

# The new Italian Seismic Hazard Model (MPS19)

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## Abstract

We describe the main structure and outcomes of the new probabilistic seismic hazard model for Italy, MPS19 [Modello di Pericolosità Sismica, 2019]. Besides to outline the probabilistic framework adopted, the multitude of new data that have been made available after the preparation of the previous MPS04, and the set of earthquake rate and ground motion models used, we give particular emphasis to the main novelties of the modeling and the MPS19 outcomes. Specifically, we (i) introduce a novel approach to estimate and to visualize the epistemic uncertainty over the whole country; (ii) assign weights to each model components (earthquake rate and ground motion models) according to a quantitative testing phase and structured experts' elicitation sessions; (iii) test (retrospectively) the MPS19 outcomes with the horizontal peak ground acceleration observed in the last decades, and the macroseismic intensities of the last centuries; (iv) introduce a pioneering approach to build MPS19\_cluster, which accounts for the effect of earthquakes that have been removed by declustering. Finally, to make the interpretation of MPS19 outcomes easier for a wide range of possible stakeholders, we represent the final result also in terms of probability to exceed 0.15 g in 50 years.

Keywords: Seismic hazard; Epistemic uncertainty; Ensemble modeling; Probabilistic methods; Italy.

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## 1. Introduction

In 2015, the Seismic Hazard Center (Centro di Pericolosità Sismica, CPS) of the Istituto Nazionale di Geofisica e Vulcanologia (INGV) was commissioned by the Italian Civil Protection Department to engage and coordinate the national scientific community with the aim of elaborating a new seismic hazard model for Italy (Modello di Pericolosità Sismica, MPS19 hereafter). MPS19 takes into account the large amount of new data and methods that have been published after the last release of MPS [MPS04; <http://zonesismiche.mi.ingv.it>; Stucchi et al., 2011], and it is developed according to the most advanced international standards in seismic hazard [Gerstenberger et al., 2020].

MPS19 will be taken into consideration to update the Italian building code [NTC, 2018], currently based on MPS04 model. For this reason, since the beginning CPS representatives met the national earthquake engineering community to discuss the most proper outcomes of the new seismic hazard model for the building code purpose. Specifically, the model has to consider only declustered seismicity, to cover the whole national territory, the reference soil is rock (i.e. EC8 site category A,  $V_{s,30} > 800\text{m/s}$ ); the hazard has to be expressed in terms of mean horizontal peak ground acceleration (PGA), peak ground velocity (PGV), and peak ground displacement (PGD), spectral acceleration (SA), velocity, displacement and macroseismic intensity (the first release described here contains only PGA and SA results); the oscillatory period of the spectral ordinates has to be in the range between 0.05 and 4 seconds, and the return periods from 30 to 5000 years.

The definition of MPS19 has engaged about one hundred of researchers from Institutes and Universities through an open call to the Italian scientific community (see Appendix A in Supplementary Material). Such a complex endeavor has been completed in 2019, after a successful interaction with a participatory review panel, which gave a final positive approval to MPS19. At the time of writing this paper, applications for building code purposes, or for any other kind of risk analysis are still under discussion. In this paper, we summarize the scientific background of MPS19, the overall structure of the model, the novelties introduced, and the main outcomes. The specific description of the components is (or will be) covered in different papers.

MPS19 is based on a Probabilistic Seismic Hazard Analysis [PSHA, Cornell, 1968], i.e., MPS19 describes in a probabilistic way the forecast of a variety of ground motion intensity measures (IMs; such as PGA and other quantities) on the Italian territory, and it is rooted in open and transparent procedures that guarantee completely reproducible outcomes; specifically, we use the open-source OpenQuake engine platform [Pagani et al., 2014] and all data will be available through a dedicated website. Before getting into the details, there are some features and novelties introduced in MPS19 that deserve to be highlighted, as they mark a difference with previous similar activities.

- MPS19 honors the hazard/risk separation principle [Jordan et al., 2014], i.e., the hazard is evaluated without being conditioned by the effects on the risk. In other words, MPS19 is only focused on providing the forecast of the IM (i.e., the likelihood of observing different levels of IM) in a time window of 50 years, describing at best the historical seismicity and ground motion; purposely, we do not evaluate the impact on the risk, and/or introduce any additional subjective “precautionary” choices that, according to the hazard/risk separation principle, should be introduced only at the risk level.
- MPS19 relies on a clear and coherent probabilistic framework [Marzocchi and Jordan, 2014; 2017; 2018], which allows a univocal distinction of aleatory variability and epistemic uncertainty, a meaningful testing phase (to find possible ontological errors), and a clear interpretation of the mean and percentiles of the hazard outcomes. Then, we use an ensemble modeling strategy to get a complete description of the seismic hazard [Marzocchi et al., 2015].
- MPS19 includes an extensive testing phase to assign weights to each component of the model, to verify the reliability of the earthquake rate models (ERMs hereafter) and of ground motion models (GMMs hereafter), and of the hazard outcomes in terms of PGA and macroseismic intensity at several sites and, on average, over the whole territory.
- Large aftershocks and foreshocks may affect significantly the seismic hazard and risk, but they are usually removed by classical declustering techniques to get a set of time-independent (Poissonian distributed) earthquakes [Cornell, 1968]. For this reason, we also estimate a version of our model introducing a pioneering attempt to correct the classical PSHA model to account for earthquake clustering [Iervolino et al., 2012; 2019; Marzocchi and Taroni, 2014].

The roadmap followed to build MPS19 consists of few basic steps which will be outlined in this paper: i) define a coherent probabilistic framework to fully describe the uncertainties (ensemble modeling), to check the consistency and relative predictive skill of each component of MPS19, and to test the consistency of the final MPS19 outcome (section 2); ii) improve the quality and accuracy of the input data (e.g. seismic catalog, seismotectonic zonation, etc.; section 3); iii) build ERMs based on these new input data, and test their consistency and relative predictive skill (section 4); iv) select the most proper GMMs for Italy, and estimate their relative predictive skill (section 5); v) weigh ERMs and GMMs through a combination of formal experts’ judgment elicitations with the results of the testing phase (section 6); vi) show the outcomes of MPS19 and the results of the testing phase (section 7);

vii) introduce pilot procedures to build a national seismic hazard model which takes into account the effects of aftershocks and foreshocks that have been removed by the declustering technique (MPS19\_cluster; section 8). Finally, in section 9, we summarize the main novelties of the model, the open questions, and future developments.

## 2. The probabilistic framework and the testing phase

PSHA combines ERMs and GMMs to estimate the hazard curve in each site, i.e., the probabilistic function of the IM in a given time interval. To acknowledge the role of pervasive uncertainties and their importance, SSHAC [1997] states that PSHA should “*represent the center, the body, and the range of technical interpretations that the larger technical community would have if they were to conduct the study.*”

Handling uncertainties of different kinds is far from being trivial. For example, in PSHA it is common to separate aleatory variability and epistemic uncertainty even though a satisfactory definition has never been provided [SSHAC, 1997; NRC, 1997]. In MPS19 we adopt a probabilistic framework which makes this separation possible and external to the model [Marzocchi and Jordan, 2014; 2017]. The framework is rooted in a univocal hierarchy of uncertainties which allows a clear interpretation of what a probability is (e.g., frequency or degree of belief), a meaningful testing of the PSHA model, and a clear interpretation of the mean hazard and of the percentiles that have been matter of intense discussion [Musson, 2005; McGuire et al., 2005; Abrahamson and Bommer, 2005]. The details of the framework can be found in Marzocchi and Jordan [2014; 2017]. Here we just summarize the most important feature in the PSHA context: the probability of exceedance (PoE), which characterizes the hazard in one site, is the frequency of an infinite sequence of yearly IM exceedances that are considered exchangeable (i.e., the sequence of exceedances can be reshuffled without losing information); the true frequency is the aleatory variability of the data-generating process and it is unknown (we could know it only if we have the “true” hazard model). We estimate this unknown aleatory variability using different models (or different parametrizations of the same model); the dispersion of these models describes the epistemic uncertainty and allows bounding where the true frequency (aleatory variability) should be. For example, every single model gives an estimation of the PoE for one selected IM, and the PoE of all models is described by a distribution; the center of this distribution is the best guess of the aleatory variability, and the dispersion mimics the epistemic uncertainty (an example of this distribution will be shown in section 6: “Weighing and merging all MPS19 components”); if the observed frequency of future data does not belong to this distribution, it means that we have found an ontological error, i.e., using the Rumsfeld’s words, “*the presence of unknown unknowns*”; this test is what Marzocchi and Jordan [2014] define as model validation (see Figure 4 in Marzocchi and Jordan [2017], for a visual representation of the problem).

This probabilistic framework allows us to use a meaningful testing phase to check the reliability of MPS19. In past common practice, the reliability of a seismic hazard model is based on the consensus of the scientific community on the model’s outcomes [Marzocchi and Zechar, 2011], even though a full consensus is impossible, and the definition of “scientific consensus” is intrinsically fuzzy [Aspinall, 2010]. The testing phase is a more objective approach that may contribute to check the reliability of a model. Owing to the length of the time window considered (50 years), seismologists cannot usually validate models using independent data, even though some efforts in this direction are ongoing, with some limitations, and usually using earthquakes from small to moderate magnitudes [e.g. Mak et al., 2014; Brooks et al, 2017; Mousavi and Beroza, 2018]. Most of the time, as for MPS19, only past data, which have been used more or less directly to build the hazard model, can be considered for testing [e.g. Albarello and D’Amico, 2008; Stirling and Gerstenberger, 2010; Mak and Schorlemmer, 2016]. Hence, instead of using the term “validation”, we prefer to use the less ambitious term “consistency” of the model’s output with data. In other words, the model should be consistent with the data that were available when the model was built. At the same time, the adopted framework outlines a clear pathway to validate the model with future observations.

We test the MPS19 outcomes, ERMs and GMMs to check their consistency with the available data, and their relative forecasting skill. To this purpose, we adopt some tests that have been developed in the Collaboratory for the Study of Earthquake Predictability (CSEP) framework [Jordan, 2006; Zechar et al., 2010a; Schorlemmer et al., 2018]. Specifically, we check the consistency of a model through the *N*- and *S*-test [Schorlemmer et al., 2007; Zechar et al., 2010b], which have been modified to account for the epistemic uncertainty [Marzocchi and Jordan, 2018]. The *N*-test compares the total number of earthquakes forecast with the observed number over the entire region. The *S*-test is similar to a classical likelihood test, but it is applied to a normalized forecast isolating the spatial

component of the forecast. A model that passes both tests explains satisfactorily the overall number of target earthquakes and their location.

The forecasting skill of the models can be quantified through one or more scoring schemes. We argue that the wide proliferation of scoring schemes in scientific literature is due to the fact that a perfect single scoring system does not exist, but each one of them ranks models according to a specific metric, or loss function. Here we use two proper scoring methods [Scherbaum et al., 2009; Zechar and Zhuang, 2014] that explore different aspects of the models: the parimutuel gambling score which tends to reward a model that produces forecasts often close to the observations, and the Log-likelihood score which rewards models that never failed badly (one model that predicts perfectly all the events but one, at which it assigns zero probability, will have the worst score possible).

To sum up, for the testing phase, we:

- test the consistency of every single ERM, comparing its rates with the observed seismicity rates of the last centuries (taken from the CPTI15 catalog [Rovida et al., 2016; 2020]). An ERM that does not pass both consistency tests is not included in MPS19.
- score ERMs and GMMs, comparing all models against each other using different metrics to rank them according to their retrospective forecasting (hereafter hindcasting) skill.
- evaluate the overall consistency of MPS19 with the past accelerometric data that are available in Italy, accounting for the epistemic uncertainty. In other words, we test if the numbers of exceedances recorded in the last decades are in agreement with what expected by MPS19.
- test MPS19 using macroseismic intensity data. Specifically, we test if the number of the macroseismic intensities 6 MCS (Mercalli–Cancani–Sieberg scale [Sieberg, 1932]) and larger observed in the last centuries is consistent with the MPS19 outcomes.

A full description of the tests is beyond the scope of this paper and has been thoroughly discussed in Gneiting and Raftery [2007], Schorlemmer et al. [2007], Zechar et al. [2010b], and Marzocchi and Jordan [2018]. In this paper we just report the main results of the testing phase.

### 3. The input data

MPS19 takes advantage of a collection of new datasets (see Table 1) that have been used to build and test ERMs and GMMs. They include, among others, an updated historical-instrumental earthquake catalog [CPTI15 v. 1.5, Rovida et al., 2016; 2020], an updated database of seismogenic sources [DISS 3.2.1, DISS Working Group, 2018], a new harmonized GPS velocity model for the Mediterranean area [Devoti et al., 2017], a new flatfile of strong motion recordings for the moment magnitude  $M_w \geq 4.0$  earthquakes for the selection and test of the GMMs [Lanzano et al., 2019; see Section 5] and two assessments of the focal mechanism of expected earthquakes [Roselli et al., 2018; Pondrelli et al., 2020]. To be used in MPS19, all datasets cover the whole Italian territory, with the exception of those referred to volcanic areas. Although it was not prescriptive, these datasets have been used by several modelers to build their ERMs.

N.	Dataset	Description/notes	References
1	Earthquake Catalog CPTI15 v 1.5	Parametric catalog for the Italian area and the period 1000-2014, $M_w \geq 4.0$	Rovida et al. [2016; 2020]
2	Earthquake catalog for the Mt Etna volcanic area	Input for CPTI15 and for the earthquake rate model of the Etna area (added to all ERMs)	Azzaro et al. [2017]
3	Instrumental catalog 1981-2015	Input for CPTI15 ( $M_w \geq 4.0$ ) and for some ERMs for low magnitudes	Gasperini et al. [2016]; Lolli et al. [2020]
4	Statistical completeness of CPTI15	Used only in MA5 ERM	Rotondi and Garavaglia [2002], (method)
5	Focal mechanism determination and deformation model	Used in MA4 and MA5 ERM	Pondrelli et al. [2020]

6	GPS velocity model for the Euro-Mediterranean Area	Main input for MG1 ERM	Devoti et al. [2017]
7	Database of Italian seismogenic sources DISS 3.2.1	Main input for MAs and MFs ERMs	DISS Working Group [2018]
8	Strong motion data flatfile	Input for the GMMs selection (see Section 5)	Lanzano et al. [2019]
9	Conversion between ground motion parameters and macroseismic intensity	Used in the MPS19 testing phase (see Section 7)	Gomez Capera et al. [2020]
10	Forecast of focal mechanism of future earthquakes	Used in MSs, MF2, and MG1	Roselli et al. [2018]

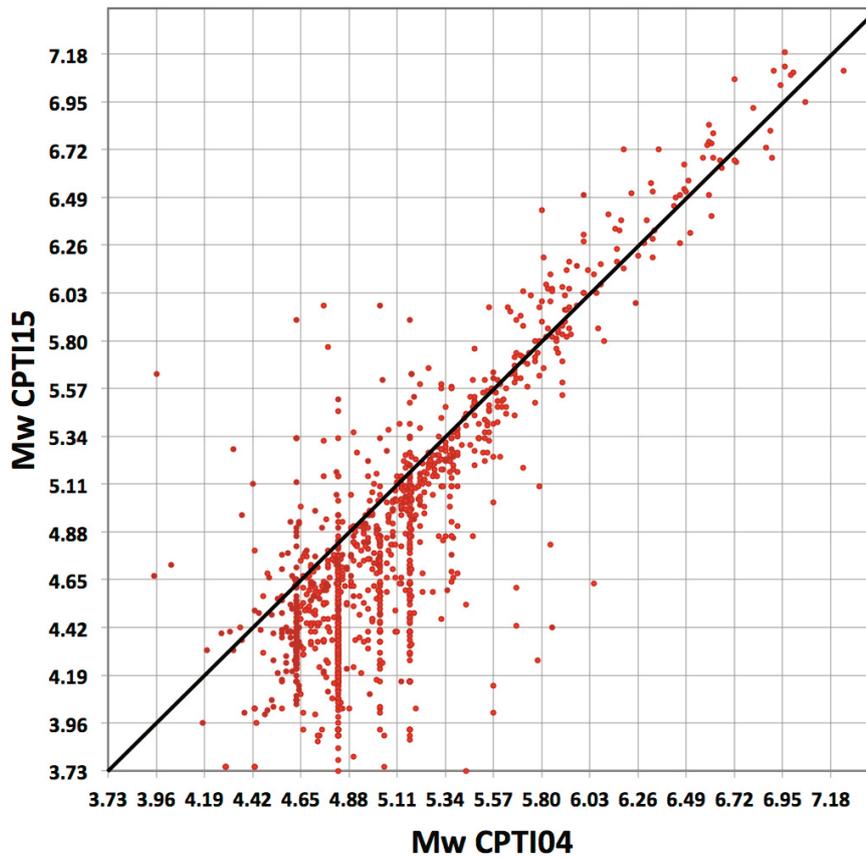
**Table 1.** Datasets, with references, used as a common input for MPS19.

One important dataset is the CPTI15 catalog, which represents the common input to almost all ERMs, but those that use strain rate field derived from geodesy as main input. CPTI15 contains the parameters of more than 4500 earthquakes, from 1000 CE to 2014, determined with homogeneous criteria merging updated macroseismic intensity data of past earthquakes and instrumental determinations of modern ones, as described in Rovida et al. [2020]. Owing to major improvements with respect to previous versions, we argue that the updated catalog turns out to be one of the most impacting factors to explain the differences of MPS19 with previous Italian PSHA models. Among the major improvements, macroseismic magnitudes (which are meant to estimate the moment magnitude) derived only from epicentral intensity are replaced in CPTI15 with magnitudes that are obtained by a new and updated intensity data distribution. The novelties of the catalog lead to a magnitude-frequency distribution that is more realistic than in previous versions of the catalog (i.e., the Gutenberg-Richter law has the b-value close to 1), and to the removal of spurious temporal patterns which characterize the time evolution of the seismicity in previous versions of the catalog. Figure 1 shows how the earthquakes magnitudes are changed between CPTI04 ([Gruppo di Lavoro, 2004] adopted in MPS04) and CPTI15; in particular, a substantial magnitude reduction is evident for the magnitude range between 4.5 and 5.8.

Furthermore, we provide additional pieces of information which have been used by several ERM modelers: i) a model for the maximum magnitude [Visini et al., 2019], ii) a determination of the seismogenic layer and of the hypocentral depths [D’Amico et al., 2019], and iii) the completeness intervals of the CPTI15 catalog and its declustered version [Meletti et al., 2019]. For the purpose of MPS19, we estimate the maximum magnitude in 18 macro-areas with homogenous tectonic features. The maximum magnitude is evaluated by comparing the observed magnitudes in the catalog (spanning approximately 1000 years) with the magnitudes calculated from geological information in the DISS 3.2.1. Within each macro-area, we initially define the maximum observed magnitude as the magnitude of the strongest event in the CPTI15 catalog, considering a minimum level of uncertainty for the magnitude evaluation of 0.2 and 0.3, respectively for the instrumental (post 1980 AD) and historical (pre 1980 AD) earthquakes. This magnitude is then compared to the magnitudes of the composite seismogenic sources of the DISS 3.2.1, and the largest between the two values is taken as representative of the maximum magnitude of each macro-area. Then, according to the procedure adopted in the European Seismic Hazard Model ESHM13 [Woessner et al., 2015], we define the lower threshold for the maximum magnitude as 6.5 for active shallow crustal regions (ASCR), 6.0 for stable continental regions (SCR), and 5.6 for Etna volcanic area. In lack of better qualitative and quantitative data coverage, we define a second and highest value of maximum magnitude, adding a uniform magnitude increment of 0.3 (except for the Etna area where we defined it to be 0), to account for epistemic uncertainty.

