



Originally published as:

Lacombe, O., Mazzoli, S., von Hagke, C., Rosenau, M., Fillon, C., Granado, P. (2019): Style of deformation and tectono-sedimentary evolution of fold-and-thrust belts and foreland basins: From nature to models. - *Tectonophysics*, 767, pp. 228163.

DOI: <http://doi.org/10.1016/j.tecto.2019.228163>

1 **Style of deformation and tectono-sedimentary evolution of fold-and-thrust**  
2 **belts and foreland basins: from nature to models**

3  
4 Olivier Lacombe

5 *Sorbonne Université, CNRS-INSU, Institut des Sciences de la Terre de Paris, ITeP UMR 7193, F-75005 Paris,*  
6 *France*

7  
8  
9 Stefano Mazzoli

10 *Università di Napoli Federico II, Scuola Politecnica e delle Scienze di Base, Dipartimento di Scienze della*  
11 *Terra, dell'Ambiente e delle Risorse (DiSTAR), Complesso di Monte S. Angelo, I-80126 Napoli, Italy*

12  
13  
14 Christoph von Hagke

15 *Institute of Geology & Palaeontology, RWTH Aachen University, 52056 Aachen, Germany*

16  
17 Matthias Rosenau

18 *Helmholtz Centre Potsdam – GFZ German Research Centre for Geoscience, D-14473 Potsdam, Germany*

19  
20  
21 Charlotte Fillon

22 *Convergent Margins to Foothills Project, Exploration Production TOTAL SA, 64018 Pau, France*

23  
24  
25 Pablo Granado

26 *Universitat de Barcelona, Institut de Recerca Geomodels,*  
27 *Departament de Dinàmica de la Terra i de l'Oceà, Martí i Franquès s/n 08028, Barcelona, Spain*

28  
29

30 **1. Fold-and-thrust belts : some recent advances in their description and their**  
31 **understanding**

32

33 Orogenic systems, including their external fold-and-thrust belts and foreland basin  
34 systems, generally evolve from the inversion and imbrication of former continental margins.  
35 Continental margins are characterized by displaying along-strike variations in the degree of  
36 inherited regional extension (i.e., from limited lithospheric stretching to full necking, leading  
37 to mantle exhumation and oceanic accretion). These differences have a fundamental impact on  
38 the pre-shortening thermal state of the lithosphere and on structural style development.

39 Indeed, one of the key processes in fold-and-thrust belts is the reactivation and inversion  
40 of pre-existing extensional faults. Inversion tectonics is widespread during the evolution of  
41 many orogens and this process can exert a strong control on the structural and mechanical  
42 evolution of fold-and-thrust belts (Lacombe and Bellahsen, 2016).

43 The presence of evaporitic sequences interacting during lithospheric stretching and  
44 subsequent thermal subsidence is also a key parameter in the structural styles and deformation  
45 distribution of thrust sheets involving inverted basins and salt structures. On the other hand,  
46 deformation can also be transferred ahead and downward of the shallow deformation front,  
47 leading to frontal imbrication of deep seated structures in cratonic forelands or the sub-thrust  
48 region of active fold-and-thrust belts.

49 Defining the correct structural style of fold-and-thrust belts and understanding the  
50 controlling factors are necessary steps towards predicting their long-and short-term evolution,  
51 with implications for crustal/lithospheric rheology, mountain building processes and seismic  
52 hazard, and for the correct assessment of their potential for hydrocarbon exploration (e.g.,  
53 Butler and Mazzoli, 2006; Lacombe et al., 2007; Poblet and Lisle, 2011; Lacombe et al., 2016).  
54 For these reasons, fold-and-thrust belts and adjacent foreland basin systems represent

55 outstanding places to investigate (active) deformation and surface processes and the way these  
56 processes interact to shape mountain belts. On a short-time scale, the pattern of deformation of  
57 fold-and-thrust belts provides information on crustal mechanics, the sequence of active faulting  
58 and its relation to large earthquakes; on a long-time scale, the structure and dynamics of the  
59 fold-and-thrust belt -foreland basin systems offers unique insights into the influence of  
60 structural, thermal and rheological inheritance, together with coupling between surface and  
61 deep processes.

62           During the last ten years, significant advances have been made in the description and  
63 understanding of fold-and-thrust belts and foreland basins. Among (many) others : better  
64 definition of structures at depth (seismic imaging, 3D visualization/ geomodelling, better  
65 appraisal of geometrical uncertainties); use of analogue and numerical modelling to constrain  
66 long-term and short-term surface and deep processes; applications of thermochronology  
67 (detrital thermochronology for sediment routing and paleo-burial estimates, coupled  
68 thermochronological and 2D/3D mechanical/kinematical modelling); recognition of the  
69 influence of salt and salt tectonics; renewed conceptualization of fold-fractures relationships  
70 and new ways to unravel paleostress history.

71

## 72 Imaging structures at depth in fold-and-thrust belts and addressing uncertainties

73           Our view of the geometry of fold-thrust belts relies on the interpretation of field data,  
74 borehole data and seismic lines. Classically, geoscientists collect a large amount of outcrop and  
75 geophysical data and interpret them to arrive at a plausible geometric or steady state model.  
76 Such a model can then be used as a basis for kinematic restorations or mechanical  
77 considerations. While this approach has been successful for many years, improving our  
78 understanding of fold-and-thrust belts as well as finding oil and gas reservoirs, a topic that  
79 becomes more and more recognized is to take the uncertainties associated with the input data

80 into account. This emerging field may yield new fundamental insights on subsurface  
81 geometries, and consequently may provide different kinematic and mechanical solutions of a  
82 study area. Field data are subject to possibly large uncertainty, due to vegetation cover or  
83 variable erodibility of different rock types (Moosdorf et al. 2018), which may lead to  
84 overrepresentation of less erodible rocks in the measurements. This is particularly an issue for  
85 the location of faults, as fault rocks are commonly easily eroded. Vegetation cover is a common  
86 problem in fold-thrust belts, as they form the foothills of orogens, an area commonly inhabited  
87 and cultivated. These uncertainties will often make possible different equally viable  
88 interpretations of field data. Some of these uncertainties can be reduced by combining them  
89 with subsurface data, i.e. borehole measurements and seismic images. However, these two data  
90 sets are similarly associated with uncertainties, and it has been shown using synthetic seismic  
91 images that interpretations may have a strong bias depending on the expertise of the interpreter  
92 (Bond et al. 2007). Even when interpreters are all familiar with the area, different interpretations  
93 of the same high-resolution seismic section may be offered, as pointed out by von Hagke &  
94 Malz (2018).

95         For future research in fold- and-thrust belts addressing these uncertainties will be an  
96 important way forward. At the moment there are two main directions. First, presenting not only  
97 one plausible model, but arriving at a wide range of possible models, ideally associated with a  
98 quantitative estimate of the respective uncertainties. This requires a probabilistic approach, and  
99 has successfully been applied to a wide range of areas (Bond, 2015, Wellmann et al. 2012,  
100 Wellmann et al., 2018). This approach at least partly circumvents the bias introduced by the  
101 subjectivity of the interpreters. However, the full complexity of geological systems often cannot  
102 yet be captured. Therefore it is essential to include as much geological information as possible,  
103 appreciating the 4-D nature of structures (e.g. Duffy et al., 2018, Hessami et al., 2001, Ruh &  
104 Vergés, 2018, von Hagke et al., 2016). To that respect, 3D structural modelling using various

105 geomodelling softwares (e.g., Caumon et al., 2009; Pellerin et al., 2015) may be of great help  
106 in testing the geometrical and kinematical compatibility of structures from various geological  
107 or seismic sections (Turrini et al., 2014), performing meaningful kinematic restorations  
108 (Durand-Riard et al., 2011) or performing seismotectonic studies at a regional scale (Turrini et  
109 al., 2015). Second, and maybe even more importantly, models of the subsurface structure of  
110 fold-thrust belts should be not only tested with respect to their kinematic plausibility, but their  
111 mechanics needs to be understood. This has exemplarily been shown for the interpretation of  
112 triangle zones (von Hagke & Malz, 2018), but is equally applicable for any structural model.  
113 Understanding the mechanics of fold-thrust belts and their detachments has progressed much  
114 in the last few years, partly because of the increasing strength of numerical models and  
115 increasing resolution of micro-structural techniques. The mechanics of shale and salt-  
116 detachments become increasingly known (see reviews by Morley et al. 2017, 2018 for shale  
117 detachments, and summary below for salt detachments). Addressing the mechanics of fold-  
118 thrust belts and their respective detachments also will partly shift focus of field-based research  
119 from the regional scale to the outcrop scale, where individual key structures have to be analyzed  
120 at high resolution. At the same time, it is necessary to explore the parameter space of what is  
121 responsible for fault weakness, the influence of evolving and transient rheologies, or the  
122 influence of mechanical stratigraphy on geometries. This is best done with a combination of  
123 numerical and analog models. However, a sound understanding of the structures in the field,  
124 using classic geological field techniques combined with digital mapping (Fernández et al.  
125 2004), drone imagery (Pavlis and Mason, 2017), LiDAR and photogrammetry studies at  
126 regional and outcrop scales will remain essential, and a cornerstone of research applied to fold-  
127 thrust belts (Tavani et al. 2014; Corradetti et al. 2017; García-Sellés et al. 2011, 2018).

128

129 Analogue and numerical modelling of tectonic processes in fold-and-thrust belts

130           Analogue and numerical modellings are rather mature geoscientific tools used to  
131 understand the dynamics of fold-and-belts, accretionary wedges and orogens for several  
132 decades. Recent methodological advances, both experimental and numerical, allows nowadays  
133 simulating key geoprocesses and their coupling at various time scales at high resolution and in  
134 3D.

135           With the advent of high-resolution monitoring techniques and accurate rheological  
136 characterization of rock analogue materials, analogue modelling has transformed from a  
137 concept-testing qualitative tool to a quantitative simulation technique (see reviews by Rosenau  
138 et al., 2017; Graveleau et al., 2013).

139           On the methodological side recent efforts in analogue modelling focus for instance on  
140 the use of image-processing software to generate 3D voxel (i.e., *volumetric pixel*) models of the  
141 internal structure of sandbox models; this technique allows producing arbitrary virtual sections  
142 (i.e. inlines, cross-lines and depth slices) through sandbox models based on the reconstruction  
143 from cross-sectional images in a similar manner to 3D seismic data (Dooley et al. 2009;  
144 Granado et al. 2017). These voxels can be converted into seg-y files to be loaded to seismic  
145 interpretation software platforms (Roma et al. 2018) to generate 3D structural static models;  
146 these models can later be populated with properties to carry out dynamic simulations (i.e., fluid  
147 injection, production, etc.). Kinematic monitoring of model surfaces and volumes, on the other  
148 hand, can be carried out using digital imaging techniques (Galland et al., 2016; Boutelier et al.  
149 2019; Toeneboehm et al. 2019; Adam et al. 2013; Poppe et al., 2019) including time adaptive  
150 imaging (Rudolf et al., 2019). Dynamic monitoring using stress sensors measuring lateral push  
151 (Ritter et al., 2018 a,b, Cruz et al., 2010; Souloumiac et al., 2012) or pressure sensors providing  
152 in-situ observations (Moulas et al., 2019) improved greatly our understanding of the force  
153 balance and work budget in experimental tectonic systems. A number of new analogue  
154 materials mimicking brittle-elastoplastic and viscoelastic behaviour of rocks in a more realistic

155 (often complex) way have been developed and characterized recently (e.g. Di Giuseppe et al.,  
156 2009, 2015; Abdelmalak et al., 2016; Brizzi et al. 2016).

157 In parallel to methodological advances, studies dedicated to critical assessment of the  
158 reproducibility of experiments (Cubas et al., 2010; Santimano et al., 2015) and their boundary  
159 conditions (Souloumiac et al., 2012) raised awareness for uncertainty in analogue modelling  
160 and the relation between intrinsic and extrinsic variability of experimental observations.  
161 Community benchmarks including material characterization (Klinkmüller et al., 2016, Rudolf  
162 et al., 2016) and comparison of analogue and numerical models (Schreurs et al., 2016, Buitert  
163 et al. 2006, 2012) has moreover helped validate established analogue and numerical modelling  
164 techniques.

