



Originally published as:

Lacombe, O., Mazzoli, S., von Hagke, C., Rosenau, M., Fillon, C., Granado, P. (2019): Style of deformation and tectono-sedimentary evolution of fold-and-thrust belts and foreland basins: From nature to models. - *Tectonophysics*, 767, pp. 228163.

DOI: <http://doi.org/10.1016/j.tecto.2019.228163>

1 **Style of deformation and tectono-sedimentary evolution of fold-and-thrust**
2 **belts and foreland basins: from nature to models**

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30 **1. Fold-and-thrust belts : some recent advances in their description and their**
31 **understanding**

32

33 Orogenic systems, including their external fold-and-thrust belts and foreland basin
34 systems, generally evolve from the inversion and imbrication of former continental margins.
35 Continental margins are characterized by displaying along-strike variations in the degree of
36 inherited regional extension (i.e., from limited lithospheric stretching to full necking, leading
37 to mantle exhumation and oceanic accretion). These differences have a fundamental impact on
38 the pre-shortening thermal state of the lithosphere and on structural style development.

39 Indeed, one of the key processes in fold-and-thrust belts is the reactivation and inversion
40 of pre-existing extensional faults. Inversion tectonics is widespread during the evolution of
41 many orogens and this process can exert a strong control on the structural and mechanical
42 evolution of fold-and-thrust belts (Lacombe and Bellahsen, 2016).

43 The presence of evaporitic sequences interacting during lithospheric stretching and
44 subsequent thermal subsidence is also a key parameter in the structural styles and deformation
45 distribution of thrust sheets involving inverted basins and salt structures. On the other hand,
46 deformation can also be transferred ahead and downward of the shallow deformation front,
47 leading to frontal imbrication of deep seated structures in cratonic forelands or the sub-thrust
48 region of active fold-and-thrust belts.

49 Defining the correct structural style of fold-and-thrust belts and understanding the
50 controlling factors are necessary steps towards predicting their long-and short-term evolution,
51 with implications for crustal/lithospheric rheology, mountain building processes and seismic
52 hazard, and for the correct assessment of their potential for hydrocarbon exploration (e.g.,
53 Butler and Mazzoli, 2006; Lacombe et al., 2007; Poblet and Lisle, 2011; Lacombe et al., 2016).
54 For these reasons, fold-and-thrust belts and adjacent foreland basin systems represent

55 outstanding places to investigate (active) deformation and surface processes and the way these
56 processes interact to shape mountain belts. On a short-time scale, the pattern of deformation of
57 fold-and-thrust belts provides information on crustal mechanics, the sequence of active faulting
58 and its relation to large earthquakes; on a long-time scale, the structure and dynamics of the
59 fold-and-thrust belt -foreland basin systems offers unique insights into the influence of
60 structural, thermal and rheological inheritance, together with coupling between surface and
61 deep processes.

62 During the last ten years, significant advances have been made in the description and
63 understanding of fold-and-thrust belts and foreland basins. Among (many) others : better
64 definition of structures at depth (seismic imaging, 3D visualization/ geomodelling, better
65 appraisal of geometrical uncertainties); use of analogue and numerical modelling to constrain
66 long-term and short-term surface and deep processes; applications of thermochronology
67 (detrital thermochronology for sediment routing and paleo-burial estimates, coupled
68 thermochronological and 2D/3D mechanical/kinematical modelling); recognition of the
69 influence of salt and salt tectonics; renewed conceptualization of fold-fractures relationships
70 and new ways to unravel paleostress history.

71

72 Imaging structures at depth in fold-and-thrust belts and addressing uncertainties

73 Our view of the geometry of fold-thrust belts relies on the interpretation of field data,
74 borehole data and seismic lines. Classically, geoscientists collect a large amount of outcrop and
75 geophysical data and interpret them to arrive at a plausible geometric or steady state model.
76 Such a model can then be used as a basis for kinematic restorations or mechanical
77 considerations. While this approach has been successful for many years, improving our
78 understanding of fold-and-thrust belts as well as finding oil and gas reservoirs, a topic that
79 becomes more and more recognized is to take the uncertainties associated with the input data