The same macro-areas defined for the maximum magnitude are used to assess the depth of the seismogenic layer. Following Boncio et al. [2009], in each macro-area we first calculate the depth distribution of the earthquakes in the instrumental catalog 1981-2015 (dataset 3 in Table 1; Lolli et al. [2020]). Then, for the events within the 5th and 95th percentiles, we evaluate the parameters of both a unimodal and a bimodal distributions, selecting the best

one according to the Akaike Information Criterion, AIC [Akaike, 1974]. Finally, we check the stability of the results using different magnitude thresholds ( $M_w \geq 2.0, 2.5, 3.0, 4.0$ ).



**Figure 1.** Plot of  $M_w$  values in CPTI04 and CPTI15 catalogs.

For seismic hazard analysis purposes, we distributed a version of the catalog in which every earthquake is marked as mainshock or aftershock, with the starting years of the complete periods for its magnitude according to the two approaches described below, and the completeness region it belongs to. Since in PSHA an earthquake catalog is declustered to get the mainshocks distributed in time according to Poisson [Lomnitz, 1966], we use GK74 algorithm [Gardner and Knopoff, 1974] that is the most effective to achieve this goal [van Stiphout et al., 2011; Luen and Stark, 2012]. The use of the general GK74 technique is justified by a recent empirical analysis of worldwide seismic sequences [Stallone and Marzocchi, 2019], which shows that earthquake clustering is similar in different crustal tectonic regions. Independently from the declustering technique used, we underline that its effect on the seismic hazard could be significantly mitigated adding a suitable correction for declustering. We discuss this point in section 8.

As regards the completeness of the catalog we use the same approaches adopted in MPS04 [MPS Working Group, 2004, Stucchi et al., 2011], which are based on the historical and statistical procedures described in Stucchi et al. [2004] and Albarello et al. [2001], respectively. To match the new features of CPTI15 with the previous historical completeness estimates, we i) redefine the  $M_w$  bins according to the new empirical conversion relation between epicentral intensity and magnitude used in CPTI15 [Rovida et al., 2016; 2020]; ii) evaluate the completeness intervals for epicentral intensity  $< 6$ , because such intensities were not included in the catalog used for MPS04; iii) revise the geometry of the regions with uniform completeness, and add one more that covers the off-shore areas. As in MPS04, we assess completeness for  $M_w$  bins of  $0.23 M_w$  units, to avoid the oversampling of some  $M_w$  intervals that contain values derived from the conversion of more than one discrete epicentral intensity values [see Stucchi et al., 2011]. As regards the statistical completeness, we apply the statistical analysis of Albarello et al. [2001] in the same regions

defined for the historical completeness with bins of 0.46  $M_w$  units to enhance the stability of the results. In this way we consider in the same class integer and intermediate values of the macroseismic intensity (e.g. epicentral intensity 6-7 together with 6). The results are then applied to bins of 0.23  $M_w$  units, consistently with the historical completeness estimates.

#### 4. The ERMs

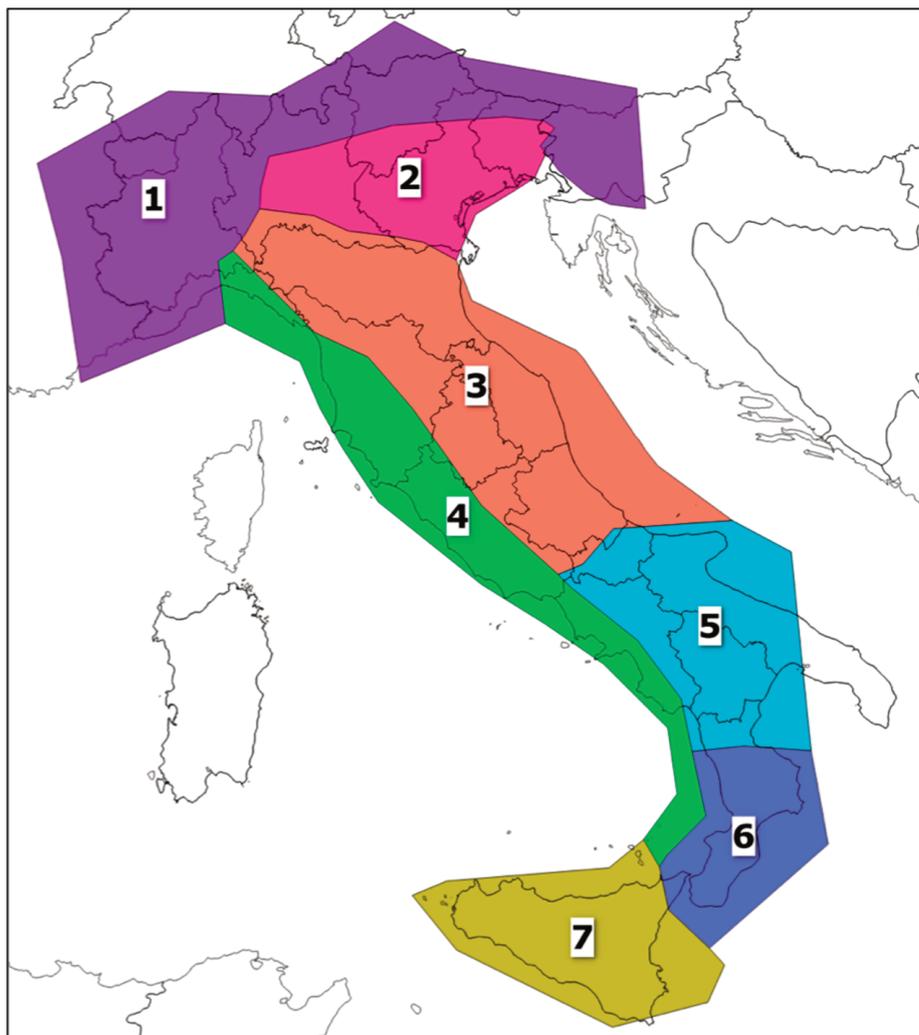
The researchers who answered to the open call spontaneously organized themselves in working groups and provided 11 independent ERMs for the whole Italian territory, and one model for the Etna volcano. An ERM forecasts the earthquake rate for different magnitudes, locations, geometry and kinematics. The models were built independently on purpose, to minimize the possibility that single models would be adjusted towards the mean consensus value, which unavoidably leads to underestimate the real uncertainty. The 11 ERMs are grouped according to the main typology of seismogenic sources: 5 area source ERMs (MA1 to MA5) based on the definition of areal sources for which seismicity rates are defined; 2 smoothed, zonation-free ERMs (MS1, MS2) that model the spatial density distributions of earthquakes in seismic catalogs; 2 fault sources plus background ERMs (MF1, MF2), based on the identification of seismogenic sources using tectonic and geophysical data; and 2 geodetically-derived, zonation-free ERMs (MG1, MG2) that model seismicity rates derived from the geodetic strain rate field. The last two models, which are mostly based on ground deformation, provide a different and independent level of information with the respect to the other ERMs because they are almost completely independent from CPTI15.

In table 2 we show the ERMs proposed and their main features. As any part of MPS19, the models have to be reproducible, and so modelers have been encouraged to publish the description of their model in scientific publications. Eventually, modelers provided codes, databases, and the outcomes of their models to make the process fully reproducible. Each ERM describes the aleatory variability of the data-generating (earthquake occurrence) process, and it explores its own epistemic uncertainty, i.e., how the model's outcome varies as a function of the variability of the parameters of the model itself, through a dedicated logic tree. The epistemic uncertainty of the final model includes the uncertainty of each ERM and the dispersion among the ERMs. A brief description of the methodology used to build each ERM and its main references are given in Appendix C of the Supplementary Material.

Model structure	Model code	Input data for earthquakes rate	Epistemic uncertainty explored	Number of branches
Seismotectonic zonation	MA1	CPTI15	Historical/statistical completeness, $M_{max}$ , GR/observed rates	8
	MA2	CPTI15	Historical/statistical completeness, Anchoring/partitioning from Macroarea to source zones	8
	MA3	CPTI15	Historical/statistical completeness, Anchoring/partitioning from Macroarea to source zones	8
	MA4	CPTI15	Historical/statistical completeness, $M_{max}$ , Anchoring/partitioning from Macroarea to source zones	20
	MA5	CPTI15		1
Fault (3D geom) + background (grid)	MF1	Slip rates and composite seismogenic source from DISS geometries	Slip rate (min-mean-max)	3
Fault + background (grid)	MF2	CPTI15, Slip rates and individual seismogenic source from DISS geometries	Historical/statistical completeness, Percentage of smoothed versus faults	6

Smoothed seismicity (grid)	MS1	CPTI15 + INGV seismic bulletin	Historical/statistical completeness, Percentage of fixed versus adaptative ratio	6
	MS2	CPTI15	Historical/statistical completeness, truncated/no-truncated GR, Isotropic/anisotropic	8
Geodetics (grid)	MG1	GPS	Modeling strain rate to seismic moment, seismogenic thickness	14
	MG2	GPS and SHmax	Smoothing factor, coupling coefficient, coupled thickness	12

**Table 2.** Main features of the ERMs; kind of seismogenic sources, input data, uncertainties explored, number of branches of an internal logic tree.



**Figure 2.** Macro-areas in which the testing phase is carried out. See text for more details.

Each ERM is deemed appropriate to be included in MPS19 if it i) passes consistency tests described in Section 2, and ii) is considered physically plausible from a group of independent experts through a formal experts' elicitation, which is described in Section 6. At first, we carry out statistical tests on a national scale, to define if an ERM is appropriate to be included in MPS19, and to measure the relative hindcasting skill which is used to weigh each single ERM (see Section 6). Following a suggestion of the participatory review panel (see Appendixes A and B in the Supplementary Material) we carry out the consistency tests of ERMs also on regions identified by a group of experts according to the following criteria: separate regions with high and low number of geological data, earthquakes, and geodetic stations; minimize the overlapping of areas with different completeness; keep the areas as large as possible avoiding marked tectonics inhomogeneity; minimize the off-shore part included in the regions. Following these criteria, we define 7 macro-areas (Figure 2).

Testing ERMs in each region allows us to check if their hindcasting skill varies with space and/or with the number of input data. If an ERM does not pass both consistency tests in one region, its overall earthquake rate is replaced in that subregion by the weighted average of the other consistent ERMs, preserving the geometries, depth, and focal mechanisms of the model. All models of Table 2 passed both consistency tests at national scale, hence they are deemed to be included in MPS19. However, as shown in Table 3, some of them failed to describe satisfactorily the data in some macro-areas of Figure 2.

Model	Macro-area
MA1	None
MA2	7
MA3	7
MA4	None
MA5	7
MF1	1,6
MF2	None
MS1	None
MS2	4
MG1	4
MG2	3,4,6

**Table 3.** Results of tests conducted on ERMs on 7 macro-areas. For each model, in the second column we list the number of macro-areas (shown in Figure 2) in which it does not pass the test; “None” means that the model passes the test in all macro-areas.

ERMs in Table 2 are defined for the entire Italian territory (excluding Etna area), and do not include seismogenic sources external to the area covered by CPTI15. For this reason, we integrate each ERM with: a) an ERM for the Etna area [adapted from Azzaro et al., 2017 and Peruzza et al., 2017]; b) an ERM for deep seismicity related to the slab activity in the southern Tyrrhenian Sea [adapted from Maesano et al., 2017] and; c) an ERM which models seismogenic sources located out of the area of the CPTI15 catalog [adapted from Woessner et al., 2015]. A brief description of the methodology used to build these 3 additional ERMs is given in Appendix D of the Supplementary Material.

Figure 3 maps the mean rates of occurrence of earthquakes with  $M_w$  larger or equal than 4.5 of the 11 national ERMs considered in MPS19 and Figure 4 shows the epistemic uncertainty of the rates. The mean rates of the 11 ERMs lie in a range from ~4 eq/yr to ~8eq/yr. Lowest and highest rates are ~3eq/yr and ~17eq/yr, respectively from a fault-based and an area-source ERM. Considering the overall weighed distribution of the rates, 2.5<sup>th</sup> - 97.5<sup>th</sup> and 16<sup>th</sup> - 84<sup>th</sup> percentiles are, respectively, ~4.4 eq/yr - ~8.6 eq/yr, and ~4.8 eq/yr - ~6.9 eq/yr.

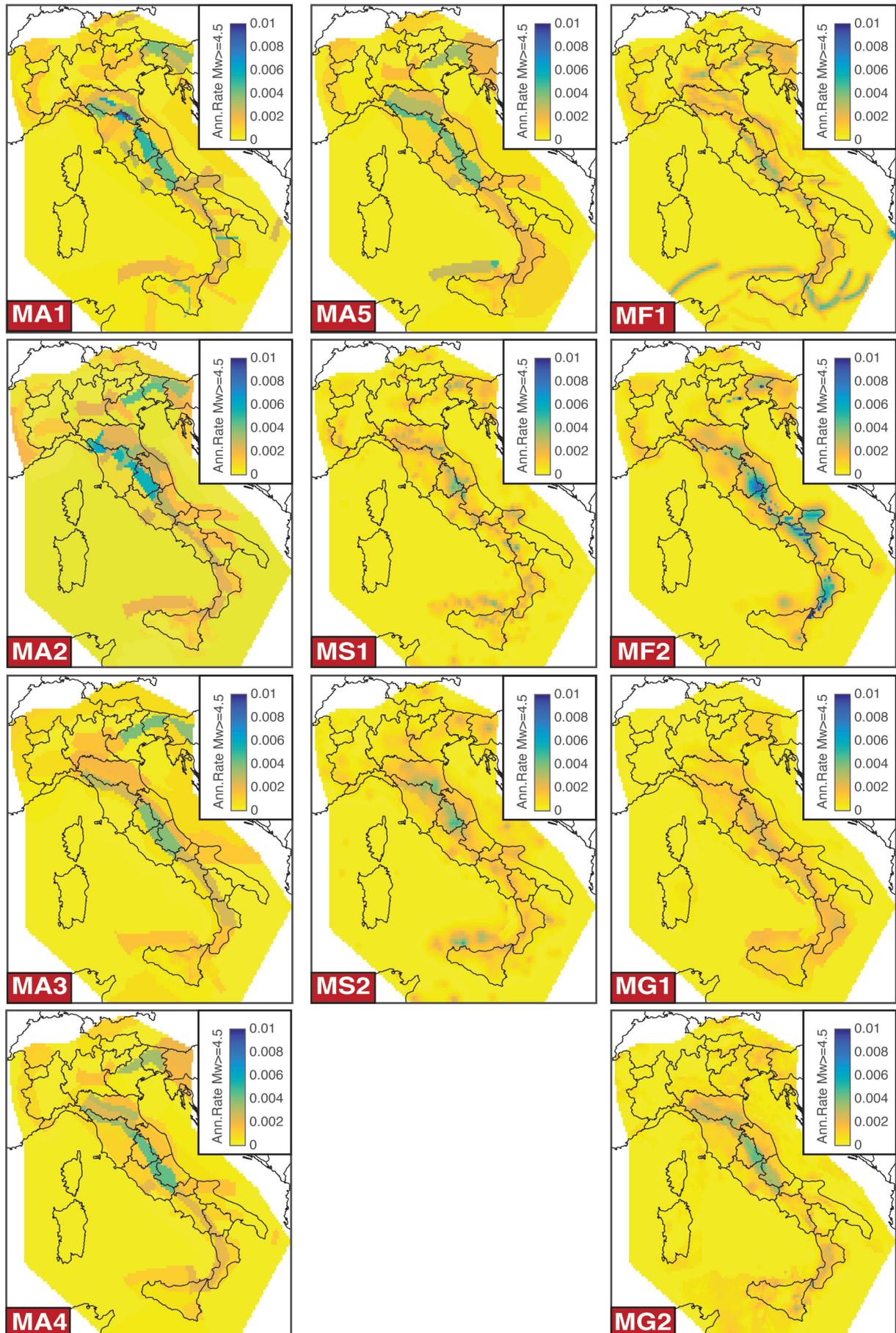
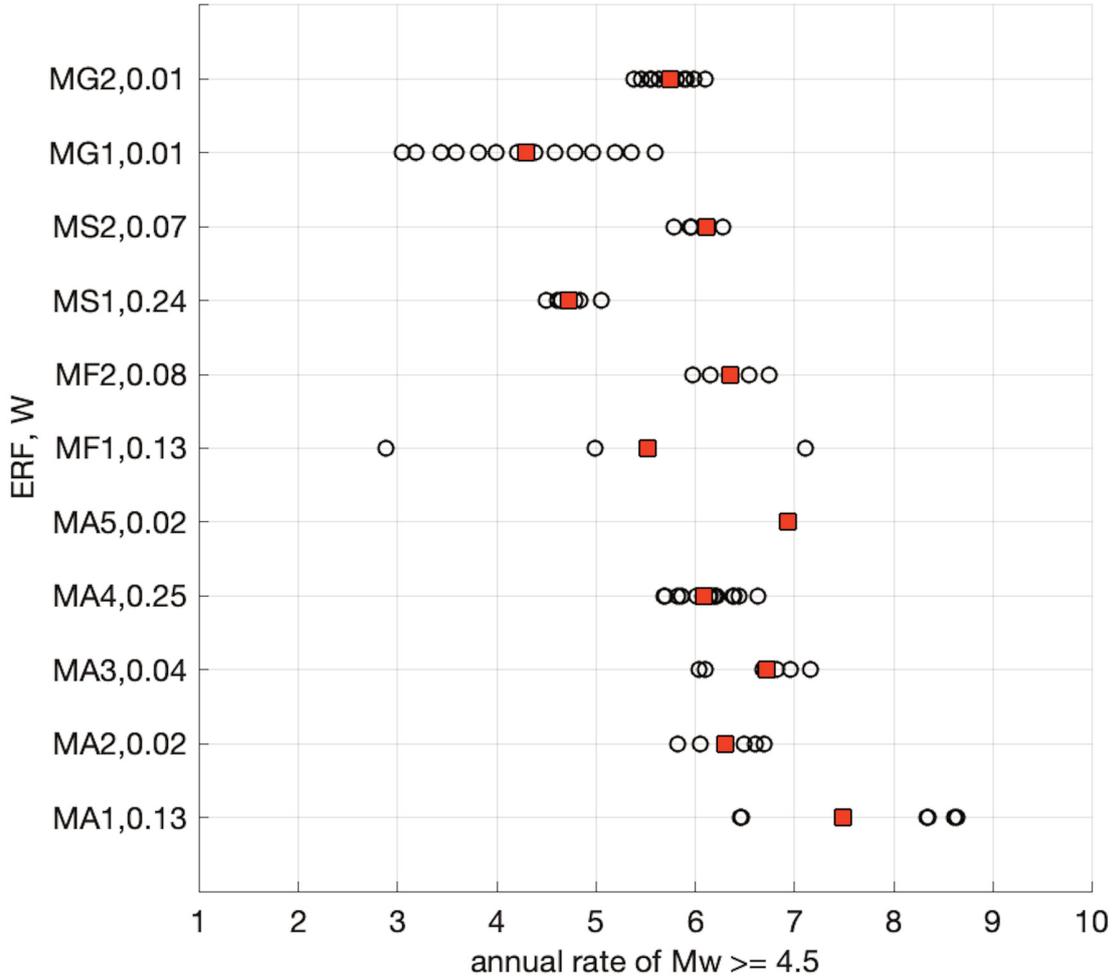


Figure 3. Mean annual earthquake rates for  $M_w \geq 4.5$  of each ERM.