165 Based on methodological advances, applications of analogue modelling developed in  
166 recent years towards ever shorter timescales: while classical sandbox modelling of orogenic  
167 systems and fold-and-thrust belts (Borderie et al., 2019, Saha et al., 2019) as well as lithospheric  
168 scale multilayer models (Munteanu et al., this issue) represent the state-of-the art in analogue  
169 modelling of tectonic processes at the million year time scale, approaches to model  
170 geomorphologic (Graveleau et al., 2006, Guerit et al., 2016) and seismic cycle time scales  
171 (Rosenau et al., 2017) have entered the stage with the perspective of understanding tectonic  
172 evolution across all relevant time-scales.

173 The deeper our understanding of shallow earth deformation processes, the more  
174 complicated their modelling, however. It is for these matters that numerical modelling is  
175 gaining importance in geoscience as applied mathematical codes aid at providing a better  
176 control on geological processes involving fluid pressure, mineral reactions, surface processes  
177 etc., which cannot, or only in a very limited way, be implemented experimentally. Moreover,  
178 modern computer processing power allows running many realizations simultaneously, testing  
179 the effect/s of different parameters in parallel resulting in a generally wider parameter space to

180 be tested numerically compared to analogue models. Finally, complex geometries as well as  
181 depth-dependent rheologies and their change with temperature (i.e. thermo-mechanical) and  
182 accumulated strain (i.e., strain-weakening, strain-hardening) can also be modelled more  
183 accurately, and their influence on the resulting realizations, numerically constrained. These are  
184 possibly the great advantages of numerical methods in comparison to traditional analogue  
185 modelling; however both techniques need to be regarded as complementary, and not mutually  
186 exclusive.

187         There are three main numerical methods currently applied to fold-and-thrust belts:  
188 discrete element methods (also referred to distinct element methods, Burbidge and Braun,  
189 2002), finite element methods (Simpson, 2011; Erdős et al. 2014a; Bauville and Schmalholz,  
190 2015), and finite difference methods (Ruh, 2019). More rarely, discrete element methods are  
191 used because of their limitations in resolution with respect to geological systems. All numerical  
192 methods are challenged by the increasing demand for 3D “cross scale” (in space and time)  
193 models requiring massively parallelized codes to be developed (Kronbichler et al., 2012; van  
194 Dinther et al., 2013; Ruh et al., 2013) implementing adaptive time-steps and meshes as well as  
195 the most efficient and reliable mathematical formulations of deformation laws (Pipping et al.,  
196 2016; Herrendörfer et al., 2018; Glerum et al., 2018).

197         Future work on numerical modelling applied to fold-and-thrust belts should address the  
198 following points: the controls of inheritance from passive margins stages, including thermal  
199 effects, fault orientation and strength in 2D/3D, as well as sedimentary basin architecture and  
200 mechanical properties of the basin-infill (i.e., changes in pore-fluid pressures), include the  
201 presence of several décollements with variable strengths, underlying basement steps, as well as  
202 the existence of previous salt structures, and *non-layer cake* stratigraphies.

203

204 Fold-and-thrust belts and the thermochronology toolbox

205           Thermochronological age dating of various minerals is a key tool in Earth sciences to  
206 quantify landscape evolution and the metamorphic and tectonic history of orogens and  
207 sedimentary basins (Wagner and Reimer, 1972; Brown et al., 1994; Reiners and Brandon,  
208 2006). Thermochronology provides information on the temperature history of minerals and  
209 rocks. Thermochronometric datasets provide cooling ages, which is the time elapsed since a  
210 mineral cooled below a certain temperature. In fold-and-thrust belts, the application of low-  
211 temperature thermochronometers is ideally suited, as the method is sensitive to the uppermost  
212 few kilometers of the crust. Commonly used techniques are fission track dating and (U-Th-  
213 Sm)/He dating on zircon and apatite. Different thermochronometers are sensitive to different  
214 cooling intervals. For instance whereas the apatite fission track system is sensitive to  
215 temperatures above approximately 110°C (depending on apatite chemistry), the apatite helium  
216 system records cooling below ~60°C (Wolf et al., 1996; Farley and Stöckli, 2002). In fold-  
217 thrust belts, analyzed samples commonly derive from foreland sediments that have been  
218 incorporated later into the orogenic wedge. This implies that dated grains may have a  
219 complicated time-temperature history, starting with cooling in the hinterland, transport through  
220 the drainage area, deposition and burial (and consequently heating) in the basin and later  
221 exhumation in the fold-and-thrust belt (Fitzgerald et al. 2019). The youngest exhumation history  
222 in the fold-and-thrust belt can be complex in itself due to repeated activity of the same fault at  
223 different times. These complex t-T paths provide some challenges to be interpreted correctly,  
224 however, by integrating thermochronological data from the orogen, the foreland as well as the  
225 fold-thrust belt it is possible to provide a comprehensive picture of the t-T history of a mountain  
226 belt. Thermochronological data from fold-and-thrust belts provide also information on the  
227 provenance history of the sediments, estimates of maximum burial, as well as timing and rates  
228 of deformation. The strength of low-temperature thermochronometry has been successfully  
229 exploited in many fold-thrust belts in the world, as for instance in the Andes (McQuarrie et al.,

230 2005; Barnes et al., 2006; Savignano et al., 2016), the European Alps (von Hagke et al., 2012,  
231 2014), and the Pyrenees (Beamud et al. 2013; Mouthereau et al., 2014, Ternois et al., 2019). It  
232 has to be noted that thermochronometric ages cannot be directly translated into time of  
233 deformation (e.g. Mora et al., 2015). Particularly for fold-and-thrust belts, where sediments are  
234 commonly dated, thermochronometric ages may be associated with uncertainties due to  
235 unknown provenance ages, hydrothermal fluxes, low crystal quality due to rounded grains,  
236 limited amount of grains that can be analyzed, or unrecognized zoning of the grains. These  
237 uncertainties provide exciting future pathways for thermochronological age dating, and future  
238 research on fold-and-thrust belts should strongly follow advances in thermochronological  
239 methods, as this may help reduce the uncertainties of proposed geological models.

240         An example for the successful combination of thermochronometry and structural  
241 geology is the newly emerging field of coupling thermochronological data with kinematic  
242 restorations. Sequential restoration of balanced cross sections is a powerful tool in structural  
243 geology, as it allows one to: (i) draw a reliable picture of the changing geometry of deforming  
244 geological structures through time; (ii) determine the original position and dip of the structures;  
245 (iii) calculate the amount of shortening; (iv) define timing of basin formation and evolution,  
246 and (v) extrapolate rates of tectonic processes, such as exhumation and erosion (Bulness and  
247 McClay, 1999). Relying on pioneering works in contractional settings (Bally et al., 1966;  
248 Dahlstrom, 1969, 1970; Mitra and Namson, 1989), kinematic restoration of balanced cross  
249 sections has been applied progressively both in extensional and inverted basin areas (Coward,  
250 1996; Bulnes and McClay, 1999). In the oil industry, this technique is routinely used to evaluate  
251 the position of source rocks and model hydrocarbon generation, expulsion and migration, as  
252 well as to analyze structural traps in terms of timing and geometry (Buchanan, 1996). However,  
253 this method alone can be inadequate in case syntectonic deposits are not present/well preserved  
254 (Almendral et al., 2014; Mora et al., 2015; Granado et al. 2016a). Coupling cross-section

255 balancing and restoration with thermochronological constraints provides the possibility to  
256 define the various stages of deformation and to quantify both their extent and timing, in the lack  
257 of a syntectonic sedimentary record or in conjunction with it (Andreucci et al., 2013; Mora et.  
258 al., 2015; Castelluccio et. al., 2016; Chapman et al., 2017). However, the shift from a  
259 temperature (i.e. related to the movement of the sample through the isotherms) to a space  
260 domain (i.e. the depth of the sample during time) requires combining kinetic and  
261 thermochronological information within a coherent model capable of taking into account  
262 variations in the distribution of the isotherms in a dynamically active scenario. Another issue  
263 to take into account is the role of topographic evolution during time and the way in which it  
264 interacts with isothermal surfaces. In fact, wavelength of the relief, exhumation rate and heat  
265 advection strongly perturb the isotherm state (Stüwe et al., 1994; Mancktelow and Grasemann,  
266 1997; Braun, 2002; Reiners and Brandon, 2006). Consequently, understanding better landscape  
267 evolution (Braun et al., 2012), drainage reorganization (Yanites et al., 2013), or rock erodibility  
268 (Moosdorf et al., 2018) is important for a correct interpretation of thermochronological data  
269 and consequently the tectonics of fold-and-thrust belts. In the last few years, successful results  
270 have been obtained with software dedicated to both inverse and forward modeling of  
271 thermochronometric data (Fillon and van der Beek, 2012; Almendral et al., 2014; Erdős et al.,  
272 2014b; Castelluccio et al., 2015; Ternois et al., 2019). Forward modeling uses  
273 thermochronometric ages calculated starting from kinematic restoration integrated with thermal  
274 parameters. The output is a model of low-temperature thermochronometric ages along a  
275 geological cross-section, which can be compared with measured apatite/zircon fission track  
276 and/or (U-Th-Sm)/He cooling ages on samples collected along the profile. The comparison with  
277 measured and modeled ages allows in turn improving the structural model through an iterative  
278 process.

279

280 Influence of salt and salt tectonics in fold-and-thrust belts

281           The critical taper theory states that the external geometry and internal deformation of  
282 fold-and-thrust belts is a function of the coefficient of friction of their basal décollement (Davis  
283 et al., 1983). This has been the basis for the current modern understanding of fold-and-thrust  
284 belts. The importance of layered evaporitic sequences -commonly referred to as *salt* for short-  
285 as preferential décollements in fold-and-thrust belts has also been recognized for a long time  
286 (Davis and Engelder, 1985). Salt is inherently weak, and contrary to most rocks whose strength  
287 increases as a function of depth, salt's strength depends on its viscosity and on the applied strain  
288 rate, meaning that at high strain rates salt will undergo brittle failure; at geological strain rates  
289 however, salt will flow (Jackson and Vendeville, 1994). For such differences in mechanical  
290 behaviour, the structural styles in salt-detached fold-and-thrust belts are markedly different to  
291 those contractional systems developed over frictional décollements (Smit et al. 2003; Granado  
292 et al. 2017). In this sense, the latter are constituted by thrust faults and related folds (Jamison,  
293 1987) which tend to display a dominant forelandward vergence largely developed in footwall  
294 thrust sequences. In their simplest scenario, salt-detached fold-belts tend to show a regular  
295 spacing of narrow symmetric anticlines separated by broader *box-like* synclines; these anticlines  
296 are commonly detachment folds, or transported detachment folds, which wavelength is  
297 controlled by the thickness of the dominant mechanical unit (Mitra, 2002). Detachment folds  
298 are usually cored by salt and can reach several kilometres of structural relief; due to crestal  
299 erosion and/or faulting, salt can eventually extrude as diapiric structures (Santolaria et al. 2014).  
300 To attain such amplitudes, detachment anticlines need severe syn-orogenic sedimentation on  
301 their immediate synclines (Izquierdo-Llavall et al. 2018; Borderie et al. 2019). Salt-cored  
302 detachment folds commonly plunge gently towards their periclinal terminations; however steep  
303 plunges can also be attained in short distances compared to their lengths depending of the  
304 distribution of underlying salt. Thrusts and reverse faults in salt-detached systems display no

305 preferred sense of transport given the low friction or even frictionless nature of saline  
306 décollements. For the same reason, salt-detached fold-and-thrust belts commonly display an  
307 extremely narrow cross-sectional taper when compared to those belts detached along frictional  
308 horizons, hence being comparatively wider, with deformation concentrated at the edges of the  
309 salt basin (Davis and Engelder, 1985; Jaumé and Lillie, 1988). Thrust salients are also typical  
310 features of salt-detached fold-and-thrust belts, being commonly related to the distribution of the  
311 underlying salt décollement and to lateral variations in décollement efficiency (Becker et al.  
312 2000; Jackson et al. 2003; Muñoz et al. 2013).

313         Some of these well-established structural templates remain true and applicable, but only  
314 when a stratigraphy of uniform thickness (i.e. *layer-cake*) is involved. However, as most fold-  
315 and-thrust belts develop from the incorporation of rifts and passive margins basins in the  
316 shortening system, contrasting structural styles will develop depending on which parts of the  
317 basins are involved. In the case of rift to passive margins salt basins (Rowan, 2014; Granado et  
318 al. 2016b; Kukla et al. 2018), significant structural complexities differing from those templates  
319 described earlier can arise. For instance, many salt-detached and salt-influenced fold-and-thrust  
320 belts display structural styles which include: i) multiple structural orientations for folds and  
321 faults; ii) strong changes in fold plunges; iii) large panels of completely overturned stratigraphy  
322 (i.e. *flaps*); iv) mechanical contacts omitting or repeating stratigraphy; v) severely deformed  
323 evaporite bodies, or their equivalent salt welds, bounding structural units of markedly different  
324 sizes and aspect ratios, and contrasting stratigraphic thicknesses and facies (i.e. *non-layer cake*  
325 stratigraphy).

