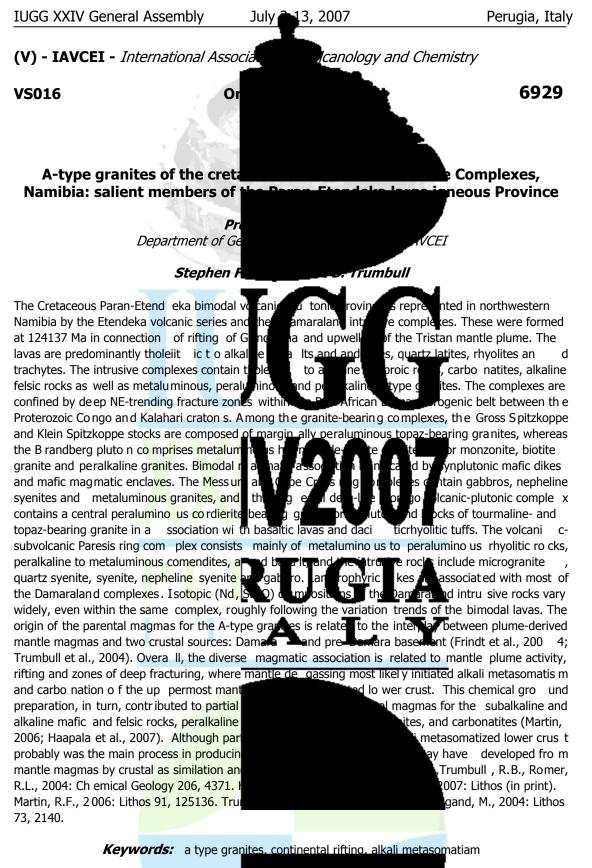
> hat un der hydrostatic lity, and that sills form id, stronger layer and a b ntrasts can co ntrol s ill The formation of a sill

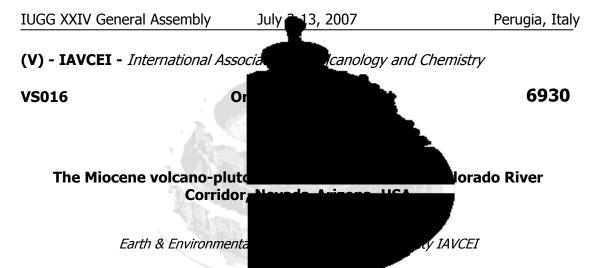
provides favourable rigidity a nisotropy for the empl accement of subseque nt sills so that laccoliths and ly by vertical expansion, ai

ai ni comparatively constant hs and plutons form is e ti οv wh hΙ CCO sca essentially determined by t he cumulative time between successive sill intrusions. The experiments also show that sill dynamics are controlled by viscous assipation of the fluid all of their length, which have trolled d mics enable sills to propagate further and thus to grow thicker than dykes of similar magmas. These dynamics enable sills to propagate faster and thus to induce no n-elastic deformations in surrounding rocks that will deviate them from the

them to feed n ew sills along other ressi ve character as well as the







J.E. Faulds, J.S. Miller, J.L. Wood , D.J. Furbish, L.L. Claiborne, B.A. Walker, D.S. Perr KF. Hody

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South of Las Vegas, Nev ada, the Co lorade Riv Arizona and Nevada and ma rks the center corridor. This tecto nic corridor underwent ra Ma, with less intense extens ion continuing volcanic sequences, spanned a longer inte al, extension reveals fossilized plumbing systems that

swarms, and hypabyssal pods and sills. We focus here on a 70 km long zone that exposes three major magmatic systems on the west side of the Searchlight pluton; and Aztec Wash and Ne intrusions; the erupted products of Search bh systems and volcanic sequences, while ver lik removal by faults. Roofs of the major plut ns i exposed, owing in part to demonstrates that plutonism at each center rema

Mountain and Searchlight, the earliest dated field relations document that it predates Azt ~16-17 Ma. The final intrusive pulse at each later stages (post-16.0 Ma) of intr usion are marked by i nput of bas alt that mingled extensively and locally mixed with granite (I ow-silica rhyoli te). A fic input in the system cryptic, manifested by widely distributed, fine-grain

The volcanic sequence and dikes in the r (trachydacite/trachyandesite) to fels ic (rhy The pluto nic portions of all centers reve replenishments of intermediate to felsic and magma within a more felsic h ost; transpo and roof and wall rocks; and accumulation magmas. The older (pre-16 Ma) portions a mushes in which details of h istories were

Mountain and Aztec Wash plutons indicate emplacem ent of replenishments as subho rizontal sheets. Earlier parts of the intrusions were probably also mostly into the weak crystal mush. Highly siliceous fractionates (~high-silica rhyolites) were extracted and emplaced at all centers as dikes, local Elemental zoning in zircons provides a reco

Mountain system through time that is con histories for zircons in the other centers (el

of the Colorado Ri ver extensional g mi<u>d-Mioce</u>ne time, from ~13-16 tism, marked by thi ck us ma g tha teep t t accompanied al million years spanning the onset of rapid extension. This system is manifested by small stock- to batholith-scale intrusions, dike

thro ugh tro ugh that separates



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it Mountain batholith; bundant hypab yssal tions between the other of lack of exposure and B-6 km; floors are poorly

ined a ctive until about 15.6-15.8 Ma; at Spirit

ton is not well dated, but d Aresults suggest an age of dik sv m. At each center, th e

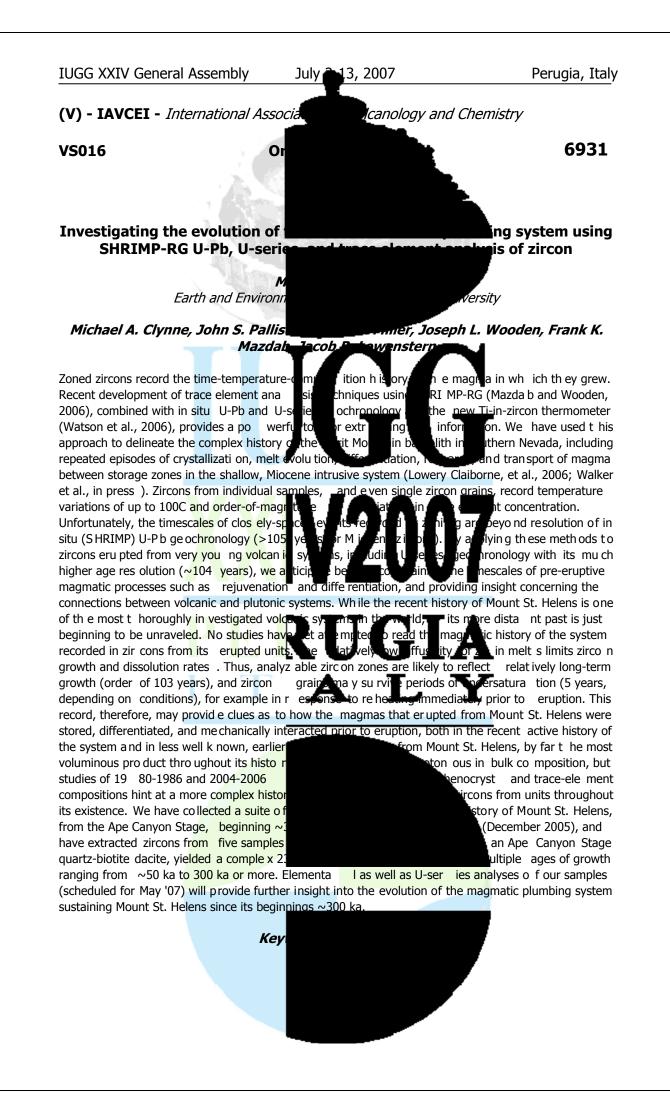
Te-16.0 Ma is largely 🖌 mafic-i miediate en aves with localized early hornblende gabbro. The dominant magmas during this stage were granite (possibly with qtz monzonite) at Spirit Mountain and quartz monzonite (trachydacite +/- trachyandesite) at Searchlig ht and Nelson. so reveal the shift from intermediate

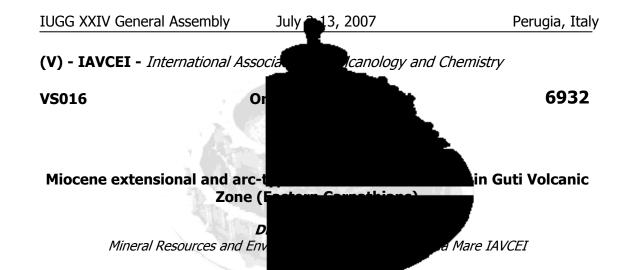
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magmatic inclusions) at ~16 Ma . d histories, with multiple nd disaggrega tion of mafic arlier magmatic produ cts the intermediate to felsic ished, long-lived crystal est port ions of Spirit

emplaced as sheets, both vertical and horizontal, ccumulations at roofs. pos ition in the Spirit b work suggests similar







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Miocene extensional and arc-type calquelline Carpathians)MARINEL KOV ACS, Alexan drine Ful 430083 Baia Mare, RomaniaGuti Volcanic Zo volcanic chain representing one of the most The volcanic chain has been built up in c Carpathian-Pannonian region. The magmatis European Plat e beneath the two microplat developed in the GVZ. A felsic rhyoliti c calc-alkalin

and associated reworked volcaniclastics developed in the south-western and southern part of the area. The felsic volcanism represents the onset of the volcanic activity in GV Z (~15.4 Ma) and has been correlated with the areal felsic extensi Transylvanian Basins (Pecskay et al. 2006) the felsic one, overlaying the ignimbrites de (13.4-7.0 Ma). It consists of predominant and associated subvolcanic intrusions. A m

intermediate volcanic act ivity in GV Z (8.1-7.0 Ma). The in termediate vol canism is typic al subductionrelated/arc-type with strong LILE and LREE enrichments, HFSE depletion and Sr-Nd isotopes negative correlation. Cr ustal assimilation involving elements geochemistry and isotopic da t 0.5123). A depleted MORB type mantle wed was asserted based mainly on HFSE geochemistry (Nb/Ta=16.1, Zr/Hf=34.6, and strong Nb depletion in the NMORB normalized diagrams). The main markle source-related proce magma genesi s was the mantle source enrichment (constrained by low U/T h(0.28) reported to high value s of Th, th e low Ce/Pb(3.5), the high Th/Ce

(0.15) and Pb/Nd (0.58) and enriched Pb isotope composition). The felsic extensional volcanism shows geochemical signatures rese mbling subdu enrichments and Eu negative anomaly, hig and Nb depletion in NMORB normalized dia the intermediate arc-type volcanics of GVZ. composition, suggesting mantle-source en magma of the felsic extensio nal volcanism extensional and arc- type is related to th Pannonian regio n. The felsic extensio na

be related to the uprise of an enriched (by Pannonian Basin. The continuous northwestward translation and rotation of the ALCAPA microplate has changed the position of the hinterland exte of the arc-type/intermediate volcanism on now GVZ explains the spatial coincidence o volcanic sources/centers). The arc- type vo of the subducted slab before the continent

blcanic Zone (Eastern re, V. Babes S tr. 62A, Baia I ene-Quaternary Inner Carpathians the Carpathian-Pannonian region. iary <u>aeotectonic</u> evolution of the bcene subduction of the ith the cia/Ti wo types of volcanism dera-related ignimbrites

> he Pannonian and ine volcanism following huch longer time interval rom basalts to rhyolitesgh Al b asalts ceased the

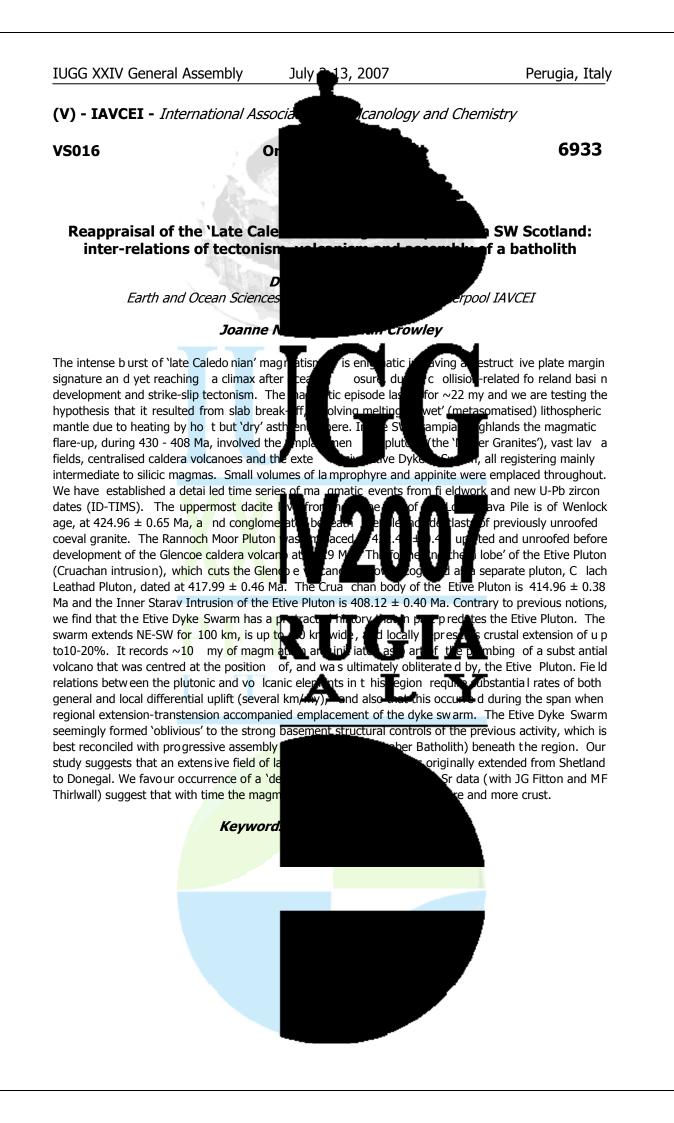
st ropy ly constrained by trace 7094 3Nd/144Nd = 0.5125-Z in ern diate/arc-type magmas solved in the arc-type **ef** (3-7 %) of subducted sediments canics: chondritic patt erns with LRE E 25), LILE and LREE enrichments ences are similar with those of hasized in the Pb isotope ents also for the parental t inct volcanism of GVZ, ion of the Carpathianns in the ALCAPA