**Figure 4.** Distribution of rates of occurrence for earthquakes with  $M_w$  larger or equal than 4.5, as evaluated by the 11 ERMs and their sub-models (open circle). Red squares are the mean rate of each ERM. The y axis reports the ERMs with their own code and final weight (see Section 6).

## 5. The GMMs

This section summarizes the activity of the selection and weight of GMMs for MPS19, that has been exhaustively treated in Lanzano et al. [2020]. The candidate GMMs potentially suitable for Italy have to satisfy the following general conditions: i) representation of the characteristics of ground motion for whole Italian territory; ii) calibration for the engineering bedrock (i.e. EC8 site category A,  $V_{s,30} > 800$  m/s) and flat topography conditions; iii) prediction of the horizontal PGA and 12 ordinates of the acceleration response spectra (5% damping), at periods  $T$  0.05, 0.1, 0.15, 0.2, 0.3, 0.4, 0.5, 0.75, 1, 2, 3, 4 seconds; iv) horizontal motion in terms of geometrical mean of the horizontal components.

We identify three main tectonic regions, which are characterized by peculiar ground motion attenuation properties: i) Active Shallow Crustal Regions (ASCRs); ii) Volcanic Zones (VZs); and iii) Subduction Zone (SZ), in the southern Tyrrhenian sea. A preliminary selection of candidate GMMs is performed over nearly one thousand models published in scientific literature, which are listed in Douglas [2019]. In order to reduce the number of available models, we adopt the rejecting criteria by Cotton et al. [2006] and Bommer et al. [2010], and then add further project-specific criteria to fulfil the requirements of MPS19 listed above [see Lanzano et al., 2020 for details]. The candidate models for ASCRs after this pre-selection are shown in Table 4, and Figure 5 displays the trellis plots of their spectral acceleration predictions, assuming rock-site conditions (EC8 site category A or  $V_{s,30} > 800$  m/s) and normal faulting.

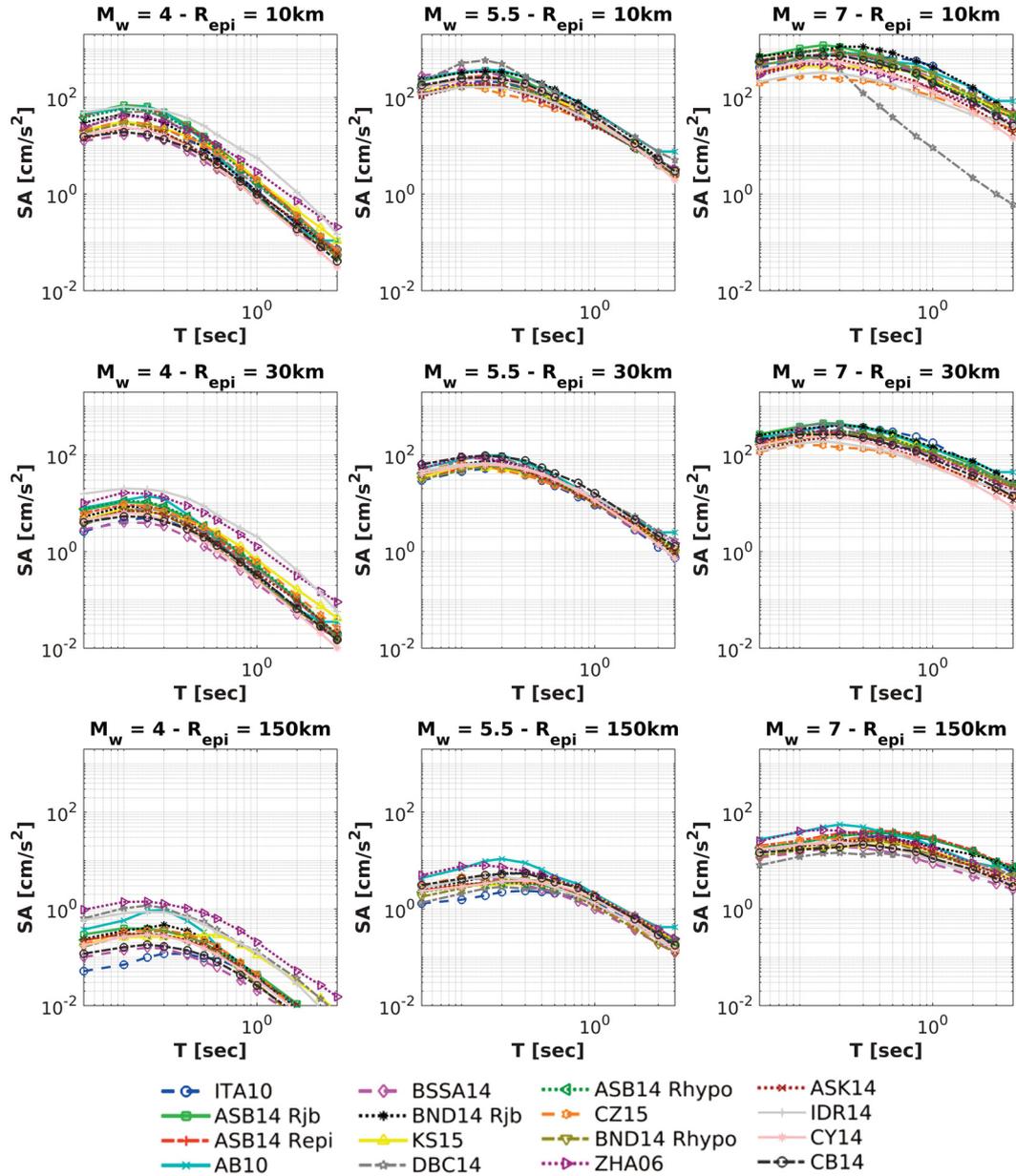


Figure 5. Trellis plot of the median values of the GMMs pre-selected for ASCRs (see Table 4 for the abbreviations).

Name	Code	Area	H comp	M range	R type	R range (km)	Period range (s)	Site-effect term	Style of faulting
Zhao et al. [2006]	ZHA06	Japan	GM	5.0 - 8.3	$R_{rup}$	0 - 300	0.05 - 5.0	5 Classes	N, R, SS
Bindi et al. [2011] <sup>1</sup>	ITA10	Italy	GM	4.1 - 6.9	$R_{JB}$	0 - 200	0.04 - 2.0	5 EC8 <sup>2</sup> Classes	N, R, SS, U
Akkar and Bommer [2010] <sup>3</sup>	AB10	Europe and Middle- East	GM	5.0 - 7.6	$R_{JB}$	0 - 100	0.05 - 3.0	3 Classes	N, R, SS

<sup>1</sup> The GMMs coefficients up to 4 s by ITA10 and BND14 have been provided by the Authors, even if they are not published.

<sup>2</sup> The EC8 site categories are defined in the Eurocode 8 part 1 (CEN, 2003) and are mainly based on the average shear wave velocity in the uppermost 30 m,  $V_{S,30}$ .

Abrahamson et al. [2014]	ASK14	Worldwide	RotD50	3.0 - 7.9	$R_{rup}$	0 - 300	0.01 - 10	$V_{S,30}$	N, R, SS
Boore et al. [2014]	BSSA14	Worldwide	RotD50	3.0 - 7.9	$R_{JB}$	0 - 400	0.01 - 10	$V_{S,30}$	N, R, SS, U
Campbell & Bozorgnia [2014]	CB14	Worldwide	RotD50	3.0 - 7.9	$R_{rup}$	0 - 300	0.01 - 10	$V_{S,30}$	N, R, SS
Chiou & Youngs [2014]	CY14	Worldwide	RotD50	3.1 - 7.9	$R_{rup}$	0 - 400	0.01 - 10	$V_{S,30}$	N, R, SS
Idriss [2014]	ID14	Worldwide	RotD50	4.5 - 7.9	$R_{rup}$	0 - 175	0.01 - 10	$V_{S,30}$	N, R, SS
Akkar et al. [2014]	ASB14	Europe and Middle -East	GM	4.7 - 7.6	$R_{JB}$ $R_{epi}$ $R_{hyp}$	0 - 200	0.01 - 4.0	$V_{S,30}$	SS, R, N
Bindi et al. [2014] <sup>1</sup>	BND14	Europe and Middle -East	GM	4.0 - 7.6	$R_{JB}$ $R_{hyp}$	0 - 300	0.02 - 3.0	<sup>4</sup> Classes, $V_{S,30}$	SS, R, N, U
Derras et al. [2014]	DBC14	Europe and Middle -East	GM	3.6 - 7.6	$R_{JB}$	0 - 550	0.01 - 4.0	$V_{S,30}$	N, R, SS
Cauzzi et al. [2015] <sup>4</sup>	CZ15	Worldwide	GM	4.5 - 7.9	$R_{rup}$	0 - 150	0.02 - 10	$V_{S,30}$	SS, N, R
Kuehn & Scherbaum [2015]	KS15	Europe and Middle -East	GM	4.0 - 7.6	$R_{JB}$	0 - 200	0.01 - 4.0	$V_{S,30}$	SS, N, R

**Table 4.** List of the candidate GMMs for the application to ASCRs for PGA<sup>5</sup> and acceleration response spectra ordinates. GM: geometrical mean; RotD50: median of orientation-independent ground motion amplitudes [Boore, 2010].  $R_{epi}$ : epicentral distance;  $R_{JB}$ : Joyner-Boore distance;  $R_{hyp}$ : hypocentral distance;  $R_{rup}$ : distance from the rupture plane. N: normal fault; R: reverse fault; SS: strike-slip fault; U: undefined style of faulting.

We evaluate the relative hindcasting skill of the pre-selected GMMs through the comparison with accelerometric records from the strong-motion databases available for Italy [ITACA, Italian ACcelerometric Archive, <http://itaca.mi.ingv.it>; Luzi et al., 2008; Pacor et al., 2011] and Europe [ESM, Engineering Strong Motion Database, <https://esm.mi.ingv.it>; Luzi et al., 2016]. The testing datasets for ASCRs and SZ are provided and described in Lanzano et al. [2020]. We verify the stability of the hindcasting skill of candidate GMMs, and examine their behavior on the basis of the requirements of MPS19, using 4 different subsets: i) the whole dataset (named All); ii) a subset of records with stations classified as EC8 site class A (named EC8-A); iii) a subset of events with  $M_w \geq 5$  and stations located at  $R_{JB} < 50$  km (named MR), since these magnitude and distance ranges are those mostly affecting the hazard in Italy; iv) a validation subset without the Italian records preceding 2009 (named After 2009), to obtain a dataset independent of that used to calibrate the ITA10 [Bindi et al., 2011] model. The last subset is the only one which allows a truly independent comparison of the GMMs; this is essential to quantify possible overfitting effects, which could favor those GMMs built using, at a different extent, the data from ITACA.

For the purpose of measuring the relative GMM hindcasting skill, we consider the proper scoring methods described in Section 2, the Log-likelihood (LLH), proposed by Scherbaum et al. [2009], and the parimutuel gambling score [Zechar and Zhuang, 2014]; we also consider an additional method, i.e. the quantile score [Gneiting and Raftery, 2007], to check the consistency of the results. The details of the scoring procedure and the scoring values for each GMM are reported and discussed in Lanzano et al. [2020]; here in Table 5 we show the average values of the scores, and the un-normalized weight for each GMM using the EC8-A subset; the un-normalized weight is the ordinal number of the score. Summing up the total un-normalized scores of EC8-A and MR subsets, the grand-total

<sup>3</sup> The AB10 predictions at 4 s (not provided by the Authors) are linearly extrapolated from those at 2 and 3 s.

<sup>4</sup> CZ15 is provided in terms of ordinates of the displacement response spectra (SD), and the predictions have been converted to SA by multiplying for  $(2\pi/T)^2$ .

<sup>5</sup> Most of the GMMs are also calibrated for PGV, with the exception of the model IDR14 and ZHA06.

score is obtained, providing a final rank of the candidate GMMs. The 3 best performing GMMs are: #1) Bindi et al. [2011], ITA10 (93 points); #2) Bindi et al. [2014], BND14  $R_{hyp}$  (82 points); #3) Cauzzi et al. [2015], CZ15 (75 points). Worthy of note, these GMMs, besides showing the best relative hindcasting skill, use different metrics for the distance (i.e.,  $R_B$ ,  $R_{hyp}$ ,  $R_{rup}$ , respectively) and are calibrated on different datasets (i.e., Italian, European, and global, respectively).

GMMs	LLH		Gambling		Quantile		Total
	mean*	UNW	mean	UNW	mean	UNW	UNW
ITA10	1.888	16	0.329	16	1.143	16	48
ASB14 Repi	1.918	15	0.015	12	1.096	13	40
BND14 Rhyp	2.004	10	0.088	15	1.113	15	40
CZ15	1.979	12	0.040	13	1.098	14	39
ASB14 Rhyp	1.926	14	0.006	10	1.093	12	36
ASB14 Rjb	1.926	13	-0.005	8	1.088	11	32
DBC14	2.095	6	0.053	14	1.041	9	29
BND14 Rjb	2.015	8	0.004	9	1.079	10	27
KS15	2.055	7	0.007	11	1.025	6	24
ASK14	1.985	11	-0.084	3	1.028	7	21
CY14	2.014	9	-0.046	4	1.032	8	21
BSSA14	2.127	5	-0.037	5	0.967	5	15
CB14	2.167	4	-0.017	7	0.965	4	15
IDR14	2.297	3	-0.036	6	0.964	3	12
AB10	2.771	2	-0.188	1	0.889	2	5
ZHA06	2.889	1	-0.129	2	0.790	1	4

\*The scores are averaged over 5 periods of the spectral acceleration (0.1, 0.2, 0.5, 1 and 2 s).

**Table 5.** Rank of candidate GMMs for the ASCR dataset EC8-A. UNW: un-normalized weight.

For the Italian subduction zone, which corresponds to the Calabrian Arc in the southern Tyrrhenian sea, we select two GMMs after the pre-selection and scoring procedures: i) Skarlatoudis et al. [2013], HEL13, and ii) Abrahamson et al. [2016], HYD15 [for details, see Lanzano et al., 2020]. For volcanic zones (i.e., Etna, Aeolian islands, and Campanian volcanoes) we use a single GMM that is the model calibrated on purpose by Lanzano and Luzi [2019], since no GMM available in the previous literature [e.g. Faccioli et al. 2010; Tusa and Langer, 2016] is able to describe satisfactorily the ground motion amplitudes observed recently in Italy (2018-2019 Mount Etna sequence, mainshock  $M_w$  4.9, and 2017 Ischia event,  $M_w$  3.9), especially in near-source conditions.

## 6. Weighing and merging all MPS19 components

It is essential to weigh each ERM and GMM to properly merge them in the final MPS19. The weighing procedure is often based on unstructured experts' judgment [Gerstenberger et al., 2020]. Here we aim at providing a more objective procedure that could make the weighing scheme more sound, transparent and reproducible.

The final weight of each component is assigned accounting for its statistical scoring (described in the sections 4 and 5), and the outcomes of an experts' elicitation session with independent national and international experts. As a matter of fact, the weight of each model (either ERM or GMM) is a combination of two different evaluations:

i) a weight based on the testing phase of the model,  $W_1$ , and ii) a weight derived by expert judgment,  $W_2$ . The final weight of that model is

$$W = W_1 \times W_2 / C \tag{1}$$

where  $C$  is a normalization factor that sets  $W$  in the range (0,1) and that the sum of all weights is 1; specifically,  $C$  is the sum of  $W_1 \times W_2$  for all models. The need of merging these two different scores is due to the fact that the number of data for the testing phase is limited and the test is made retrospectively (so models that tend to overfit the data are more rewarded than others in the testing phase), and that the earthquake catalog could not cover the whole seismic cycle; in this case, the experts' opinion is expected to provide an independent additional information for weighing models [Delavaud et al., 2012]. Note that equation (1) provides zero weight to models that are either not able to explain the data, or unrealistic, e.g. based on assumptions that are blatantly wrong.

In Tables 6 and 7 we show the weights of the ERMs and GMMs, which take into account the statistical scoring and the experts' judgment. In Appendix E of the Supplementary Material we report the description of the experts' elicitation sessions.

Model	Testing phase weight $W_1$	Experts' elicitation weight $W_2$	Final weight $W$
MA1	0.1439	0.0955	0.13
MA2	0.0303	0.0539	0.02
MA3	0.0833	0.0539	0.04
MA4	0.1136	0.2240	0.25
MA5	0.0833	0.0198	0.02
MF1	0.1136	0.1172	0.13
MF2	0.1136	0.0756	0.08
MS1	0.1136	0.2193	0.24
MS2	0.1439	0.0501	0.07
MG1	0.0303	0.0529	0.01
MG2	0.0303	0.0378	0.01

Table 6. ERMs weights.

GMMs	$W_1$	$W_2$	$W$
<b>ASCRs</b>			
ITA10 [Bindi et al., 2011]	0.37	0.423	0.45
BND14 Rhyp [Bindi et al., 2014]	0.33	0.328	0.32
CZ15 [Cauzzi et al., 2015]	0.3	0.259	0.23
<b>SZ</b>			
HEL13 [Skarlatoudis et al., 2013]	0.53	0.5	0.53
HYD15 [Abrahamson et al., 2016]	0.47	0.5	0.47
<b>VZs (depth <math>\leq</math> 5 km)</b>			
LL19 [Lanzano & Luzi, 2019]	-	1.0	1.0

Table 7. GMMs weights.

To summarize, the weighing procedure adopted is as much objective and auditable as possible; it considers independent information, such as the hindcasting performance of the models and the experts' judgment on the scientific reliability of the models. Nonetheless, the procedure has also some limitations, which are mostly due to the finite number of data, and the impossibility to carry out more specific tests at regional scale, or tests that depend on the earthquake magnitude. Here we report a list of these limitations.