326         All these complex structural styles have been sometimes explained by invoking several  
327 deformation phases, strike-slip tectonics, or even the gravitational emplacement of thrust  
328 sheets. It has not been until recently that early salt tectonics processes inherited from the rift to  
329 passive margin stages have been taken into consideration. As salt is the weakest stratigraphic

330 units involved in the developing fold-and-thrust belt, shortening would be first focused on those  
331 inflated salt structures, influencing the formation and orientation of structures immediately  
332 around them (Duffy et al. 2018; Snidero et al. 2019). Diapirs and salt walls can be squeezed  
333 depending on their orientation in respect to shortening, to eventually neck-off forming sub-  
334 vertical welds (i.e. secondary welds); with ongoing shortening, sub-vertical welds may become  
335 reactivated as thrusts, reverse or transpressive faults, also depending on their orientation with  
336 respect to the direction of shortening (Rowan and Vendeville, 2006; Duffy et al. 2018; Granado  
337 et al. 2018; Roma et al. 2018). Depocenters related to early salt structures such as rollovers will  
338 become inverted, and along with minibasins if present, will be incorporated and transported  
339 along the weak basal décollement, imbricated by later thrusts, and undergo rotations along  
340 vertical and/or horizontal axes (López-Mir et al. 2014; Saura et al. 2016; Granado et al. 2018;  
341 Snidero et al. 2019). All these processes will be strongly influenced by syn-orogenic erosion  
342 and sedimentation (Izquierdo-Llavall et al. 2018).

343 To summarize, the structural styles of contractional salt-detached and salt-influenced  
344 fold-and-thrust belts are largely dependent on the relative thickness of salt and its overburden,  
345 the lateral distribution of salt previous to shortening, and whether early salt structures and  
346 related depocentres are present (Hudec and Jackson, 2007).

347

#### 348 Fracture analysis and paleostress/paleopressure reconstruction in fold-and-thrust belts

349 The reconstruction of the past kinematic and tectonic history in fold-and thrust belts  
350 requires constraints on the evolution through space and time of both stress and strain which  
351 affected sedimentary (and basement) rocks. Fractures are the most common response of brittlely  
352 deformed rocks submitted to tectonic stresses and are therefore classical and reliable  
353 palaeostress indicators (Lacombe et al., 2011). In addition, in carbonate rocks of low matrix  
354 permeability, the characteristics of the fracture network play a fundamental role in hydrocarbon

355 migration and reservoir quality (Casini et al., 2011). A good understanding of the mechanical  
356 and chronological development of the meso-scale fracture network is therefore key for tectonic  
357 analysis as well as natural resources exploration, and waste repositories studies

358         A recent step forward in the understanding of fracture occurrence in fold-and-thrust belts  
359 is related to the recognition, in addition to fold-related meso-structures, of the widespread  
360 occurrence of pre-folding fracture sets. These fracture sets may have originated from far-field  
361 earlier tectonic events unrelated to the phase of thrusting and folding or from foreland flexuring  
362 and along-foredeep stretching (Tavani et al., 2015) and even from differential compaction  
363 controlled by deep-seated faulting before the foreland domain has become part of the fold-and-  
364 thrust belt (Tavani et al., 2018). This points toward the need to carefully consider pre-existing  
365 fractures, possibly unrelated to folding, to build realistic conceptual fold–fracture models.  
366 Moreover, a blind spot for most fracture analyses has been for long the lack of constraints on  
367 the absolute timing of fracture development. This absolute timing is never resolved through  
368 field-based geometrical relationships, so the relevance and meaning of some fracture sets with  
369 respect to regional deformation can be disputable, especially in regions that underwent  
370 polyphase tectonics. Recent developments in absolute dating of calcite cements of veins within  
371 folds using U/Pb technique (Parrish et al., 2018; Hansman et al., 2018; Beaudoin et al., 2018)  
372 have brought the proof of the time relevance of meso-scale structures to regional fold-and-thrust  
373 structures, hence have helped refine the tectonic history.

374         Providing constraints on paleostress orientations and magnitudes and how they evolved  
375 during geological history is a challenging but important task that can lead to major breakthrough  
376 in the appraisal of long-term mechanical and paleohydrological behaviour of the upper crust.  
377 Improvement and application of new paleopiezometric techniques (e.g., calcite twinning and  
378 stylolite roughness paleopiezometry) calibrated in the diagenetic conditions of pressure and  
379 temperature have recently helped refine the tectonic and paleostress history in the Apennine

380 fold-and-thrust belt (Beaudoin et al., 2016). Stress quantification provides insights into the  
381 burial history independently of any assumption on the past geothermal gradient, the overall  
382 long-term mechanical behaviour of the crust, the degree of coupling between the cover and the  
383 basement and the way orogenic stresses are transmitted from the plate boundary to the (far)  
384 foreland (Beaudoin and Lacombe, 2018).

385 Last but not least, reconstruction of fluid (over)pressure and its evolution in fold-and-  
386 thrust belts and sedimentary basins is of prime interest for both academy (e.g., fault  
387 reactivation) and industry (hydrocarbon generation and migration). The use of hydrocarbon-  
388 bearing fluid inclusions, when developing contemporaneously with aqueous inclusions,  
389 provides a direct access to the pore-fluid temperature and pressure of cemented fractures or host  
390 rocks at the time of cementation and hydrocarbon trapping, in line with the tectonic evolution  
391 (Roure et al., 2010). Alternatively, the combination of the calcite twinning paleopiezometer  
392 with fracture analysis and rock mechanics tests has led to pioneering reconstructions of fluid  
393 (over)pressure evolution during the different stages of foreland shortening (e.g., Sevier-  
394 Laramide foreland, Amrouch et al., 2011; Beaudoin et al., 2014).

395

## 396 **2. Content of the Special Issue**

397 This special issue of Tectonophysics, sponsored by the International Lithosphere  
398 Program, presents a new collection of 27 papers dealing with different aspects of fold-and-  
399 thrust belts and foreland basins evolution, such as structural geology, geomorphology,  
400 exhumation, sediment transport, dating, seismicity, surface processes and basin dynamics  
401 during pre-and syn-collision stages, analogue or numerical modelling approaches, in addition  
402 to regional case studies. Some of these contributions were presented as part of a session devoted  
403 to this topic at the 2018 European Geosciences Union General Assembly in Vienna (Austria).  
404 The aim of this session was to assemble a broad group of Earth scientists interested in fold-and-

405 thrust belts and peripheral basins spanning a broad array of tectonic settings, geographical  
406 locations, and geological times. This volume presents a collection of some of the diverse  
407 research that is being carried out on this topic. We believe that these studies contribute to a  
408 better understanding not only of fold-and-thrust belts in particular, but also of orogenic  
409 processes and of the rheology of the continental lithosphere in general, and that this volume  
410 will help promote new contacts between interdisciplinary earth scientists.

411

412 \* Structural inheritance and inversion tectonics in fold-and-thrust belts

413 The role of inherited structures for the geometry of later tectonic events has been recognized  
414 since a long time. However, a dynamic understanding of the role of inherited structures on fold-  
415 thrust belt geometries is often lacking. **Granado & Ruh** address this research gap by modeling  
416 the role of inversion of a half-graben during shortening using finite differences. In their  
417 experiments they test how the strength of the inherited fault as well as different fluid pressure  
418 in the syn-rift sediments and strength of the upper décollement are reflected in the structural  
419 style of foreland fold-and-thrust belts. A straight forward result is that a weak inherited fault  
420 will be easily reactivated, influencing the kinematics and consequently structural style. Weak  
421 syn-rift sediments favor hanging-wall bypass thrusting. Generally the strength ratio of upper  
422 and lower décollement is key to understanding fold-and-thrust belts. The results of numerical  
423 models can be applied to natural examples. The authors selected the Helvetic nappes of the  
424 European Alps, and the Malargüe fold-thrust belt and the Salta Rift System, both in the  
425 Argentinian Andes as case studies for their models. They can show a first order geometric  
426 comparison of nature and experiments, indicating a quantitative comparison of the rheologies  
427 used in the model with paleo-rheologies in nature is possible.

428 **Espurt et al.** combined field geological, structural, paleo-temperature and sub-surface data  
429 together with deep geophysical data to build a new 210km-long crustal scale balanced cross-

430 section across the Central Pyrenean belt. Along the section, the belt corresponds to the inversion  
431 of the Mesozoic Pyrenean Rift system, which consisted in a hyper-extended relay zone of two  
432 metamorphic zones with exhumation of continental lithospheric mantle. Comparison between  
433 present-day crustal geometry and sequentially retro-deformed stages (lower Santonian, upper  
434 Jurassic and lower-middle Triassic) of this section shows that the Pyrenees were superimposed  
435 onto a complex structural template affected by the Variscan orogeny and subsequent Permian  
436 rifting, that in turn controlled subsequently the geometry of the Mesozoic rifting and the  
437 building of the upper Cretaceous-lower Miocene Pyrenean orogen. This study puts emphasis  
438 on the long-term influence of inherited tectonic crustal fabric in the evolution of orogens.

439 Using a kinematic forward lithosphere deformation model (RIFTER), **Gómez-Romeu et al.**  
440 produce flexural isostatically compensated, balanced geological sections across the Western  
441 Pyrenees. The tectonic evolution of both the original rift and the subsequent orogeny are  
442 investigated in order to obtain new insights into the role of extensional structural inheritance in  
443 the development of collisional orogens. The proposed model shows how, following an  
444 extensional stage characterized by a hyper-extended rift system that also led to mantle  
445 exhumation, Pyrenean shortening included two main stages. A pre-collisional stage produced  
446 inversion of the hyper-extended rift system, whereas a syn-collisional stage involved the  
447 southern proximal rift domain, resulting in the formation of the Axial Zone and the Iberian pro-  
448 foreland.

449 **Munteanu et al.** present results from state-of-the art lithospheric scale analogue modelling of  
450 compressional systems. Based on a systematic series of analog models the authors investigate  
451 the structural imprint of local crustal weaknesses on the deformation front in contractional  
452 settings. The model setup features brittle (granular) and viscous (fluid) multilayers similar to a  
453 continental lithosphere. Because the models rest on a dense low viscosity fluid (similar to the  
454 asthenosphere) they are isostatically balanced and applied lateral forces are transmitted

455 throughout the model domain. In the model, the crustal weaknesses are effectively localizing  
456 early stage deformation. Depending on the size, shape, number and location of the weakness  
457 zones either laterally continuous but bent structures evolve nucleating at the weak spots and  
458 growing laterally into stronger lithosphere or discontinuities (transfer zones) form separating  
459 strong and weak lithospheric domains. This study highlights that crustal weaknesses inherited  
460 for example from an earlier rifting phase needs to be considered as a source of along-strike  
461 irregularity of orogenic structures.

462 **Martins-Ferreira** integrates seismic interpretation (calibrated by well logs) and field data to  
463 investigate the relationships between fault reactivation and thrusting in the cratonic region of  
464 central Brazil. The study area is located between the opposed-verging Brasília and Araçuaí  
465 Neoproterozoic belts, which formed during the initial stages of West Gondwana amalgamation.  
466 Seismic interpretation suggests a tight relationship between shallow folds and thrusts and rift  
467 inversion. Reverse-slip reactivation of inherited normal faults exerts a clear control on thrust  
468 ramp nucleation. Widespread thrusting occurs over buried rifts characterized by multiple  
469 inverted faults, whereas sedimentary successions overlying non-rifted basement are not  
470 significantly affected by thrusting and folding. Furthermore, the development of salients and  
471 recesses characterizing the Brasiliano-Pan Africano orogen appears to be controlled by the  
472 morphology of buried rifts.

473 **Tavani et al.** analyzed meso-scale fractures and faults exposed in the Triassic to Miocene  
474 sedimentary succession of the Lurestan region (Zagros Belt). The authors document  
475 development of syn-sedimentary extensional fractures formed during Early Jurassic rifting then  
476 in response to foreland flexuring and along-foredeep stretching during Late Cretaceous-Eocene  
477 and late Miocene-Pliocene pulses of convergence. These repeated extensional episodes  
478 produced oblique-slip reactivation of inherited basement structures, above which differential  
479 compaction and subsidence prevailed during tectonically quiescent periods. The fractures

480 related to differential compaction add to the complexity of the ‘tectonic’ pattern formed during  
481 true tectonic (extensional and compressional) pulses, leading to a complex articulated fracture  
482 network. This study emphasizes how meso-scale structures may be relevant to regional-scale  
483 tectonics and highlights the role of tectonic inheritance in fold-and-thrust belts at all scales.

