



## Tiefer See – a key site for lake sediment research in NE Germany

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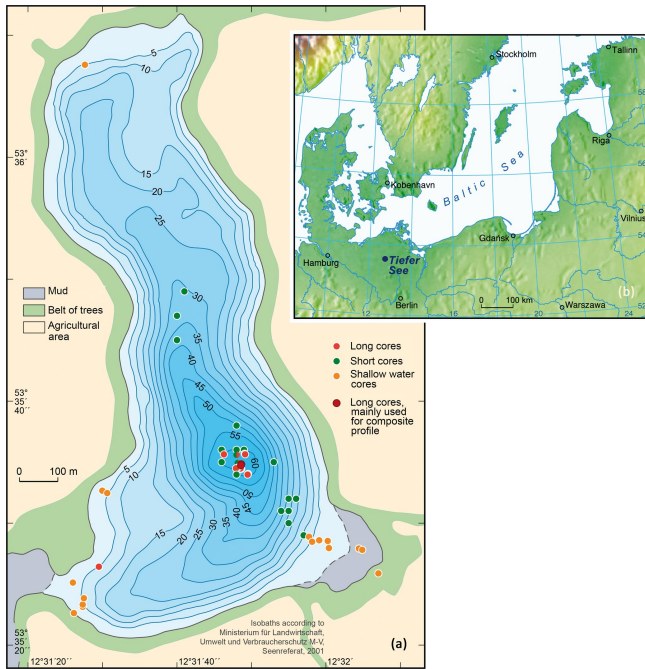
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**Abstract:** Tiefer See formed in a subglacial gully system at the end of the last glaciation in the northeast German lowlands. The lake has been selected as a focus site within the TERENO (Terrestrial Environmental Observatory) NE German observatory because it forms annual laminations (calcite varves) providing detailed information of past climate and environmental changes. Our research integrates palaeolimnology and limnology by combining high-resolution analyses of the sediment record with a comprehensive monitoring of the lake and its sedimentation processes since 2012. This allows evaluation of the observed effects of ongoing climate change in the context of the long-term history of the lake. The lacustrine sediment profile comprises the last 13 000 years and is dated by a multiple dating approach. The sedimentation is dominated by biochemical calcite formation and algal blooms. Detrital material from the catchment forms only a minor component even during times of increased human impact. Repeated changes between well-varved, poorly varved and homogeneous sediment intervals indicate that sedimentation processes in the lake are particularly sensitive to changes in lake circulation. The research at Tiefer See is embedded in ICLEA (<https://www.iclea.de>, last access: 2 August 2019) and BaltRap (<https://www.io-warnemuende.de/projekt/167/baltrap.html>, last access: 2 August 2019) projects.

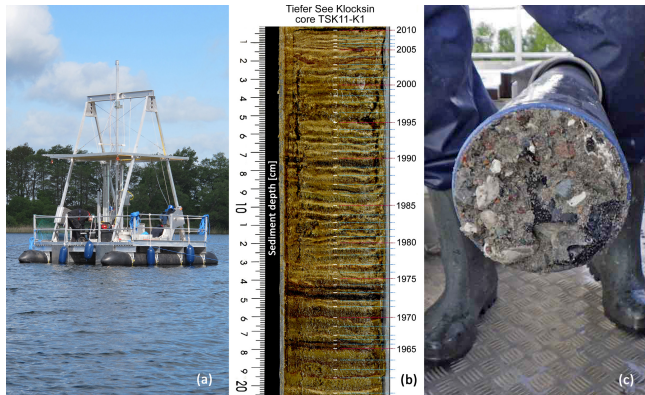
### 1 Geomorphology and Geology

Tiefer See is part of the Klocksin lake chain, a subglacial gully system in a morainic terrain located in the natural park Nossentiner/Schwinzer Heide at 53°35.5' N, 12°31.8' E. The Klocksin lake chain extends from NNE to SSW and includes the four lakes Flacher See (64.4 m a.s.l.), Tiefer See (62.9 m a.s.l.), Hofsee (62.7 m a.s.l.) and Bergsee (62.6 m a.s.l.). The catchment geology of the northern part of the lake chain including Tiefer See is formed by tills while the southern part consists of glaci-fluvial sands. Numerous er-

atics along the shoreline of Tiefer See indicate the proximity of terminal moraines which formed during the Pomeranian ice advance W2 at around 16 kyr BP. Today, Tiefer See is connected to Hofsee in the south while the connection to Flacher See in the north has been channelised in a tunnel after construction of a railway dam between the two lakes. Tiefer See has a surface area of ca. 0.75 km<sup>2</sup> and the catchment area is about 5.5 km<sup>2</sup>. With a maximum depth of 62 to 63 m Tiefer See is the deepest lake of the Klocksin lake chain (Fig. 1). It has no major inflow and outflow. The present-day lake water is mesotrophic and the circulation mode is