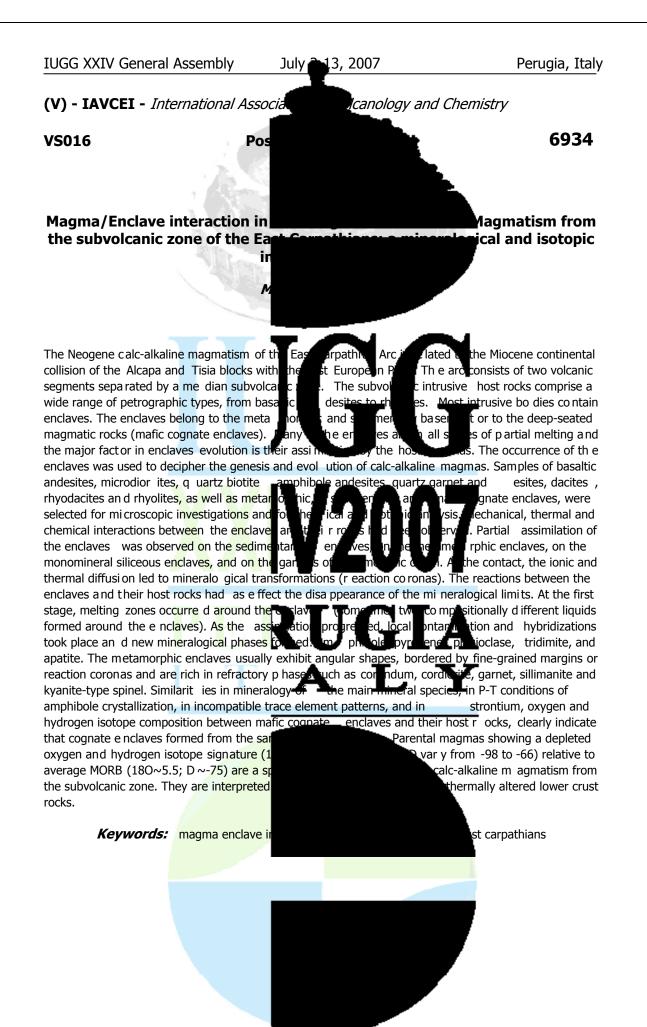
microplate (Pannonian Basin / back-arc site) related to the subducti on roll-back. Magma genesis could subduction components) asteno spheric mantle belo w

of

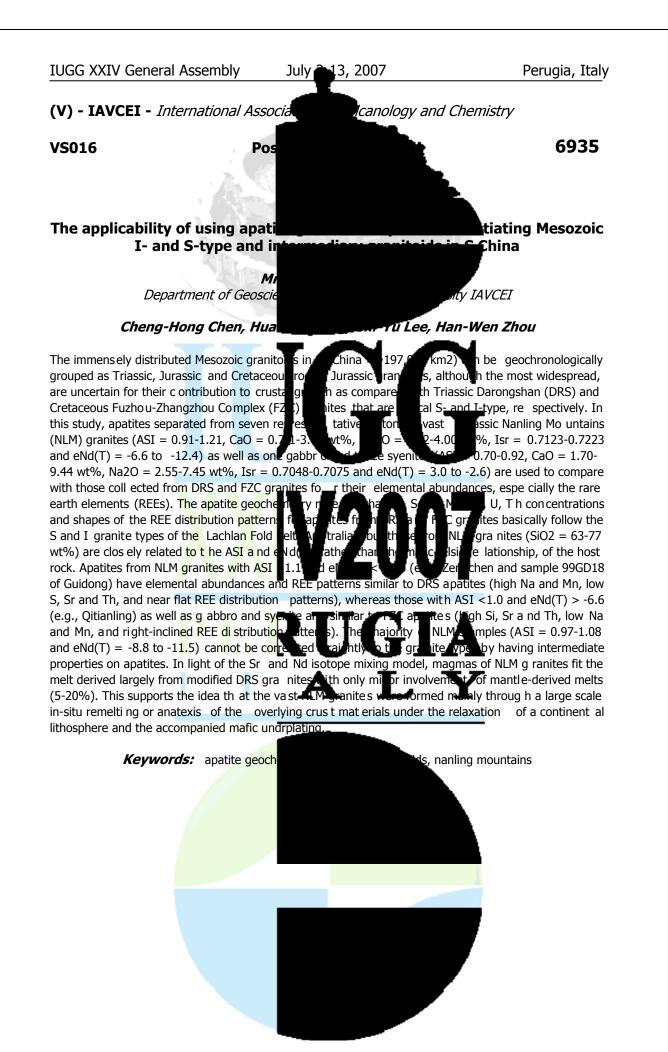
uction front. The onset LCAPA) in what means er locations (overlapped uring the complete sinking lians arc.







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## IUGG XXIV General Assembly July 13, 2007

Perugia, Italy

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**VS017** 

## Symposium Pedagogical and didactical met geopark concepts in demonstra

Convener : Dr. Ulrike Martin

Methods in ed ucation to demonstrate the living environments should be a to ol to le there is a good connection between the

science education. Through the creation of a world p features, label led UNESCO Geopark, UNE CO environment and enhancing sustainable eco hoi tool for a better understanding of the ge increasing public awareness for a balanced multi level programmes, e.g. to demonstrate the feeling of responsibility for our environ her transfer scienti fic information about volcani pr popolation. Contributors are encouraged to consider

aspects of the complexity understanding volcanoes and their hazards canology and Chemistry

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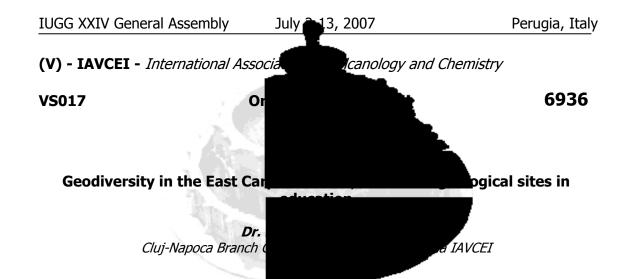
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cation and

ween animal, plant and nonwir onment. In this point of view veveloped under the UNESCO and e arth f natural s with significant geological als of nserving a healthy parks an designed to become a se use of the Earth's crust, by na nkind and t he earth. Especially featu e necessary to transfer thods are n ecessary to dhood lcar azards to the gener al inductions are invited on all of pedagogical and di dactical methods to transfer knowledge about



Recent efforts to record , upgrade and e Carpathians have produced significant result many sites or geotopes possessing high sci rock types (e.g. garnet beari ng and esites, deposits, mud stones); fossils (e.g. foramini (e.g. salt domes, salt mines, cross-bedding lakes (salty lake, volcanic lake); landscapes (e.g.

mofettes, CO2, H2S exhalations, caverns with intense native sulfur and alums depositions on walls). Preliminary evaluation has identified severa on their unique characteristics. Furt her del areas as poten tial Geoparks. The use of v East Carpathian Arc in education is materia joint project between schools from several science school teachers is needed in order to determine how best the geological sites could be used in

earth science education. The expected impact to be achieved is to raise students interest in the study phenomena by understanding the mechan the objectives to be studied and to enrich t are expected t o change their attitude t owa students from different countries is viewed. The decisions on selecting the objectives to be studied, to enrich existing biography through fieldwork, to

(photos, shor t films), t o dis seminate the teachers and upo n scho ols co uld mate geoconservation organized in the following

on uments in the East ng geologi cal sites h as iden tified ogical sites comprise of: different basalts. piroclastic deposits, flish eleton edimentary structures folds ust and normal faults); bes, keys, cliffs, caves);

and a wide range of postvolcanic phenome na (e.g. mineral and thermal water springs , dry and wet



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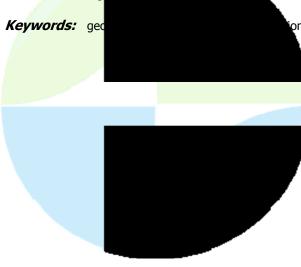
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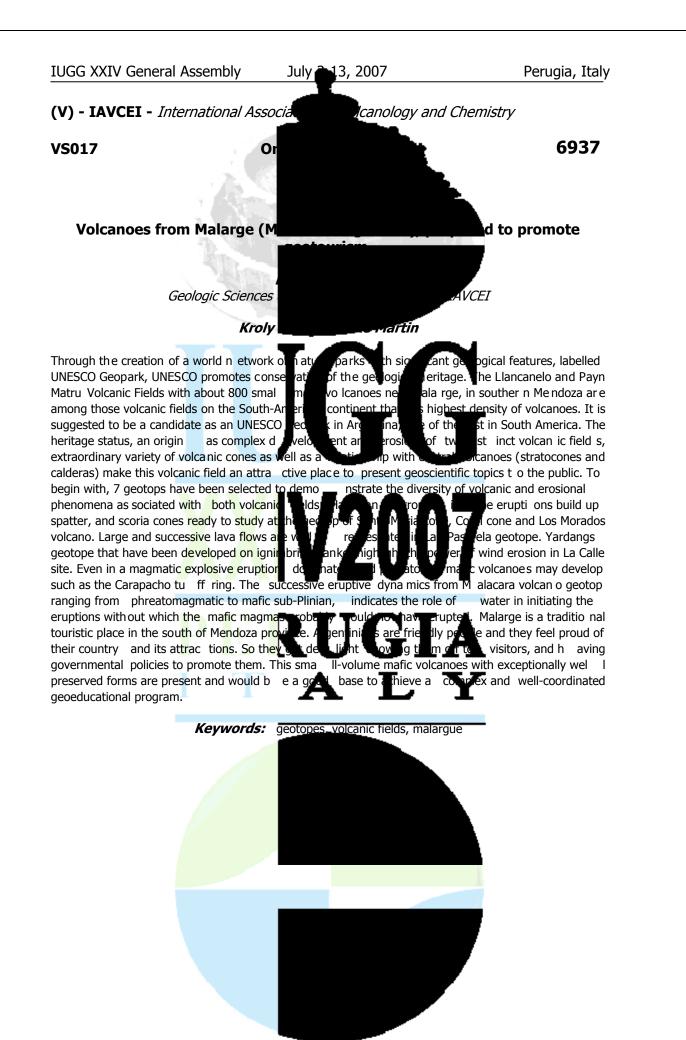
alues. These

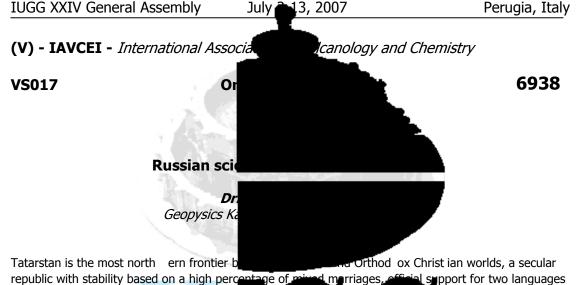
al heritage sites based but in order to promote c phenomena from the bmenius School Project a etween geoscientists and

a ke the maware of the value of of E n Science. Yet, students ce. By selecting similar me tow ds scie phenomena in the participating countries, the conver gent/divergent elements of the sam e phenomena in different vo lcanic environments can be followed by the students. Tr and of kno wledge between participa tudents a encouraged to make

elaborate texts oriented towards the geological explanation of processes, to release multimedia products e lessons. The pro jects impact upo n c ourses on geodiversity an d







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and cultural tolerance. Kazan, the capital largest port on Volga river, re ferred to as "t Tatarstan has four million inhabitants. As over 80 million people. The beautiful city Ka the most remarkable and ancient towns at t monuments of Kazan, together with its bea centre as a Ru ssian Treasure. The history of the

N.N.Zinin ob tained aniline (1842), K.K. Klauss propounded the theory of composition of during the round-the-world expedition of 1 further succes sful development of scier Tatarstan, Kazan is offered the national pro the Internet of Technologies "GeoNa"

enthusiasm, p ride, grandeur), which including: or riginal designs building "GeoNa" - "Lobachevsky surface", 59 floors, height 2 15 m (with a spike 302 m), the general a rea in 148,000 sq. meters, a modern complex of conference halls (up to Computer center, 3D Plan etarium, training a cognitive system "Spheres of Know ledge", botanical and landscape oas es, business-hotel, where will be hosted conferences, the congresses, fundamental scientific researches, educational arter creation level. Expositions of museum -geopark "Travel o c sensual demonstration of fundamental processes inside the Earth: movement of continents and tectonic

flashes and the polar lights, destructive ea of the Earth, t sunami a nd to rnado, fo rm deposits, paleontology, occurrence of the f (http://www.geona.ksu.ru) will enable sci advanced achievements of a science, infor with foreign colleagues in sphere of the hi centers.

itv of million p eople and the ver o of five seas" ortation hub, 05yr<u>), situat</u>e st rive g tifi c a nd : /ersitv

ut the side of Ireland or Portugal, an serves a regio n encompassing 0 km East from Mos cow, is one of ga, Russia. The ancient sphere, made it historic the history of Russian

science, social thought and culture: N.I. Lobachevsky - the founder of non-Euclidean geometry (1826);



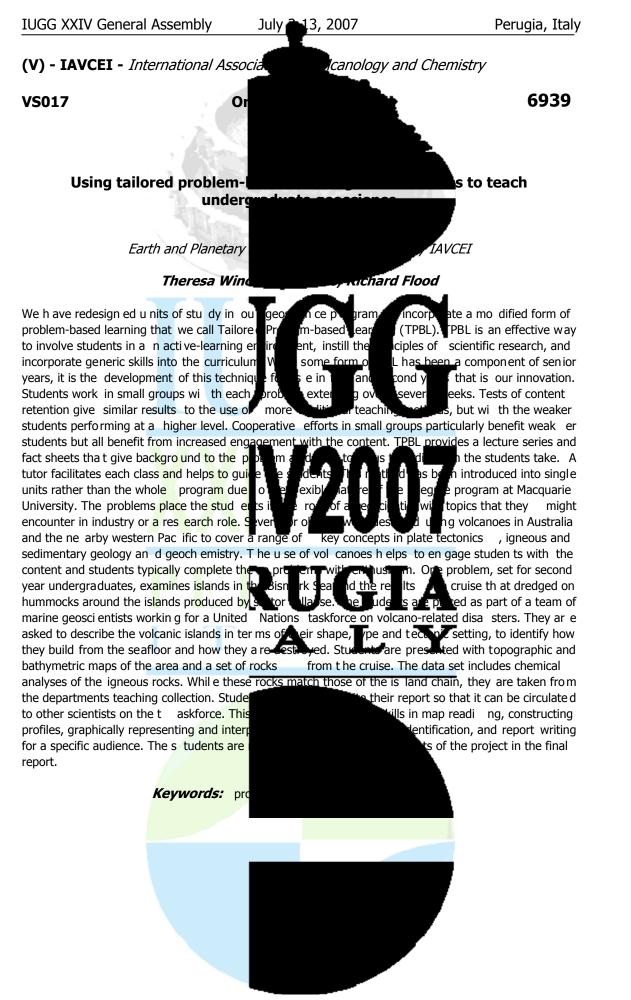
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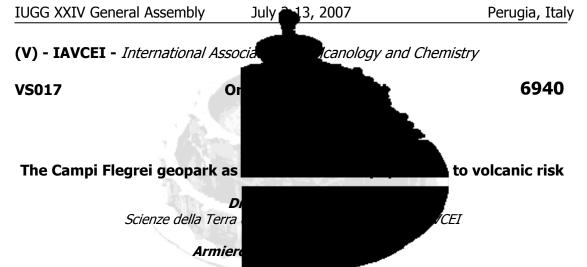
discoverer of ruthenium (1844), A.M. Butler ov scovered the Antarctic d M.P.Lazarev. For the deration, the Republic e Science, Education and developed wisdom 's

t he<u>I</u>nternet of Technologies, ve r eum of natural sciences, vate ment (5 million liters ), ourist action a world organ izational Fart h″ wil be presented of 3D and plates, rotation of a liquid ter restrial core and generation of a magnetic field, magnetic storms, solar tic eruptions of volcano, global warming

> inerals, generation of oil and gas Center "GeoNa" luman. an univer sities to join to scientific co mmunications ects with world scientific

Keywords: scientific education center





The Campi Fle grei area combines fascinati pageolog customs and t he outstanding ancient Greek and volcanic products are visible from a small to geological sites of particular importance in educational value, most of which are part d archaeological, ecological, historical or cultur highest volcanic risk area of t he world, but that share the territory. All the previously dooted for

people to come in touch with the volcanoes. Main aim of this project is to promote the knowledge of the weight of the active volcanoes in the territ first step to re duce the volcanic risk is to cultural revitalization as a whole. The re consequence of both being aware of the v not worth only the violence of the present itinerary from Cuma to Solfatara, passing through Averno and Monte Nuovo Volcanoes. Starting from a

awakening of younger generations to the in

th its unique ilizati lan. nal scale Thi their scient ological he reove gre ī ardly

ural long history, individual his territo ry all kind o f Wit h a comprises a good number of ualit y, rarity, aesthetic appeal or e, b <u>ut their</u> intere st may also b e in ic field is one of the habitants of the towns a very suitable area for

a volcanological geopark, in which geolog ical and archaeological records can concur to help young