80 into account. This emerging field may yield new fundamental insights on subsurface
81 geometries, and consequently may provide different kinematic and mechanical solutions of a
82 study area. Field data are subject to possibly large uncertainty, due to vegetation cover or
83 variable erodibility of different rock types (Moosdorf et al. 2018), which may lead to
84 overrepresentation of less erodible rocks in the measurements. This is particularly an issue for
85 the location of faults, as fault rocks are commonly easily eroded. Vegetation cover is a common
86 problem in fold-thrust belts, as they form the foothills of orogens, an area commonly inhabited
87 and cultivated. These uncertainties will often make possible different equally viable
88 interpretations of field data. Some of these uncertainties can be reduced by combining them
89 with subsurface data, i.e. borehole measurements and seismic images. However, these two data
90 sets are similarly associated with uncertainties, and it has been shown using synthetic seismic
91 images that interpretations may have a strong bias depending on the expertise of the interpreter
92 (Bond et al. 2007). Even when interpreters are all familiar with the area, different interpretations
93 of the same high-resolution seismic section may be offered, as pointed out by von Hagke &
94 Malz (2018).

95 For future research in fold- and-thrust belts addressing these uncertainties will be an
96 important way forward. At the moment there are two main directions. First, presenting not only
97 one plausible model, but arriving at a wide range of possible models, ideally associated with a
98 quantitative estimate of the respective uncertainties. This requires a probabilistic approach, and
99 has successfully been applied to a wide range of areas (Bond, 2015, Wellmann et al. 2012,
100 Wellmann et al., 2018). This approach at least partly circumvents the bias introduced by the
101 subjectivity of the interpreters. However, the full complexity of geological systems often cannot
102 yet be captured. Therefore it is essential to include as much geological information as possible,
103 appreciating the 4-D nature of structures (e.g. Duffy et al., 2018, Hessami et al., 2001, Ruh &
104 Vergés, 2018, von Hagke et al., 2016). To that respect, 3D structural modelling using various

105 geomodelling softwares (e.g., Caumon et al., 2009; Pellerin et al., 2015) may be of great help
106 in testing the geometrical and kinematical compatibility of structures from various geological
107 or seismic sections (Turrini et al., 2014), performing meaningful kinematic restorations
108 (Durand-Riard et al., 2011) or performing seismotectonic studies at a regional scale (Turrini et
109 al., 2015). Second, and maybe even more importantly, models of the subsurface structure of
110 fold-thrust belts should be not only tested with respect to their kinematic plausibility, but their
111 mechanics needs to be understood. This has exemplarily been shown for the interpretation of
112 triangle zones (von Hagke & Malz, 2018), but is equally applicable for any structural model.
113 Understanding the mechanics of fold-thrust belts and their detachments has progressed much
114 in the last few years, partly because of the increasing strength of numerical models and
115 increasing resolution of micro-structural techniques. The mechanics of shale and salt-
116 detachments become increasingly known (see reviews by Morley et al. 2017, 2018 for shale
117 detachments, and summary below for salt detachments). Addressing the mechanics of fold-
118 thrust belts and their respective detachments also will partly shift focus of field-based research
119 from the regional scale to the outcrop scale, where individual key structures have to be analyzed
120 at high resolution. At the same time, it is necessary to explore the parameter space of what is
121 responsible for fault weakness, the influence of evolving and transient rheologies, or the
122 influence of mechanical stratigraphy on geometries. This is best done with a combination of
123 numerical and analog models. However, a sound understanding of the structures in the field,
124 using classic geological field techniques combined with digital mapping (Fernández et al.
125 2004), drone imagery (Pavlis and Mason, 2017), LiDAR and photogrammetry studies at
126 regional and outcrop scales will remain essential, and a cornerstone of research applied to fold-
127 thrust belts (Tavani et al. 2014; Corradetti et al. 2017; García-Sellés et al. 2011, 2018).

128

129 Analogue and numerical modelling of tectonic processes in fold-and-thrust belts

130 Analogue and numerical modellings are rather mature geoscientific tools used to
131 understand the dynamics of fold-and-belts, accretionary wedges and orogens for several
132 decades. Recent methodological advances, both experimental and numerical, allows nowadays
133 simulating key geoprocesses and their coupling at various time scales at high resolution and in
134 3D.