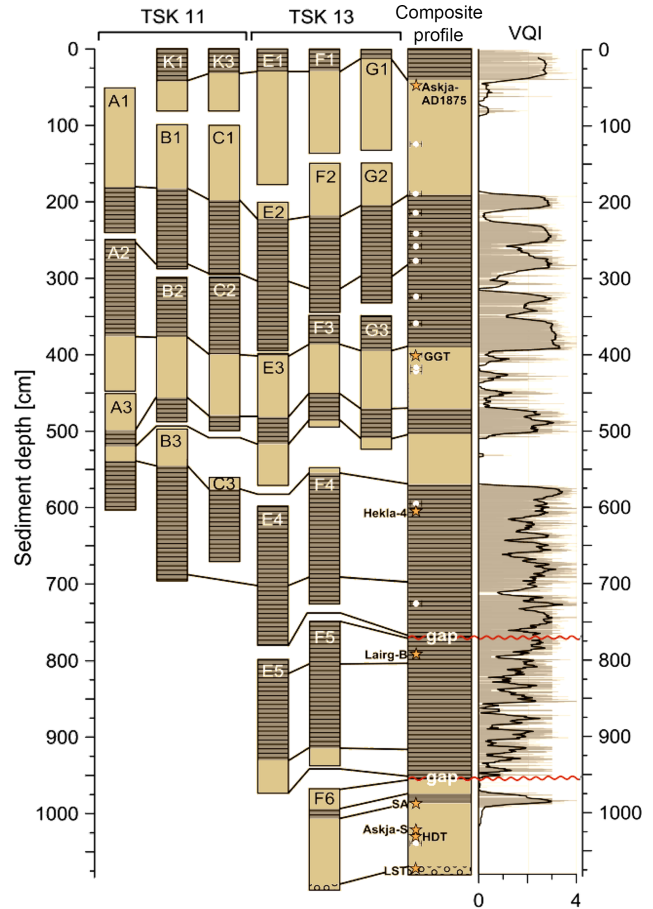


**Figure 1.** (a) Lake bathymetry and core locations (red: long cores; green: surface cores; orange: shallow water cores). (b) Tiefer See location.



**Figure 2.** (a) GFZ lake sediment coring device (<https://www.UWITEC.at>, last access: 2 August 2019); (b) subrecent varves; (c) basal gravel and sand.

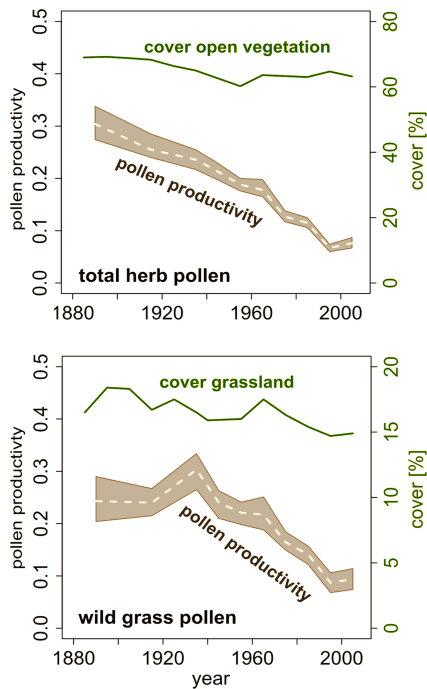
either mono- or dimictic, depending on the formation of a winter ice cover (Kienel et al., 2013). The climate is warm-temperate at the transition from oceanic to continental conditions. Mean monthly temperatures range from 0 °C in January to 17–18 °C in July with maxima up to 30 °C and minima down to –5 °C. Mean monthly precipitation varies between ca. 40 mm during winter and ca. 60 mm in summer with a mean annual precipitation of 560–570 mm. Although the catchment is mainly used for agriculture, the direct shoreline of the lake is covered by a fringe of trees without any buildings and roads.