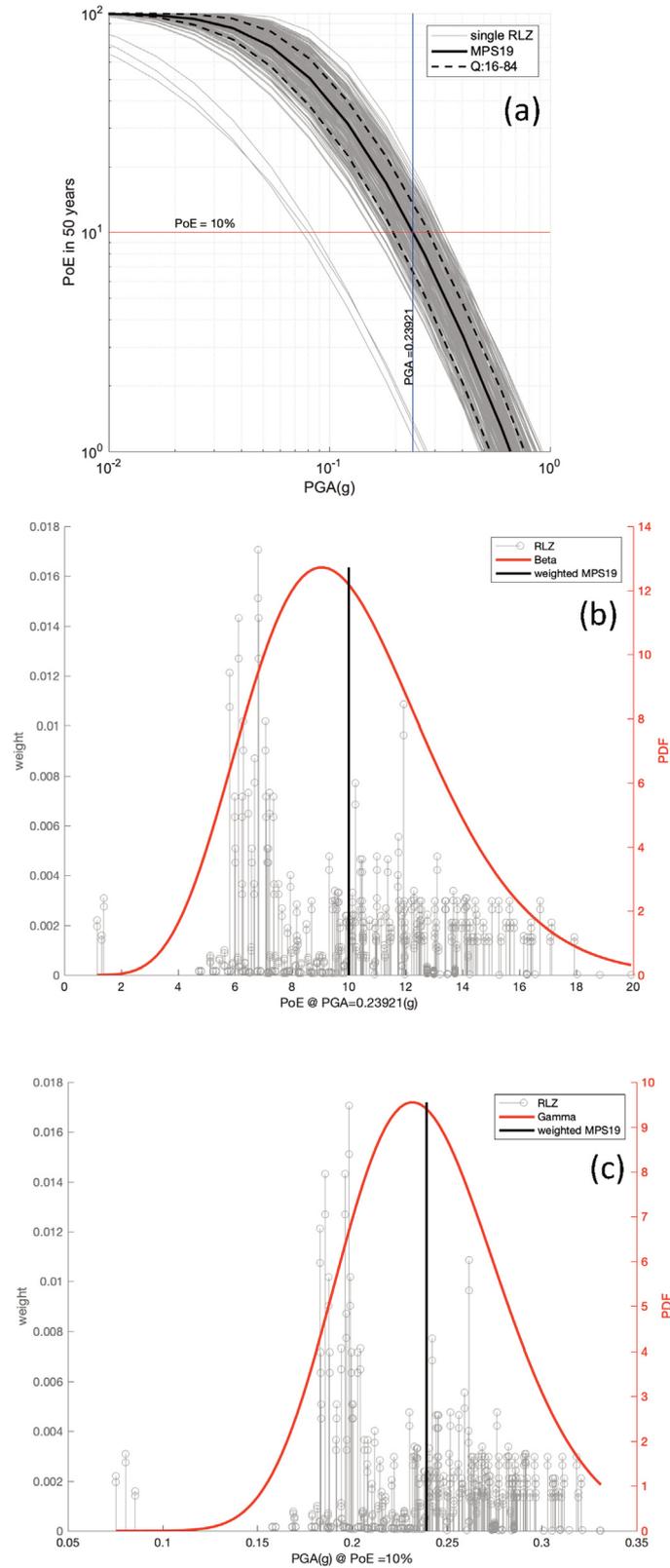
- i. The statistical procedure to select and to score ERMs and partially GMMs does not take into account that some models were calibrated with the same data used for testing, while others do not. This means that the latter are probably underscored (and underweighted).
- ii. The weight of each ERM and GMM is constant for the whole country, whereas some ERMs are not used in some sub-regions where they failed the tests.
- iii. Owing to the limited number of data, the testing procedure is limited only to some specific aspects of the earthquake occurrence and ground motion processes. For example, the data available do not allow us to test ERMs for M7+ earthquakes, or to compare GMMs performance in the very near field (less than 10 km) which are very relevant for seismic hazard.
- iv. The procedure carries on an unavoidable subjectivity in merging the different scoring metrics of the testing phase and the scoring of the experts' judgment. Other procedures to merge these scores are possible.
- v. The weights based on the testing phase do not take into account that some ERMs may perform poorly because of uncertainties in the oldest earthquakes of the seismic catalog, such as large uncertain location of off-shore and/or very old earthquakes, and the magnitude of historical events which may be affected by cumulative damages caused by seismic sequences. This is unavoidable, but it highlights again the need to not weigh a model only using its hindcasting performance.

Once the ERMs and GMMs weights have been defined, the seismic hazard at each site is estimated by  $\mathcal{N}$  hazard curves, where  $\mathcal{N}$  is the number of ERMs and GMMs combinations. Worthy of mention, this set of hazard curves is expected to sample the epistemic uncertainty, and a higher number of models cannot reduce it. In fact, even though models are developed independently from one another, almost all of them use the same amount of information [Marzocchi and Jordan, 2014].

If we dissect the bunch of curves for one specific IM value, we get a set of estimations of the PoE for that specific IM value (Figure 6). We model them using a Beta distribution [Marzocchi et al., 2015], which describes where the true PoE is expected to be [Marzocchi and Jordan, 2014; 2017]. The choice of using the Beta distribution is subjective and other ones may be used [see Gelman et al., 2003], keeping the same probabilistic framework adopted here. That said, we argue that this choice is certainly less subjective than describing the hazard using the mean alone, which would imply the use of a Dirac distribution, or using the percentiles of the distribution, which implicitly means to impose a nonparametric distribution. The Beta distribution usually fits well aleatory variables with a unimodal distribution and bounded between 0 and 1; in case the data show more modes [for examples 2 modes as in Figure 6 of Marzocchi et al., 2015], we can still use the Beta distribution if we think the bimodality is only due to the fact that the models used are just exploring two extreme scenarios, and a proper epistemic uncertainty has to fill the gap between the modes. To sum up, no matter what distribution we decide to use, this choice is certainly less critical than using one single value for the hazard.

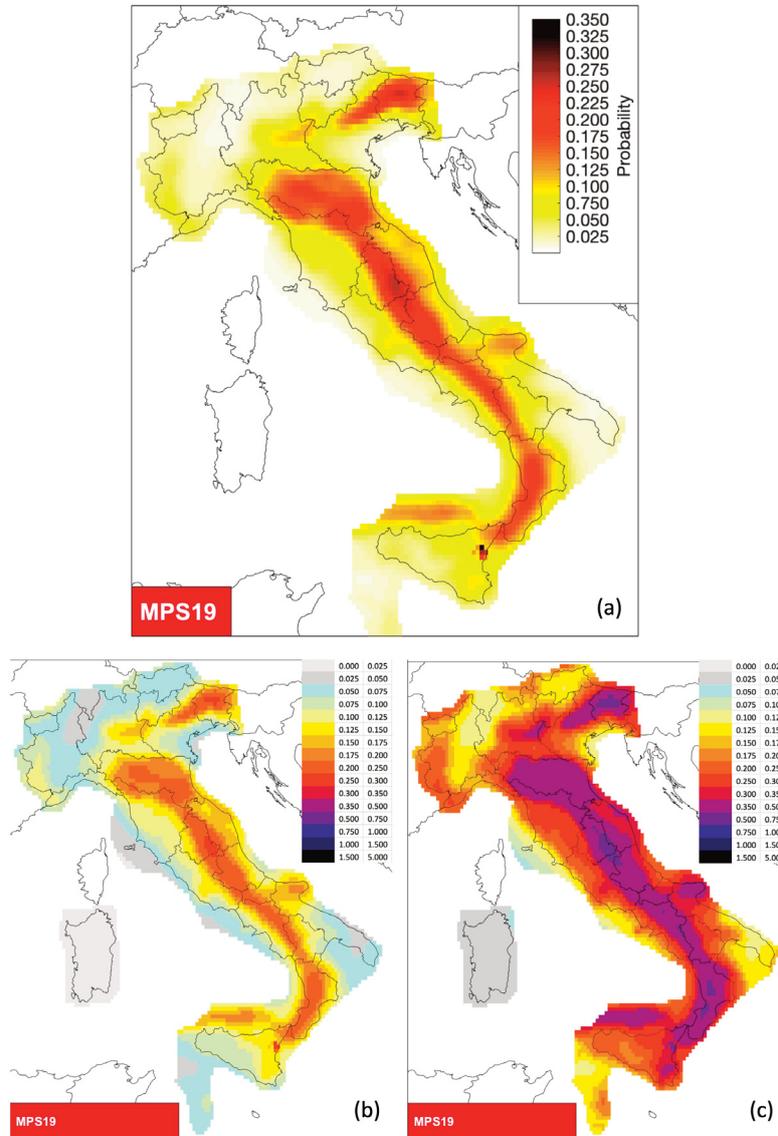
## 7. The MPS19 outcomes and consistency test

Regional seismic hazard is usually presented as a map showing the PGA values that have 10% PoE in 50 years. This representation is of particular interest to the engineering community which establishes this hazard level to be of interest for the building code. However, this representation is often misinterpreted by laymen and some end-users, which assume that this map is reporting the maximum PGA that can occur in each area. In fact, although it is quite obvious to observe PGA exceedances just close to the epicenter of a large earthquake, this observation is often deemed as a proof that the probabilistic model is wrong. For this reason, we propose here an additional graphical representation of MPS19 that may be easier interpreted by the non-engineering community. Specifically, we show the PoE of PGA = 0.15 g in the whole Italian territory for a time interval of 50 years (Figure 7a). The choice



**Figure 6.** Seismic hazard for L'Aquila. (a) Hazard curves of all the realizations of MPS19 and weighted mean hazard curve along with 16<sup>th</sup> and 84<sup>th</sup> percentiles; the horizontal red line marks the 10% PoE in 50 years, and the blue vertical line identifies the PGA corresponding to the mean hazard at 10% PoE in 50 years ( $PGA^*$ ). (b) The Beta distribution (red line) of the PoE corresponding to  $PGA^*$ ; each vertical line shows the weight of the outcome of one hazard model. (c) The Gamma distribution (red line) of the PGA corresponding to PoE of 10% in 50 years.

of 0.15 g has been adopted as a criterion for significant shaking, and other values can be used [WGCEP, 1995]. Figures 7b and 7c show the more classical representation of MPS19, in terms of PGA with 10% and 2% PoE in 50 years.

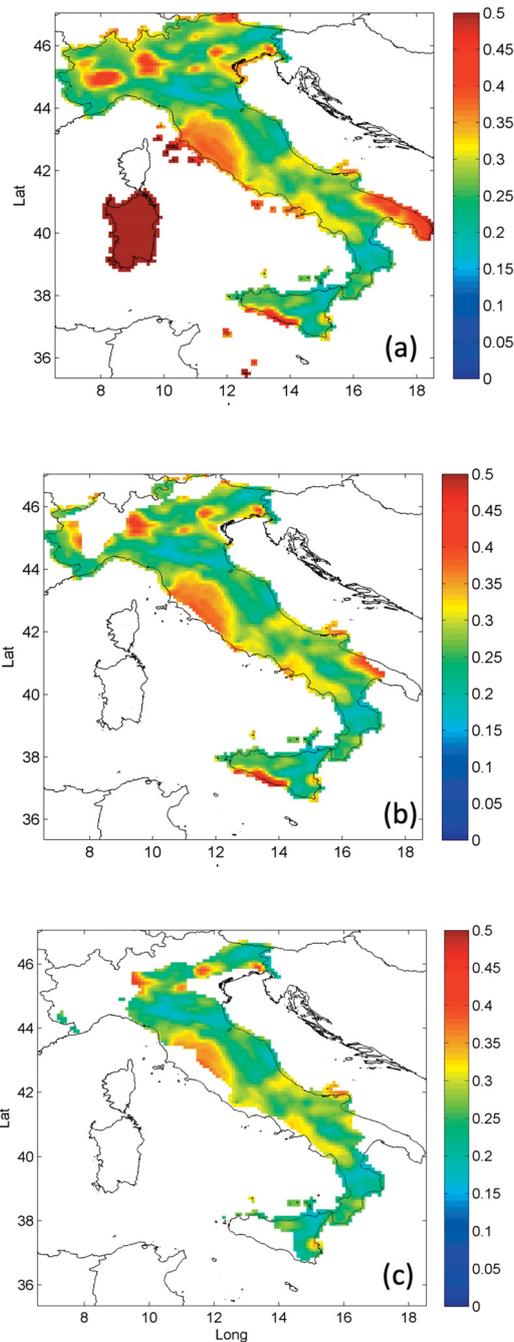


**Figure 7.** The outcomes of MPS19: (a) PoE for PGA=0.15 g in 50 years; (b) PGA with 10% probability to be exceeded in 50 years; (c) PGA with 2% probability to be exceeded in 50 years.

One of the most innovative aspects of MPS19 concerns the quantification of the epistemic uncertainties of the final hazard estimates. As mentioned above, for the probabilistic scheme used, the exceedance probability at one site for one specific IM is not a single number, but it is a Beta distribution, the dispersion of which describes the epistemic uncertainty on the true exceedance frequency (random variability). One way to represent epistemic uncertainties on the whole national territory is given by the mapping of the CoV, i.e. coefficient of variation. CoV is given by

$$\text{CoV} = \frac{\sigma}{\mu} \quad (2)$$

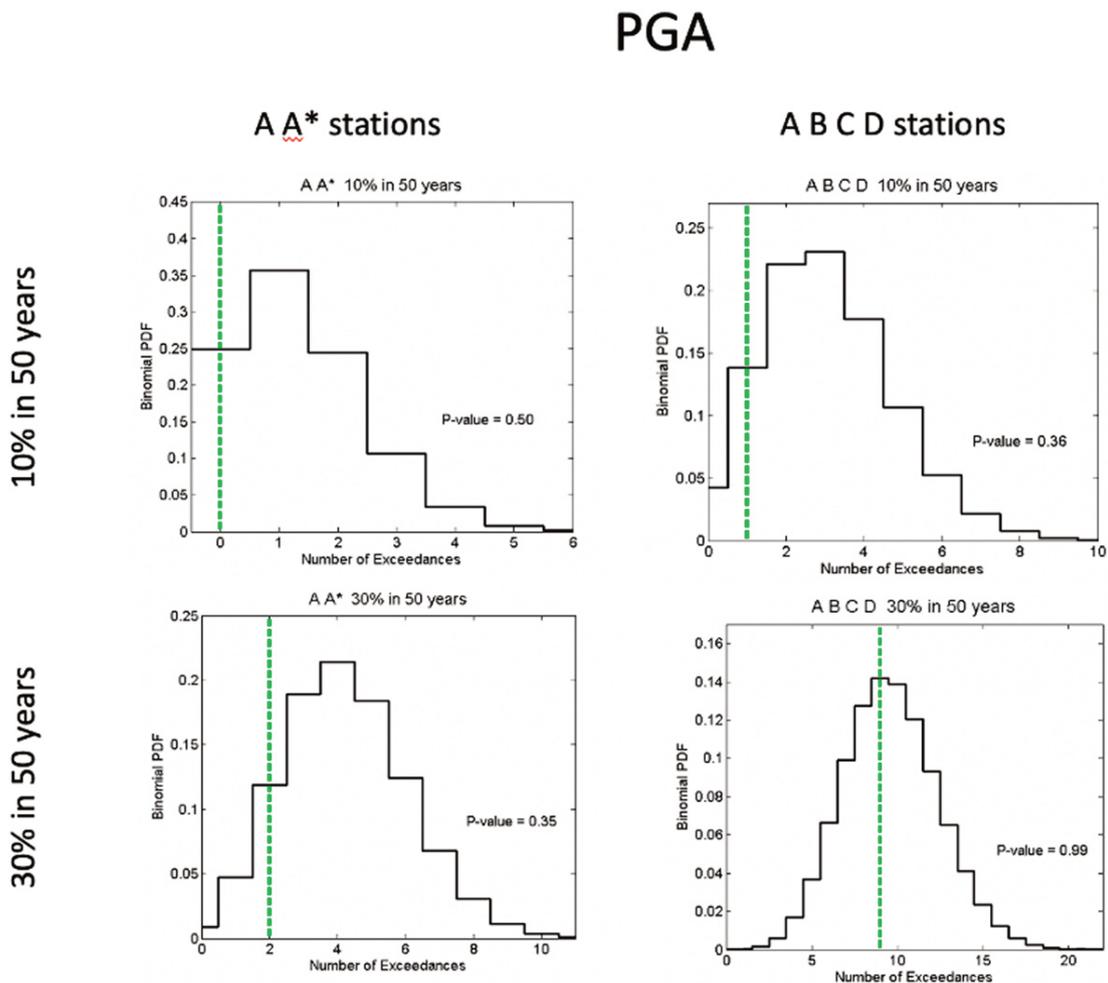
where  $\sigma$  is the weighted standard deviation of all the hazard estimates, and  $\mu$  is the weighted average value among all estimates (these parameters are used to calculate the parameters of the Beta distribution). A CoV of 0.3 indicates that the standard deviation between all the models used (including the uncertainties within each model) is 30% of the average. Figure 8 shows the CoV map for the whole national territory relating to PGA values which have the exceedance probability of 10% in 50 years (Figure 7b). From the figure it can be seen that the highest values of the CoV are for areas with low seismic hazard. When we consider only areas with high hazard (i.e. those with a higher  $\mu$ , panel (c)), the CoV mainly varies between 0.15 and 0.30. The fact that the CoV is higher for low hazard areas is due to the fact that in those areas the information available are relatively scarce, and so it is understandable to have greater uncertainties.



**Figure 8.** CoV (equation 2) map for the 10% PoE in 50 years PGA (Figure 7b). (a) CoV for the whole territory; (b) CoV for the sites of Figure 7b with PGA above the 25<sup>th</sup> percentile; (c) CoV for the sites of Figure 7b with PGA above the 50<sup>th</sup> percentile.

It is worth remarking that this result poses a great challenge to the use of MPS19 (as well as of any other model built on the same amount of information) for the definition of the building code. As a matter of fact, the current Italian building code is based on one hazard input having 4 decimal digits and it does not consider epistemic uncertainty; *de facto*, the latter is handled by engineers introducing some arbitrary conservative parameters. Figure 8 shows that this request of precision does not match the current epistemic uncertainty in PSHA. In practice this means that even in the most optimistic (and likely unrealistic) option where MPS19 is “right”, the mean values of future PSHA models considering more data are expected to vary according to the range given by the CoV. In essence the CoV represents the lower bound of the expected future variations of the mean hazard.