135 With the advent of high-resolution monitoring techniques and accurate rheological
136 characterization of rock analogue materials, analogue modelling has transformed from a
137 concept-testing qualitative tool to a quantitative simulation technique (see reviews by Rosenau
138 et al., 2017; Graveleau et al., 2013).

139 On the methodological side recent efforts in analogue modelling focus for instance on
140 the use of image-processing software to generate 3D voxel (i.e., *volumetric pixel*) models of the
141 internal structure of sandbox models; this technique allows producing arbitrary virtual sections
142 (i.e. inlines, cross-lines and depth slices) through sandbox models based on the reconstruction
143 from cross-sectional images in a similar manner to 3D seismic data (Dooley et al. 2009;
144 Granado et al. 2017). These voxels can be converted into seg-y files to be loaded to seismic
145 interpretation software platforms (Roma et al. 2018) to generate 3D structural static models;
146 these models can later be populated with properties to carry out dynamic simulations (i.e., fluid
147 injection, production, etc.). Kinematic monitoring of model surfaces and volumes, on the other
148 hand, can be carried out using digital imaging techniques (Galland et al., 2016; Boutelier et al.
149 2019; Toeneboehm et al. 2019; Adam et al. 2013; Poppe et al., 2019) including time adaptive
150 imaging (Rudolf et al., 2019). Dynamic monitoring using stress sensors measuring lateral push
151 (Ritter et al., 2018 a,b, Cruz et al., 2010; Souloumiac et al., 2012) or pressure sensors providing
152 in-situ observations (Moulas et al., 2019) improved greatly our understanding of the force
153 balance and work budget in experimental tectonic systems. A number of new analogue
154 materials mimicking brittle-elastoplastic and viscoelastic behaviour of rocks in a more realistic

155 (often complex) way have been developed and characterized recently (e.g. Di Giuseppe et al.,
156 2009, 2015; Abdelmalak et al., 2016; Brizzi et al. 2016).

157 In parallel to methodological advances, studies dedicated to critical assessment of the
158 reproducibility of experiments (Cubas et al., 2010; Santimano et al., 2015) and their boundary
159 conditions (Souloumiac et al., 2012) raised awareness for uncertainty in analogue modelling
160 and the relation between intrinsic and extrinsic variability of experimental observations.
161 Community benchmarks including material characterization (Klinkmüller et al., 2016, Rudolf
162 et al., 2016) and comparison of analogue and numerical models (Schreurs et al., 2016, Buitter
163 et al. 2006, 2012) has moreover helped validate established analogue and numerical modelling
164 techniques.

165 Based on methodological advances, applications of analogue modelling developed in
166 recent years towards ever shorter timescales: while classical sandbox modelling of orogenic
167 systems and fold-and-thrust belts (Borderie et al., 2019, Saha et al., 2019) as well as lithospheric
168 scale multilayer models (Munteanu et al., this issue) represent the state-of-the art in analogue
169 modelling of tectonic processes at the million year time scale, approaches to model
170 geomorphologic (Graveleau et al., 2006, Guerit et al., 2016) and seismic cycle time scales
171 (Rosenau et al., 2017) have entered the stage with the perspective of understanding tectonic
172 evolution across all relevant time-scales.

173 The deeper our understanding of shallow earth deformation processes, the more
174 complicated their modelling, however. It is for these matters that numerical modelling is
175 gaining importance in geoscience as applied mathematical codes aid at providing a better
176 control on geological processes involving fluid pressure, mineral reactions, surface processes
177 etc., which cannot, or only in a very limited way, be implemented experimentally. Moreover,
178 modern computer processing power allows running many realizations simultaneously, testing
179 the effect/s of different parameters in parallel resulting in a generally wider parameter space to

180 be tested numerically compared to analogue models. Finally, complex geometries as well as
181 depth-dependent rheologies and their change with temperature (i.e. thermo-mechanical) and
182 accumulated strain (i.e., strain-weakening, strain-hardening) can also be modelled more
183 accurately, and their influence on the resulting realizations, numerically constrained. These are
184 possibly the great advantages of numerical methods in comparison to traditional analogue
185 modelling; however both techniques need to be regarded as complementary, and not mutually
186 exclusive.