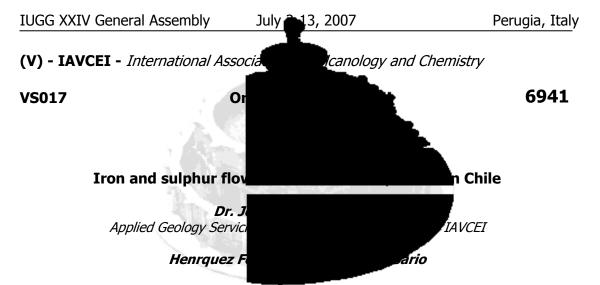
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is well known that the ipate in the territorys com es as a direct t such a valuable area is ork we here propose an

new detailed geological survey of the area, whichmad e us also possible to identify many sites with a strong pedagogical and educational value, eosites of Cuma, Averno, Monte Nuovo and Solfatara. All these area s ar lge, their i nterest is als o her archaeological and historical, as testified by do lected and investigated archaeological and historical, as testified by the numerous accient doments collected and investigated during the research. Each geosite satisfied the criteria adopted for the most recent italian geosites, i.e. n n mej representativity, scientific interest, rarity, landscan value, edu **T**cessibility, preservation ationa I valu Cebparks. Whin each g eosite at least and vulnerability, which correspond to the criteria Europea two possible itineraries with many stops were planned. At each st op posters illustrating both the geological and archeological valuables can be <u>set\_for</u> self-guided to urs. We hypo thesize a Western Campi Flegrei Geopark linking the four geo from one to another ar e guaranteed by present roads and, where possible, by ro ient times for military aims. This geopark could represent a st ep forward fo e geological h eritage and the

Keywords: ca



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Although silicate lavas formed by silicate minerals are activity, other magmatic phe nomena also g type. In addition to the geotectonic and petr Andes of n orthern Chile, have allow the ge sulphur, as iron oxide lava flows at El Laco y flows at Lastar ria volcano (5.700 m asl, 25 on the flanks of El Laco, a 30 km2, Pliocene El Laco volcano is similar to other volcanic Comple

lavas that are partially hydrothermally documented in Chile in Creta ceous formations hosting the iron ore belt of the Coastal C ordillera to the south of the Atacama Desert. With similar I in Sweden (Kirunavaara) USA (Missouri), b preserved, used as reference to explain an typical central Andean volcano with consta 50 to 350 m long sulphur flows as well a activity of this volcano. Their origin was c

apparently more common in andesitic cent pristine featur es showed by the iron oxide a respectively, make them world class example unique site of esthetic, touris t and scientific value. Detailed mapping is being carried out in order to

features to pahoehoe basaltic lavas. They look fresh and resently formed, being one of the last eff usive aused by removilisati on of prec ipitated sulphur deposits, Etio ro

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most comm terial produced by volcanic carbon atite and sulphur as of oxic hiperarid conditions of the central of volcanic products of iron and m asl, 2348S/6730W) and sulphu r xid e d osits.

altered. T he occu rrence of iron

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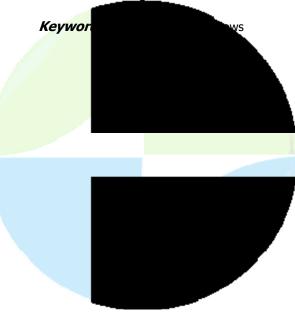


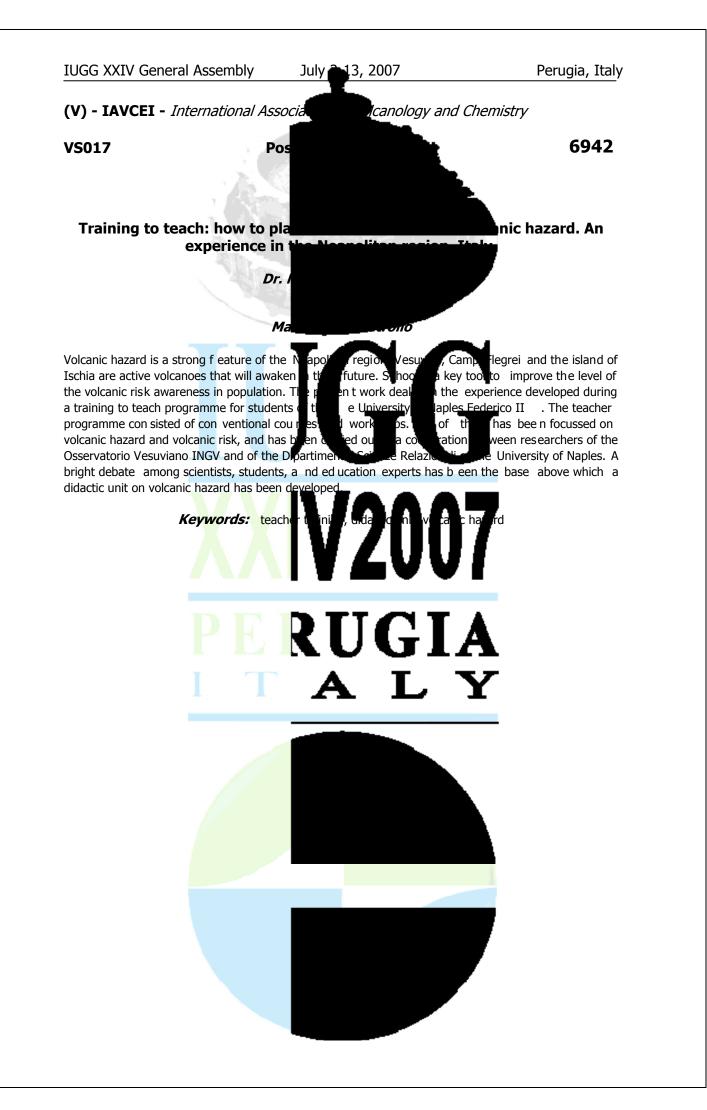
sits have been mapped ept for these iron units, hly compound of silicate oxide magmas is also d of Precambrian ages he yo ungest and best ther hand, Lastarria is a shows unique large scale

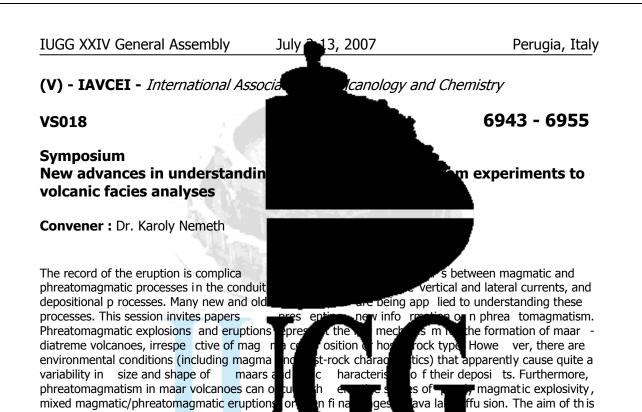
notably similar surficial

n other volcano provinces. The nd Lastarria volcanoes, ng and preservation as from

catalogue the main features, morphologic evolution Tetending them to b on of their rigin source е es of this kind in different n State declared as natural geology monuments by the occurrer volcanic environments are expected to be discovered. This is a c ontribution of F ondecyt Project No. 1070428.









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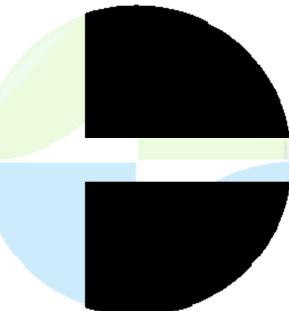
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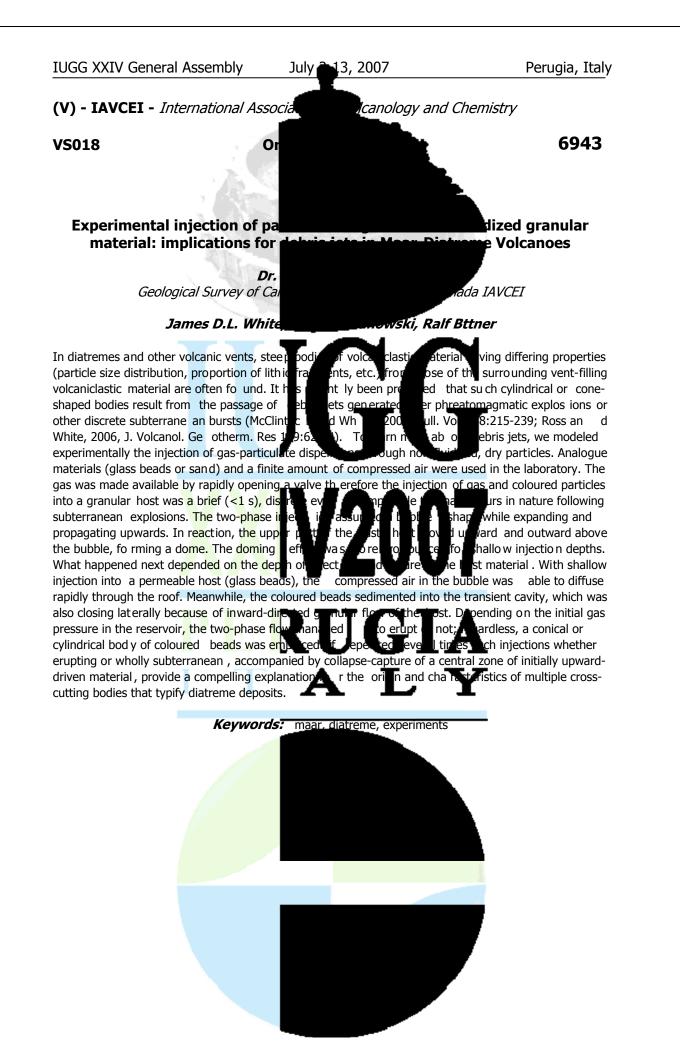
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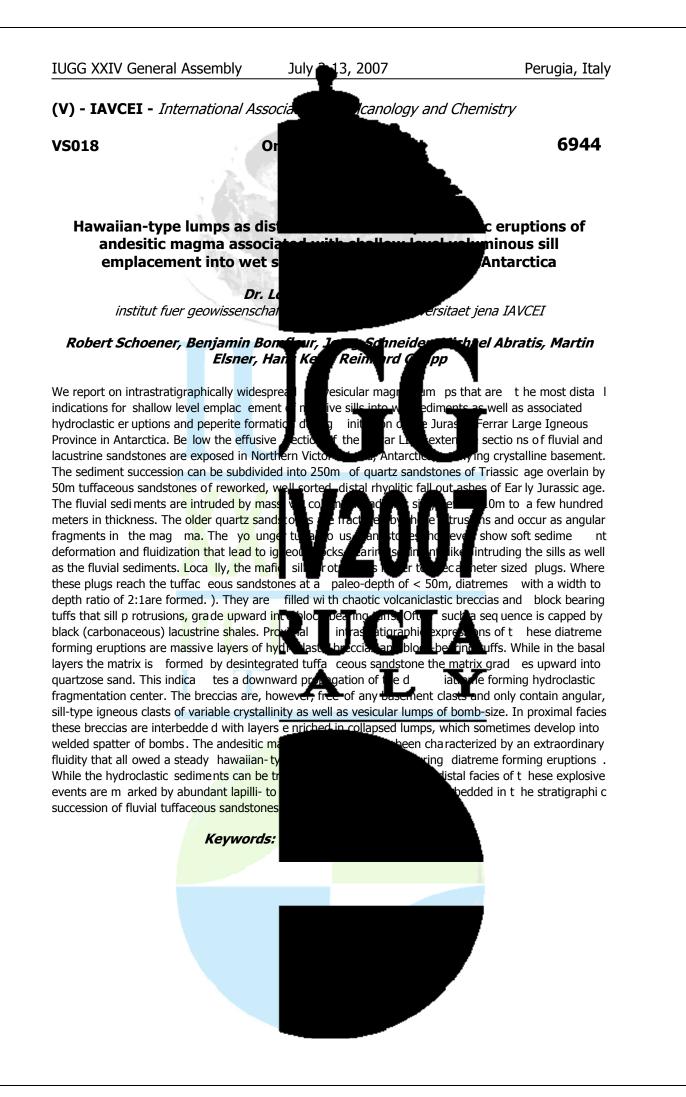
uch processes lead to

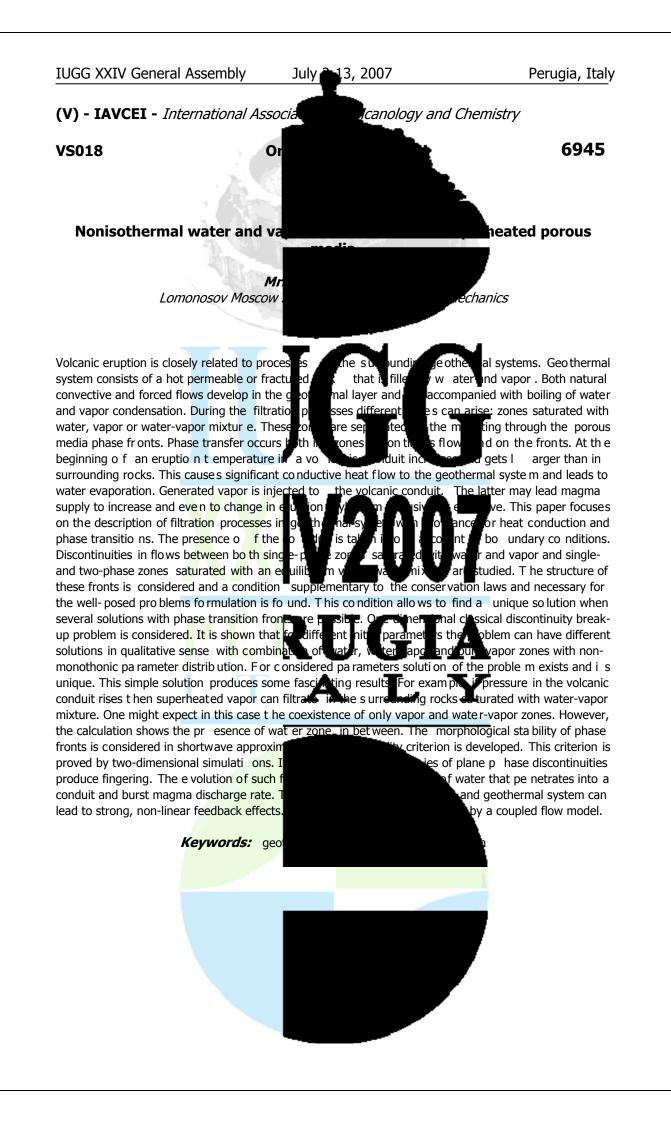
session is the discussion of physical process