**Figure 3.** Schematic profiles of Tiefer See sediment cores and the respective composite profile with well-varved (dark) and non-varved (light) sections; white dots indicate AMS 14C dates, and orange stars show positions of cryptotephras (data from Wulf et al., 2016).

## 2 The sediment profile

In addition to several surface sediment cores from the entirety of the lake basin, a total of nine overlapping sediment cores have been obtained from the deepest part of the lake (Fig. 1a) with the GFZ piston coring device (Fig. 2a) during three field campaigns in 2011, 2013 and 2019. A composite profile has been established from six cores which includes two minor gaps below 7.5 m depth in the early Holocene (Dräger et al., 2017). These gaps have been closed with the new cores taken in 2019. The composite profile is about 11 m long and depicts a main lithological change from basal sands and gravel (Fig. 2c) to massive fine-grained and calcite-rich lacustrine sediments at about 10.7 m depth. At 10.4 m sediment depth (Fig. 3) a rapid transition to organic-rich sediments occurred. The organic-rich sediments vary in structure from finely laminated to poorly laminated and massive. The finely laminated sediments represent three different micro-facies types of varves, (1) Ca-rhodochrosite varves, (2) or-

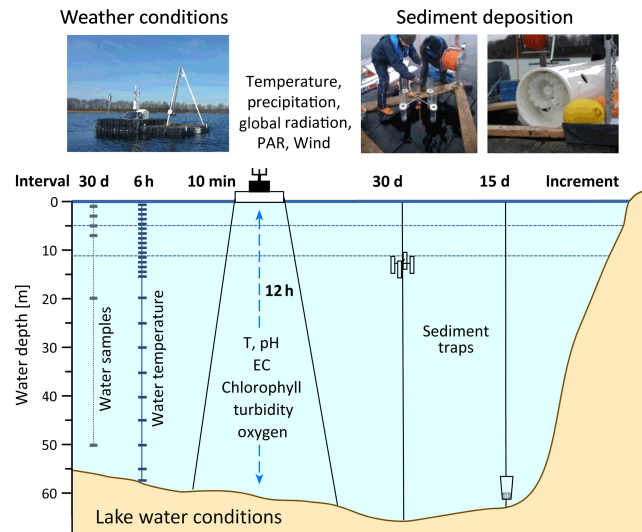


**Figure 4.** Changes in the cover of open vegetation and grasslands in the surroundings of Tiefer See, and changes in the mean pollen productivity of herbs and grasses since 1880 CE (after Theuerkauf et al., 2015).

ganic varves and (3) calcite varves (Dräger et al., 2017). The most recent period of varve preservation began in 1924 CE (Kienel et al., 2017) with the formation of calcite varves (Fig. 2b). The alternation of varved and poorly or non-varved sediment intervals is also reflected in stable carbon isotope data of organic matter which is interpreted as variations between predominantly anoxic and oxic lake bottom waters (Dräger et al., 2019). The frequency and duration of poorly varved intervals increased during the late Holocene and has been related to either climatic changes (increased windiness) or human impact in the catchment (Dräger et al., 2017).

### 3 The age–depth model

Dating is crucial for robust reconstruction of past climate and environment changes. Therefore, a multiple dating approach has been applied for the sediment record from Tiefer See including varve counting, radiocarbon dating, tephrochronology and  $^{10}\text{Be}$  analyses. Due to the minor gaps in the lower part of the sediment profile, a robust and continuous chronology could be established only for the last  $6030 \pm 85$  yrs BP (ca. 7.5 m sediment depth). Varves have been counted using a petrographic microscope and non-varved intervals have been interpolated using sedimentation rate estimates based on microstructural analyses on thin sections. The resulting varve chronology has been confirmed by 12 radiocarbon dates of organic macroremains and by three cryptotephra



**Figure 5.** Sketch of installed lake sediment monitoring equipment in Tiefer See.

horizons: (1) the Askja 1875 CE, (2) the Glen Garry ( $2088 \pm 122$  yrs BP) and (3) the Hekla 4 tephra ( $4293 \pm 43$  yrs BP). The identification of another five Holocene cryptotephra of Icelandic origin in the lower part of the sediment profile (Hässeldalen, Askja-S, Saksunarvatn, Lairg B; Wulf et al., 2016) provides a good estimate of the early Holocene age model.  $^{10}\text{Be}$  data further confirmed the chronology and allowed an independent synchronisation with the varved sediment record from Lake Czechowskie a few hundred kilometres further east (Czymzik et al., 2015, 2018). The base of the profile is dated into the late Allerød through the detection of glass shards from the Laacher See eruption (12 880 yrs BP).