As regards the test of MPS19 against observations, in Figure 9 we show the results of the consistency test of MPS19 compared to the PGA recorded ed by different accelerometric stations in recent decades (from the ITACA database). In particular, we consider accelerometric stations that have been operating for at least 30 years. If the period is longer, the last 30 years are considered (until 31 December 2018). In this way all stations have observations for the same length of the time interval considered. The fact that the time intervals are temporally staggered is irrelevant because of the stationarity of the Poisson process. Owing to the PGA values of MPS19 refer to a 50-yr time window, we rescale the PoE to 30-yr time window. This is done maintaining the same annual exceedance rate of PGA; the 10% threshold of PoE in 50 years (which becomes about 6% in 30 years) is chosen for the obvious practical interest in terms of building code; the 30% threshold in 50 years (which becomes about 19% in 30 years) is used to lower the expected PGA values in order to have more exceedances and therefore to empower the test.



**Figure 9.** Results of the consistency test of MPS19 with the observed PGA values in the last decades for different PoE (10% and 30% in 50 years). The meaning of sites A, A\*, B, C, and D is reported in the text.

We consider accelerometric stations on rock (EC8 A site category) and on all types of terrain (EC8 A, B, C and D) with flat topographical surface (i.e., T1 category according to NTC18 Italian building code; NTC [2018]). In case the station is not located on rock, the PGA observation is scaled down by the weighted average of the site coefficients of the 3 GMMs selected for ASCRs according to their weights. For rock sites we consider all stations, either with a good site characterization (type A), or with more uncertain site characterization (named type A\*). When considering all types of terrain, we use only the stations with a good site characterization. The IM observed at different sites is correlated [Park et al., 2007], and this may be an issue for testing [Iervolino et al., 2017]. We adopt a pragmatic approach choosing sites that have a minimum distance of 30 km from one another. In this case, although part of the IM correlation remains independently from the distance [Park et al., 2007], we may assume that, for small PoE, exceedances are independent, i.e., the observation of an exceedance  $E_Y$  in one site  $Y$  does not modify the probability to observe an exceedance  $E_X$  in the nearby site  $X$ ,  $P(E_X|E_Y) = P(E_X) \equiv \text{PoE}$ . Hence, the exceedances follow a binomial distribution with the parameter equal to the selected PoE. Worthy of note, this assumption makes the test of PGA conservative: if the null hypothesis (the exceedances are generated by a process which is well described by MPS19) is not rejected, it will not be rejected also relaxing the assumption, i.e., using a more exact distribution for the number of exceedances, which is overdispersed with respect to the binomial distribution [see Figure 2 in Iervolino et al., 2019]. In the analysis reported in Figure 9, we show the results of the test when the epistemic uncertainty of the PoE at each site is not included, which means to not consider possible common mode errors [Marzocchi and Jordan, 2018]. This also makes the test very conservative for our purposes, i.e., a non-rejection means that it will be also a non-rejection if uncertainties are included.

MPS19 is also tested against the macroseismic data observed in Italy as reported in the database DBMI15 v1.5 [Locati et al., 2016]. In this way, we can use a much longer time period of observations, i.e., several centuries. Since MPS19 does not currently produce estimates in macroseismic intensity, the first step is to convert the PGA exceedance rates  $\lambda(x)$  where  $x$  is the PGA, in terms of rates of exceedance of macroseismic intensities  $\lambda(I)$ . This is done using the procedure suggested by D'Amico and Albarello [2008]

$$\lambda(I) = \sum_{i=1}^N \lambda(x_i) P(I|x_i) \quad (3)$$

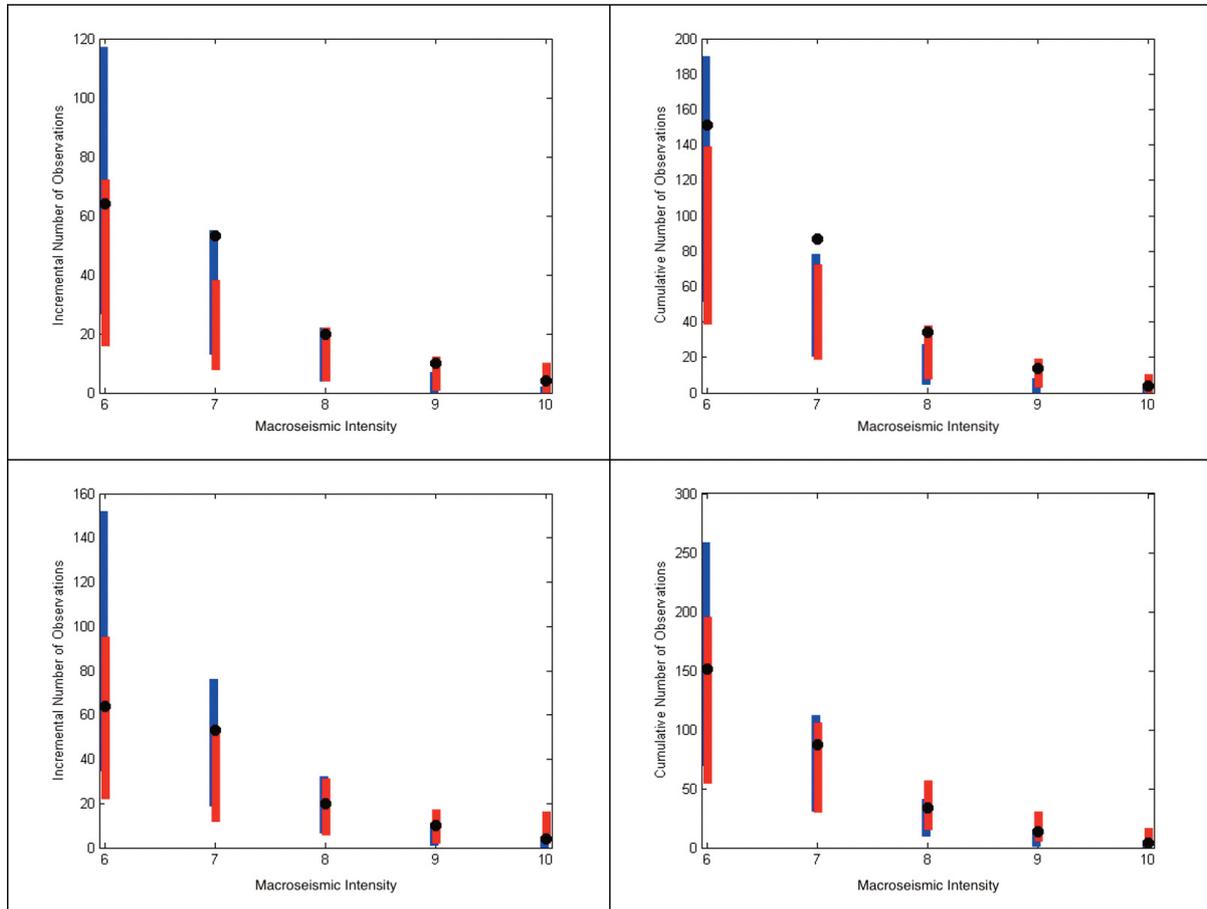
where  $P(I|x_i)$  is the empirical function for converting PGA into macroseismic intensity, and  $N$  is the number of PGA levels considered in the hazard curves of the model. Obviously, the passage introduces a further source of possible ontological error, as the empirical function for transforming PGA into macroseismic intensity is subject to high uncertainty. For this test, we account for these uncertainties using two functions estimated by Faenza and Michelini [2010; 2011], and Gomez Capera et al. [2020].

We select 27 sites that have a long seismic history for the analysis. The number of observations at each site takes into account the completeness of the data and considers only macroseismic intensities equal to 6 or higher which have been generated by mainshocks (macroseismic intensities associated to foreshocks and aftershocks are removed). To avoid the spatial correlation of macroseismic intensities, the distance between any couple of sites is always larger than 30 km (see the discussion above for testing the PGA exceedances), and we keep only one intensity (the maximum observed) for each earthquake; being generated by mainshocks following the Poisson distribution, we assume that these macroseismic observations are independent and follow the same distribution. This assumption facilitates the consistency test phase since we can test with aggregate observations. The procedure is similar to that described for the PGA exceedances. In this case, the epistemic uncertainty is taken into account, that is,  $P(\lambda_{tot})$  is the pdf Gamma of the parameter  $\lambda_{tot}$ , which is estimated by the different (weighted) rates of all the models. Since there is considerable uncertainty and variability on the type of soil where macroseismic intensities have been observed, the test is carried out for two different types of soil: A and C.

Since all the GMMs used in MPS19 assume a linear behavior of soils, to estimate the PGA for type C soils the hazard curve calculated for type A soils is translated by a factor derived for Italy from the GMM by Bindi et al. [2011]. Note that the test for type A soils represents an underestimation of the number of exceedances because it does not take into account any site amplification.

Figure 10 displays the test plots for each macroseismic intensity value. These results show that MPS19 is compatible with the overall number of macroseismic intensities observed at the selected localities in the last centuries. As for the testing of PGA, we underline that any violation of the assumption on the independence of observations does not affect the conclusions of this test (MPS19 is not rejected). In Appendix F of the Supplementary

Material we add the same test for individual sites. The results show a general good agreement in all sites, with some exceptions (mostly when considering type A soils) where the number of observations for intensity 7 or less are not in agreement with the model.



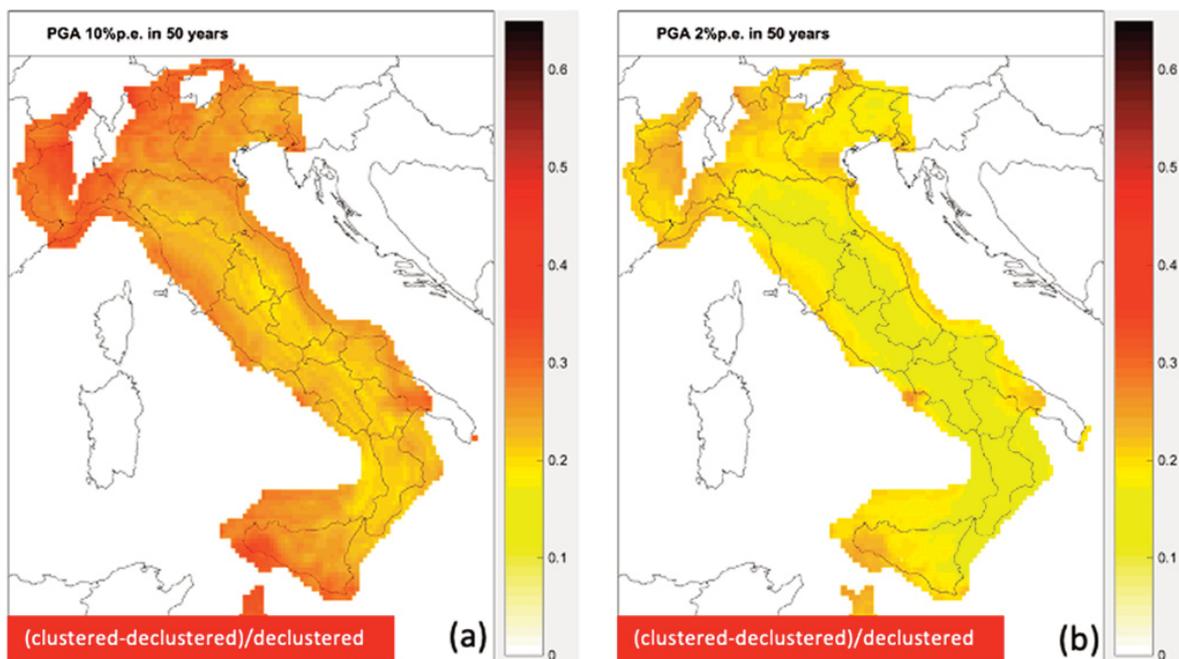
**Figure 10.** Results of the consistency test of MPS19 with the observed macroseismic intensities (black dots) of the last centuries for different soils. The upper row is for A sites; the lower row is for C sites. The left panels report the incremental number of observations for different macroseismic intensity, and the right panels the cumulative versions; the blue [Faenza and Michellini, 2010] and red bars [Gomez Capera et al., 2018] are the forecasts of MPS19 using two different transformation laws between PGA and macroseismic intensities (see text for more details).

### 8. A pilot modeling: MPS19\_cluster

One of the basic assumptions in PSHA is that the earthquake occurrence has to be stationary in time and Poisson distributed [Cornell, 1968]. To satisfy this assumption seismic catalogs are commonly declustered, i.e., clusters of seismicity, which represent a clear deviation from the Poisson distribution, have to be removed. The basic rationale behind this choice is that by keeping the largest events we remove earthquakes that have a negligible contribution in terms of IMs and seismic hazard [Lomnitz, 1966]. However, recent earthquakes sequences, such as the Canterbury sequence 2010-2011 in New Zealand, Kumamoto sequence 2016 in Japan, the central Italy sequence 2016-2017, have raised doubts on the validity of this assumption, in particular for a complete risk analysis. For this reason, scientists are beginning to estimate the effects of the declustering procedure on PSHA.

As previously said, different declustering procedures exist, which may lead to important variability of the declustered seismicity rate [vanStiphout et al., 2011]. We think that such a variability cannot be considered as epistemic, because most of the declustering techniques do not do what they were planned for, i.e., leaving a mainshocks distribution similar to Poisson. So, even assuming that a true Poisson mainshocks distribution really exists (yet we are not aware of any convincing physical reason that allows us to postulate the existence of such unknown sharp distinction between mainshocks and all other earthquakes), the epistemic uncertainty cannot be estimated by the variability generated by the current declustering techniques.

In MPS19 we use GK74 declustering technique, because it is the most efficient procedure to get a mainshocks distribution similar to Poisson distribution (see Section 3). No matter of the declustering method used, this operation implies that strong ground shaking caused by aftershocks and foreshocks is not considered in classical PSHA [Boyd, 2012; Iervolino et al., 2012; 2019; Marzocchi and Taroni, 2014] and risk [Papadopoulos et al., 2020]. For this reason, we build a version of MPS19 which includes the effects of earthquake clustering (hereafter MPS19\_cluster) and allows us to quantify its impact in terms of seismic hazard. MPS19\_cluster is rooted in two different strategies, which have been described by Iervolino et al. [2012] and Marzocchi and Taroni [2014]. For the sake of example, in Figure 11 we show the MPS19\_cluster obtained by applying only the procedure described by Marzocchi and Taroni [2014] on the mean ERM used in MPS19. The correction is magnitude-dependent in order to restore the magnitude-frequency distribution of the earthquakes before declustering, and it is applied to each spatial bin.



**Figure 11.** Variations of the (a) 10% PoE in 50 years, and (b) 2% PoE in 50 years PGA when the effect of clustered earthquakes are included to the maps of Figure 7b and 7c, respectively, according to the procedure described by Marzocchi and Taroni [2014]. The maps show only the sites with a PGA above the 25<sup>th</sup> percentile.

This figure intends to give a flavor of what the impact may be in the PGA estimated for 10% and 2% PoE in 50 years of MPS19 (Figure 7b and 7c respectively); from the figure we notice that the mean hazard could increase of about 20-25% in the regions of higher hazard for 10% PoE in 50 years with some increases of more than 30%. For the 2% PoE in 50 years, the increase is less strong being about 10-15% with some increases of about 25%.

A similar work is carried out also using the procedure described by Iervolino et al. [2012]; the outcomes will be described in a separate paper [Chioccarelli et al., 2021].

## 9. Final remarks, highlights, and open questions

MPS19 uses all new available data at the best and implements the most advanced practice introducing several novelties that are worth being remarked:

- MPS19 honors the hazard/risk separation principle, i.e., it does not include any subjective “precautionary” choices that should be introduced only at the risk level. In other words, all choices made in MPS19 are based on scientific thoughts, independently from the impact on the seismic hazard.
- MPS19 is modular, transparent and reproducible.
- MPS19 shows the aleatory variability and the epistemic uncertainty in the final output, making the interpretation of the dispersion of the hazard curves more clear, and providing the interpretation for any possible end-user.
- MPS19 has been checked through an extensive testing phase to assign weights to each component of the model, and to verify the reliability of the ERMs, GMMs and of the hazard outcomes in terms of PGA and macroseismic intensity at several sites and, on average, over the whole territory.
- We introduce MPS19\_cluster, which is a pioneering attempt to correct the classical PSHA model to account for the large foreshocks and aftershocks that have been removed by the declustering technique.

At the same time, the work that has been made for MPS19 have highlighted also some challenges that should be addressed in the future. Some of them are general, others are specific for the Italian territory:

- MPS19\_cluster shows the impact of removing earthquake clusters in terms of seismic hazard. Other methods based on synthetic catalogs may offer some additional interesting new perspectives in this field [e.g., UCERF3; Field et al., 2017].
- One of the basic assumptions in PSHA is that seismicity is stationary in time once the short-term clustering has been removed by declustering. In other words, time dependency on longer time scales is not included in this kind of modeling. This assumption is often taken for granted and not carefully tested. We think that in the future more efforts will be necessary to verify this assumption.
- The magnitude-frequency distribution has a large impact on PSHA. The common assumption of using a Gutenberg-Richter-like distribution over wide areas may be questionable, and currently we do not have enough data to test this assumption, at least for what concerns long-term PSHA purposes.
- GMMs are still mostly empirical and very likely not able to capture the details of IM variability at the spatial scale of interest for PSHA. Moving from general GMMs to more specific models which account for the regional geology may lead to significant variations of PSHA.

As a final remark, we argue that the procedures adopted in MPS19 make the comparison with the previous model MPS04 challenging. In particular, the elements of substantial difference are: i) MPS04 represents the median of the branches, MPS19 the average; ii) MPS04 represents the maximum horizontal component of the shaking, whereas MPS19 the geometric mean; iii) MPS04 made extensive use of expert judgment essentially aimed at introducing precautionary factors, whereas MPS19 keeps this judgment to a minimum, avoiding the use of precautionary components; iv) the basic data are completely different and up to date (more than 15 years have passed); v) MPS19 is based on a fully revised seismic catalog and magnitude estimated with more advanced procedures; vi) some earthquakes in MPS04 (based on the CPTI04 catalog) were aftershocks that have been removed in MPS19 by the declustering method; vii) MPS04 and MPS19 use completely different GMMs; viii) accelerometric databases used to define GMMs contain marked differences on the soil classification for the seismic stations.

Owing to these reasons, a detailed appraisal of the differences will be addressed in a separate future paper.

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**SUPPLEMENTARY MATERIAL TO**

**THE NEW ITALIAN SEISMIC HAZARD MODEL (MPS19)**

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## **APPENDIX A: List of participants, review panel, and external experts which contribute to MPS19**

Besides the authors of this paper the following researchers contributed at different levels to the preparation of MPS19.

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## **APPENDIX B. Iterations with the participatory review panel**

From the beginning, the work carried out in the preparation phase of MPS19 has been revised by the experts identified by National Civil Protection Department (DPC). In particular, from 2015 to 2017, the revision was made by the National Commission for major risks prediction and prevention (section Seismic Risk; MRC hereafter), the Operational Committee of DPC. At the beginning of 2018 a new Commission has been appointed; in order to not interrupt the review process, DPC named a participatory review panel with 5 out of 12 previous members of MRC (listed in Appendix A). The 5 reviewers followed the work carried out by the MPS19 Working Group until the end of the project.