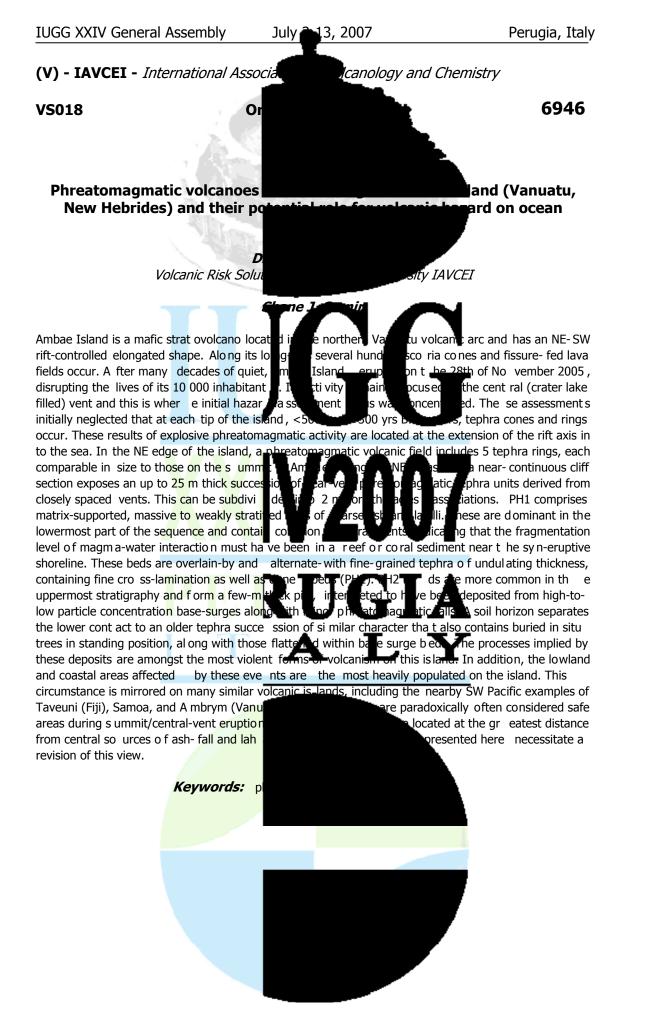
eruption variability, and the environmental factors the



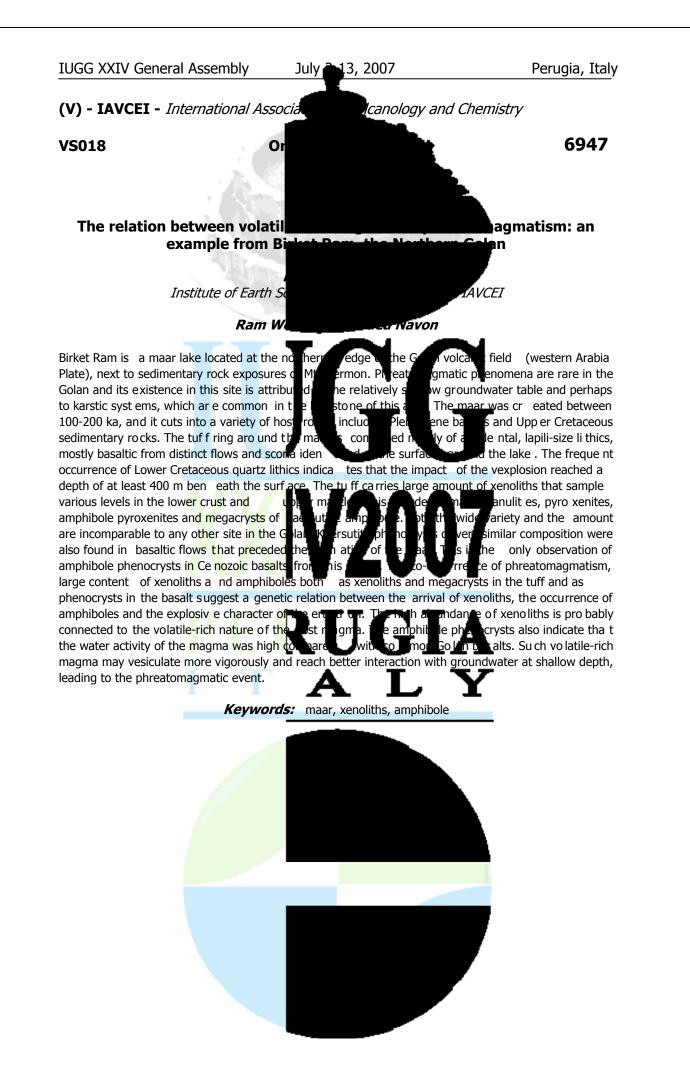




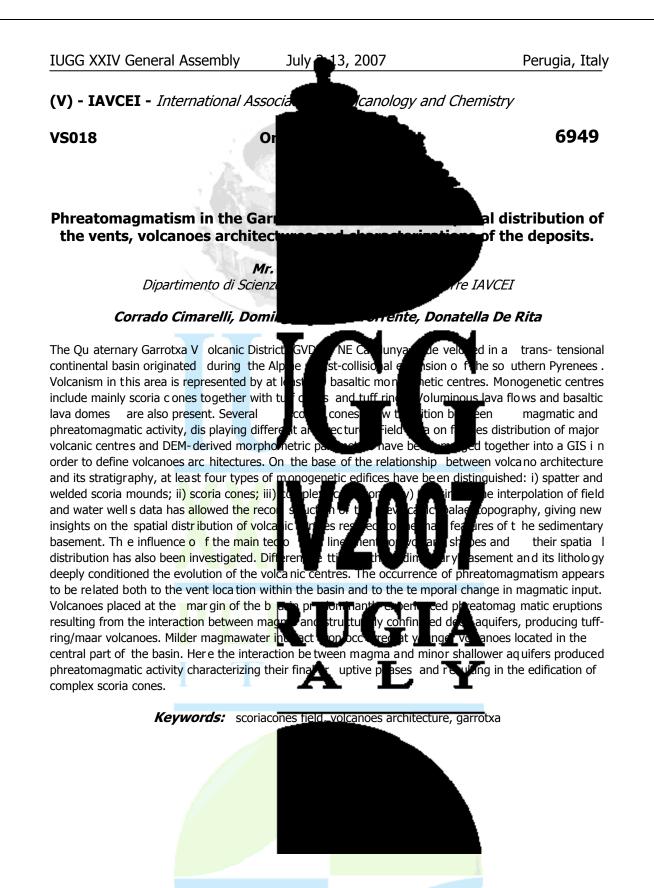


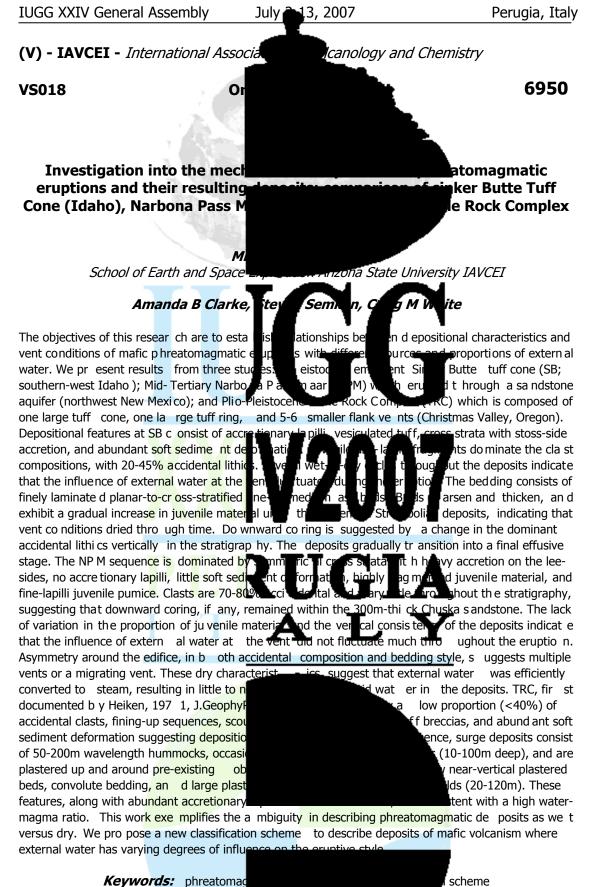


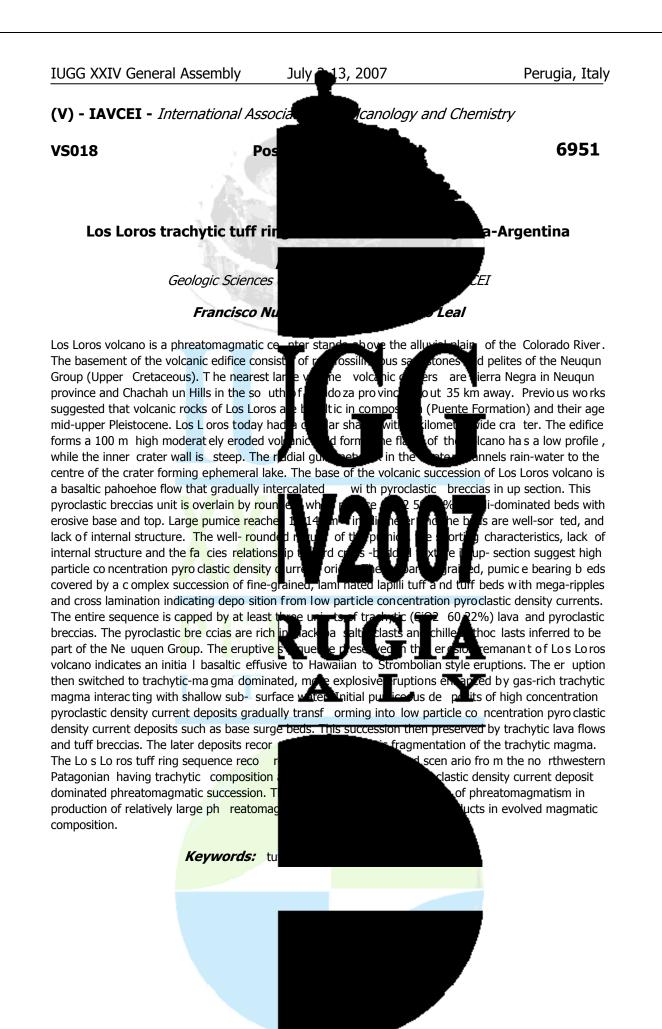
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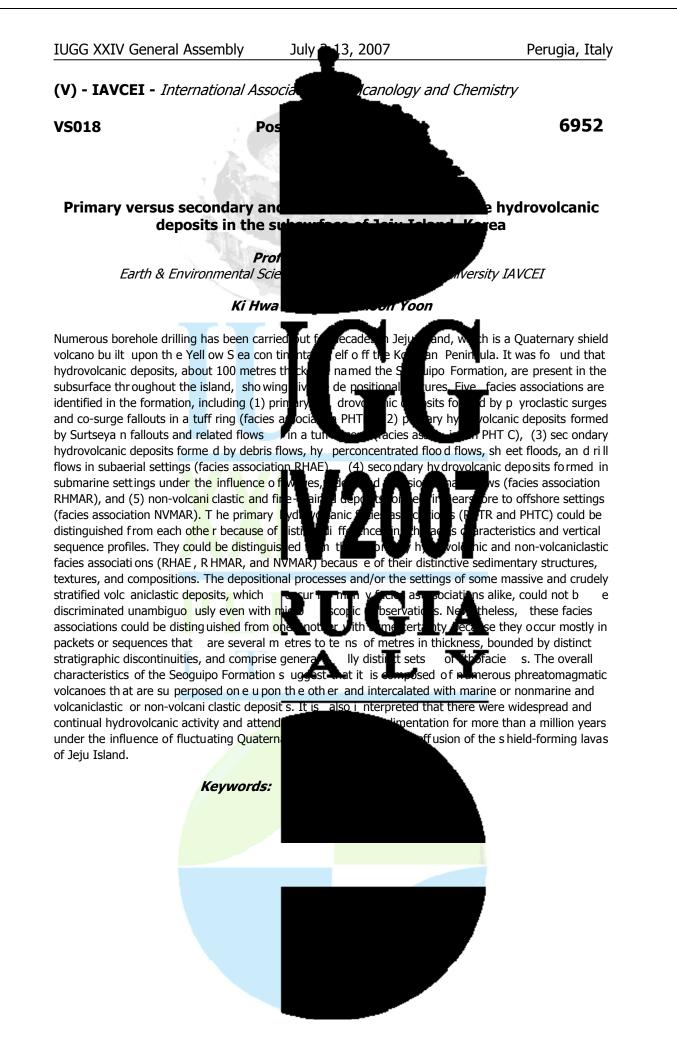


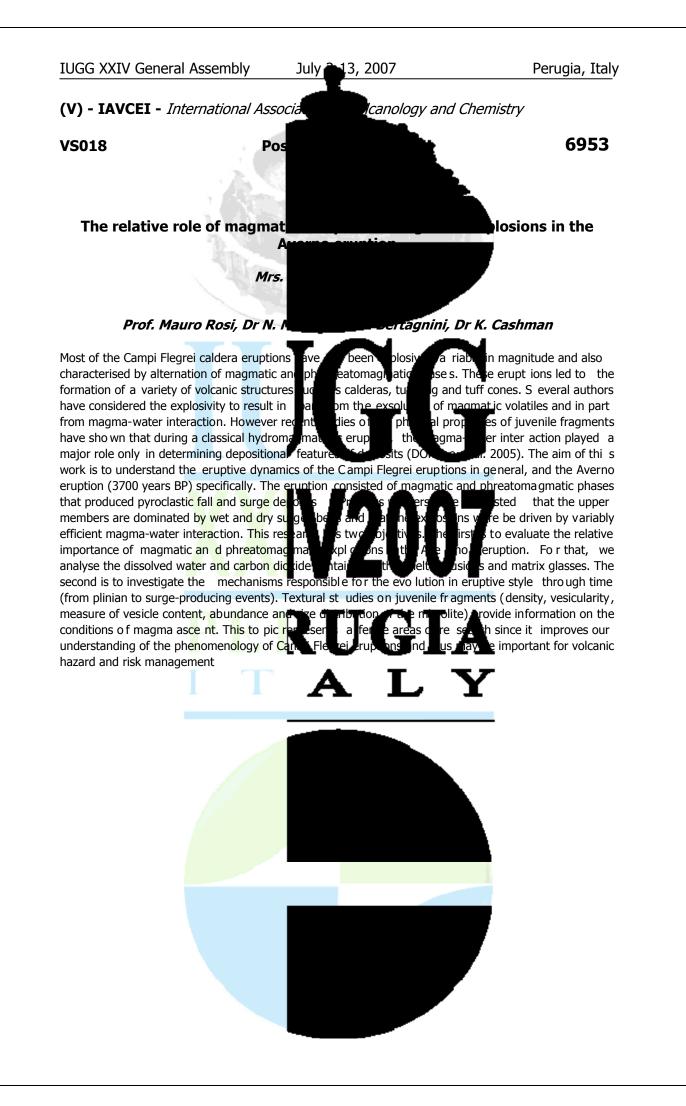


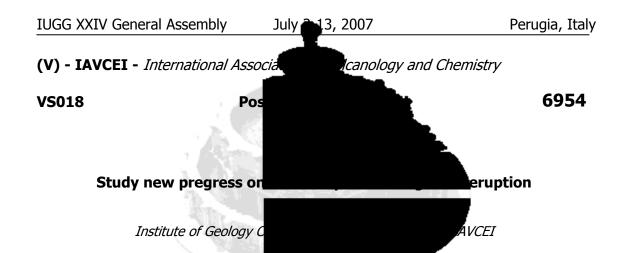








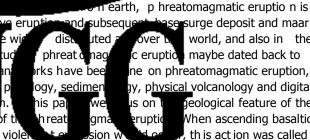




As an es sential and import ant type of characterized by groundwater-related explosive eruption lakes. Base su rge deposit an d maar lakes a

Northeast Chin a and the southern Chin a. 1921, and in the following over 80 years, r using various of methods of volcanic geology modeling, to discuss its origin and mechanis base surge deposit and dynamic mechanism magma meets with ground (surface) water, viole

the base surg e, it has long been treated as sedi distributed around maar, different from the surge caused by phre atomagmatic er upti large-scale and low-angle cr oss-bedding, s lapilli. In order to explain the dynamic med a simple model in this paper in light of the the larger the radius of maar, the larger the exp



sessurge deposit and maar world, and also in the c eruption maybe dated back to ne on phreatomagmatic eruption, y, p<u>hysical vol</u>canology and digital jeological feature of the When ascending basaltic , th is act ion was called

phreatomagmatic eruption. The main product of this kind of eruption are maars and base surge. As to