187 There are three main numerical methods currently applied to fold-and-thrust belts:
188 discrete element methods (also referred to distinct element methods, Burbidge and Braun,
189 2002), finite element methods (Simpson, 2011; Erdős et al. 2014a; Bauville and Schmalholz,
190 2015), and finite difference methods (Ruh, 2019). More rarely, discrete element methods are
191 used because of their limitations in resolution with respect to geological systems. All numerical
192 methods are challenged by the increasing demand for 3D “cross scale” (in space and time)
193 models requiring massively parallelized codes to be developed (Kronbichler et al., 2012; van
194 Dinther et al., 2013; Ruh et al., 2013) implementing adaptive time-steps and meshes as well as
195 the most efficient and reliable mathematical formulations of deformation laws (Pipping et al.,
196 2016; Herrendörfer et al., 2018; Glerum et al., 2018).

197 Future work on numerical modelling applied to fold-and-thrust belts should address the
198 following points: the controls of inheritance from passive margins stages, including thermal
199 effects, fault orientation and strength in 2D/3D, as well as sedimentary basin architecture and
200 mechanical properties of the basin-infill (i.e., changes in pore-fluid pressures), include the
201 presence of several décollements with variable strengths, underlying basement steps, as well as
202 the existence of previous salt structures, and *non-layer cake* stratigraphies.

203

204 Fold-and-thrust belts and the thermochronology toolbox

205 Thermochronological age dating of various minerals is a key tool in Earth sciences to
206 quantify landscape evolution and the metamorphic and tectonic history of orogens and
207 sedimentary basins (Wagner and Reimer, 1972; Brown et al., 1994; Reiners and Brandon,
208 2006). Thermochronology provides information on the temperature history of minerals and
209 rocks. Thermochronometric datasets provide cooling ages, which is the time elapsed since a
210 mineral cooled below a certain temperature. In fold-and-thrust belts, the application of low-
211 temperature thermochronometers is ideally suited, as the method is sensitive to the uppermost
212 few kilometers of the crust. Commonly used techniques are fission track dating and (U-Th-
213 Sm)/He dating on zircon and apatite. Different thermochronometers are sensitive to different
214 cooling intervals. For instance whereas the apatite fission track system is sensitive to
215 temperatures above approximately 110°C (depending on apatite chemistry), the apatite helium
216 system records cooling below ~60°C (Wolf et al., 1996; Farley and Stöckli, 2002). In fold-
217 thrust belts, analyzed samples commonly derive from foreland sediments that have been
218 incorporated later into the orogenic wedge. This implies that dated grains may have a
219 complicated time-temperature history, starting with cooling in the hinterland, transport through
220 the drainage area, deposition and burial (and consequently heating) in the basin and later
221 exhumation in the fold-and-thrust belt (Fitzgerald et al. 2019). The youngest exhumation history
222 in the fold-and-thrust belt can be complex in itself due to repeated activity of the same fault at
223 different times. These complex t-T paths provide some challenges to be interpreted correctly,
224 however, by integrating thermochronological data from the orogen, the foreland as well as the
225 fold-thrust belt it is possible to provide a comprehensive picture of the t-T history of a mountain
226 belt. Thermochronological data from fold-and-thrust belts provide also information on the
227 provenance history of the sediments, estimates of maximum burial, as well as timing and rates
228 of deformation. The strength of low-temperature thermochronometry has been successfully
229 exploited in many fold-thrust belts in the world, as for instance in the Andes (McQuarrie et al.,

230 2005; Barnes et al., 2006; Savignano et al., 2016), the European Alps (von Hagke et al., 2012,
231 2014), and the Pyrenees (Beamud et al. 2013; Mouthereau et al., 2014, Ternois et al., 2019). It
232 has to be noted that thermochronometric ages cannot be directly translated into time of
233 deformation (e.g. Mora et al., 2015). Particularly for fold-and-thrust belts, where sediments are
234 commonly dated, thermochronometric ages may be associated with uncertainties due to
235 unknown provenance ages, hydrothermal fluxes, low crystal quality due to rounded grains,
236 limited amount of grains that can be analyzed, or unrecognized zoning of the grains. These
237 uncertainties provide exciting future pathways for thermochronological age dating, and future
238 research on fold-and-thrust belts should strongly follow advances in thermochronological
239 methods, as this may help reduce the uncertainties of proposed geological models.