The interaction was realized through periodic meetings, production of internal reports, elaborations, graphs. The reviewers asked details on the advancements with respect to construction of the model, the testing phase, the weights assigned to ERMs e GMMs, and so on.

The reviewers asked for several improvements of the model during its elaboration. Among them, one of the most important is the introduction of macroareas for testing and weighing the ERMs; the rationale is that some ERM could be more efficient in some regions and less in others. Then they proposed to evaluate the stability of our results comparing the model outcomes with the ones produced by an additional ERM built with the same operating choices of the European SHARE model (Woessner et al., 2015). The reviewers also asked to develop a new GMPE that adopted the so-called “backbone” approach; this new equation was then not implemented in MPS19 model, because not yet published in peer review journal, but we verify that the outcomes were comparable with MPS19. Since the testing phase is one of the most important points in our project, the reviewers asked to implement more the testing phase, for example evaluating the reliability of some crucial assumptions, such as the stationarity of seismicity and the suitability of the Gutenberg-Richter law, in some regions were the data showed some apparent discrepancies from these assumptions.

The reviewers shared their final evaluation on MPS19 model in a document, which was submitted to DPC. The main points of the evaluation are reported in the following, in its Italian formulation and then translated in English by automatic translator (Google Translate), in order to not introduce any subjective consideration. The few changes introduced by us are reported in italics.

*L'operazione di test [...], a rigore non è definibile come "validazione" - che è una procedura statistica ben codificata - visto che il learning set di dati e il voting set di dati non sono indipendenti. Questo confronto, tuttavia, fornisce un importante controllo di coerenza (sanity check): se i risultati di un modello sono in disaccordo con la sismicità storica, potranno credibilmente essere in accordo con quella futura? Da questo punto di vista, MPS19 costituisce un indubbio progresso non solo rispetto a MPS04 ma anche, a conoscenza del GDL, a tutte le procedure PSHA che sono state adottate per derivare la mappe di pericolosità nazionali attualmente esistenti al mondo.*

*Un ulteriore elemento significativo della robustezza del modello finale è dato dalla sostanziale coerenza tra tassi di eccedenza di intensità macrosismiche osservate e quelle calcolate dal modello. Questo confronto, pur con tutte le limitazioni del caso dovute all'incertezza sui dati osservati e alla completezza degli stessi, è importante, perché i dati di intensità macrosismica non sono stati usati in modo esplicito durante lo sviluppo del modello MPS19, se si eccettua l'utilizzo per la stima della magnitudo di eventi pre-strumentali. Inoltre, ancorché non richiesto dal GDL e non discusso nella bozza di rapporto del 27 febbraio 2019, è importante segnalare che l'adeguatezza del modello MPS19 è stata anche testata mediante un confronto, presentato nella riunione del 27 febbraio 2019, tra le stime dei tassi di eccedenza di PGA misurate in stazioni accelerometriche operative da almeno 30 anni e quelle predette dal modello. Anche in questo caso le stime del modello MPS19 dei tassi di eccedenza delle PGA sono risultate compatibili con quelle osservate.*

*Per quanto riguarda i risultati pratici, il pregio fondamentale di MPS19 è il mostrare chiaramente le incertezze delle stime di pericolosità, che nel caso del modello ensemble hanno un ragionevole grado di credibilità. Poiché l'incertezza costituisce un aspetto fondamentale per il normatore, il GDL suggerisce di presentare le mappe fianco a fianco con media e media+deviazione standard delle accelerazioni spettrali a 0.5, 1 e 3 Hz, corrispondenti ai periodi di ritorno più rappresentativi (ad esempio, 475 anni, 975 anni e 2475 anni). Allo stesso modo, il GDL suggerisce di presentare per i suddetti periodi di ritorno le mappe con la stima di PGA di MPS19 fianco a fianco a quelle di MPS04 utilizzando la stessa scala di colori, in modo da garantire una comprensione immediata delle differenze.*

*Infine, si esprime soddisfazione per l'onestà intellettuale portata dai ricercatori nell'interazione con il GDL e la trasparenza con cui MPS19 è stato sviluppato, caratteristiche che costituiscono la migliore garanzia per ulteriori avanzamenti. Gli aggiornamenti futuri delle mappe di pericolosità derivate da MPS 18 troveranno dei fondamenti robusti su cui basarsi.*

The test operation [...], strictly speaking, cannot be defined as "validation" - which is a well-coded statistical procedure - since the learning data set and the voting data set are not independent. This comparison, however, provides an important sanity check: if the results of a model disagree with historical seismicity, can they credibly agree with the future one? From this point of view, MPS19 represents an undoubted progress not only compared to MPS04 but also, to the knowledge of the GDL (*the reviewers' panel*), to all the PSHA procedures that have been adopted to derive the national hazard maps currently existing in the world.

A further significant element of the robustness of the final model is the substantial coherence between the observed macroseismic intensity excess rates and those calculated by the model. This comparison, even with all the limitations of the case due to the uncertainty on the observed data and the completeness of the same, is important, because the data of macroseismic intensity were not used explicitly during the development of the MPS19 model, except for the 'use for estimating the magnitude of pre-instrumental events. Furthermore, although not required by the GDL and not discussed in the draft report of 27 February 2019, it is important to point out that the adequacy of the MPS19 model was also tested by means of a comparison, presented at the meeting of 27 February 2019, between the *PGA exceedances* rate measured in accelerometric stations operating for at least 30 years and those predicted by the model. Also in this case, the estimates of the MPS19 model of the *PGA exceedances* rate were found to be compatible with those observed.

As far as practical results are concerned, the fundamental advantage of MPS19 is that it clearly shows the uncertainties of the hazard estimates, which in the case of the ensemble model have a reasonable degree of credibility. Since uncertainty is a fundamental aspect for the regulator, the GDL suggests presenting the maps side by side with mean and mean + standard deviation of the spectral accelerations at 0.5, 1 and 3 Hz, corresponding to the most representative return periods (for example , 475 years, 975 years and 2475 years). Similarly, the GDL suggests presenting for the aforementioned return periods the maps with the *PGA* estimate of MPS19 side by side with those of MPS04 using the same color scale, in order to ensure an immediate understanding of the differences.

Finally, satisfaction is expressed for the intellectual honesty brought by the researchers in the interaction with the GDL and the transparency with which MPS19 was developed, characteristics that constitute the best guarantee for further progress. Future updates of MPS19-derived hazard maps will find robust foundations to build upon.

## **APPENDIX C. Brief description of ERM**

### **1 Area source models (MA)**

#### **1.1 MA1**

*(contribution by M. Santulin, A. Tamaro, D. Slejko, F. Sani, L. Martelli, M. Bonini, G. Corti, M. E. Poli, A. Zanferrari, A. Marchesini, M. Buseti, M. Dal Cin, D. Spallarossa, S. Barani, D. Scafidi, G. Barreca, C. Monaco, A. Rebez)*

The seismotectonic zoning model (SZ) of MA1 was built by integrating new and updated seismotectonic data into the Meletti et al. (2008) zoning scheme. Where the difference between the new SZs and those of the Meletti et al. (2008) were found negligible, in terms of geographical boundaries and seismotectonic characteristics, the SZ9 zones were used. Individual seismicity rates have been computed for each zone, using the declustered CPTI15 catalogue (Rovida et al., 2016), with the completeness provided by Meletti et al. (2019), using two different approaches: (i) zones were grouped into some macroareas to evaluate the b-value and then this b-value is used to calculate seismicity rates into the zones according to a Gutenberg-Richter magnitude-frequency distribution (the zones are kept independent for the a-value computation) and; (ii) observed seismicity rates for each zone. MA1 explores uncertainties on maximum magnitude, completeness approaches (historical and statistical) and approach for seismicity rates (GR and observed).

#### **1.2 MA2 and MA3**

*(contribution by R. Basili, P. Burrato, U. Fracassi, V. Kastelic, G. Tarabusi, M. M. Tiberti, G. Valensise, P. Vannoli)*

The two models, MA2 and MA3, are here described together as they use the same approach to evaluate seismicity rates but different definition of area sources. Both models used a two-level area source model. MA2 and MA3 defined a 2D subdivision of the crustal space in large regions consistent with geotectonic properties inspired by concepts and definitions of plate tectonics (Bates and Jackson, 1980), and already used in earthquake hazards studies (Delavaud et al., 2012; Selva et al., 2016; Basili et al., 2020). In addition to these large regions, MA2 and MA3 defined two different seismotectonic zoning, that were topologically coherent subdivisions of the former, considering geotectonic properties at the local scale, including the presence of volcanoes. For the seismicity annual rate calculations, MA2 and MA3 used large regions to estimate b and the corner moment of a tapered-Pareto frequency-moment. Seismotectonic zone were then used to estimate the annual rate (a-value) by anchoring the theoretical FMD to the observations from the CPTI15 earthquake catalog. MA2 and MA3 considered uncertainties on completeness studies, and partitioning rates approaches (from superzones to zones).

### **1.3 MA4**

*(contribution by C. Meletti, F. Visini, B. Pace, V. D'Amico, A. Rovida, S. Pondrelli)*

SZ of MA4 represents an update the previous model SZ9 (Meletti et al., 2008), which is the reference source model for MPS04. MA4 is consistent with the general background delineated by the geodynamic model proposed for MPS04 and incorporate all recent advances on the understanding of the active tectonics of the Italian territory and on the distribution of seismogenic sources (e.g., DISS Working Group, 2018). To estimate the expected seismicity rates, firstly zones were grouped into macroareas, according to their tectonic regime. Assuming that the distribution of the earthquake sizes follows a truncated Gutenberg-Richter model,  $b$  and  $a$  parameters were estimated applying the Weichert (1980) approach to each macroarea. Then, 4 approaches to distribute activity rates from a macroarea to the zones included, were used. In addition, Weichert (1980) was also applied to the single zones, but upper and lower bounds to  $b$  was introduced. MA4 explored uncertainties on completeness studies (historical and statistical), maximum magnitude and activity rates evaluations (from macroareas to zone).

### **1.4 MA5**

*(contribution by R. Rotondi, E. Varini)*

MA5 uses the same SZ of MA4, but it applied a different approach, in respect of MA4, to estimate the seismicity rates. The MA5 estimated jointly the completeness time and the seismicity rate of the complete part of data sets associated with each seismogenic zone and constituted by events of magnitude exceeding a varying threshold (Rotondi and Garavaglia, 2002). Assuming that the main earthquakes follow a homogeneous Poisson process with constant rate, and that the beginning of the complete part of the catalog corresponds to a change in the Poisson parameter, the problem was translated into a change-point problem which combines stochastic simulation methods with a Bayesian hierarchical model (Pievatolo and Rotondi, 2008). Assuming that the process has two rates, one in the incomplete part and another in the complete one, following the Bayesian approach, both the two rates and the time when the process changes rate were considered as random variables, and the method computed, in addition to their estimates, the entire probability distribution of these variables as well.

## **2 Fault-based models (MF)**

### **2.1 MF1**

*(contribution by R. Basili, P. Burrato, U. Fracassi, V. Kastelic, F. E. Maesano, , G. Tarabusi, M. M. Tiberti, G. Valensise, P. Vannoli)*

MF1 is a fault-based seismicity model using exclusively geological information taken from the DISS 3.2.1. Other datasets, such as earthquake catalogues, focal mechanisms, stress orientations, and crustal models were used to constrain various parameters of the model but not the seismicity rates. Although the model was based on fault information only, an empirical proximity function has been specifically

developed to predict off-fault seismicity, i.e. the occurrence of earthquakes of any size away from the faults. These characteristics ensured the independence of the seismicity rates predicted by this model from the earthquake catalogue, thereby providing a valid alternative to explore the epistemic uncertainty in PSHA. The basic idea of this model was that potentially seismogenic faults can provide a representation of the seismicity on a time frame much longer than that of any earthquake catalogue. The seismic moment rate  $\dot{M}_s$  of a seismogenic fault is derived from the geologic moment rate, by multiplying the shear modulus, fault length and width, slip rate, and a coefficient (often referred to as seismic efficiency) that determines how much of the geologic rate is converted into the seismic rate. Using the moment conservation principle, the seismic moment rate of any seismogenic fault was related to the annual earthquake rate of a Truncated GR magnitude-frequency distribution using the formulations by Kagan (2002a; 2002b) and assuming a unique beta equal to 0.667 for all the faults. To evaluate the off-fault seismicity rates, MF1 followed an empirical approach aimed to capture the natural distribution of the observed earthquakes around the faults and the respective location uncertainties. To this aim, an empirical proximity function representing the frequency at which an earthquake has occurred at any distance from a mapped fault was developed. Then, the earthquakes predicted by the calculated magnitude-frequency distributions, as previously described, are symmetrically spread around proportionally to the magnitude and distance pairs dictated by the empirical proximity function.

## **2.2 MF2**

*(contribution by M. Murru, R. Console, G. Falcone, M. Taroni)*

The MF2 model combines the seismic rates obtained by three different databases in order to cover all the Italian territory: seismogenic sources, the historical seismic catalogue CPTI15 and the Instrumental catalogue provided by Meletti et al. (2019). For the historical catalog, two different criteria of completeness were used. In this approach we adopted over all Italy a tapered Gutenberg-Richter relation with a unique b-value (0.99) and a unique corner magnitude (7.3 Mw) estimated by the catalogs.

As regards the seismogenic sources, MF2 considered a revised group of 126 individual sources from DISS Working Group (2018). The seismic moment rate of each individual source was computed using the seismic-moment conservation principle (Field et al., 1999), that allows to derive seismic moment rate from long-term slip rate and geometry of the fault source. Successively, balancing the seismic moment rate over a tapered Gutenberg-Richter magnitude-frequency distribution, from magnitude greater or equal than Mw 4.5, seismicity rates of occurrence were evaluated for each individual source. Moreover, using the smoothed-seismicity approach proposed by Frankel (1995), with the correction for time varying completeness magnitude (Hiemer et al., 2014), having as input the combined historical and instrumental catalogue, rates of occurrences for Mw greater or equal than Mw 4.5 were computed over a grid. Finally, for cells of the grid intersected by a portion of the surface projection of an individual fault, the rates of occurrence evaluated by the smoothed seismicity approach, were combined with the rates estimated for that individual source, in a quantity proportional to the area of the intersection. Considering the three

different combinations of the seismogenic source rate and seismic catalog rate, we obtained a total of six different ensemble models.

### **3 Smoothing seismicity models (MS)**

#### **3.1 MS1**

*(contribution by A. Akinici, M. Moschetti, M. Taroni)*

MS1 model is developed using smoothed seismicity methods for long-term earthquake rate space-time. A forecast of the probabilities of ( $M \geq 4.5$ ) earthquakes is calculated from ensemble smoothed seismicity models derived from longer Italian historical (Catalogue Parametrico dei Terremoti Italiani, CPTI15, 1000-2014) and shorter instrumental (1981-2016) earthquake catalogs. Two different smoothed seismicity methods are adopted for the MS1 model construction following: 1) the well-known and widely applied fixed smoothing (Frankel, 1995) such that the kernels used to smooth catalog-derived seismicity rates were invariant to spatial variations in seismicity rate and 2) the most recent adaptive smoothing method (Helmstetter et al., 2007) which determines the smoothing distance for each earthquake from its distance to neighboring earthquake epicenters. The optimized fixed smoothing distances and adaptive neighbor numbers are identified ranking the ability of all of the trial smoothed seismicity models to forecast the location of earthquakes in the recent part of the earthquake catalog with a likelihood parameter. Finally, MS1 is created as an ensemble model that combines two different equally weighted smoothing models (adaptive and fixed) from the two earthquake catalogs (historical and instrumental), through a logic tree approach to improve the forecast capability. MS1 takes into account and considers the uncertainties in the model constructions using different percentage combinations of the models obtained from two different smoothing approaches (for details see, Akinici et al., 2018).

#### **3.2 MS2**

*(contribution by C. G. Lai, E. Zuccolo)*

MS2 followed the Woo (1996) approach and proposed a zone-free method solely based on the use of the earthquake catalogue. Proxy seismogenic sources were created from the epicentral locations of the events that are smoothed according with their fractal distribution in space.

A square sub-grid of point sources (approximate dimensions  $0.1^\circ \times 0.1^\circ$  with step 0.2 km) was defined around each node of the a grid and the activity rates of each point source were calculated from the density and proximity of events lying within the considered magnitude range. The contribution of each earthquake to the seismicity of the region was smeared over all sub-grid points falling within an epicentral distance that depends on the magnitude of the event itself.

Finally, the activity rates referred to each node of the grid were computed by summing the activity rates of the point sources multiplied by the area of the sub-cells.

To consider the uncertainties associated with some model assumptions (namely kernel shape, kernel directionality, completeness periods), the MS2 model was created by producing 8 models that were integrated, with proper weights, within a logic tree framework.

## **4 Geodetic models (MG)**

### **4.1 MG1**

*(contribution by N. D'Agostino)*

MG1 is a geodetically-derived seismicity model based on the analysis of 919 GPS horizontal interseismic velocities and on the hypothesis that future seismic moment release will scale linearly with the rate of present-day strain accumulation (Ward, 1998; Ward, 2007). The strain rate tensor field has been a) calculated on a regular  $0.1^\circ \times 0.1^\circ$  grid using the VISR software (Shen et al., 2015) taking into account the variable station spacing for the optimal smoothing parameters, and b) converted in rates of seismic moment accumulation following Kostrov (1974). The rate of seismic moment accumulation density was finally converted to earthquake rate density (or earthquake potential) under the assumption that seismic moment distributes into earthquake sizes that follow a tapered Gutenberg- Richter distribution of given  $b$ -value and corner magnitude (Kagan, 2002b). The effect of uncertainties and spatial variations in the seismogenic thickness and  $b$ -values has been introduced by considering alternative realizations of the seismicity model and considering the 50th, 5th and 95th percentiles as representative of the median and associated uncertainties of the final seismicity model. To account for the effects of declustering and aseismic component, the forecasted seismicity rate has been scaled to the seismic moment release of the declustered CPTI15 catalogue in the interval 1787-2015. Following the approach described above the seismicity rates forecasted by MG1 satisfy the overall 1787-2015 CPTI15 historical moment release while respecting the spatial variations shown by the geodetic strain rate field.

### **4.2 MG2**

*(contribution by M. M. C. Carafa, P. Bird, V. Kastelic, G. Valensise)*

MG2 uses GPS interseismic horizontal velocities and stress-azimuth data to determine both the interseismic and long-term strain-rates and velocities on a finite element grid with NeoKinema code version 5.2 (Bird and Liu 2007, Bird 2009, Bird and Carafa 2016). Short term geodetic and non-tectonic (e.g. landslide movements) transients are modeled as principal factors in the covariance matrix (Carafa and Bird 2016) before running NeoKinema, and the strain rate in Calabria is calculated only for the upper plate under the assumption of a locked subduction (Carafa et al., 2018). The methodology to compute models of the long-term earthquake rates density for shallow earthquakes from the long-term strain rate follows Bird and Liu (2007), Bird et al. (2010), and Bird and Kreemer (2015). The parameters of the Tapered Gutenberg-Richter frequency-magnitude relations and seismic coupling were computed at the regional scale of the broader study area. The seismic moment rate for each cell of the grid is calculated in a two-step procedure: first, the interseismic strain-rates of the portion of the finite elements falling in the

individual grid cell are converted to long-term tectonic moment rates; then the long-term tectonic moment-rates are converted to seismic moment rates.