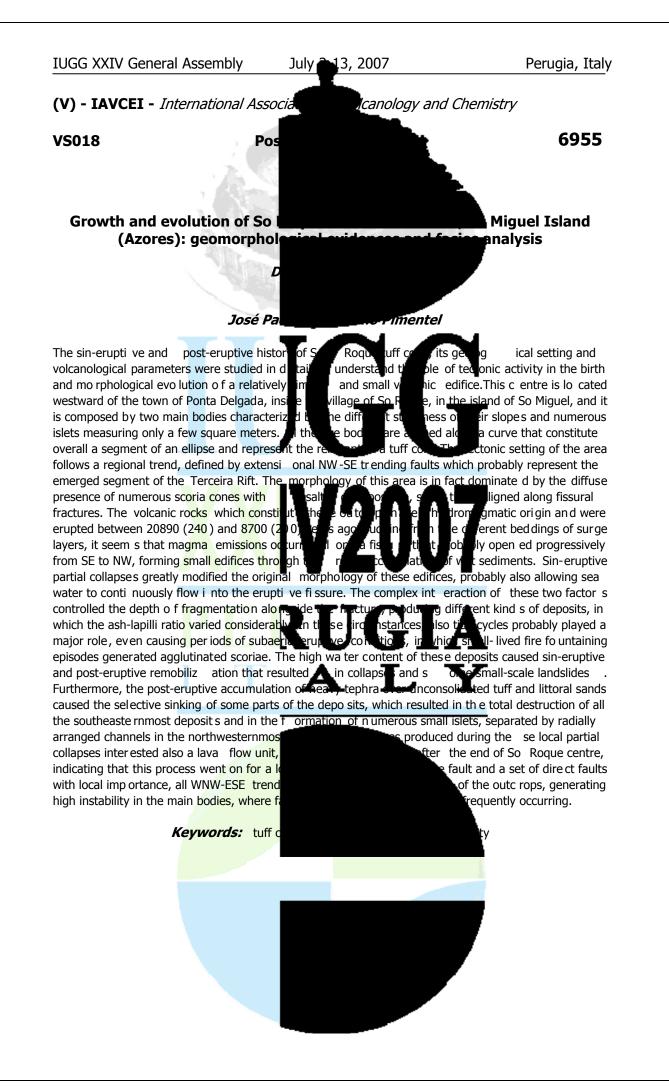


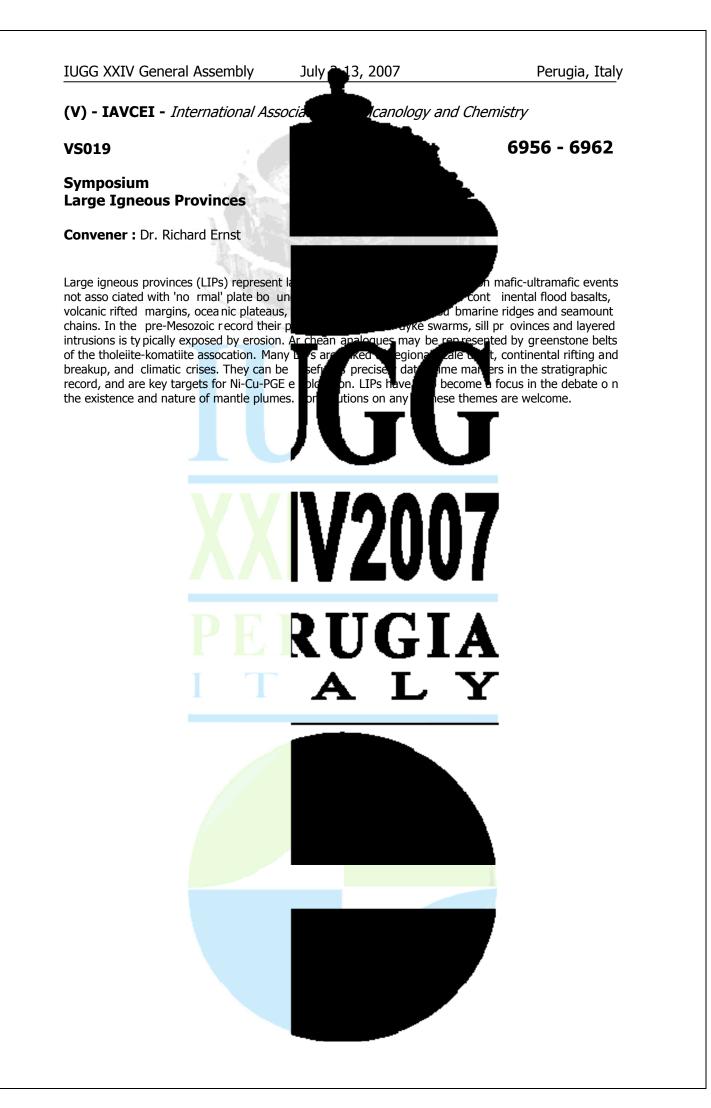
base surge is al phenomena of base ail field work, such a s istal facies accretionary thoroughly, we propose can be drawn as follows:

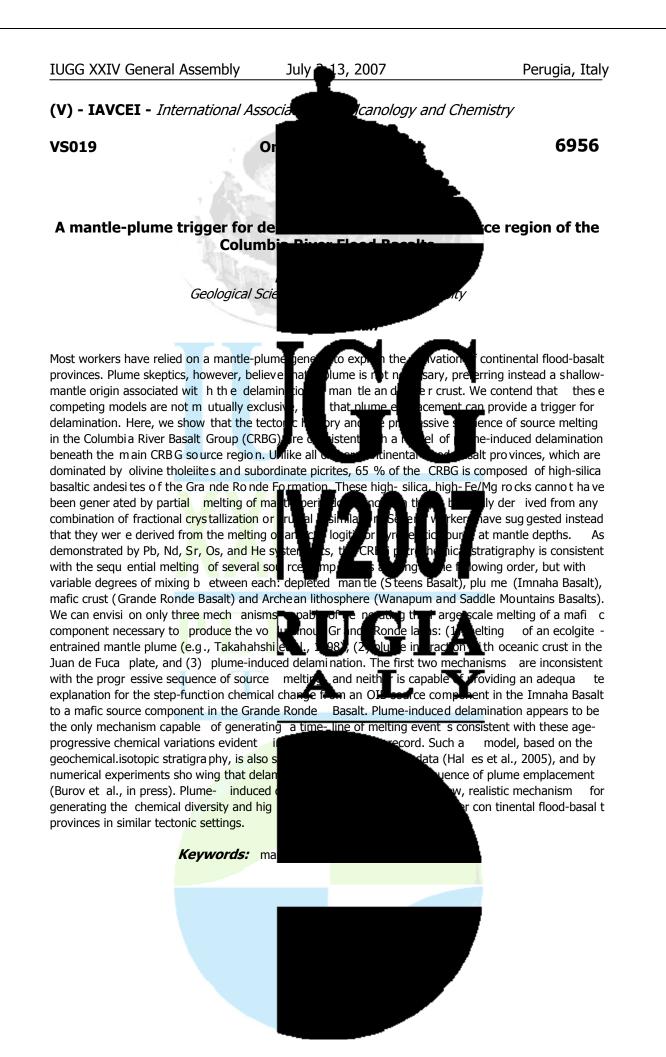
provided that the radius of m aar and depth of explosive point are limited, then the larger the area of contact surface between magma and group xplosive energy will be; if the explosive energy and area of explosive point the r us of ma ar, the greater the **l**arge d radius of maar the depth of explosive point can be infer he , the depth of explosive point can be inferred; when the explosive one gy and radius of maar are qualified, the depth of explosive point decreases with increasing of the area of contact surface between sive are s, undoub magma and groundwater. As for the maximum st dly it should occur on the surface of the overlying formation.

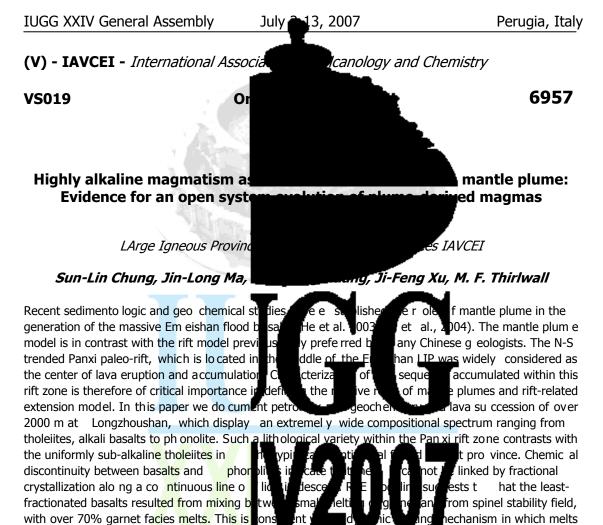










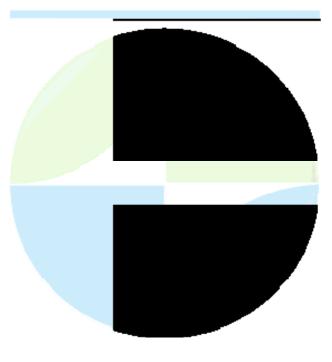


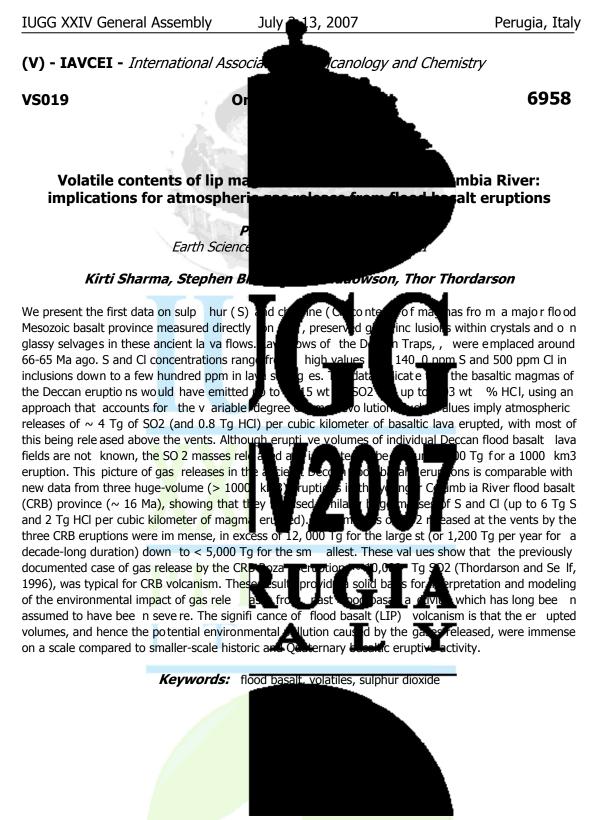
resulted from poolin g of melts from a meltin g column. Phonolite series may have formed via partial melting of underplated magmas. Crystal fractionation is coupled with crustal assimilation as indicated by

the correlation between S iO2 and radiogen indicates a stable, shallow-level magma chapter be related to the combined effect of lithosph a conserve of the abunda asing marma samply. The study menotive y.

the abundant evolved lava s supply. The volcanism could

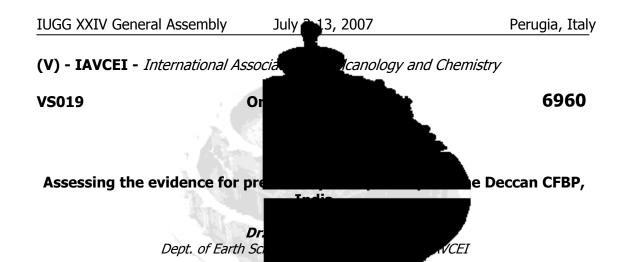
Keywords: basalt phonolit, plume riftin eraction, en visionalip











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The timing an d duration of surface uplift constraints on mantle co nvection pro cesse regional uplift and associated erosion prior to between ero sion, sedimentation, tectonics uplift has o ccurred, and fo r constraining the describe the emplacement of LIPs without explaining the o bservations of surface up absence of regional uplift, ero sional, or sediment

models (e.g. Sheth, 2005). White et al. (1987) and White and McKenzie (1989) argue that rifting above an incubating plume head plume triggered the rapid eruption of the Deccan volcanic province (DVP). If the models of Griffiths and Campbell (1991 then the maxi mum transient uplift preced Kutch - Cambay - western Narmada-Son ri to the east and south. To test these mode at key localities around the periphery of th

when conside red together, record the environments. The DVP succession youngs southward, with the earliest lavas preserved in the Narmada Rajpipla area (c. 22N), and the youngest invoking a southward migrating locus of vole plume head. In the no rth, early lavas lie between the tilted sedimentary succession and rerlying lava erosion having occurred in the very earliest phases

west Indian coast (i.e. offshore of the west Cauvery Basins on the east coast, are all c sedimentation during the late Campania successions indicate a significant Late Cret entirely consistent with regional uplift effect

nces provide importan t me m el predicts pre-volcanic wever, it is the inter-relationships vid e evidence of whether or not ift. <u>Those</u> mo dels attempting to en hav e difficulties in olumes n. Ne neless, any appar ent to su bstantiate suc h

VOI ffe developm ent of pre-, syn-, and early-stage volcanic

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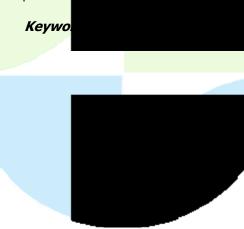
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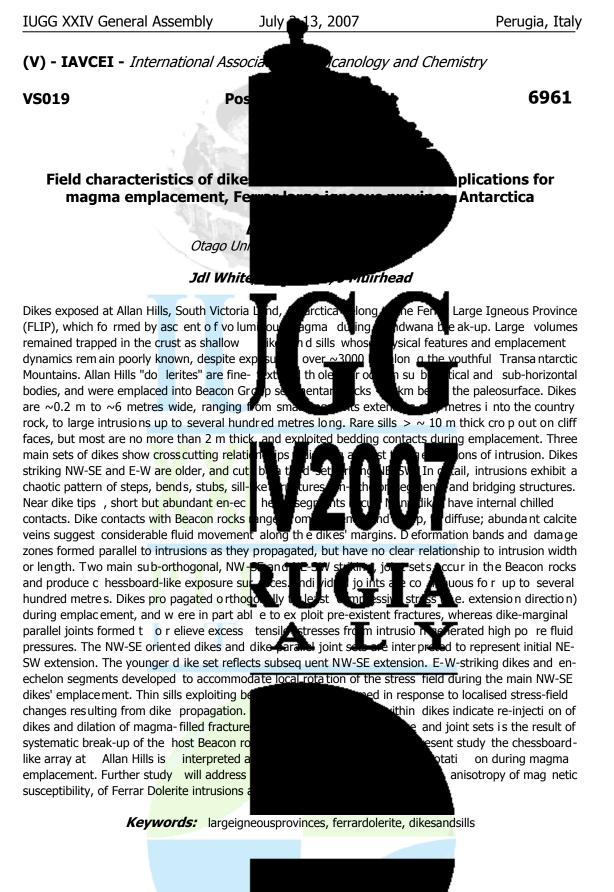
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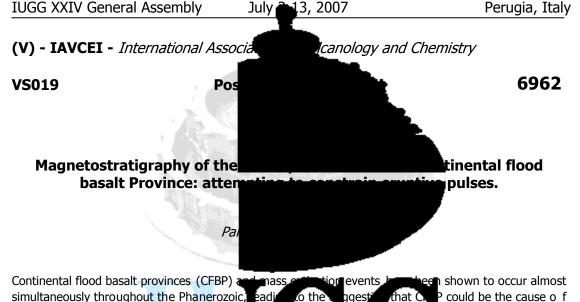
re correct for the DVP , d have occurred in the diminishing uplift further basement-basalt contact nt geological setting and,

ese data support models d over a (proto-Reunion) g pa te Cretaceo us marine brm v ur bin limestone fluvial sandstone s uccessions and, in some instances, thick conglomerate sheets intervene evidence of significant . These pr contrast, finer sediments ef Deco aption. B accumulated in shallow lakes around the per iphery of the most distal lav a fields, consistent with their remoteness from any locus of uplift. Offshore, the sedimentary records of the Kutch Basin of the northvalley), and the Krishna-Godavari and

fiv e-fold increase in the rate of kett et al., 2001). These inental hinterlands which is







the extinctions. The volumes of SO2 erupted enough volumes are injecte d into the str perturbations significant enough to cause million years a nd are us ually composed of individual sheet lobes. If the lava had erupt probable that this would not have been high enoug

is unlikely that the lava was erupting constantly as major red soil -like horizons, thought to represent long periods of quiescence, are found between some majority of the lavas were er upted by a fe years of a CFBPs life is spent dormant. Re Cretaceous-Tertiary mass ext inction; KTB) produced by these very large eruptions. using geo magnetic secular v ariation as a d ent cores from 70 sites along the Naudes Nek traverse of the Karoo CFBP to study the magnetostratigraphy of its southern part. We chose this CFBP because it is thought to have been larger than the Deccan

o the ggest h ought b be re and conve ctinct<u>io ns co</u>u n 1-2 kn a co l af nt ra 'major