240 An example for the successful combination of thermochronometry and structural
241 geology is the newly emerging field of coupling thermochronological data with kinematic
242 restorations. Sequential restoration of balanced cross sections is a powerful tool in structural
243 geology, as it allows one to: (i) draw a reliable picture of the changing geometry of deforming
244 geological structures through time; (ii) determine the original position and dip of the structures;
245 (iii) calculate the amount of shortening; (iv) define timing of basin formation and evolution,
246 and (v) extrapolate rates of tectonic processes, such as exhumation and erosion (Bulness and
247 McClay, 1999). Relying on pioneering works in contractional settings (Bally et al., 1966;
248 Dahlstrom, 1969, 1970; Mitra and Namson, 1989), kinematic restoration of balanced cross
249 sections has been applied progressively both in extensional and inverted basin areas (Coward,
250 1996; Bulnes and McClay, 1999). In the oil industry, this technique is routinely used to evaluate
251 the position of source rocks and model hydrocarbon generation, expulsion and migration, as
252 well as to analyze structural traps in terms of timing and geometry (Buchanan, 1996). However,
253 this method alone can be inadequate in case syntectonic deposits are not present/well preserved
254 (Almendral et al., 2014; Mora et al., 2015; Granado et al. 2016a). Coupling cross-section

255 balancing and restoration with thermochronological constraints provides the possibility to
256 define the various stages of deformation and to quantify both their extent and timing, in the lack
257 of a syntectonic sedimentary record or in conjunction with it (Andreucci et al., 2013; Mora et.
258 al., 2015; Castelluccio et. al., 2016; Chapman et al., 2017). However, the shift from a
259 temperature (i.e. related to the movement of the sample through the isotherms) to a space
260 domain (i.e. the depth of the sample during time) requires combining kinetic and
261 thermochronological information within a coherent model capable of taking into account
262 variations in the distribution of the isotherms in a dynamically active scenario. Another issue
263 to take into account is the role of topographic evolution during time and the way in which it
264 interacts with isothermal surfaces. In fact, wavelength of the relief, exhumation rate and heat
265 advection strongly perturb the isotherm state (Stüwe et al., 1994; Mancktelow and Grasemann,
266 1997; Braun, 2002; Reiners and Brandon, 2006). Consequently, understanding better landscape
267 evolution (Braun et al., 2012), drainage reorganization (Yanites et al., 2013), or rock erodibility
268 (Moosdorf et al., 2018) is important for a correct interpretation of thermochronological data
269 and consequently the tectonics of fold-and-thrust belts. In the last few years, successful results
270 have been obtained with software dedicated to both inverse and forward modeling of
271 thermochronometric data (Fillon and van der Beek, 2012; Almendral et al., 2014; Erdős et al.,
272 2014b; Castelluccio et al., 2015; Ternois et al., 2019). Forward modeling uses
273 thermochronometric ages calculated starting from kinematic restoration integrated with thermal
274 parameters. The output is a model of low-temperature thermochronometric ages along a
275 geological cross-section, which can be compared with measured apatite/zircon fission track
276 and/or (U-Th-Sm)/He cooling ages on samples collected along the profile. The comparison with
277 measured and modeled ages allows in turn improving the structural model through an iterative
278 process.