484 \* Deformation of basement rocks in fold-and-thrust belts

485 The paper by **Searle et al.** provides a review of the geology of the Caledonian Moine Thrust  
486 zone in the Loch Eriboll region, NW Scotland. In addition to already published maps, the  
487 authors use new maps, balanced and restored cross sections from fieldwork at Loch Eriboll  
488 together with a cross-section from the Moine thrust hinterland to infer geological processes  
489 including sequence of thrusting, shortening estimates and regional tectonic implications for this  
490 famous fold-and-thrust belt in the NW Scottish Highlands. The authors suggest in a rather  
491 provocative way that two major crustal-scale thrusts that extend down into the upper mantle  
492 imaged on seismic profiles across the foreland, the Outer Isles and Flannan thrusts, are unrelated  
493 spatially or temporally to the Moine thrust sequence and speculate on the regional significance  
494 of these thrusts. Finally, the authors reflect on the metamorphic sequence overlying the Moine  
495 Thrust and draw parallels to Himalayan style orogenesis.

496 **Bellahsen et al.** use the case study of the Bielsa basement unit in the Axial Zone of the Pyrenees  
497 to provide new insights into the modes and style of upper crustal shortening in orogens. In the  
498 study area, distributed strain associated with widespread, minor shear zones appears to predate  
499 strain localization along major crustal ramps (as constrained by zircon fission-track data). The  
500 authors suggest that chemical weakening – in the form of feldspar sericitization – exerts a major  
501 control on basement rheology during the early stage of distributed shortening. Sericitization is  
502 widespread, and occurs not only in ultramylonites, ultra-cataclasites and phyllonites, but also  
503 in un-deformed granodiorites. Based on these observations, the authors propose that the strength

504 of the upper crust was very low at the onset of shortening, due to a high thermal gradient and  
505 fluid circulation that induced large-scale sericitization in greenschist facies conditions.

506 \* Geomorphic signatures of tectonics and surface processes in fold-and-thrust belts

507 **Obaid & Allen** focus on the Zagros fold-thrust belt, which forms one of the best-exposed  
508 examples of and active fold-thrust belts in the world. Therefore, it is perfectly suited to  
509 investigate the role of climate on landscape evolution in an actively deforming system. The  
510 authors use geomorphic indices such as normalized channel steepness index or integrated relief  
511 and hypsometric index to test how geomorphology is a sensitive indicator of tectonic processes.  
512 The geomorphic indices show differences between different regions, correlating with climatic  
513 differences. A possible interpretation is that wetter conditions retard plateau growth, whereas  
514 dry climate allows for plateau growth, as the river draining the area have lower stream power  
515 (which is partly also a function of rock erodibility). This does not imply that climate drives  
516 landscape evolution. Instead, a positive feedback exists, where tectonics forms topographic  
517 barriers that control precipitation. Despite local differences, there is topographic similarity  
518 along five relatively evenly distributed swath profiles across the Zagros. Similar strain rates or  
519 similar overall shortening may explain this topographic similarity across the belt.

520 **Sanchez Nassif et al.** combine forward modelling and structural reconstruction to link tectonic  
521 steps to erosion rates from the Argentinian Precordillera Jachal section. Their workflow  
522 successfully articulates at each time steps forward kinematic modelling and alpha calculation  
523 following the Coulomb wedge analysis to deduce the cumulative erosion budget. Their  
524 methodology allows to obtain the sequence of deformation as well as the associated erosion  
525 rates on the basis of geological field evidences such as tectonic features and preserved  
526 sedimentary structures, without any additional t-T data. Applied to the example of the Jachal  
527 section, the authors validate their analysis with the comparison to the published section and

528 thermochronological data, showing a two-stepped evolution of the erosion rates (1 and 1.3  
529 km/Myr) associated to the activity of the Niquivil fault.

530 \*Active tectonics and seismicity in fold-and-thrust belts

531 The Hengchun Peninsula at the southern tip of Taiwan is of global importance, as it is one of  
532 the few locations in the world where an accretionary prism is exposed on land. Furthermore,  
533 the area is seismically very active and constraining the geometries of the structures is important  
534 for geohazard assessment. **Deffontaines et al.** address this using a new high-resolution digital  
535 terrain model and SAR images. The authors conclude that the peninsula consists of two ramp  
536 structures at depth, and that the Hengchun Fault, interpreted as the major tectonic lineament of  
537 the peninsula, extends towards the offshore. This study stimulates the debate on the geometry  
538 of the structures of the Hengchun Peninsula and in other systems, showing the high  
539 uncertainties at depth despite extensive data coverage. The study particularly highlights the  
540 need for further research, as a nuclear power plant is located close to the Hengchun Fault.

541 **Mescua et al.** integrate field and wellbore data to discuss the stress field in the frontal sector of  
542 the Malargüe fold-and-thrust belt (Andes of Argentina). Surface observations indicate N-S  
543 thrusts and active NW to WNW and ESE strike-slip faults in the study area. Inversion of fault  
544 kinematic indicators, combined with borehole breakout data and a mini-frac test, constrain the  
545 Quaternary to recent stress state, which is characterized by a subhorizontal, E-W oriented  
546 maximum principal stress, and by intermediate and minimum stresses with similar magnitudes  
547 that are locally interchanged, producing a setting in which reverse and strike-slip faults are  
548 alternatively active. The implications of the recognized structures for earthquake hazard are  
549 examined.

550 **Rivas et al.** analysed seismic activity in the Precordillera in the Andean backarc region of  
551 Argentina where a long record of large and damaging earthquakes in the last century exists. In

552 the northern part, which is poorly known either seismically or in terms of style of deformation,  
553 the authors determine seismic locations, seismic moments, moment magnitudes, focal depths  
554 and focal mechanisms for local earthquakes over the periods 2000–2002 and 2007–2010.  
555 Overall the results agree with the E-W compression and shortening in the Andean northern  
556 Precordilleran backarc region. The results further provide new constraints on the patterns of  
557 earthquake distribution and on the Andean backarc crustal deformation.

558

559 \* Exhumation and sediment routing in fold-and-thrust belts

560 **Odlum et al.** investigate the geological history of the South-Eastern Pyrenees, from the syn-  
561 rift period to the syn-orogenic period. With the combination of several detrital proxies such as  
562 U/Pb and (U-Th)/He on Zircons as well as U/Pb on rutiles and with an extensive mapping of  
563 the bedrock ages dated with the same techniques, the authors retrace the source of the sediments  
564 deposited from early Cretaceous to Oligocene times. By doing so, they provide a geological  
565 scenario of hinterland exhumation, foreland basin evolution, and sediment routing system of  
566 the south Eastern Pyrenees.

567 **Buford Parks & McQuarrie** present a study on the Central Andes highlighting the  
568 importance of thermal, flexural, and kinematic models for understanding the evolution of fold-  
569 and-thrust belts and orogens. The authors present a compilation of thermochronological data  
570 and use them as independent tests for different structural models. The authors show that with  
571 this method it is possible to gain insights on the sequence and rates of deformation and the  
572 plausibility of different geometric models can be tested. Coupling with flexural models shows  
573 that an additional geodynamic driver for uplift may be present, such as mantle delamination,  
574 isostatic attainment, or lower crustal flow. Apart from insights on the geologic evolution of the  
575 Central Andes, this contribution provides a tool for assessing kinematic models in fold-and-  
576 thrust belts in general.

577 Based on the integration of field geology, seismic interpretation, apatite fission track and (U-  
578 Th)/He (AHe) dating, as well analogue modelling, **Chang et al.** investigate the tectonic  
579 evolution of the Kalpin fold-and-thrust belt. As this mountain belt accommodates crustal  
580 shortening between the Tianshan and the Tarim Basin, unravelling its architecture, modes and  
581 timing of development may significantly improve our understanding of the tectonic evolution  
582 of central Asia. The timing of thrusting, constrained by low-temperature thermochronometry,  
583 implies a southward propagation of Cenozoic deformation. Combined inverse and forward  
584 thermal modelling is used to obtain information on the cooling and deformation of specific  
585 thrust sheets, while sandbox models are used to simulate and confirm the thrust sequence. The  
586 results are effectively used to provide new insights into the timing of deformation at the  
587 northern margin of the Tarim Basin.

588 \* Salt processes in fold-and-thrust belts

589 Based on a new regional balanced cross section, **Espurt et al.** present and discuss a new  
590 interpretation of the Provence fold-and-thrust belt. The system is described as a Mesozoic  
591 halokinetic salt province above a basement with strong structural inheritance, which has been  
592 subsequently shortened during the Pyrenean and Alpine orogenies. They analysed the geometry  
593 and timing of deformation and provide estimates of the amount of pre-orogenic contraction.  
594 This paper highlights the major role of halokinetic processes and shows that a significant  
595 amount of folding in fold-thrust belts can result from an early halokinetic fold system developed  
596 during the pre-contractional passive margin evolution.

597

598 On the basis of field data, **Snidero et al.** describe and interpret the stratigraphic and structural  
599 relationships between the Hormuz salt and its overburden around the Darmadan anticline in the  
600 eastern Fars region of the Zagros fold-and-thrust belt (Iran). Their model describes for the first  
601 time Late Jurassic-Early Cretaceous halokinetic sequences indicative of passive diapirism,

602 followed by early squeezing and tilting of the diapir's flanks during the Campanian-  
603 Maastrichtian. They furthermore show that second order structural features indicate secondary  
604 welding during upper Miocene times. Their work supports that the present structural trends of  
605 the Zagros fold-and-thrust belt in the eastern Fars region are the result of the reactivation of  
606 pre-existing salt structures.

607 \* Case studies

608 Andean fold-and-thrust belts and basins

609 **Barrionuevo et al.** report on the structural kinematics of the Malargüe fold-and-thrust belt, and  
610 the control imposed by the local stress field on magmatic and hydrocarbon fluid migration.  
611 They propose that the structural framework controlling the magmatic activity corresponds to  
612 inverted Mesozoic normal faults and Cenozoic thrusts, with oblique structures showing strike-  
613 slip kinematics. Miocene dykes and sills were emplaced in relation to strike-slip and reverse  
614 faults, respectively. Structural analysis suggests local switches from compressional to strike-  
615 slip/compressional likely related to the similar values of the minimum ( $\sigma_3$ ) and intermediate  
616 ( $\sigma_2$ ) principal stress with an E-W oriented maximum principal stress ( $\sigma_1$ ), favouring the  
617 emplacement of igneous intrusions and hydrocarbon migration through both thrusts and sub-  
618 vertical strike-slip faults.

619 **Ronda et al.** present a series of balanced and sequentially restored geological sections across a  
620 segment of the Southern Patagonian Andes between 46 and 48 °S. Opening of the Austral-  
621 Magallanes basin and the Rocas Verdes back-arc oceanic basin as a result of Jurassic extension  
622 were followed by Cretaceous to Cenozoic shortening. A regional westward-dipping listric  
623 detachment, interpreted as formed during Jurassic extension, was reactivated during Andean  
624 shortening as the main thrust detachment for the basement structures and the fold and thrust  
625 belt. Based on the occurrence of angular unconformities and growth strata, three main

626 shortening stages have been recognized by the authors: (i) an early (Late Cretaceous) stage  
627 associated with positive inversion and closure of the northernmost Rocas Verdes basin; (ii) a  
628 Miocene stage (between 18 Ma and 10 Ma), which produced most of the shortening; and (iii) a  
629 late (younger than 7 Ma) stage involving out-of-sequence thrusting in interior of the belt,  
630 possibly associated with the onset of glaciations and glacial erosion.

631 Peri-Mediterranean fold-and-thrust belts and basins

632 **Masrouhi et al.** provide an extensive tectono-sedimentary study of the Southern Atlas front in  
633 Tunisia. By a combination of field observations, well correlation and seismic profiles  
634 interpretation, they propose a series of restored cross-sections across the particular area of the  
635 Chott basin. The authors conclude on a mixed thick- and thin-skinned structural style with an  
636 efficient décollement level defined by the Triassic evaporites. From their reconstruction, they  
637 also find a decreasing amount of shortening from West to East, of a relatively small amount  
638 (from 7 % to 1 %) that is explained by the inversion pattern. Masroushi et al documents that  
639 this area, and the Chott basin in particular, is a large Mesozoic roll-over structure that was  
640 reactivated during Alpine orogeny.