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could be the cause o f m ost important factor: if large lint o sulphate aerosols, cl imate <u>CEBP</u> are e mplaced in c. 1 ade up of hundreds of e 1 million years, it is ate change. However, it

lava flows. It\_is therefore pos sible that the ajority of the 1 million th was coeval with the tified erup tive packages

as much smaller than the KTB

d magnetostratigraphy, e have taken o rientated

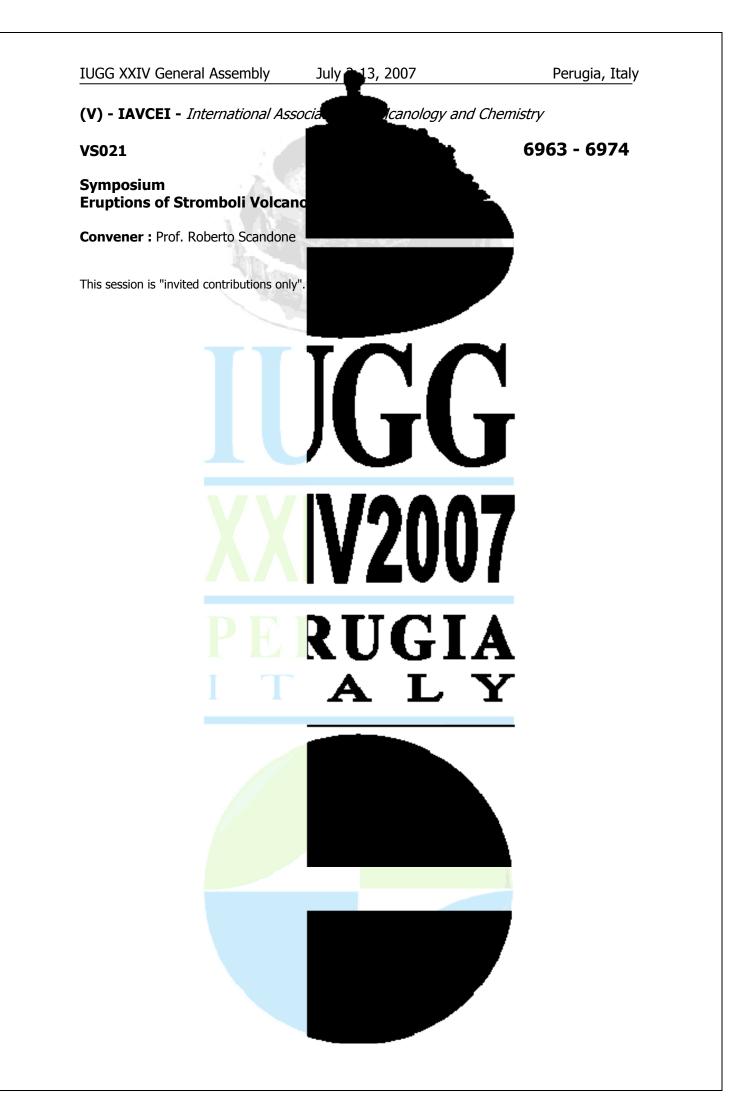
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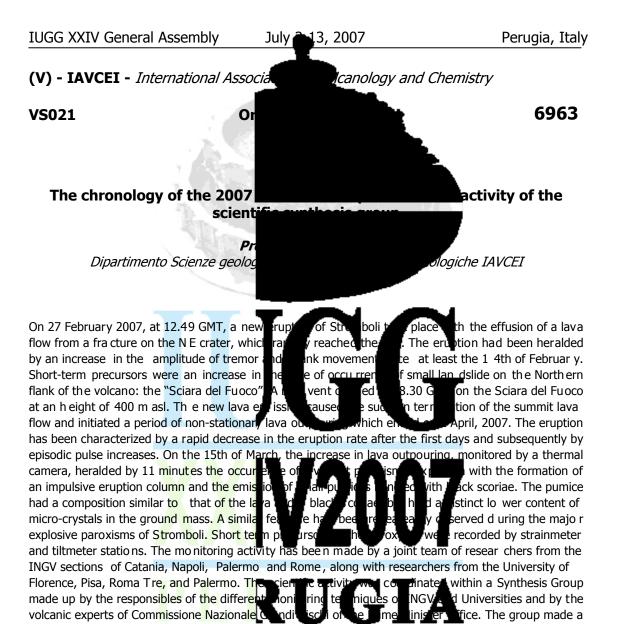
in turn will help us to

Traps, yet the extinction associated with extinction. We suggest that detailed mag nd hen Karoo CFBP will shed light on the reasons h s lle further understand CFBPs in general and their impact on the global climate. We will present the first results of this new study at the meeting.

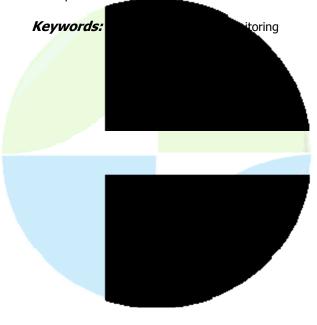
*Keywords:* karoo, paleomagnetism, climate

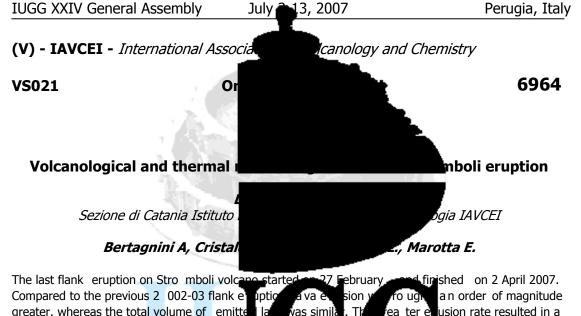






daily evaluation of the state of the volcano and transmitted its recommendations to the Dipartimento di Protezione Civile (DPC). Several prevention measures were addited by DPC, the main of which were the evacuation of the coast zone when strong acceleration of the being a del Fusco slope motion (occurred twice) could led to a dangerous tsunami by flank collapse and two days before the 15 March paroxysm when access was prohibited to the part of the <u>volcano above 400m asl</u>.





Compared to the previous 2 002-03 flank e liptic greater, whereas the total volume of emitted la much shorter duration, and this conf irms the st on the NE flank of the NE-Cr ater and the erett sea. Late on the first day, the three initial flows at about 400 m a.s.l., at the eastern margin of the this vent formed a lava bench, several terms of me

started on 27 February tic cua va e sion v a was simila . Th y supply of th a va formed th ped at ne to Sciara Fud me. Swith which s

rough an order of magnitude rea ter enusion rate resulted in a lea no. An eruptive fissure opened branches that rapidly reached the nt open on the Sciara del Fuoco In a fee ays, a lava flow fed b y uifi comy modified the coastline.

Between 4 and 9 March, important changes occurred in the summit area, w ith significant widening of the crater rim due to crater collapses and ash explosions. The 400-m-vent stopped for a few hours on 9 March, when another vent opened at about 10 march the provide the NE-Crater, almost in

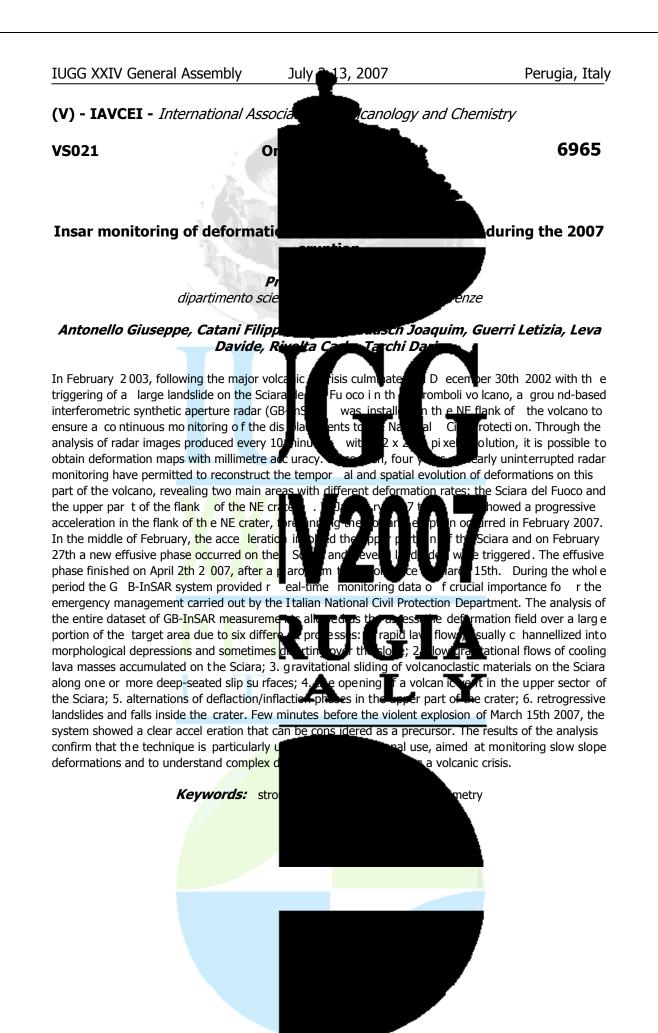
the same position as one of the vents of the 2 02 24 hours, and when it closed, the 400 m vent p 2007, while the effusion from the 400 m vent v was similar to the 5 April 2003 paroxysm, and v number of monitoring systems, including therman

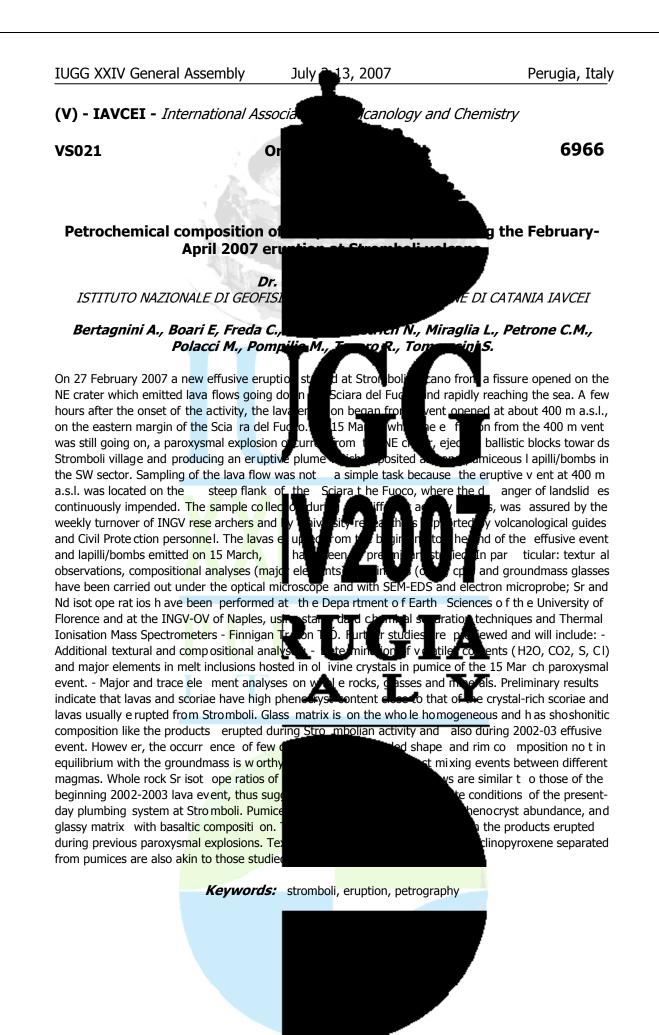
curred in the summit area explosions. The 400-m-vent 0 minus and the portfol 02 of current in the 500ozer operad in unit of a still reagoin, am or recommission, am or

pped for a few hours on 9 e NE-Crater, almost in was active for less than to the sea. On 15 March sion occurred. This event monitoring web cams. A

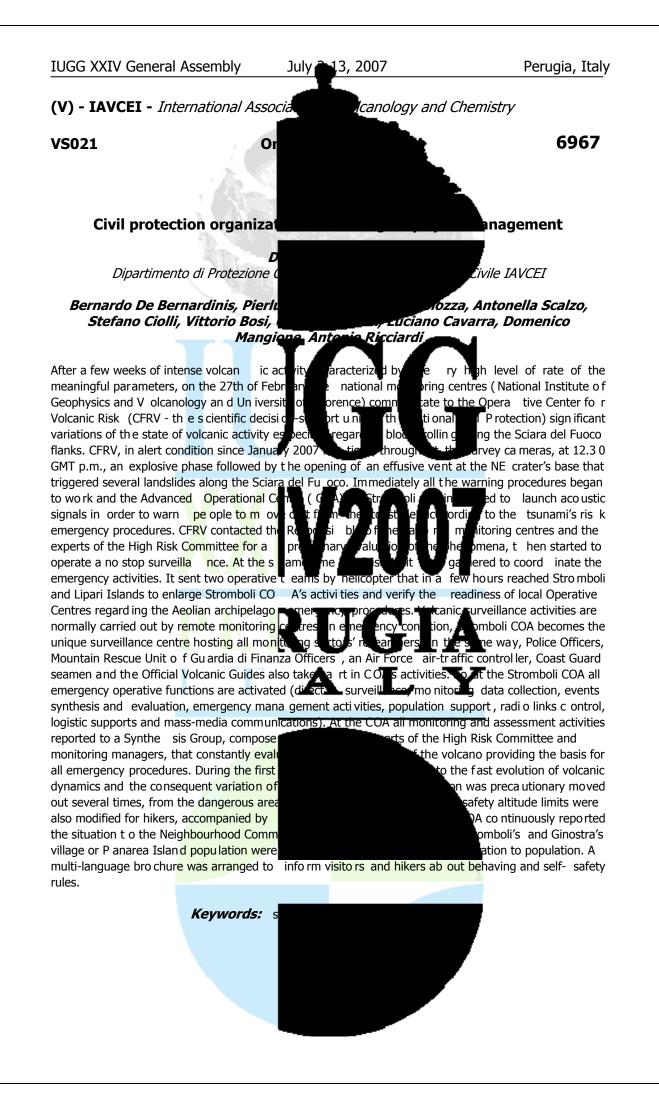
number of monitoring systems, including thermar and visual web cameras, direct field surveys, and daily helicopter flights using a portable thermal camera, enabled characterizing the different phases of the eruption.