## **APPENDIX D. ERM for Etna, Tyrrhenian subduction, and sources outside Italy.**

### **ERM for Etna**

*(contribution by R. Azzaro, G. Barberi, S. D'Amico, T. Tuvè, R. Gee, B. Pace, L. Peruzza)*

The ERM for Etna has been adapted from Azzaro et al. (2017) and Peruzza et al. (2017). The ERM includes a rather high degree of detail and complexity, motivated by the great availability of geological, seismological and geodetic data for the area. The Etna original ERM consists of 3 versions, in which all these complexities are explored, and the hazard estimates are elaborated taking into account the topography that influences the site-source distances. For the purposes of application to the MPS19 project, the time-dependent model is discarded, and the focus is exclusively on the Poisson model. Furthermore, the topography is not incorporated, due to homogeneity with the ERM on a national scale, but the sources are all modeled with respect to the sea level. The final ERM consists of a combination of individual sources (seismogenic faults) and gridded seismicity. The individual sources estimate the highest seismicity using a historical approach, and therefore with the rates estimated by maximum magnitudes observed and relative average times of recurrence; the gridded seismicity estimates the lower magnitude seismicity ( $3.5 < M < 4.5$ ) with point sources with variable depth.

### **ERM for subduction**

*(contribution by R. Basili, F. E. Maesano, M. M. Tiberti, F. Visini)*

The ERF input for the subduction of the Calabrian Arc uses the slab geometry from Maesano et al. (2017), specifically created for the tsunami hazard model NEAMTHM18 (Basili et al., 2020) and MPS19. We consider a shallowest portion of the slab, corresponding to the surface of the slab interface between 18 and 40 km of depth (the limits correspond respectively to the main change of slope of the slab below the accretion prism and at the intersection with the Moho of the upper plate), and a deepest portion that corresponds to the actual intraslab, between 20 and 440 km depth. For the intraslab, we resampled the three-dimensional geometry of the intraslab with a 10-km-spaced grid of points. The slab thickness was taken equal to the crustal thickness (seafloor to Moho) measured in the undeformed part of the subduction (Ionian Sea) which then constrained other seismogenic parameters such as the expected nodal planes that were assigned to the grid points to have earthquake ruptures at  $45^\circ$  angle with respect to the local slab dip and their size confined within the subduction body. (4) the expected seismicity rates were calculated according to the Weichert (1980) approach using, as input data, the deep earthquakes in the CPTI15 catalog, the statistical approach for estimating completeness, and a magnitude-frequency distribution of the type Tapered, with corner magnitude equal to  $M_w 7.5$ . The total magnitude-frequency distribution the seismicity rate was uniformly distributed on the grid points. It is worth noting that two of the eleven ERM, specifically MF1 and MG2, had been able to separate seismicity due to the interface from the crustal tectonic seismicity, therefore we added only to them the interface ERM. Both the two MF1 and MG2 ERF have then failed the regional statistical tests for this particular subregion, and their earthquake

rates were replaced in that subregion by the weighted average of the other consistent ERMs (according to the procedure described in the section 4 of the main text). The intraslab ERM was added to all the ERF inputs.

### **ERM for external sources**

*(contribution by F. Visini, C. Meletti)*

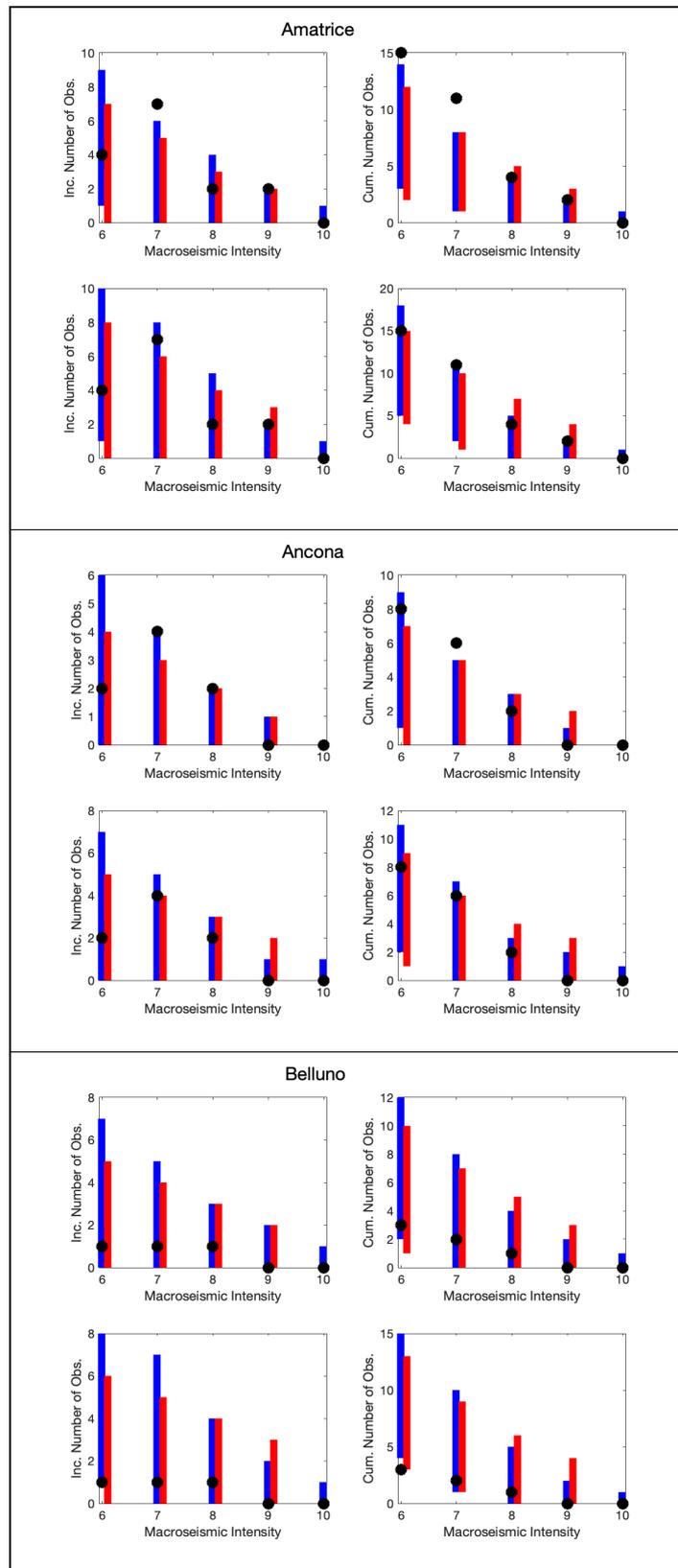
To parametrize sources located outside reliability area of the CPTI15 catalog, and to minimize edge effects, we consider the ESHM European hazard model (Woessner et al., 2015), which, by extension, methodologies and typologies of ERMs, is deemed useful for the purpose. The ESHM model consists of three types of seismicity models: areas; faults and background (FSBG); smoothed seismicity and faults (SEIFA). To avoid as much as possible the overlapping of MPS19 sources with the external ones, the three types of seismogenic sources of ESHM sources have been treated. The area source model and the FSBG are resampled according to a regular grid spacing  $0.2^\circ \times 0.2^\circ$ . To each point of this grid, we assign a weighted seismicity rate of the three models, in accordance to the weights of ESHM for the return periods between 475 and 2475 years, and respectively for area, FSBG and SEIFA of 0.5, 0.2.0.3. For each of the 11 ERM, only the points located outside the area covered by the ERM and within 100 km from the outer edge, are included.

## **APPENDIX E. Description of the experts' elicitation session.**

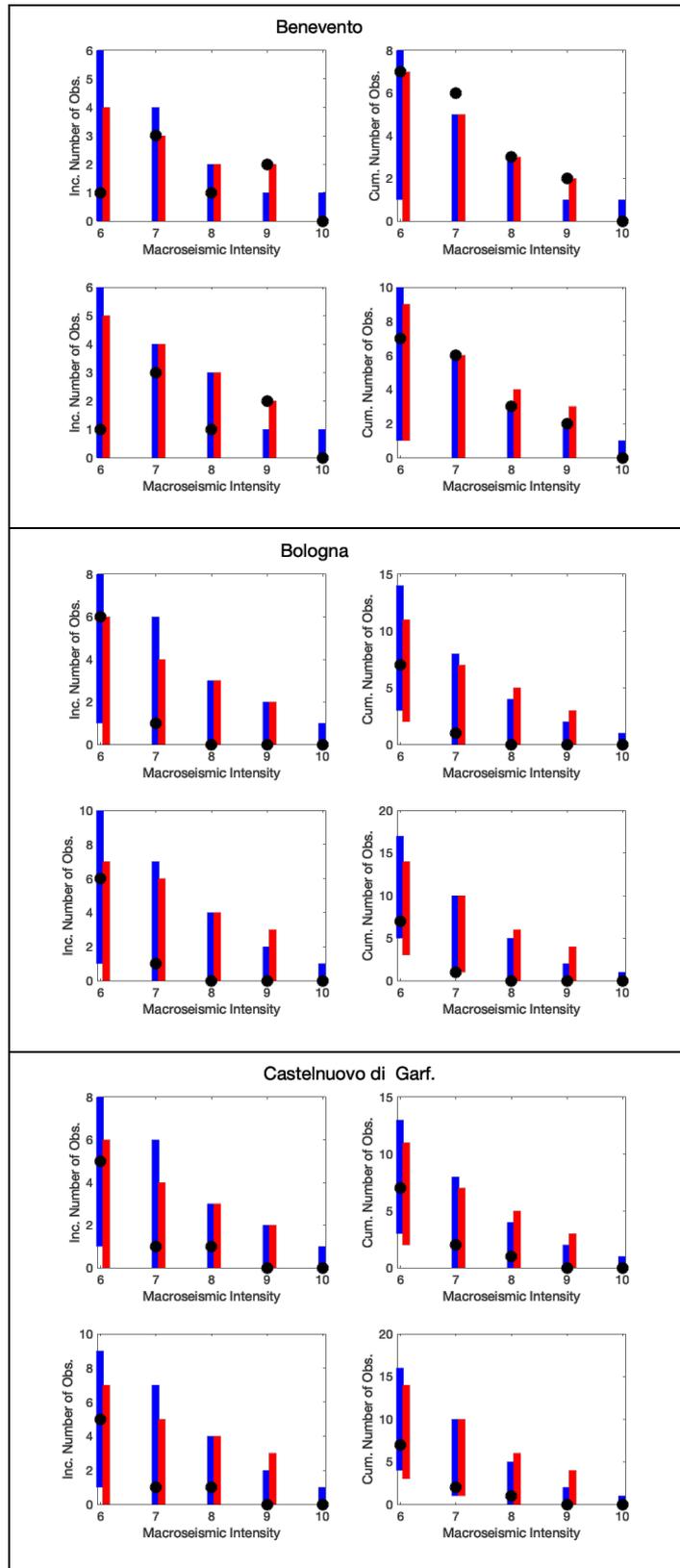
The elicitation sessions (one for ERMs and one for GMMs) are carried out by a group of experts (GoE) which have not been involved in the preparation of the models: Warner Marzocchi, Carlo Meletti (only for GMM), Dario Albarello, Iunio Iervolino, Marco Pagani, plus two additional external experts, Danijel Schorlemmer from GFZ, and Laurentiu Danciu from ETH. To avoid overlapping with the results of the scoring statistics, and to avoid well-known cognitive biases, such as anchoring and confirmation (Kahneman 2011), GoE members are not allowed to see the statistical scoring (possible anchoring), but only to evaluate the scientific coherence of the models. The confirmation bias is likely pervasive on PSHA, i.e., the human tendency to favor models that confirms one's preexisting beliefs of how the seismic hazard should be – e.g., the last seismic hazard model – while giving disproportionately less consideration to alternative possibilities. To minimize this possibility, experts' are asked to evaluate the reliability of each ERM and GMM without knowing at their overall impact in terms of seismic hazard.

For the elicitation sessions we use one procedure that is named Analytic Hierarchy Process (AHP; Saaty, 1980). AHP is a multi-criteria decision-making method that is useful to make decisions under complex problems. The hierarchy process breaks down the complex decisions into a series of pairwise comparisons, synthesizes the results, and then helps to take into account both subjective and objective aspects of the decision. Additionally, the process incorporates a useful technique for checking the consistency of each expert judgments, thus reducing bias and incoherent results. Specifically, in two separate sessions experts are elicited on the prioritization of ERM and GMM compared pairwise, and the results provide a sort of ranking of the scientific credibility of all ERMs and GMMs. Results are shown in the manuscript, and more details are in the final report of MPS19 (in italian).

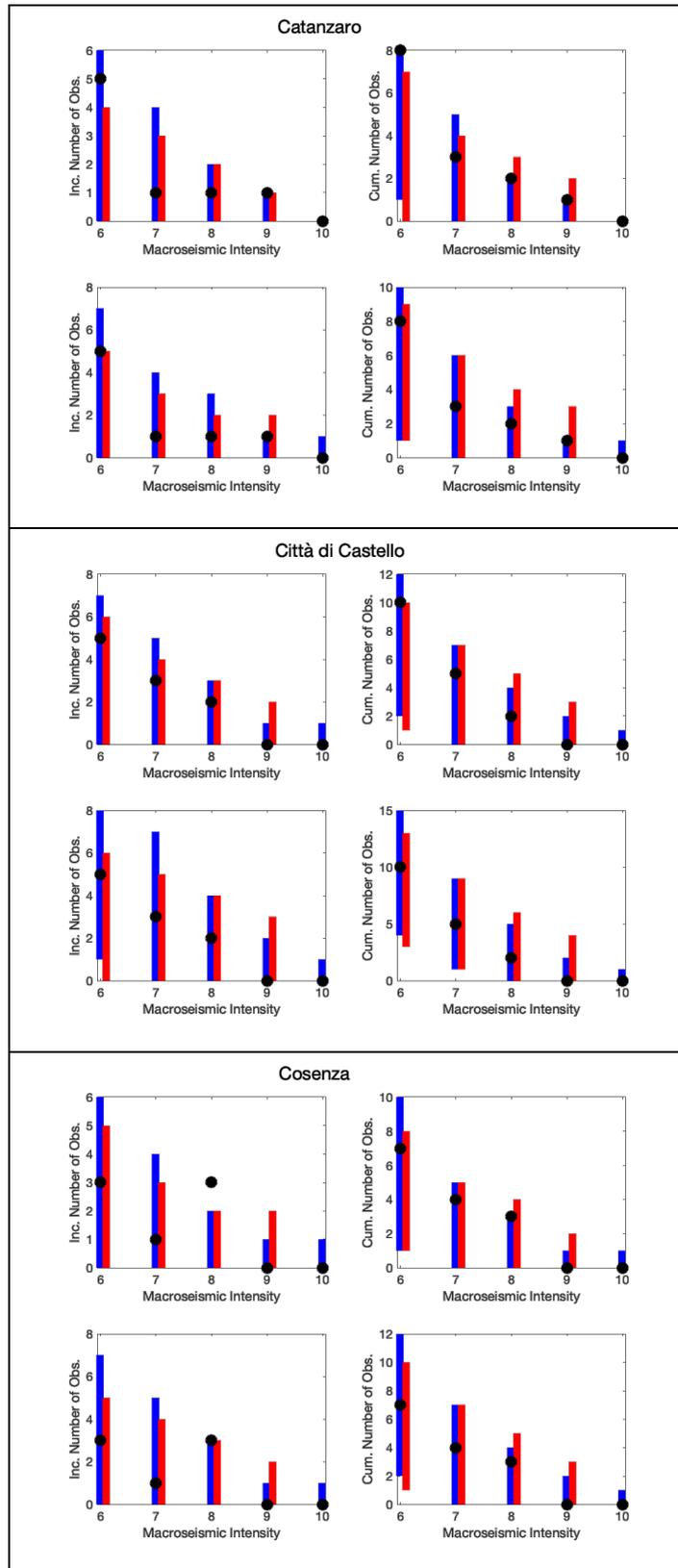
## APPENDIX F. Test of the macroseismic intensity at different sites



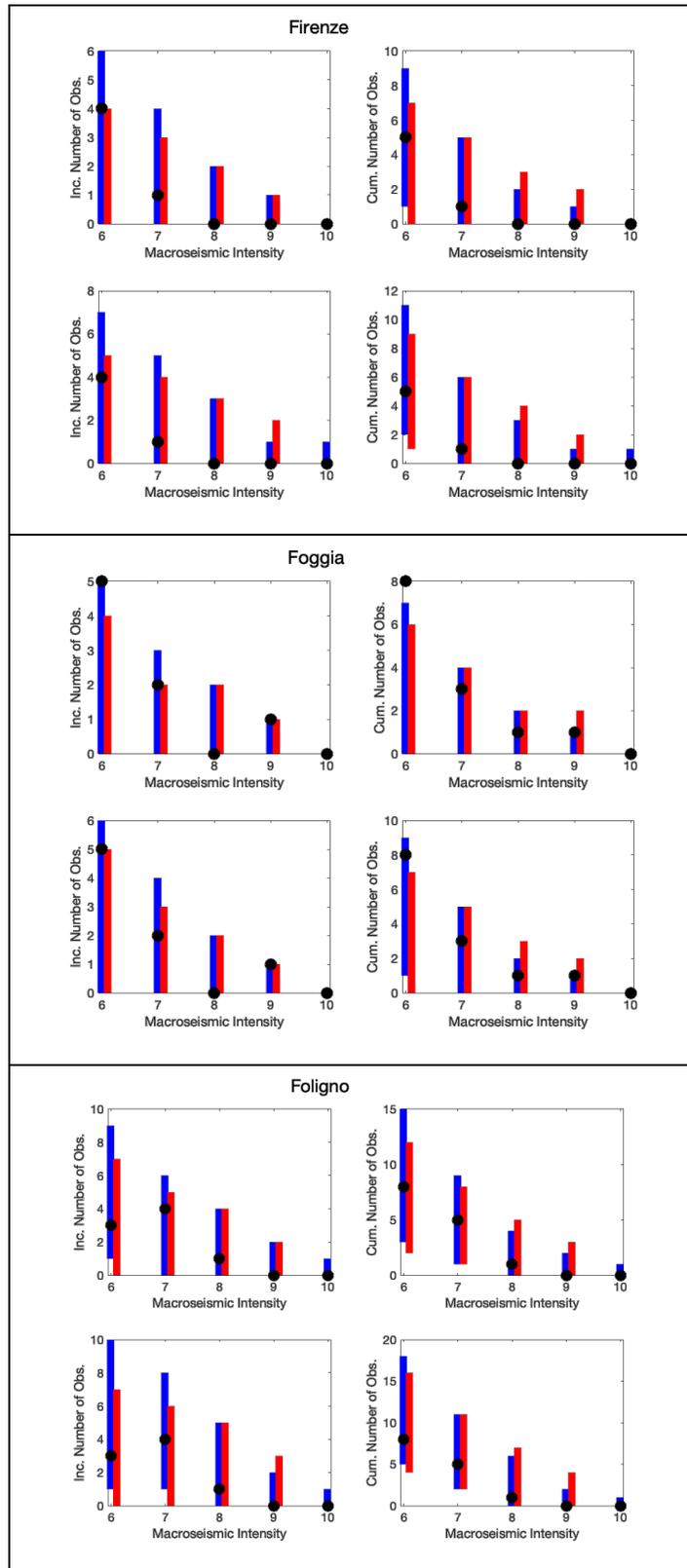
As for Figure 10 in the text but relative to Amatrice (top), Ancona (center), Belluno (below)