279

280 Influence of salt and salt tectonics in fold-and-thrust belts

281 The critical taper theory states that the external geometry and internal deformation of
282 fold-and-thrust belts is a function of the coefficient of friction of their basal décollement (Davis
283 et al., 1983). This has been the basis for the current modern understanding of fold-and-thrust
284 belts. The importance of layered evaporitic sequences -commonly referred to as *salt* for short-
285 as preferential décollements in fold-and-thrust belts has also been recognized for a long time
286 (Davis and Engelder, 1985). Salt is inherently weak, and contrary to most rocks whose strength
287 increases as a function of depth, salt's strength depends on its viscosity and on the applied strain
288 rate, meaning that at high strain rates salt will undergo brittle failure; at geological strain rates
289 however, salt will flow (Jackson and Vendeville, 1994). For such differences in mechanical
290 behaviour, the structural styles in salt-detached fold-and-thrust belts are markedly different to
291 those contractional systems developed over frictional décollements (Smit et al. 2003; Granado
292 et al. 2017). In this sense, the latter are constituted by thrust faults and related folds (Jamison,
293 1987) which tend to display a dominant forelandward vergence largely developed in footwall
294 thrust sequences. In their simplest scenario, salt-detached fold-belts tend to show a regular
295 spacing of narrow symmetric anticlines separated by broader *box-like* synclines; these anticlines
296 are commonly detachment folds, or transported detachment folds, which wavelength is
297 controlled by the thickness of the dominant mechanical unit (Mitra, 2002). Detachment folds
298 are usually cored by salt and can reach several kilometres of structural relief; due to crestal
299 erosion and/or faulting, salt can eventually extrude as diapiric structures (Santolaria et al. 2014).
300 To attain such amplitudes, detachment anticlines need severe syn-orogenic sedimentation on
301 their immediate synclines (Izquierdo-Llavall et al. 2018; Borderie et al. 2019). Salt-cored
302 detachment folds commonly plunge gently towards their periclinal terminations; however steep
303 plunges can also be attained in short distances compared to their lengths depending of the
304 distribution of underlying salt. Thrusts and reverse faults in salt-detached systems display no

305 preferred sense of transport given the low friction or even frictionless nature of saline
306 décollements. For the same reason, salt-detached fold-and-thrust belts commonly display an
307 extremely narrow cross-sectional taper when compared to those belts detached along frictional
308 horizons, hence being comparatively wider, with deformation concentrated at the edges of the
309 salt basin (Davis and Engelder, 1985; Jaumé and Lillie, 1988). Thrust salients are also typical
310 features of salt-detached fold-and-thrust belts, being commonly related to the distribution of the
311 underlying salt décollement and to lateral variations in décollement efficiency (Becker et al.
312 2000; Jackson et al. 2003; Muñoz et al. 2013).

313 Some of these well-established structural templates remain true and applicable, but only
314 when a stratigraphy of uniform thickness (i.e. *layer-cake*) is involved. However, as most fold-
315 and-thrust belts develop from the incorporation of rifts and passive margins basins in the
316 shortening system, contrasting structural styles will develop depending on which parts of the
317 basins are involved. In the case of rift to passive margins salt basins (Rowan, 2014; Granado et
318 al. 2016b; Kukla et al. 2018), significant structural complexities differing from those templates
319 described earlier can arise. For instance, many salt-detached and salt-influenced fold-and-thrust
320 belts display structural styles which include: i) multiple structural orientations for folds and
321 faults; ii) strong changes in fold plunges; iii) large panels of completely overturned stratigraphy
322 (i.e. *flaps*); iv) mechanical contacts omitting or repeating stratigraphy; v) severely deformed
323 evaporite bodies, or their equivalent salt welds, bounding structural units of markedly different
324 sizes and aspect ratios, and contrasting stratigraphic thicknesses and facies (i.e. *non-layer cake*
325 stratigraphy).

326 All these complex structural styles have been sometimes explained by invoking several
327 deformation phases, strike-slip tectonics, or even the gravitational emplacement of thrust
328 sheets. It has not been until recently that early salt tectonics processes inherited from the rift to
329 passive margin stages have been taken into consideration. As salt is the weakest stratigraphic

330 units involved in the developing fold-and-thrust belt, shortening would be first focused on those
331 inflated salt structures, influencing the formation and orientation of structures immediately
332 around them (Duffy et al. 2018; Snidero et al. 2019). Diapirs and salt walls can be squeezed
333 depending on their orientation in respect to shortening, to eventually neck-off forming sub-
334 vertical welds (i.e. secondary welds); with ongoing shortening, sub-vertical welds may become
335 reactivated as thrusts, reverse or transpressive faults, also depending on their orientation with
336 respect to the direction of shortening (Rowan and Vendeville, 2006; Duffy et al. 2018; Granado
337 et al. 2018; Roma et al. 2018). Depocenters related to early salt structures such as rollovers will
338 become inverted, and along with minibasins if present, will be incorporated and transported
339 along the weak basal décollement, imbricated by later thrusts, and undergo rotations along
340 vertical and/or horizontal axes (López-Mir et al. 2014; Saura et al. 2016; Granado et al. 2018;
341 Snidero et al. 2019). All these processes will be strongly influenced by syn-orogenic erosion
342 and sedimentation (Izquierdo-Llavall et al. 2018).