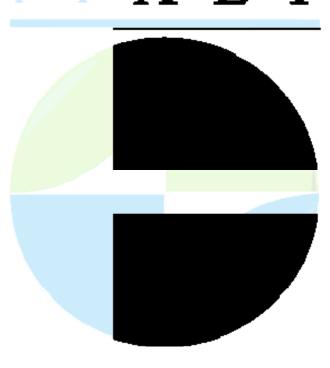


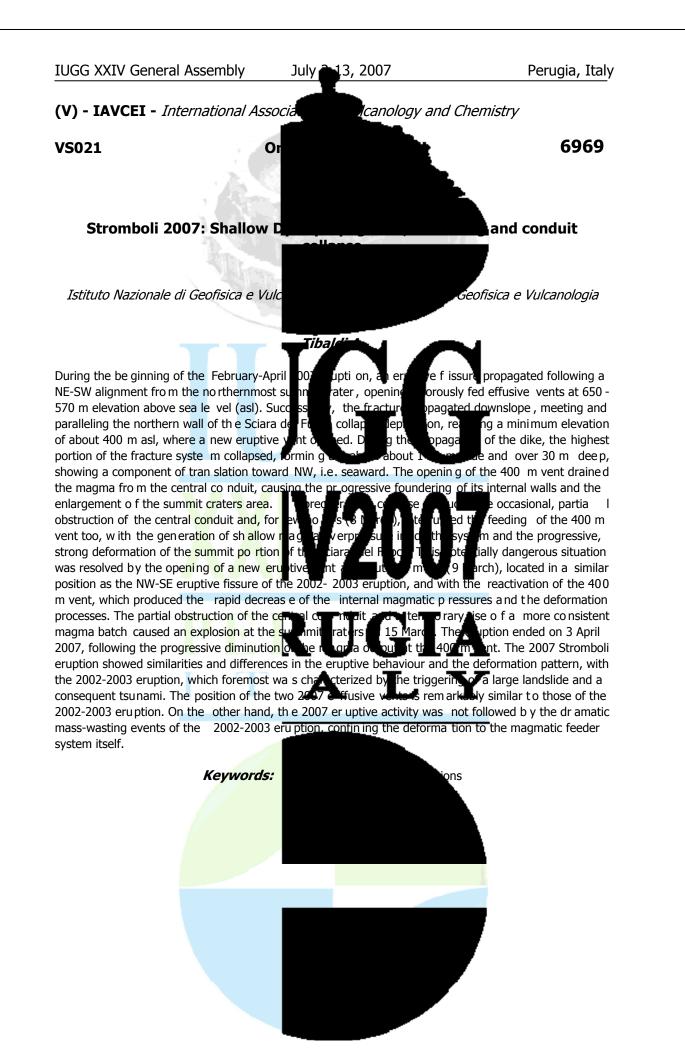


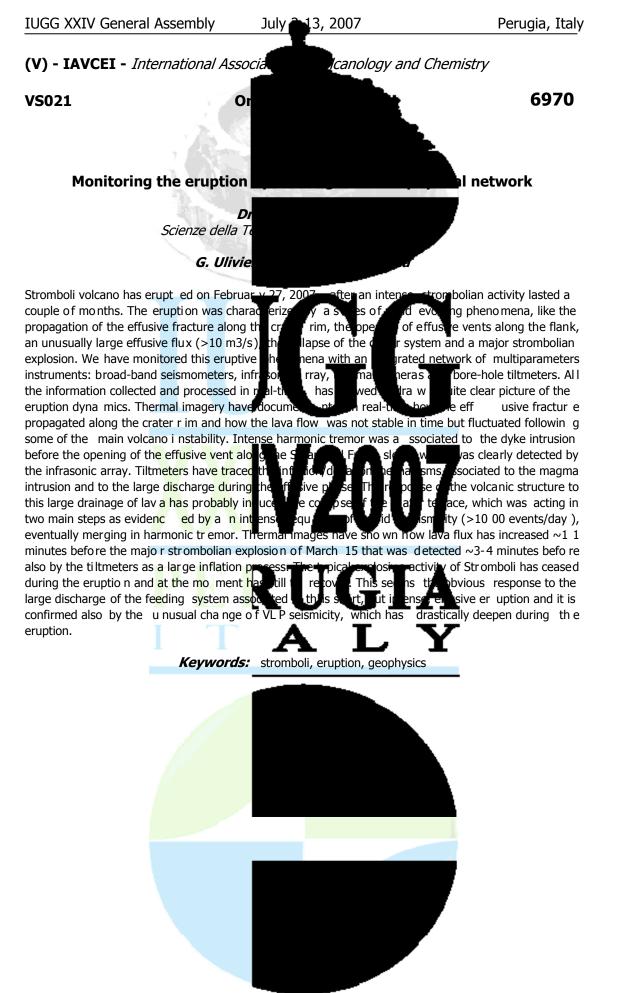
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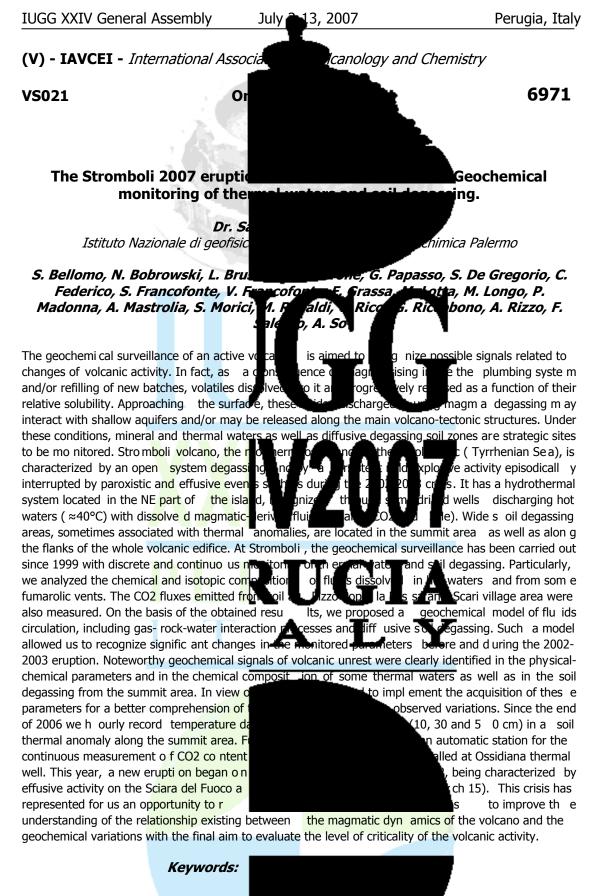