641 In their paper, **Khomsi et al.** provide a review of the known structural architecture of the Atlas  
642 and Tell fold-and-thrust belts. Whereas the major tectonic steps accounting for the development  
643 of the eastern Maghreb structures are rather well constrained by seismic data and structural  
644 analyses, the overall configuration at depth, as well as the deep architecture of the underlying  
645 basement remain poorly understood because only few wells yet penetrated deeper than the  
646 Triassic series in the Atlas and adjacent foreland basin. The paper is intended at giving an  
647 overview on the deep structural features affecting the Pan-African basement in the eastern  
648 Maghreb by means of tentative regional cross-sections, in order to stimulate exploration of deep  
649 hydrocarbon plays with particular focus on the pre-Triassic series.

650 Based on seismic profiles and field observations, **Balestra et al.** propose a 3D numerical model  
651 of the Apennine-Maghrebian chain in the Mt. Kumeta and Mt. Rocca Busambra in Sicily. Their  
652 model builds up on cross-section restoration based on interpreted seismic profiles. They are  
653 incorporated to the 3D geomechanical model to provide a full and comprehensive view of the  
654 studied area, especially useful when seismic data is of poor quality. By applying that technique,  
655 the authors revisit the geological cross-sections by testing two-end member scenarios with a  
656 one- or two-stepped structural sequence. They find that the along-strike variation in structural  
657 style observed in the Trapanese unit is controlled by structural inheritance rather than by a  
658 polyphase history.

659 **Vitale et al.** present a review on the stratigraphy, petrology, deformation and metamorphism of  
660 the oceanic, sedimentary, magmatic and metamorphic successions from northern Calabria to  
661 the Campania region of southern Italy (i.e., the Southern Ligurian Domain). These rock units  
662 underwent subduction (Eocene) to be subsequently exhumed and exposed (Tortonian). Deep-  
663 basin successions in the easternmost sector (close to the continental margin of Adria) were  
664 obducted and frontally accreted (Aquitainian-Burdigalian). The tectonic transport for the  
665 obducted successions was dominantly to SE, however new data provides a mean eastward-  
666 directed transport during tectonic exhumation. Petro-chemical comparison between the  
667 Southern Ligurian Domain mafic rocks with the corresponding rocks in the Alps, Corsica and  
668 northern Apennines suggests an ocean continental transition setting for the Ligurian Domain.  
669 Early orogenic stages of the southern Apennines-northern Calabria system were characterized  
670 by a complex kinematic evolution of the subduction system, including the migration of the basal  
671 and roof décollements within the subduction channel.

672 **Oliva-Urcia et al.** present a new magnetostratigraphic section from the Southern Pyrenees to  
673 date the Oligocene-Miocene Pyrenean succession. They detail a 5-km section of continental  
674 syn-tectonic sediments and link their magneto-stratigraphy results to the activity of the

675 Gavarnie and Guargua thrust sheets, active from 31 to 24 and from 24 to 21 Ma, respectively.  
676 They also refine the dating on the latest deformation phase, of ~5 Myr younger than previously  
677 published and the dating the Upper Riglos thrust system at 21 Ma. Finally, they also derive the  
678 sediment accumulation rates through time that they correlate to first- and second order tectonic  
679 activities, as well as to the signal of Ebro basin closure at 36 Ma.

#### 680 Fold-and-thrust belts and basins of eastern Europe

681 **Tomek et al.** investigate the tectonics of the northeastern Variscan belt, in particular the  
682 tectonic history of the Moravosilesian Culm Basin. Using structural, paleomagnetic and  
683 magnetic anisotropy data, the authors argue that the basin underwent a change in deformation  
684 history from early compression to late strike-slip dominated tectonics. Such a switch in  
685 kinematics is important, as it provides the opportunity to test how earlier structures are  
686 influencing later deformation phases. Similarly, this study shows that from studying foreland  
687 fold-thrust belts at the regional scale, inferences can be drawn for the entire orogen and its plate  
688 tectonic context. The authors conclude that, as opposed to some previous studies, the strike-slip  
689 deformation did not play a major role during continent-continent collision in the area.

690 In order to analyse the timing of late-stage Neo-Tethys subduction and subsequent continent-  
691 continent collision in central Turkey, **Gülyüz et al.** integrate field mapping, low-temperature  
692 thermochronometry and structural analyses – including inversion of fault slip data – on the  
693 Upper Cretaceous to Eocene infill of the Haymana basin. The results of paleostress analysis  
694 point out that this basin, located at the junction between the Izmir-Ankara-Erzincan and the  
695 Intra-Tauride suture zones, underwent initial N–S to NNE-SSW extension until the middle  
696 Paleocene, followed by N–S synsedimentary shortening and coeval E-W directed extension  
697 (possibly reaching the middle Miocene). Apatite (U-Th)/He cooling ages indicate that  
698 exhumation of the southeast portion of the basin started in the early Oligocene, while the  
699 northwest part was exhumed during the early Miocene. The differential uplift and unroofing of

700 the basin fill is tentatively related to the progressive NW-ward activation of a major fault  
701 bounding the basin to the north, within the framework of the evolution of the whole basin from  
702 an extensional forearc depocentre (Late Cretaceous to early Paleocene) to a foreland basin  
703 during the subsequent collision between Taurides and Pontides.

704

705

706 Acknowledgements.

707 The Guest Editor team would like to acknowledge the work of the reviewers who have played  
708 an extremely important part in maintaining a high level of rigor to the contributions. Journal  
709 Editor-in-Chief Rob Govers is thanked for his support.

710

711 References :

712

713 Abdelmalak, M.M., Bulois, C., Mourgues, R., Galland, O., Legland, J.B. & Gruber, C., 2016. Description of new  
714 dry granular materials of variable cohesion and friction coefficient: Implications for laboratory modeling of  
715 the brittle crust. *Tectonophysics*, 684, 39-51

716 Adam, J., Klinkmüller, M., Schreurs, G. & Wieneke, B., 2013. Quantitative 3D strain analysis in analogue  
717 experiments simulating tectonic deformation: Integration of X-ray computed tomography and digital volume  
718 correlation techniques, *Journal of Structural Geology*, 55, 127-149

719 Almendral, A., Robles, W., Parra, M., Mora, A., Ketcham, R.A., & Raghieb, M., 2015. FetKin: Coupling kinematic  
720 restorations and temperature to predict thrusting, exhumation histories, and thermochronometric ages. *AAPG  
721 Bulletin*, 99, 8, p.1557-1573.

722 Amrouch K., Beaudoin N., Lacombe O., Bellahsen N. & Daniel J.M., 2011, Paleostress magnitudes in folded  
723 sedimentary rocks. *Geophysical Research Letters*, 38, L17301

724 Andreucci B., Castelluccio A., Jankowski L., Mazzoli S., Szaniawski R. & Zattin M., 2013. Burial and exhumation  
725 history of the Polish Outer Carpathians: Discriminating the role of thrusting and post-thrusting extension.  
726 *Tectonophysics*, 608, 866-883.

727 Balestra M., Corrado S., Aldega L., Rudkiewicz J.-L. & Sassi W., 2019. 3D structural modeling and restoration of  
728 the Apennine-Maghrebian chain in Sicily: Application for non-cylindrical fold-and-thrust belts.  
729 *Tectonophysics*, 761, 86-107 (this issue)

730 Bally, A.W., Gordy, P.L., & Stewart, G.A., 1966. Structure, seismic data, and orogenic evolution of southern  
731 Canadian Rocky Mountains. *Bulletin of Canadian Petroleum Geology*, 14, 3, 337-381.

732 Barnes, J.B., Ehlers, T.A., Mcquarrie, N., O'sullivan, P.B. & Pelletier, J., 2006. Eocene to recent variations in  
733 erosion across the central Andean fold-thrust belt, northern Bolivia: Implications for plateau evolution. *Earth  
734 and Planetary Science Letters* 248, 118-133.

735 Barrionuevo M., Giambiagi L., Mescua J., Suriano J., de la Cal H., Soto J.L. & Lossada A.C., 2019. Miocene  
736 deformation in the orogenic front of the Malargüe fold-and-thrust belt (35°30'- 36° S): controls on the migration  
737 of magmatic and hydrocarbon fluids. *Tectonophysics*, in press (this issue)

738 Bauville, A., Schmalholz, S. M. 2015. Transition from thin- to thick-skinned tectonics and consequences for nappe  
739 formation: Numerical simulations and applications to the Helvetic nappe system, Switzerland. *Tectonophysics*  
740 665, 101-117.

741 Beamud, E., Muñoz, J. A., Fitzgerald, P. G., Baldwin, S. L., Garcés, M., Cabrera, L. Metcalf, J. R. 2011.  
742 Magnetostratigraphy and detrital apatite fission track thermochronology in syntectonic conglomerates:  
743 constraints on the exhumation of the South-Central Pyrenees. *Basin Research*, 23, 309-331

744 Beaudoin N. & Lacombe O., 2018. Recent and future trends in paleopiezometry in the diagenetic domain: insights  
745 into the tectonic paleostress and burial depth history of fold-and-thrust belts and sedimentary basins. *J. Struct.  
746 Geol.* 114, 357-365

747 Beaudoin N., Lacombe O., Bellahsen N., Amrouch K. & Daniel J.M., 2014. Evolution of fluid pressure during  
748 folding and basin contraction in overpressured reservoirs: insights from the Madison-Phosphoria carbonate  
749 formations in the Bighorn basin (Wyoming, USA). *Marine and Petroleum Geology*, 55, 214-229

750 Beaudoin, N., Koehn D., Lacombe O., Lecouty A, Billi A., Aharonov., E. & Parlangeau C., 2016. Fingerprinting  
751 stress: stylolite and calcite twinning paleopiezometry revealing the complexity of stress distribution during  
752 folding – the case of the Monte Nero anticline in the Apennines, Italy. *Tectonics*, 35, 1687-1712

753 Beaudoin N., Lacombe O., Roberts N.M.W. & Koehn D., 2018. U-Pb dating of calcite veins reveals complex  
754 stress evolution and thrust sequence in the Bighorn Basin, USA. *Geology*, 46(11), 1015-1018

755 Becker, A. 2000. The Jura Mountains: An active foreland fold-and-thrust belt?. *Tectonophysics* 321, 381-406.

756 Bellahsen N., Bayet L., Denele Y., Waldner M., Airaghi L., Rosenberg C., Dubacq B., Mouthereau F., Bernet M.,  
757 Pik R., Lahfid A. & Vacherat A., 2019. Shortening of the Axial Zone, Pyrenees: shortening sequence, upper  
758 crustal mylonites and crustal strength. *Tectonophysics*, 766, 433-452 (this issue)

759 Bond, C.E., Gibbs, A.D., Shipton, Z.K. & Jones, S., 2007. What do you think this is? "Conceptual uncertainty" in  
760 geoscience interpretation. *GSA Today* 17, 4–10.

761 Bond, C. E., 2015. Uncertainty in structural interpretation: Lessons to be learnt. *Journal of Structural Geology*, 74,  
762 185-200.

763 Borderie, S., Vendeville, B.C., Graveleau, F., Witt, C., Dubois, P., Baby, P. & Calderon, Y. 2019. Analogue  
764 modeling of large-transport thrust faults in evaporites-floored basins: Example from the Chazuta Thrust in the  
765 Huallaga Basin, Peru. *Journal of Structural Geology* 123, 1-17.

766 Boutelier D., Schrank C. & Regenauer-Lieb K., 2019. 2D finite displacements and strain from particle imaging  
767 velocimetry (PIV) analysis of tectonic analogue models with TecPIV. *Solid Earth*, 10, 1123-1139

768 Boyer, S.R. & Elliot, D. 1982. Thrust systems. *AAPG Bulletin*, 66 (9), 1196-1230.

769 Braun, J., 2002. Quantifying the effect of recent relief changes on age–elevation relationships. *Earth and Planetary*  
770 *Science Letters*, 200, 3, 331-343.

771 Braun, J., Van Der Beek, P., Valla, P., Robert, X., Herman, F., Glotzbach, C., Pedersen, V., Perry, C., Simon-  
772 Labric, T. & Prigent, C., 2012. Quantifying rates of landscape evolution and tectonic processes by  
773 thermochronology and numerical modeling of crustal heat transport using PECUBE. *Tectonophysics* 524–525,  
774 1-28.