343 To summarize, the structural styles of contractional salt-detached and salt-influenced
344 fold-and-thrust belts are largely dependent on the relative thickness of salt and its overburden,
345 the lateral distribution of salt previous to shortening, and whether early salt structures and
346 related depocentres are present (Hudec and Jackson, 2007).

347

348 Fracture analysis and paleostress/paleopressure reconstruction in fold-and-thrust belts

349 The reconstruction of the past kinematic and tectonic history in fold-and thrust belts
350 requires constraints on the evolution through space and time of both stress and strain which
351 affected sedimentary (and basement) rocks. Fractures are the most common response of brittlely
352 deformed rocks submitted to tectonic stresses and are therefore classical and reliable
353 palaeostress indicators (Lacombe et al., 2011). In addition, in carbonate rocks of low matrix
354 permeability, the characteristics of the fracture network play a fundamental role in hydrocarbon

355 migration and reservoir quality (Casini et al., 2011). A good understanding of the mechanical
356 and chronological development of the meso-scale fracture network is therefore key for tectonic
357 analysis as well as natural resources exploration, and waste repositories studies

358 A recent step forward in the understanding of fracture occurrence in fold-and-thrust belts
359 is related to the recognition, in addition to fold-related meso-structures, of the widespread
360 occurrence of pre-folding fracture sets. These fracture sets may have originated from far-field
361 earlier tectonic events unrelated to the phase of thrusting and folding or from foreland flexuring
362 and along-foredeep stretching (Tavani et al., 2015) and even from differential compaction
363 controlled by deep-seated faulting before the foreland domain has become part of the fold-and-
364 thrust belt (Tavani et al., 2018). This points toward the need to carefully consider pre-existing
365 fractures, possibly unrelated to folding, to build realistic conceptual fold–fracture models.
366 Moreover, a blind spot for most fracture analyses has been for long the lack of constraints on
367 the absolute timing of fracture development. This absolute timing is never resolved through
368 field-based geometrical relationships, so the relevance and meaning of some fracture sets with
369 respect to regional deformation can be disputable, especially in regions that underwent
370 polyphase tectonics. Recent developments in absolute dating of calcite cements of veins within
371 folds using U/Pb technique (Parrish et al., 2018; Hansman et al., 2018; Beaudoin et al., 2018)
372 have brought the proof of the time relevance of meso-scale structures to regional fold-and-thrust
373 structures, hence have helped refine the tectonic history.

374 Providing constraints on paleostress orientations and magnitudes and how they evolved
375 during geological history is a challenging but important task that can lead to major breakthrough
376 in the appraisal of long-term mechanical and paleohydrological behaviour of the upper crust.
377 Improvement and application of new paleopiezometric techniques (e.g., calcite twinning and
378 stylolite roughness paleopiezometry) calibrated in the diagenetic conditions of pressure and
379 temperature have recently helped refine the tectonic and paleostress history in the Apennine

380 fold-and-thrust belt (Beaudoin et al., 2016). Stress quantification provides insights into the
381 burial history independently of any assumption on the past geothermal gradient, the overall
382 long-term mechanical behaviour of the crust, the degree of coupling between the cover and the
383 basement and the way orogenic stresses are transmitted from the plate boundary to the (far)
384 foreland (Beaudoin and Lacombe, 2018).