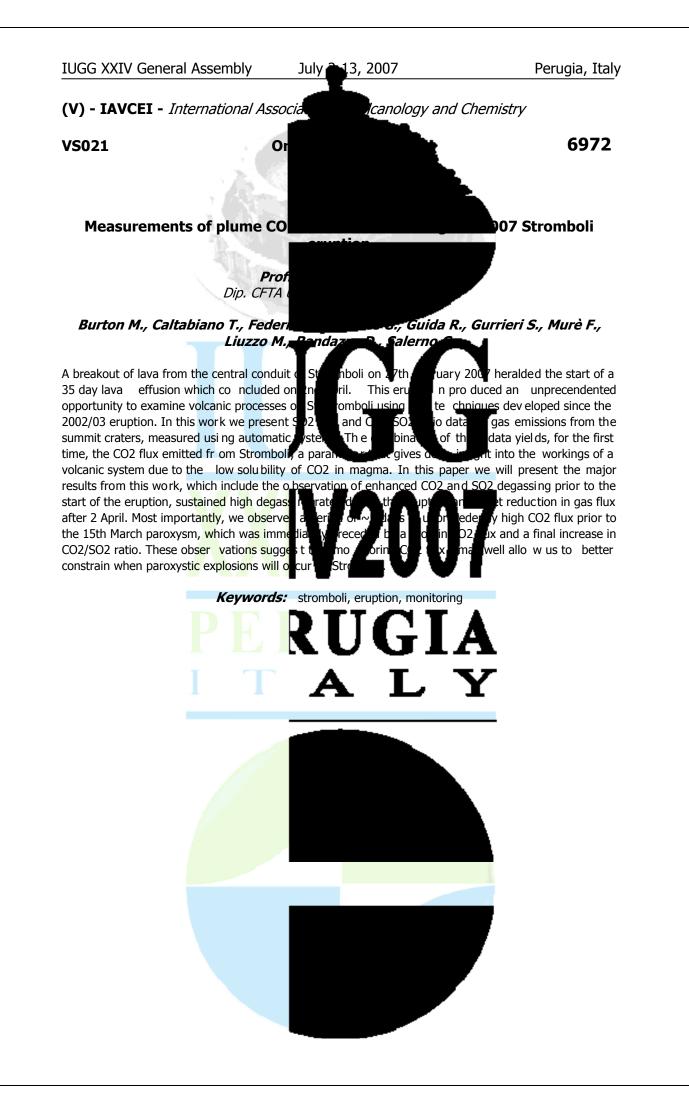


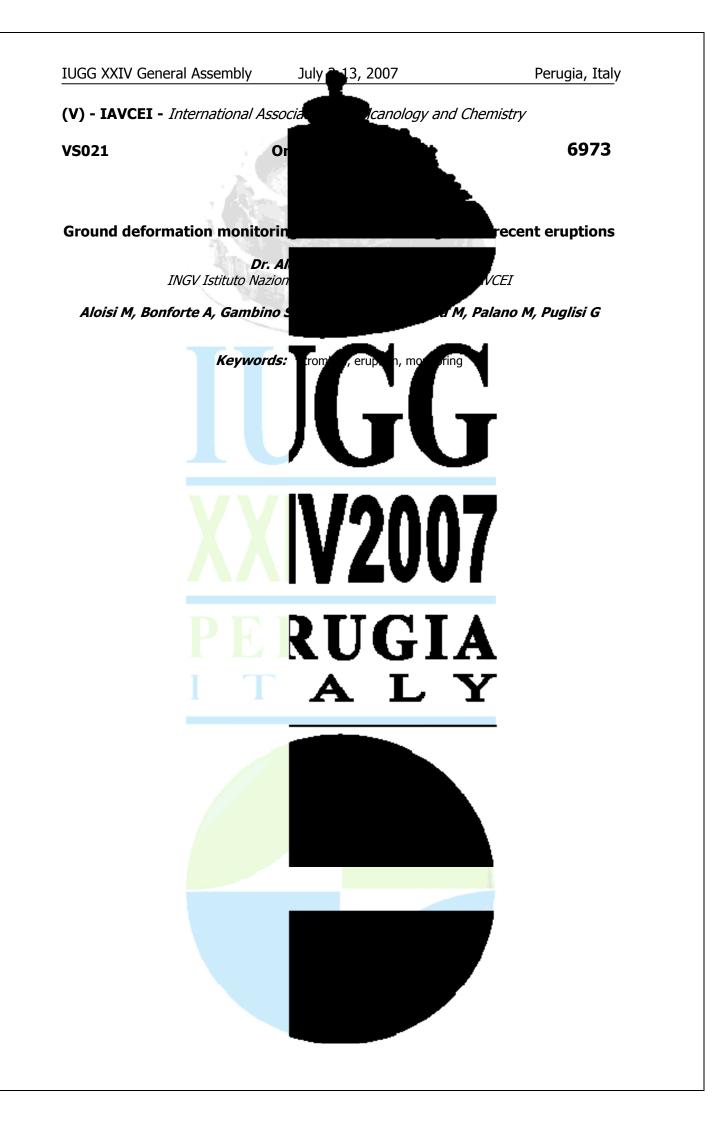


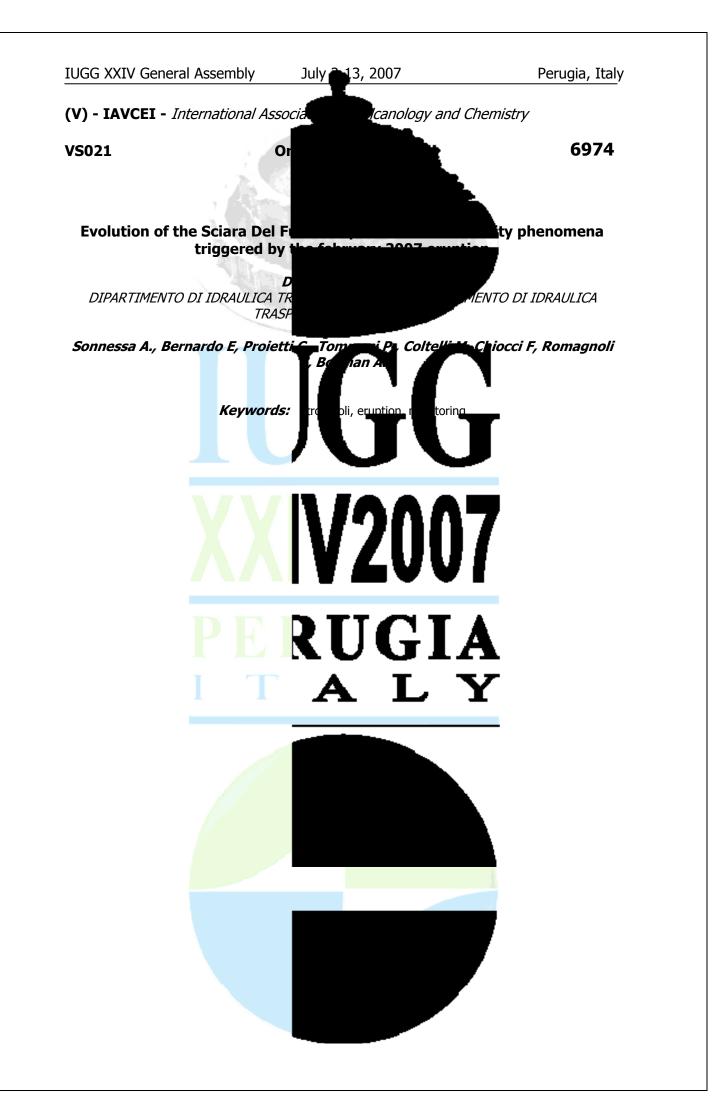


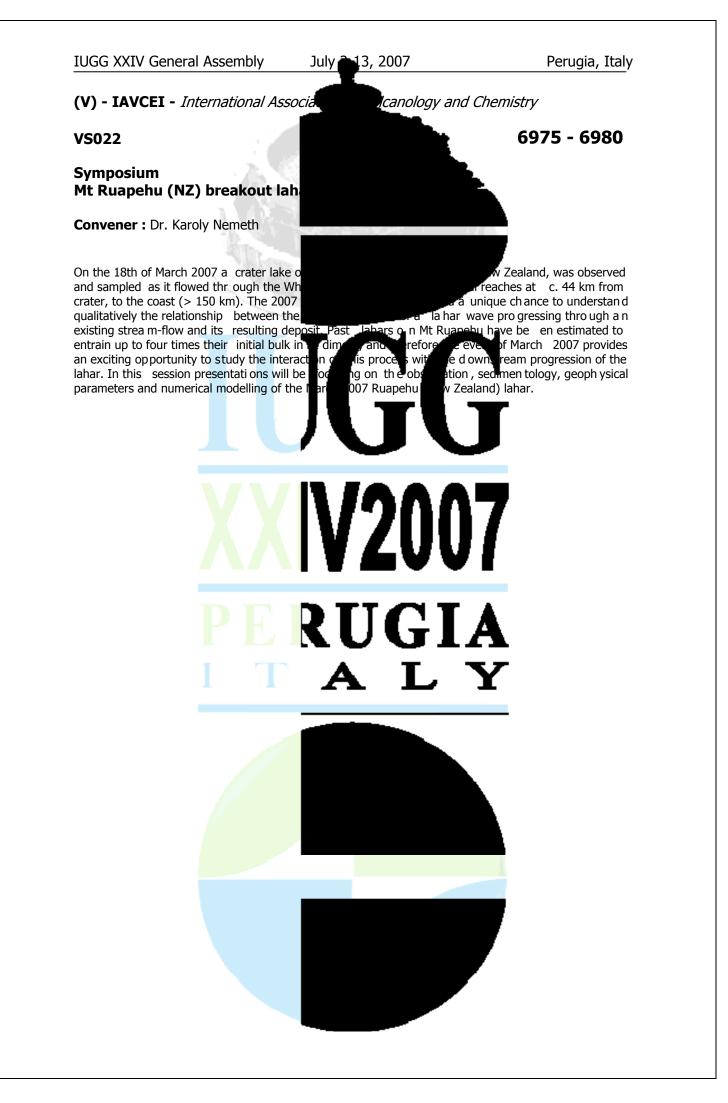
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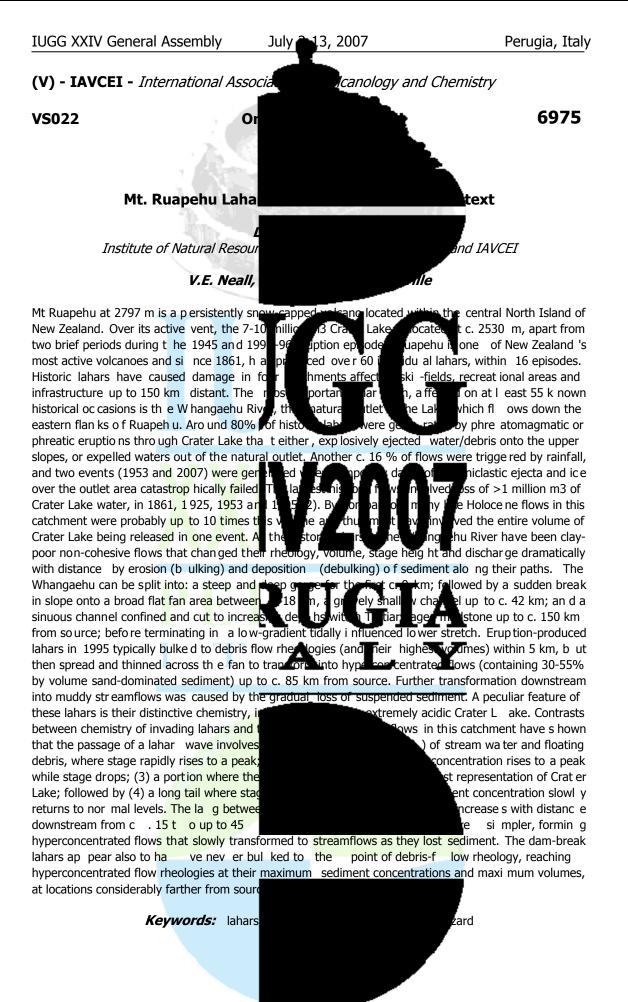


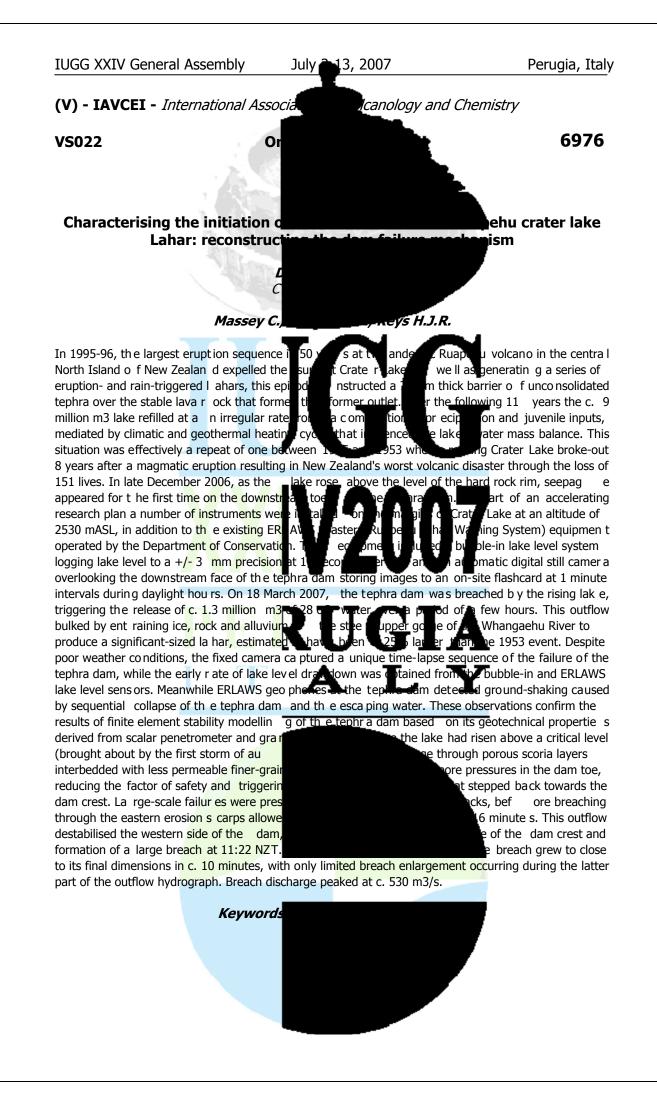


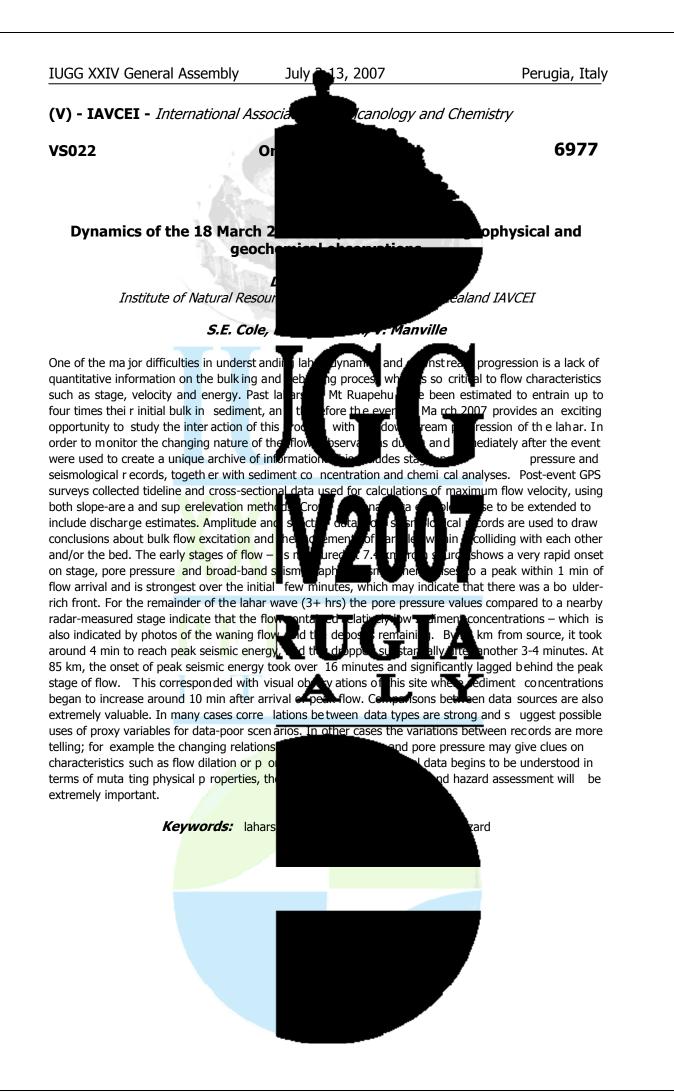


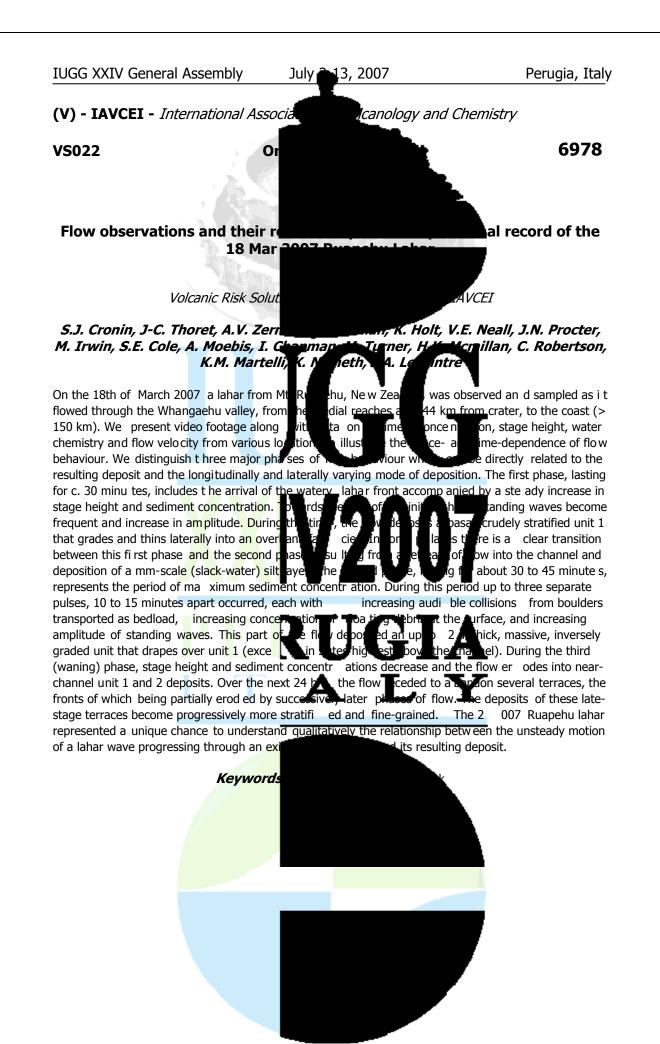


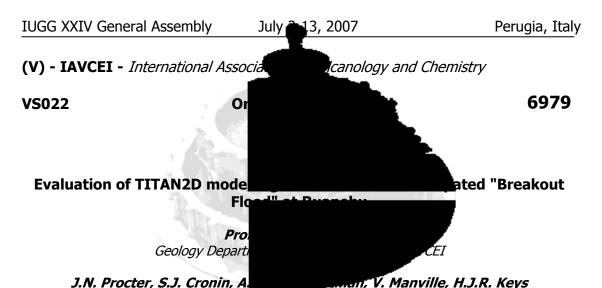












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## J.N. Procter, S.J. Cronin, A.

Following the 1996 eruptions of Ruapehu, t recognised, because an unst able 7-9 m pile stable outlet of Crater Lake. As the refilling increasingly demanded answers on the rang particular focus was a protection structure ( downstream of Crater Lake in order to stop The bund effectiveness was tested by running a rail

versions of the Titan2D mass flow modelling code. Titan-2D is a depth averaged, "shallow water" flow model, simulating either a dry granular flow, the Geophysical Mass Flow Modelling grou is that it solves the movement of a granul terrains by using an adaptive grid. We pres March 2007 b reak-out lahar to details from h th accurate in predicating both the inundatio the upper flow channel down to the bund. The velocities and discharges predicted at the bund were up

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or a two-phase viscous fluid + aranular flow developed by la

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ahar wer e immediately ad accumulated over the former hed behind this debris, authorities hars should the dam collapse. Of the Wi aehu River at c. 9.5 km nd in e Tongar iro catchment. channel reach using two

ve feature of this code ed volume over natural arios, run before the 18 models were extremely r at vario us points along to 20% higher than those of the 2007 flow, but matched closely those of the largest 1995 event in this le lated with those in the March

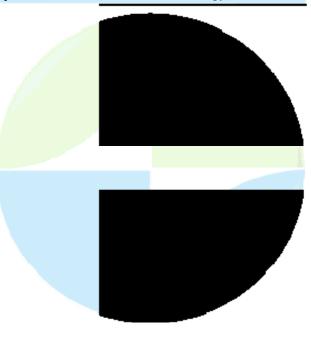
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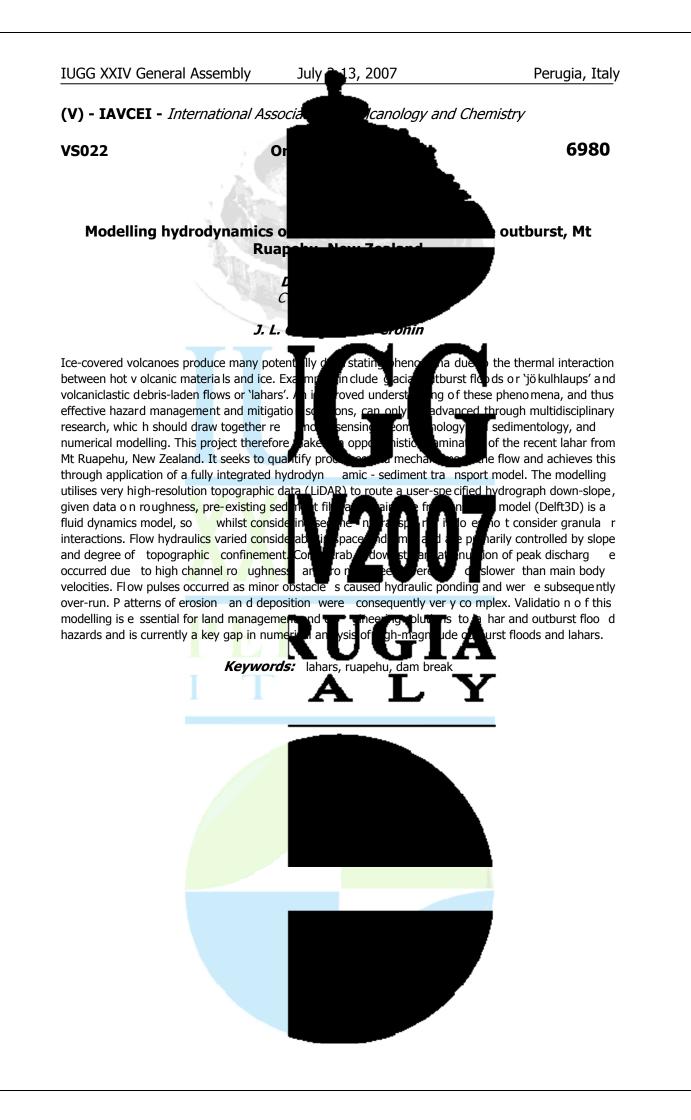
ts: In addition to these

area. Predicted paths of flow, including sid 2007 event. In addition the simulations s hydraulic ponding that correspond well to a pre-event simulations, new high-fid elity simulations using pre-event LiDAR data and actual volumes of the 2007 event will be presented and discussed.

Keywords: lahars, volcanosedimentology, volcano hazard

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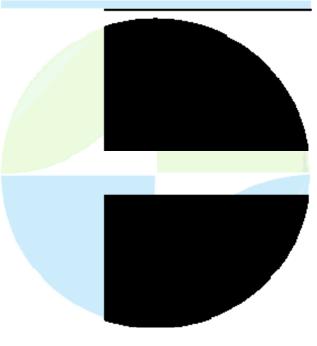
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