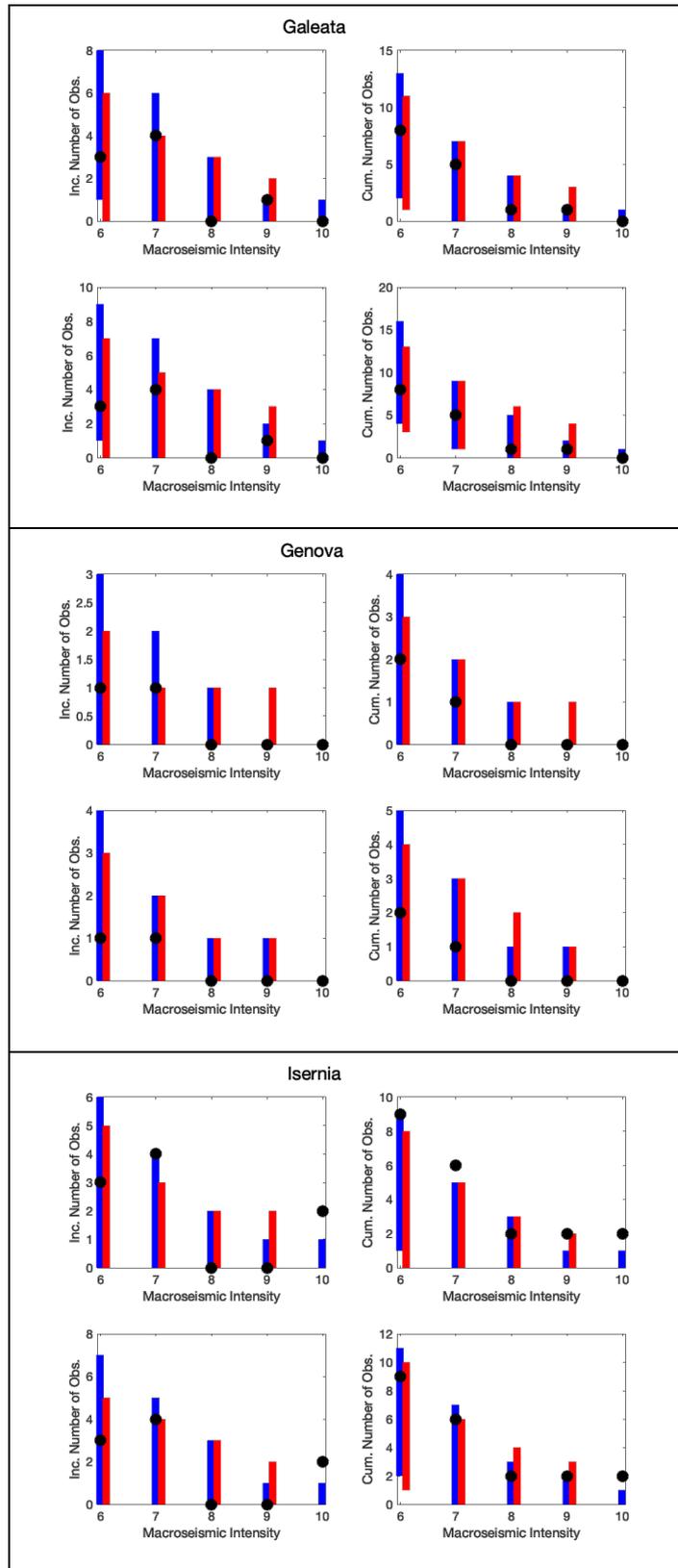
As for Figure 10 in the text but relative to Benevento (top), Bologna (center), Castelnuovo di Garfagnana (below)



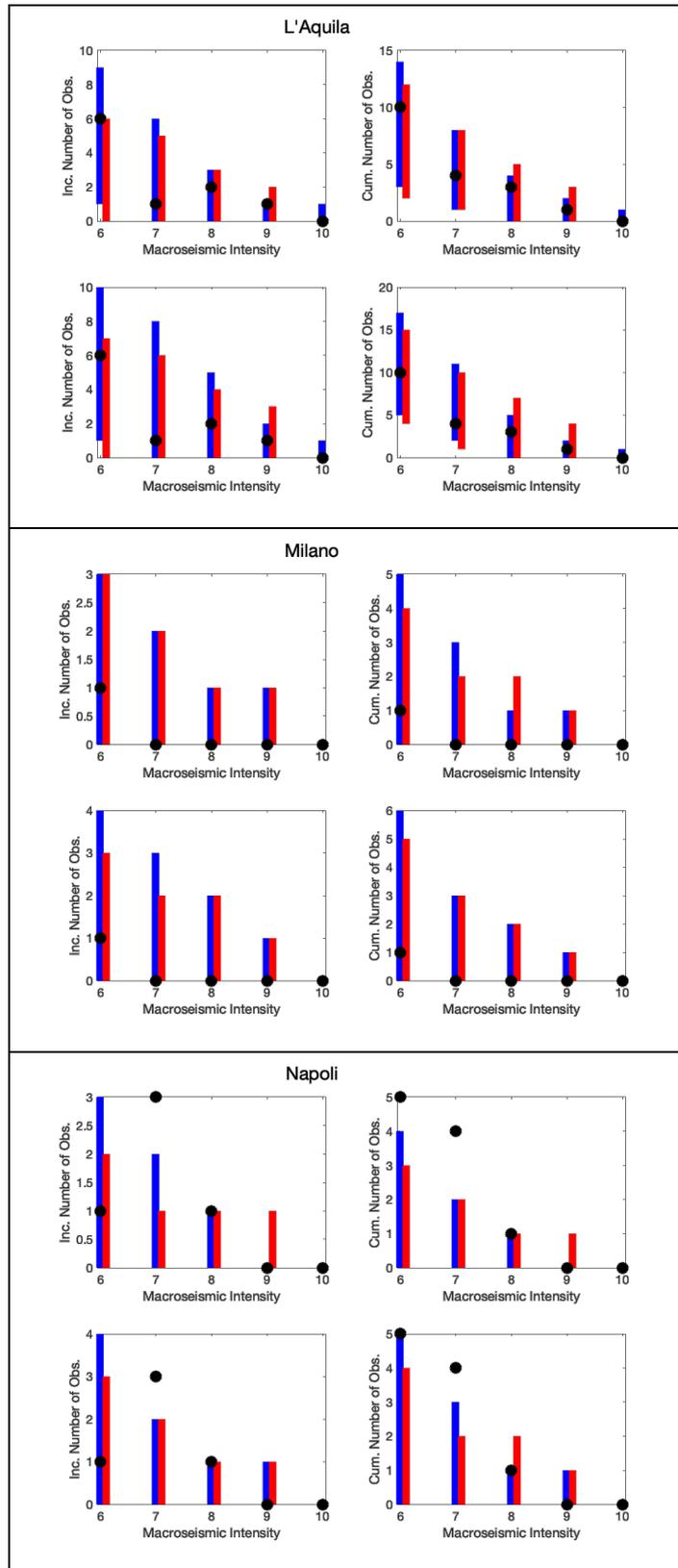
As for Figure 10 in the text but relative to Catanzaro (top), Città di Castello (center), Cosenza (below)



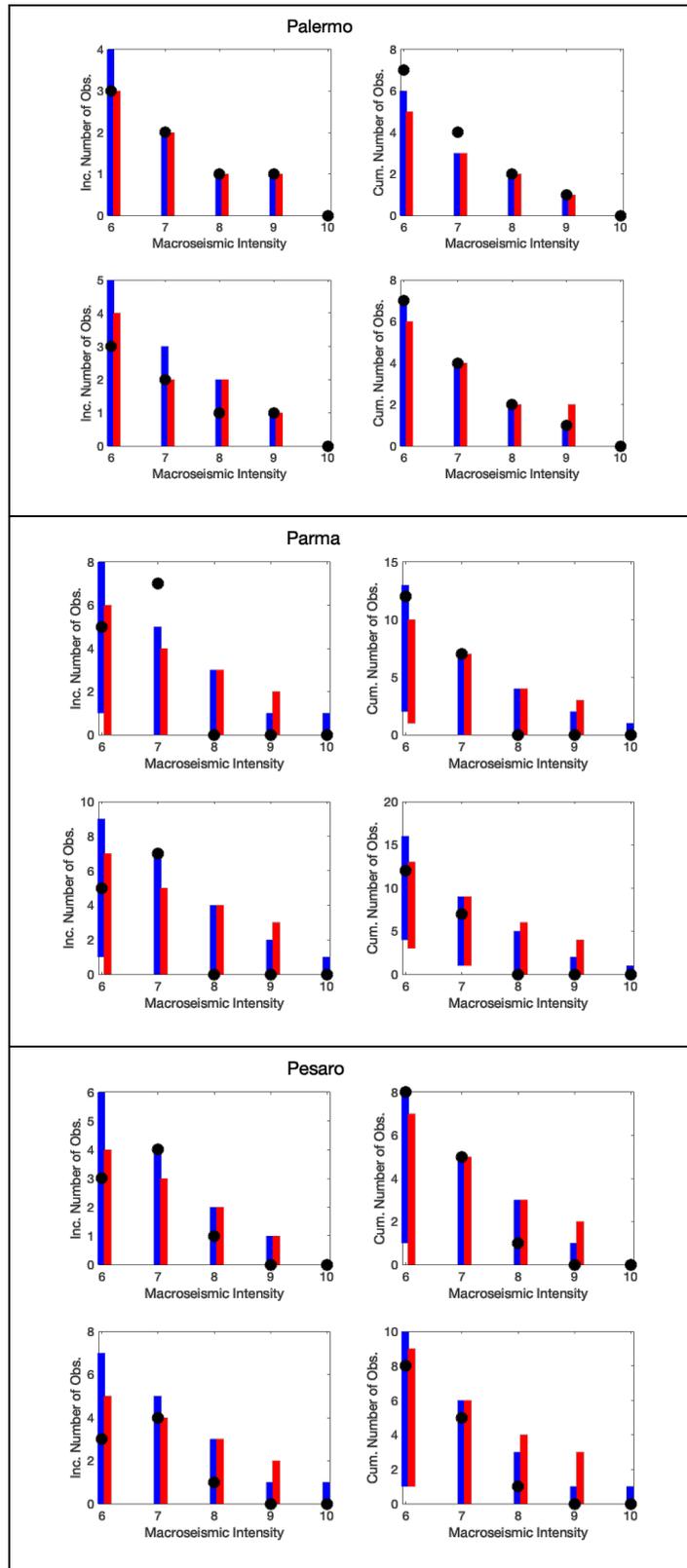
As for Figure 10 in the text but relative to Firenze (top), Foggia (center), Foligno (below)



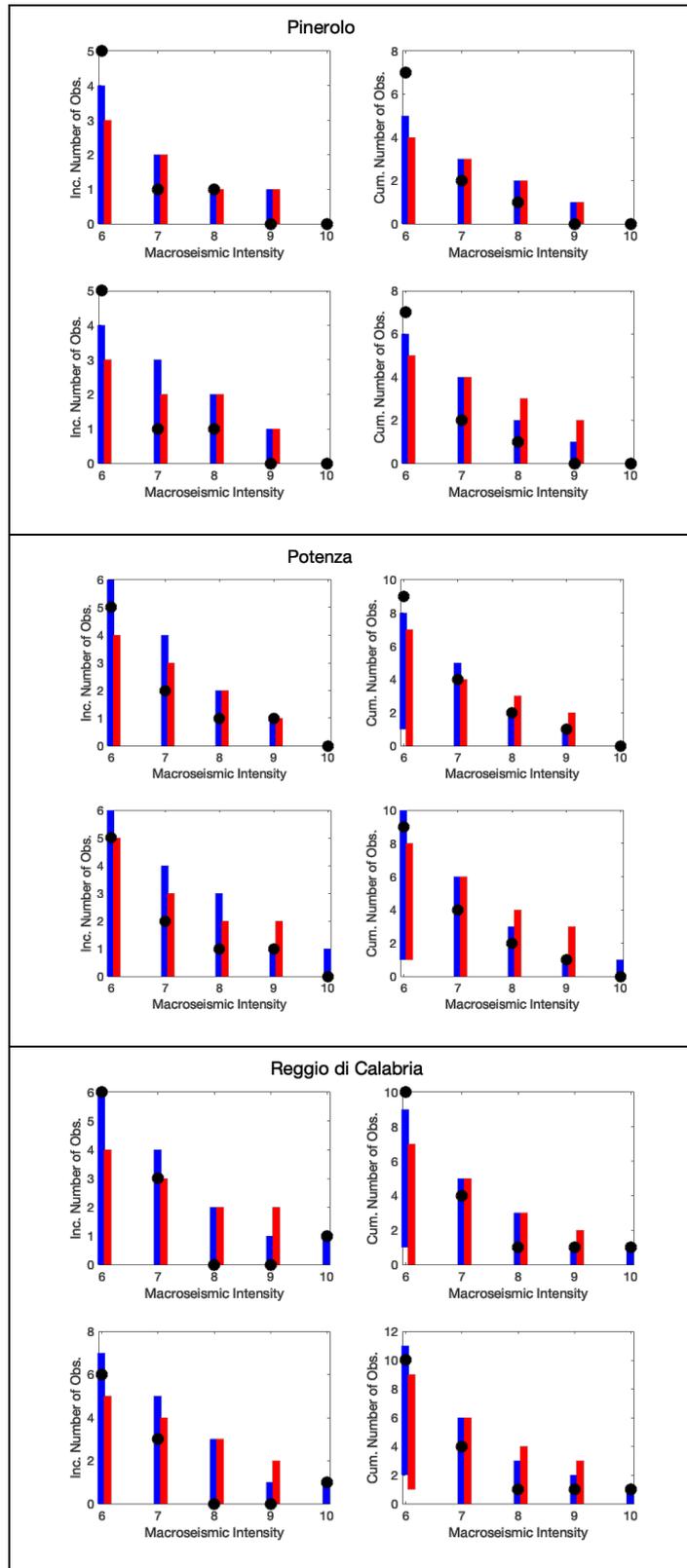
As for Figure 10 in the text but relative to Galeata (top), Genova (center), Isernia (below)



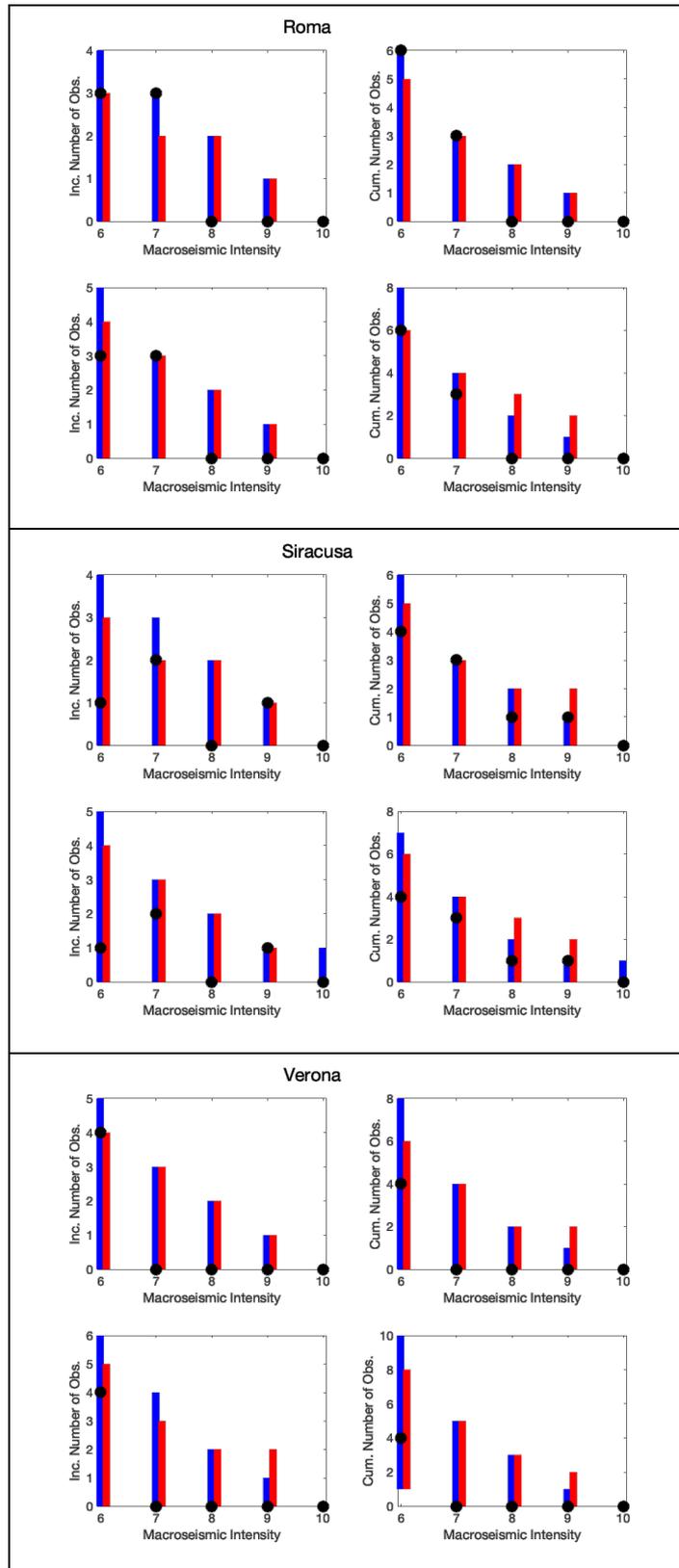
As for Figure 10 in the text but relative to L'Aquila (top), Milano (center), Napoli (below)



As for Figure 10 in the text but relative to Palermo (top), Parma (center), Pesaro (below)



As for Figure 10 in the text but relative to Pinerolo (top), Potenza (center), Reggio di Calabria (below)



As for Figure 10 in the text but relative to Roma (top), Siracusa (center), Verona (below)

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