775 Brizzi, S., Funicello, F., Corbi, F., Di Giuseppe, E. & Mojoli, G., 2016: Salt matters: How salt affects the  
776 rheological and physical properties of gelatine for analogue modelling, *Tectonophysics*, 679, pp. 88-101

777 Brown, R.W., Summerfield, M.A. & Gleadow, A.J.W., 1994. "Apatite fission track analysis: Its potential for the  
778 estimation of denudation rates and implications for models of long-term landscape development," in *Process*  
779 *models and theoretical geomorphology*, ed. M.J. Kirkby. Wiley, Chichester, 24-53.

780 Buchanan, J.G., 1996. The application of cross-section construction and validation within exploration and  
781 production: a discussion. *Geological Society, London, Special Publications*, 99, 1, 41-50.

782 Buford Parks V.M. & McQuarrie N., 2019. Kinematic, Flexural, and Thermal Modelling in the Central Andes:  
783 Unravelling age and signal of deformation, exhumation, and uplift. *Tectonophysics*, 766, 302-325 (this issue)

784 Buiter, S. J. H., 2012, A review of brittle compressional wedge models, *Tectonophysics*, 530, 1-17

- 785 Butier, S. J. H., Babeyko A.Y., Ellis S., Gerya T.V., Kaus B.J.P., Kellner A., Schreurs G. & Yamada Y., 2006.  
786 The numerical sandbox: comparison of model results for a shortening and an extension experiment, Geological  
787 Society, London, Special Publications, 253, 29-64
- 788 Bulnes, M. & McClay, K., 1999. Benefits and limitations of different 2D algorithms used in cross-section  
789 restoration of inverted extensional faults: application to physical experiments: *Tectonophysics*, 312, 2, 175-  
790 189.
- Burbidge, D.R. & Braun, J. 2002. Numerical models of the evolution of accretionary wedges and fold-and-thrust  
belts using the distinct-element method. *Geophysical Journal International*, 148 (3), 542–561.
- Butler, R. W. H. & Mazzoli, S., 2006. Styles of continental contraction: A review and introduction, in Mazzoli,  
S., and Butler, R. W. H., *Styles of Continental Contraction: Geological Society of America Special Paper 414*,  
p. 1–10
- Casini, G., Gillespie, P.A., Vergés, J., Romaine, I., Fernández, N., Casciello, E., Hunt, D.W., 2011. Sub-seismic  
fractures in foreland fold and thrust belts: insight from the Lurestan Province, Zagros Mountains, Iran.  
*Petroleum Geoscience*. 17, 3, 263–282
- Castelluccio, A., Andreucci, B., Zattin, M., Ketcham, R.A., Jankowski, L., Mazzoli, S. & Szaniawski, R., 2015.  
Coupling sequential restoration of balanced cross sections and low-temperature thermochronometry: The case  
study of the Western Carpathians. *Lithosphere*, 7, 4, 367-378.
- Castelluccio A., Mazzoli S., Andreucci B., Jankowski L., Szaniawski R. & Zattin M., 2016. Building and  
exhumation of the Western Carpathians: New constraints from sequentially restored, balanced cross sections  
integrated with low-temperature thermochronometry. *Tectonics*, 35, 2698–2733.
- Caumon G., Collon-Drouaillet P., Le Carlier de Veslud C., Viseur C. & Sausse J., 2009. Surface-based 3D  
modeling of geological structures. *Mathematical Geosciences*, 41, 8, 927-945
- Chang J., Li D., Min K., Qiu N., Xiao Y., Wu H. & Liu N., 2019. Cenozoic deformation of the Kalpin fold-and-  
thrust belt (southern Chinese Tian Shan): new insights from low-temperature thermochronology and sandbox  
modeling. *Tectonophysics*, 766, 416-432 (this issue)
- Chapman, J.B., Carrapa, B., Ballato, P., DeCelles, P.G., Worthington, J., Oimahmadov, I., Gadoev, M. &  
Ketcham, R., 2017. Intracontinental subduction beneath the Pamir Mountains: Constraints from  
thermokinematic modeling of shortening in the Tajik fold-and-thrust belt. *Geological Society of America  
Bulletin*, 129, 11-12, 1450-1471.

- Corradetti, A., Tavani, S., Russo, M., Granado, P., Arbues, P. 2017. Quantitative analysis of folds by means of orthorectified photogrammetric 3D models: a case study from Mt Catria, Northern Apennines, Italy. *The Photogrammetric Record*, 32 (160), 480-496.
- Coward, M.P. 1996. Balancing sections through inverted basins. In: *Modern Developments in Structural Interpretation, Validation and Modelling* (Buchanan, P.G., Nieuwland, D.A. eds.), Geological Society, London, Special Publication 99, 51-77.
- Cruz, L., Malinski, J., Wilson, A., Take, W.A. & Hilley G., 2010. Erosional control of the kinematics and geometry of fold-and-thrust belts imaged in a physical and numerical sandbox. *Journal of Geophysical Research*, 115, B09404
- Cubas, N., Maillot B. & Barnes C., 2010. Statistical analysis of an experimental compressional sand wedge. *Journal of Structural Geology*, 32, 818-831
- Dahlstrom, C.D.A., 1969. Balanced cross sections. *Canadian Journal of Earth Sciences*, 6, 4, 743-757.
- Dahlstrom, C.D., 1970. Structural geology in the eastern margin of the Canadian Rocky Mountains. *Bulletin of Canadian Petroleum Geology*, 18, 3, 332-406.
- Davis, D., Supper, J. & Dahlen, F.A., 1983. Mechanics of fold-and-thrust belts and accretionary wedges: *Journal of Geophysical Research*, 88, B2, 1153-1172
- Davis, D.M. & Engelder, T. 1985. The role of salt in fold-and-thrust belts: *Tectonophysics*, 119, 67-88.
- Deffontaines B., Chang K.-J., Lee C.-T., Magalhaes S. & Serries G., 2019. Neotectonics of the Southern Hengchun Peninsula (Taiwan): Inputs from high resolution UAS Digital Terrain Model, updated geological mapping and PSInSAR techniques. *Tectonophysics*, in press (this issue)
- Di Giuseppe, E., Funicello, F., Corbi, F., Ranalli, G. & Mojoli, G., 2009. Gelatins as rock analogs: A systematic study of their rheological and physical properties, *Tectonophysics*, 473 (3-4), 391-403
- Di Giuseppe, Corbi E.F., Funicello F., Massmeyer A., Santimano T.N., Rosenau M. & Davaille A., 2015. Characterization of Carbopol hydrogel rheology for experimental tectonics and geodynamics, *Tectonophysics*, 642, 29-45
- Dooley, T.P., Jackson, M.P.A., Hudec, M.R. 2009. Inflation and deflation of deeply buried salt stocks during lateral shortening. *Journal of Structural Geology*, 31, 582-600.

- Duffy, O., Dooley, T.P., Hudec, M.R., Jackson, M.P.A., Fernandez, N., Jackson, C. A-L. & Soto, I. 2018. Structural evolution of salt-influenced fold-and-thrust belts: A synthesis and new insights from basins containing isolated salt diapirs. *Journal of Structural Geology*, 114, 206-221.
- Durand-Riard P., Salles L., Ford M., Caumon G. & Pellerin J., 2011. Understanding the evolution of syn-depositional folds: Coupling decompaction and 3D sequential restoration. *Marine and Petroleum Geology*, 28, 8, 1530-1539
- Erdős, Z., Huismans, R. S., Beek, P., Thieulot, C. 2014a. Extensional inheritance and surface processes as controlling factors of mountain belt structure, *J. Geophys. Res. Solid Earth*, 119, 9042– 9061
- Erdős, Z., Beek, P. & Huismans, R.S., 2014b. Evaluating balanced section restoration with thermochronology data: A case study from the Central Pyrenees. *Tectonics* 33, 617-634
- 791 Espurt N., Angrand P., Teixell A., Labaume P., Ford M., de Saint Blanquat M. & Chevrot S., 2019. Crustal-scale  
792 balanced cross-section and restorations of the Central Pyrenean belt (Nestes-Cinca transect): Highlighting the  
793 structural control of Variscan belt and Permian-Mesozoic rift systems on mountain building. *Tectonophysics*,  
794 764, 25-45 (this issue)
- 795 Espurt N., Wattellier F., Philip J., Hippolyte J.-C., Bellier O. & Bestani L., 2019. Mesozoic halokinesis and  
796 basement inheritance in the eastern Provence fold-thrust belt, SE France. *Tectonophysics*, 766, 60-80 (this  
797 issue)
- 798 Farley, K.A. & Stöckli, D.F., 2002. "(U-Th)/He dating of phosphates: Apatite, monazite, and xenotime," in  
799 *Reviews of Mineralogy and Geochemistry, Phosphates: Geochemical, Geobiological, and Materials*  
800 *Importance*, eds. M.J. Kohn, J. Rakovan & J.M. Hughes, 559-577.
- 801 Fernández, O., Muñoz, J.A., Arbués, P., Falivene, O., Marzo, M. 2004. Three-dimensional reconstruction of  
802 geological surfaces: An example of growth strata and turbidite systems from the Ainsa basin (Pyrenees, Spain).  
803 *AAPG Bulletin*, 88 (8), 1049-1068.
- 804 Fitzgerald, P.G., Malusà, M.G. & Muñoz, J.A. 2019. Detrital thermochronology using conglomerates and cobbles.  
805 In: *Fission-Track Thermochronology and its Application to Geology* (Malusà, Marco G., & Fitzgerald, Paul  
806 G., eds.). Springer International Publishing AG, Springer, Cham, pp.295-314
- 807 Fillon, C. & van der Beek, P., 2012. Post-orogenic evolution of the southern Pyrenees: constraints from inverse  
808 thermo-kinematic modelling of low-temperature thermochronology data. *Basin Research*, 24, 418-436

809 Galland, O., Bertelsen, H.S. Gulstrand, F., Girod, L., Johannessen, R.F., Bjugger, F., Burchardt, S. & Mair,  
810 K.,2016. Application of open-source photogrammetric software MicMac for monitoring surface deformation  
811 in laboratory models. *Journal of Geophysical Research*,121, 4, 2852-2872

812 García-Sellés, D., Gratacós, O., Granado, P., Carrera, N., Muñoz, J.A., Sarmiento, S., Lakshmikantha, M.R.,  
813 Cordova, J.C. 2018. Fracture analog of the sub-Andean Devonian of southern Bolivia: Lidar applied to  
814 Abra Del Condor, In G. Zamora, K. R. McClay, and V. A. Ramos, eds., *Petroleum basins and hydrocarbon*  
815 *potential of the Andes of Peru and Bolivia: AAPG Memoir 117*, 569–604.

816 García-Sellés, D., Falivene, O., Arbués, P., Gratacos, O., Tavani, S., Muñoz, J.A. 2011. Supervised identification  
817 and reconstruction of near-planar geological surfaces from terrestrial laser scanning. *Computer and*  
818 *Geosciences*, 37, 1584-1594.

819 Glerum, A., Thieulot, C., Fraters, M., Blom, C. & Spakman, W., 2018. Nonlinear viscoplasticity in ASPECT:  
820 benchmarking and applications to subduction, *Solid Earth*, 9, 267-294.

821 Gómez-Romeu J., Masini E., Tugend J., Ducoux M. & Kuszniir N., 2019. Role of rift structural inheritance in  
822 orogeny highlighted by the Western Pyrenees case-study. *Tectonophysics*, 766, 131-150 (this issue)

823 Granado, P., Ferrer, O., Muñoz, J.A., Thöny, W. & Strauss, P. 2017. Basin inversion in tectonic wedges: a  
824 comparative approach from analogue modelling and the Alpine-Carpathian fold-and-thrust belt.  
825 *Tectonophysics*, 704-705, 50-68

826 Granado, P., Roca, E., Strauss, P., Pelz, K. & Muñoz, J.A. 2018. Structural styles in fold-and-thrust belts involving  
827 early salt structures: The Northern Calcareous Alps (Austria). *Geology*, 47, 51-54

828 Granado, P., Thöny, W., Carrera, N., Gratzner, O., Strauss, P., Muñoz, J.A. 2016a. Basement-involved reactivation  
829 in fol- and-thrust belts: the Alpine-Carpathian Junction (Austria). *Geological Magazine*, 153 (5-6), 1100-1135.