### 4 Pollen analyses

The seasonal resolution of the sediments allows for tracing historical land-use changes in great detail (Theuerkauf et al., 2015). Varve dating enabled us to accurately estimate pollen deposition of herbs and trees over the past 150 years, which has then been compared to the contemporaneous land cover documented in archival data (Fig. 4; Theuerkauf et al., 2015). Interestingly, the strong decline in herb pollen deposition since the 1950s did not reflect a change in the cover of arable land and grassland. Instead, pollen deposition of herbs and grasses was controlled by changes in pollen productivity in response to intensified land management. Today, grassland is mown earlier and more frequently, giving the grasses less time to flower and produce pollen than previously. This study demonstrated that the modern vegetation is not always a reliable reference for the past. Therefore, the ROPES approach without pollen productivity as a parameter has been developed for the long pollen record of the last 6000 years from Tiefer See (Theuerkauf and Couwenberg, 2018).

Pollen analyses in combination with sedimentological parameters as well as macrofossil and microfossil analysis from 20 sediment cores along the shore of Tiefer See and in adjacent peatlands were applied for Holocene lake level reconstruction. We found Holocene lake level fluctuations with an amplitude of about 10 m, which exceeds by far recently observed variations. Based on radiocarbon age modelling, the lowest lake levels occurred during the early Holocene and maxima in the medieval period. These changes are only partly related to climatic changes but we also observe a link to changes in forest cover. In particular during medieval times, large-scale forest opening likely increased groundwater formation and thus contributed to lake level increases.

## 5 Lake and sedimentation monitoring

A comprehensive lake monitoring instrumentation (Fig. 5) has been installed on Tiefer See since 2012 in order to decipher the response of the lake to ongoing climate change (Heinrich et al., 2018) and to study in detail the processes that transform weather signal into the sediment record. The instrumentation includes a permanent weather station on the lake measuring temperature, relative humidity, solar radiation, rainfall and wind speed in 10 min intervals. Attached to the weather station is an automatic probe that measures water parameters from the epilimnion to the lake bottom every 12 h. These measurements are complemented by temperature loggers permanently installed in 1 m apart down to 15 m water depth and 5 m apart down to 55 m depth. In addition, lake water is sampled monthly at water depths of 1, 3, 5, 7, 20 and 50 m for chemical analyses and stable oxygen and deuterium isotopes. Sediments are trapped in monthly intervals in two four-cylinder traps, one directly below the productive zone in the epilimnion and one in the hypolimnion. In the hypolimnion is a sequential trap with 15 d sample intervals installed. First results of the water isotope analyses revealed a seasonal 1.5‰ enrichment of  $\delta^{18}\text{O}$  proving the expected evaporative loss of surface water of up to 10% during the summer, which corresponds to the observed lake level drop of 50 cm. Interestingly, the deep water also generally has more positive  $\delta^{18}\text{O}$  values than the annual rainwater mean collected on the weather station. The isotopic enrichment of the deep water indicates a long water residence time in the lake, which is likely due to the lack of major in- and outflows and low groundwater inflow. The intensive monitoring also allowed the investigation of the effects of the 2018 summer heatwave on the lake hydrology (Heinrich et al., 2019).

**Data availability.** No data sets were used in this article.

**Author contributions.** AB, MJS and MT wrote the text. SP, BB, MJS and MT participated in the fieldwork and took the photographs in Fig. 2. All authors contributed to the discussion and interpreta-

tion of the presented research results. Additional information and supplementary data about this project are published (Czymzik et al., 2015, 2018; Dräger et al., 2017, 2019; Heinrich et al., 2018, 2019; Kienel et al., 2013, 2017; Theuerkauf et al., 2015, Theuerkauf and Couwenberg, 2018; Wulf et al., 2016).

**Competing interests.** The authors declare that they have no conflict of interest.

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