830 Granado, P., Urgeles, R., Sábat, F., Albert-Villanueva, E., Roca, E., Muñoz, J.A., Mazzucca, N. & Gambini, R.  
831 2016b. Geodynamical framework and hydrocarbon plays of a salt giant: the North Western Mediterranean  
832 Basin. *Petroleum Geoscience*, 22, 309-321

833 Granado P. & Ruh J.B., 2019. Numerical modelling of inversion tectonics in fold-and-thrust belts. *Tectonophysics*,  
834 763, 14-29 (this issue)

835 Graveleau, F., Hurtrez, J.-E., Dominguez, S. & Malavieille J., 2011. A new experimental material for modeling  
836 interactions between tectonics and surface processes. *Tectonophysics* 513 (1–4), 68–87.

837 Graveleau, F., Malavieille J. & Dominguez S., 2012. Experimental modelling of orogenic wedges: A review.  
838 *Tectonophysics* 538-540, 1–66

839 Guerit, L., Dominguez, S., Malavieille, J. & Castellort, S., 2016. Deformation of an experimental drainage  
840 network in oblique collision. *Tectonophysics*, 693, 210-222

841 Gülyüz E., Özkaptan M., Kaymakci N., Persano C. & Stuart F.M., 2019. Kinematic and Thermal Evolution of the  
842 Haymana Basin, a fore-arc to foreland basin in Central Anatolia (Turkey). *Tectonophysics*, 766, 326-339 (this  
843 issue)

844 Hansman, R.J., Albert, R., Gerdes, A., and Ring, U., 2018, Absolute ages of multiple generations of brittle  
845 structures by U-Pb dating of calcite. *Geology*, 46, 207–210

846 Herrendörfer, R., Gerya, T. & Dinther, Y., 2018. An invariant rate- and state- dependent friction formulation for  
847 viscoelastoplastic earthquake cycle simulations. *Journal of Geophysical Research: Solid*  
848 *Earth*, 123, 5018– 5051

849 Hessami, K., Koyi, H. A., Talbot, C. J., Tabasi, H., & Shabanian, E., 2001. Progressive unconformities within an  
850 evolving foreland fold–thrust belt, Zagros Mountains. *Journal of the Geological Society*, 158(6), 969-981.

851 Hoth, S., Kukowski, N. & Oncken, O., 2006. Influence of erosion on the kinematics of bivergent orogens: Results  
852 from scaled sandbox simulations – In: Willett, S.D., Hovius, N., Brandon, M.T., Fischer, D. (Eds.), *Tectonics,*  
853 *Climate, and Landscape evolution: Geol. Soc. Amer. Spec. Pap. 398, Penrose Conference Series*, 201-225

854 Hudec, M.R. & Jackson, M.P.A. 2007. Terra Infirma: understanding salt tectonics. *Earth-Science Reviews*, 82, 1-  
855 28.

856 Izquierdo-Llavall, E., Roca, E., Xie, H., Pla, O., Muñoz, J.A., Rowan, M.G., Yuan, N. & Huang, S. 2018. Influence  
857 of overlapping décollements, syntectonic sedimentation, and structural inheritance in the evolution of a  
858 contractional system: The central Kuqa fold-and-thrust Belt (Tian Shan Mountains, NW China). *Tectonics*, 37  
859 (8), 2608-2632.

860 Jackson, M. P. A., Warin, O. N., Woad, G. M. & Hudec, M. R. 2003. Neoproterozoic allochthonous salt tectonics  
861 during the Lufilian orogeny in the Katangan Copperbelt, central Africa. *Geological Society of America*  
862 *Bulletin*, 115, 314–330.

863 Jackson, M.P.A. & Vendeville, B.C. 1994. Regional extension as geologic trigger for diapirism. *GSA Bull.*, 106,  
864 57-73.

865 Jamison, W. 1987. Geometric analysis of fold development in overthrust terranes: *Journal of Structural Geology*,  
866 9, 207-219.

867 Jaumé, S.C. & Lillie, R.J. 1988. Mechanics of the Salt Range – Potwar Plateau, Pakistan: a fold-and-thrust belt  
868 underlain by evaporites. *Tectonics*, 7, 57-71.

869 Khomsi S., Roure R., Khelil M., Mezni R. & Echihi O., 2019. A review of the crustal architecture and related pre-  
870 salt oil/gas objectives of the eastern Maghreb Atlas and Tell: Need for deep seismic reflection profiling.  
871 *Tectonophysics*, 766, 232-248 (this issue)

872 Kronbichler, M., Heister, T. & Bangerth, W., 2012. High accuracy mantle convection simulation through modern  
873 numerical methods. *Geophysical Journal International*, 191, 12–29

874 Kukla, P.A., Strozyk, F. & Mohriak, W.U. 2018. South Atlantic salt basins – Witnesses of complex passive margin  
875 evolution. *Gondwana Research*, 53, 41-57.

876 Lacombe, O., Lavé, J., Roure, F. M., & Vergés, J. (Eds.), 2007. Thrust belts and foreland basins: From fold  
877 kinematics to hydrocarbon systems. Springer Science & Business Media.

878 Lacombe O., Ruh J., Brown D. & Nilfouroushan F., 2016. Introduction to the Special Issue : “Tectonic evolution  
879 and mechanics of basement-involved fold-and-thrust belts”, *Geological Magazine*, 153, 5-6, 759-762

880 Lacombe O. & Bellahsen N., 2016. Thick-skinned tectonics and basement-involved fold-thrust belts. Insights from  
881 selected Cenozoic orogens. *Geological Magazine*, 153, 5-6, 763-810

882 Lacombe O., Bellahsen N. & Mouthereau F., 2011, Fracture patterns in the Zagros Simply Folded Belt (Fars):  
883 New constraints on early collisional tectonic history and role of basement faults. In “Geodynamic evolution of  
884 the Zagros”, O. Lacombe, B. Grasemann and G. Simpson eds, *Geological Magazine*, 148, 940-963

885 López-Mir, B., Muñoz, J.A. & García-Senz, J. 2014. Restoration of basins driven by extension and salt tectonics :  
886 Exaple from the Cotiella Basin in the Central Pyrenees. *Journal of Structural Geology*, 69, 147-162.

887 Mancktelow, N.S. & Grasemann, B., 1997. Time-dependent effects of heat advection and topography on cooling  
888 histories during erosion. *Tectonophysics*, 270, 3, 167-195.

889 Martins-Ferreira M.A.C., 2019. Effects of initial rift inversion over fold-and-thrust development in a cratonic far-  
890 foreland setting. *Tectonophysics*, 757, 88-107 (this issue)

891 Masrouhi M., Gharbi M., Bellier O. & Ben Youssef M., 2019. The Southern Atlas Front in Tunisia and its foreland  
892 basin: Structural style and regional-scale deformation. *Tectonophysics*, 764, 1-24 (this issue)

893 McQuarrie, N., Horton, B.K., Zandt, G., Beck, S. & Decelles, P.G., 2005. Lithospheric evolution of the Andean  
894 fold–thrust belt, Bolivia, and the origin of the central Andean plateau. *Tectonophysics* 399, 15-37.

895 Mescua J.F., Barrionuevo M., Giambiagi L., Suriano J., Spagnotto S., Stahlschmidt E., de la Cal H., Soto J.L. &  
896 Mazzitelli M., 2019. Stress field and active faults in the orogenic front of the Andes in the Malargüe fold-and-  
897 thrust belt (35°-36°S). *Tectonophysics*, 766, 179-193 (this issue)

898 Mitra, S. 2002. Structural models of faulted detachment folds. *AAPG Bulletin* 86 (9), 1673–1694.

899 Mitra, S. & Namson, J.S., 1989. Equal-area balancing. *American Journal of Science*, 289, 5, 563-599.

900 Moosdorf, N., Cohen, S. & Von Hagke, C., 2018. A global erodibility index to represent sediment production  
901 potential of different rock types. *Applied Geography* 101, 36-44.

902 Mora, A., 2015. Petroleum systems of the Eastern Cordillera, foothill basins, and associated Llanos basin: Impacts  
903 on the prediction of large scale foreland and foothill petroleum accumulations: *AAPG Bulletin*, 99, 8, 1401-  
904 1406.

905 Morley, C. K., von Hagke, C., Hansberry, R. L., Collins, A. S., Kanitpanyacharoen, W. & King, R., 2017. Review  
906 of major shale-dominated detachment and thrust characteristics in the diagenetic zone: Part I, meso-and macro-  
907 scopic scale. *Earth-Science Reviews*, 173, 168-228.

908 Morley, C. K., von Hagke, C., Hansberry, R., Collins, A., Kanitpanyacharoen, W., & King, R., 2018. Review of  
909 major shale-dominated detachment and thrust characteristics in the diagenetic zone: Part II, rock mechanics  
910 and microscopic scale. *Earth-Science Reviews*, 176, 19-50.

911 Moulas, E., Sokoutis, D. & Willingshofer, E., 2019. Pressure build-up and stress variations within the Earth's crust  
912 in the light of analogue models. *Scientific Reports*, 2310, 9

913 Mouthereau F., Filleaudeau P.-Y., Vacherat A., Pik R, Lacombe O., Fellin M.G., Castelltort S., Christophoul F. &  
914 Masini E., 2014. Placing limits to shortening evolution in the Pyrenees: role of margin architecture and  
915 implications for the Iberia/Europe convergence. *Tectonics*, 33, 12, 2283-2314

916 Muñoz, J. A., Beamud, E., Fernández, O., Arbués, P., Dinarès-Turell, J. & Poblet, J. 2013. The Ainsa Fold and  
917 Thrust Oblique Zone of the Central Pyrenees: kinematics of a curved contractional system from paleomagnetic  
918 and structural data. *Tectonics* 32(5), 1142-1175.

919 Munteanu I, Willingshofer E., Matenco L., Sokoutis D., Dinu C. & Cloetingh S.A.P.L., 2019. Far-field strain  
920 transmission and contractional step-overs. *Tectonophysics*, 766, 194-204 (this issue)

921 Obaid A.K. & Allen M.B., 2019. Landscape expressions of tectonics in the Zagros fold-and-thrust belt.  
922 *Tectonophysics*, 766, 20-30 (this issue)

923 Odlum M.L., Stockli, D.F., Capaldi T.N., Thomson K.D., Clark J., Puigdefàbregas C. & Fildani A., 2019. Tectonic  
924 and sediment provenance evolution of the South Eastern Pyrenean foreland basins during rift margin inversion  
925 and orogenic uplift. *Tectonophysics*, in press (this issue)

926 Oliva-Urcia B., Beamud E., Arenas C., Pueyo E.L., Garcès M., Soto R, Valero L. & Pérez-Rivarés F.J., 2019.  
927 Dating the northern deposits of the Ebro foreland basin; implications for the kinematics of the SW Pyrenean  
928 front. *Tectonophysics*, 765, 11-34 (this issue)

929 Parrish, R.R., Parrish, C.M. & Lasalle, S., 2018, Vein calcite dating reveals Pyrenean orogen as cause of Paleogene  
930 deformation in southern England: *Journal of the Geological Society*, 175, 425–442

931 Pavlis, T.L., Mason, K.A. 2017. The new world of 3D geologic mapping. *GSA Today*, 27

932 Pellerin J., Caumon G., Julio C., Mejia-Herrera P. & Botella A., 2015. Elements for measuring the complexity of  
933 3D structural models: Connectivity and geometry. *Computers & Geosciences*, 76, 130-140

934 Pipping, E., Kornhuber R., Rosenau M. & Oncken O., 2016. On the efficient and reliable numerical solution of  
935 rate-and-state friction problems. *Geophysical Journal International*, 204 (3), 1858-1866

936 Poblet, J. & Lisle, R. J., 2011. Kinematic evolution and structural styles of fold-and-thrust belts. In *Kinematic  
937 Evolution and Structural Styles of Fold-and-Thrust Belts* (eds J. Poblet & R. J. Lisle), pp. 1–24. Geological  
938 Society of London, Special Publication no. 349

939 Poppe, S et al., 2019. An Inside Perspective on Magma Intrusion: Quantifying 3D Displacement and Strain in  
940 Laboratory Experiments by Dynamic X-Ray Computed, *Frontiers in Earth Science* 7, 62

941 Reiners, P.W. & Brandon, M.T., 2006, Using thermochronology to understand orogenic erosion. *Annual Review  
942 of Earth and Planetary Sciences*, 34, 419-466.

943 Ritter M.C., Santimano T., Rosenau M., Leever K. & Oncken O., 2018. Sandbox rheometry: Co-evolution of stress  
944 and strain in Riedel– and Critical Wedge–experiments. *Tectonophysics*, 722, 400-409

945 Ritter, M. C., Rosenau, M. & Oncken O., 2018. Growing faults in the lab: Insights into the scale dependence of  
946 the fault zone evolution process. *Tectonics*, 37, 140–153.

947 Rivas C., Ortiz G., Alvarado P., Podesta M. & Martin A., 2019. Modern crustal seismicity in the northern Andean  
948 Precordillera, Argentina. *Tectonophysics*, 762, 144-158 (this issue)

949 Roma, M., Ferrer, O., McClay, K., Muñoz, J.A., Roca, E., Gratacós, O. & Cabello, P. 2018. Weld kinematics of  
950 syn-rift salt during basement-involved extension and subsequent inversion : Results from analogue models.  
951 *Geologica Acta*, 16 (4), 391-410.

952 Ronda G., Ghiglione M.C., Barberón V., Coutand I. & Tobal J., 2019. Mesozoic – Cenozoic evolution of the  
953 Southern Patagonian Andes fold and thrust belt (47°–48°S): Influence of the Rocas Verdes basin inversion and  
954 onset of Patagonian glaciations. *Tectonophysics*, 765, 83-101 (this issue)

955 Rosenau, M., Corbi F. & Dominguez S., 2017. Analogue earthquakes and seismic cycles: Experimental modelling  
956 across timescales, *Solid Earth*, 8, 3, 597-635

957 Roure F., Andriessen P., Callot J.-P., Ferket H., Gonzalez E., Guilhaumou N., Hardebol N., Lacombe O.,  
958 Malandain J., Mougín P., Muska K., Ortuno S., Sassi W., Swennen R. & Vilasi N., 2010, The use of paleo-

959 thermo-barometers and coupled thermal, fluid flow and pore fluid pressure modelling for hydrocarbon and  
960 reservoir prediction in fold-and-thrust belts. In “Hydrocarbons in Contractual Belts”, Geol. Soc. London,  
961 Spec. Publ., 348, 87-114

962 Rowan, M.G. 2014. Passive-margin salt basins: hyperextension, evaporite deposition, and salt tectonics. *Basin*  
963 *Research*, 26, 154-182.

964 Rowan, M.G. & Vendeville, B.C., 2006, Foldbelts with early salt withdrawal and diapirism: Physical model and  
965 examples from the northern Gulf of Mexico and the Flinders Ranges, Australia. *Marine and Petroleum*  
966 *Geology*, 23, 871-891.

967 Rudolf, M., Rosenau M., Ziegenhagen T., Ludwikowski V., Schucht T., Nagel H. & Oncken O., 2019. Smart  
968 Speed Imaging in Digital Image Correlation: Application to Seismotectonic Scale Modeling, *Frontiers in Earth*  
969 *Science* 6, 248

970 Ruh, J. B., Gerya, T. & Burg, J.P., 2013. High-resolution 3D numerical modeling of thrust wedges: Influence of  
971 décollement strength on transfer zones, *Geochem. Geophys. Geosyst.*, 14, 1131– 1155,

972 Ruh, J.B. 2019. Effets of fault-weakening processes on oblique intracontinental rifting and subsequent tectonic  
973 inversion. *American Journal of Science*, 319, 315-338.

974 Ruh, J. B., & Vergés, J., 2018. Effects of reactivated extensional basement faults on structural evolution of fold-  
975 and-thrust belts: Insights from numerical modelling applied to the Kopet Dagh Mountains. *Tectonophysics*,  
976 746, 493-511.

977 Saha, P., Bose, S. & Mandal N., 2016. Sandbox modelling of sequential thrusting in a mechanically two-layered  
978 system and its implications in fold-and-thrust belts, *Journal of Geodynamics*, 100, 104-114

979 Santimano, T., Rosenau M. & Oncken O., 2015. Intrinsic versus extrinsic variability of analogue sand-box  
980 experiments – Insights from statistical analysis of repeated accretionary sand wedge experiments, *Journal of*  
981 *Structural Geology*, 75, 80-100

982 Sanchez Nassif F., Canelo H., Davila F. & Ezpeleta M., 2019. Constraining erosion rates in thrust belts: Insights  
983 from kinematic modeling of the Argentine Precordillera, Jachal section. *Tectonophysics*, 758, 1-11 (this issue)

984 Santolaria, P., Casas, A.M. & Soto, R., 2015. Anisotropy of magnetic susceptibility as a proxy to assess internal  
985 deformation in diapirs: case study of the Naval salt wall (Southern Pyrenees). *Geophysical Journal*  
986 *International*, 202, 1207-1222.

987 Santolaria, P., Casas-Sainz, A.M., Soto, R., Pinto, V., Casas, A., 2014. The Naval diapir (southern Pyrenees):  
988 Geometry of a salt wall associated with thrusting at an oblique ramp. *Tectonophysics*, 637, 30-44.

989 Saura, E., Ardèvol i Oró, L., Teixell, A. & Vergés, J. 2016. Rising and falling diapirs, shifting depocenters, and  
990 flap overturning in the Cretaceous Sopeira and Sant Gervàs subbasins (Ribagorça Basin, southern Pyrenees).  
991 *Tectonics*, 35 (3), 638-662.

992 Savignano, E., Mazzoli, S., Arce, M., Franchini, M., Gautheron, C., Paolini, M. & Zattin, M., 2016. (Un) Coupled  
993 thrust belt-foreland deformation in the northern Patagonian Andes: new insights from the Esquel-Gastre sector  
994 (41° 30'–43° S). *Tectonics*, 35, 2636-2656

995 Schreurs, G. et al., 2016. Benchmarking analogue models of brittle thrust wedges, *Journal of Structural Geology*,  
996 92, 116-139

997 Searle M., Cornish S.B., Heard A., Charles J.-H. & Branch J., 2019. Structure of the Northern Moine thrust zone,  
998 Loch Eriboll, Scottish Caledonides. *Tectonophysics*, 752, 35-51 (this issue)

999 Simpson, G. 2011. Mechanics of non-critical fold-thrust belts based on finite element models. *Tectonophysics*  
1000 499, 142-155.

1001 Smit, J.H.W., Brun, J.P. & Sokoutis, D. 2003. Deformation of brittle-ductile thrust wedges in experiments and  
1002 nature. *Journal of Geophysical Research*, 108, B10, 2480

1003 Snidero M., Muñoz J.A., Carrera N., Butillé M., Mencos J., Motamedi H., Piryaei A. & Sàbat F., 2019. Temporal  
1004 evolution of the Darmadan salt diapir, eastern Fars region, Iran. *Tectonophysics*, 766, 115-130 (this issue)

1005 Soto, R., Beamud, E., Roca, E., Carola, E. & Almar, Y. 2016. Distinguishing the effect of diapir growth on  
1006 magnetic fabrics of syn-diapiric overburden rocks: Basque–Cantabrian basin, Northern Spain. *Terra nova*, 29,  
1007 191-201.

1008 Souloumiac, P., Maillot B. & Leroy Y.M., 2012. Bias due to side wall friction in sand box experiments. *Journal*  
1009 *of Structural Geology*, 35, 90-101

1010 Stüwe, K., White, L. & Brown, R., 1994. The influence of eroding topography on steady-state isotherms.  
1011 Application to fission track analysis. *Earth and Planetary Science Letters*, 124, 1-4, 63-74.

1012 Tavani S., Corradetti A., Sabbatino M., Morsalnejad, D. & Mazzoli S., 2018. The Meso-Cenozoic fracture pattern  
1013 of the Lurestan region, Iran: The role of rifting, convergence, and differential compaction in the development  
1014 of pre-orogenic oblique fractures in the Zagros Belt. *Tectonophysics*, 749, 104-119 (this issue)

1015 Tavani S., Storti F., Lacombe O., Corradetti A., Muñoz J.A. & Mazzoli S., 2015. A review of deformation pattern  
1016 templates in foreland basin systems and fold-and-thrust belts: implications for the state of stress in the frontal  
1017 regions of thrust wedges. *Earth-Science Reviews*, 141, 82-104

1018 Tavani, S., Granado, P., Corradetti, A., Girundo, M., Iannace, A., Arbués, P., Muñoz, J.A., Mazzoli, S. 2014.  
1019 Building a virtual outcrop, extracting geological information from it, and sharing the results in Google Earth  
1020 via OpenPlot and Photoscan: An example from the Khaviz Anticline (Iran). *Computers & Geosciences*, 63, 44-  
1021 53

1022 Ternois, S., Odlum, M., Ford, M., Pik, R., Stockli, D., Tibari, B., Vacherat A. & Bernard V.,  
1023 2019. Thermochronological evidence of early orogenesis, eastern Pyrenees,  
1024 France. *Tectonics*, 38, 1308– 1336.

1025 Tomek F., Vacek F., Žák J., Petronis M.S., Verner K. & Foucher M.S., 2019. Polykinematic foreland basins  
1026 initiated during orthogonal convergence and terminated by orogen-oblique strike-slip faulting: An example  
1027 from the northeastern Variscan belt. *Tectonophysics*, 766, 379-397 (this issue)

1028 Toeneboehn, K., Cooke, M.L., Bemis S.P. & Fendick A.M., 2019. Stereovision combined with particle tracking  
1029 velocimetry reveals advection and uplift within a restraining bend simulating the Denali Fault, *Frontiers in*  
1030 *Earth Sciences*, 6, 152

1031 Turrini C., Lacombe O. & Roure F., 2014. Present-day 3D structural architecture of the Pô Valley basin, northern  
1032 Italy. *Marine and Petroleum Geology*, 56, 266-289

1033 Turrini C., Angeloni P., Lacombe O., Ponton M. & Roure F., 2015. 3D seismotectonics in the Po Valley basin,  
1034 Northern Italy. *Tectonophysics*, 661, 156-179

1035 van Dinther, Y., Gerya, T., Dalguer, L., Corbi, F., Funiciello, F. & Mai, P., 2013. The seismic cycle at subduction  
1036 thrusts: 2. Dynamic implications of geodynamic simulations validated with laboratory models, *Journal of*  
1037 *Geophysical Research. Solid Earth*, 118(4), 1502–1525.

1038 Vitale S., Ciarcia S., Fedele L., D'Assisi Tramparulo F., 2019. The Ligurian oceanic successions in southern Italy:  
1039 The key to decrypting the first orogenic stages of the southern Apennines-Calabria chain system.  
1040 *Tectonophysics*, 750, 243-261 (this issue)

1041 von Hagke, C., Cederbom, C.E., Oncken, O., Stöckli, D.F., Rahn, M.K. & Schlunegger, F., 2012. Linking the  
1042 northern Alps with their foreland: The latest exhumation history resolved by low-temperature  
1043 thermochronology. *Tectonics* 31, TC5010.

1044 von Hagke, C., Oncken, O., Ortner, H., Cederbom, C.E. & Aichholzer, S., 2014. Late Miocene to present  
1045 deformation and erosion of the Central Alps — Evidence for steady state mountain building from  
1046 thermokinematic data. *Tectonophysics* 632, 250-260.

1047 von Hagke, C., & Malz, A., 2018. Triangle zones–Geometry, kinematics, mechanics, and the need for appreciation  
1048 of uncertainties. *Earth-science reviews*, 177, 24-42.

1049 von Hagke, C., Philippon, M., Avouac, J.-P. & Gurnis, M., 2016. Origin and time evolution of subduction polarity  
1050 reversal from plate kinematics of Southeast Asia, *Geology*, 44, 659-662

1051 Wagner, G.A. & Reimer, G.M., 1972. Fission track tectonics: the tectonic interpretation of fission track apatite  
1052 ages. *Earth and Planetary Science Letters* 14, 263-268.

1053 Wellmann, J. F. & Regenauer-Lieb K., 2012. Uncertainties have a meaning: Information entropy as a quality  
1054 measure for 3-D geological models. *Tectonophysics* 526, 207-216.

1055 Wellmann, F. & Caumon G., 2018. 3-D Structural geological models: Concepts, methods, and uncertainties.  
1056 *Advances in Geophysics*, 59, 1-121.

1057 Wolf, R.A., Farley, K.A. & Silver, L.T., 1996. Helium diffusion and low-temperature thermochronometry of  
1058 apatite. *Geochimica et Cosmochimica Acta* 60, 4231-4240.

1059 Yanites, B.J., Ehlers, T.A., Becker, J.K., Schnellmann, M. & Heuberger, S., 2013. High magnitude and rapid  
1060 incision from river capture: Rhine River, Switzerland. *Journal of Geophysical Research: Earth Surface* 118,  
1061 1060-1084.

1062