385 Last but not least, reconstruction of fluid (over)pressure and its evolution in fold-and-
386 thrust belts and sedimentary basins is of prime interest for both academy (e.g., fault
387 reactivation) and industry (hydrocarbon generation and migration). The use of hydrocarbon-
388 bearing fluid inclusions, when developing contemporaneously with aqueous inclusions,
389 provides a direct access to the pore-fluid temperature and pressure of cemented fractures or host
390 rocks at the time of cementation and hydrocarbon trapping, in line with the tectonic evolution
391 (Roure et al., 2010). Alternatively, the combination of the calcite twinning paleopiezometer
392 with fracture analysis and rock mechanics tests has led to pioneering reconstructions of fluid
393 (over)pressure evolution during the different stages of foreland shortening (e.g., Sevier-
394 Laramide foreland, Amrouch et al., 2011; Beaudoin et al., 2014).

395

396 **2. Content of the Special Issue**

397 This special issue of Tectonophysics, sponsored by the International Lithosphere
398 Program, presents a new collection of 27 papers dealing with different aspects of fold-and-
399 thrust belts and foreland basins evolution, such as structural geology, geomorphology,
400 exhumation, sediment transport, dating, seismicity, surface processes and basin dynamics
401 during pre-and syn-collision stages, analogue or numerical modelling approaches, in addition
402 to regional case studies. Some of these contributions were presented as part of a session devoted
403 to this topic at the 2018 European Geosciences Union General Assembly in Vienna (Austria).
404 The aim of this session was to assemble a broad group of Earth scientists interested in fold-and-

405 thrust belts and peripheral basins spanning a broad array of tectonic settings, geographical
406 locations, and geological times. This volume presents a collection of some of the diverse
407 research that is being carried out on this topic. We believe that these studies contribute to a
408 better understanding not only of fold-and-thrust belts in particular, but also of orogenic
409 processes and of the rheology of the continental lithosphere in general, and that this volume
410 will help promote new contacts between interdisciplinary earth scientists.

411

412 * Structural inheritance and inversion tectonics in fold-and-thrust belts

413 The role of inherited structures for the geometry of later tectonic events has been recognized
414 since a long time. However, a dynamic understanding of the role of inherited structures on fold-
415 thrust belt geometries is often lacking. **Granado & Ruh** address this research gap by modeling
416 the role of inversion of a half-graben during shortening using finite differences. In their
417 experiments they test how the strength of the inherited fault as well as different fluid pressure
418 in the syn-rift sediments and strength of the upper décollement are reflected in the structural
419 style of foreland fold-and-thrust belts. A straight forward result is that a weak inherited fault
420 will be easily reactivated, influencing the kinematics and consequently structural style. Weak
421 syn-rift sediments favor hanging-wall bypass thrusting. Generally the strength ratio of upper
422 and lower décollement is key to understanding fold-and-thrust belts. The results of numerical
423 models can be applied to natural examples. The authors selected the Helvetic nappes of the
424 European Alps, and the Malargüe fold-thrust belt and the Salta Rift System, both in the
425 Argentinian Andes as case studies for their models. They can show a first order geometric
426 comparison of nature and experiments, indicating a quantitative comparison of the rheologies
427 used in the model with paleo-rheologies in nature is possible.

428 **Espurt et al.** combined field geological, structural, paleo-temperature and sub-surface data
429 together with deep geophysical data to build a new 210km-long crustal scale balanced cross-

430 section across the Central Pyrenean belt. Along the section, the belt corresponds to the inversion
431 of the Mesozoic Pyrenean Rift system, which consisted in a hyper-extended relay zone of two
432 metamorphic zones with exhumation of continental lithospheric mantle. Comparison between
433 present-day crustal geometry and sequentially retro-deformed stages (lower Santonian, upper
434 Jurassic and lower-middle Triassic) of this section shows that the Pyrenees were superimposed
435 onto a complex structural template affected by the Variscan orogeny and subsequent Permian
436 rifting, that in turn controlled subsequently the geometry of the Mesozoic rifting and the
437 building of the upper Cretaceous-lower Miocene Pyrenean orogen. This study puts emphasis
438 on the long-term influence of inherited tectonic crustal fabric in the evolution of orogens.

439 Using a kinematic forward lithosphere deformation model (RIFTER), **Gómez-Romeu et al.**
440 produce flexural isostatically compensated, balanced geological sections across the Western
441 Pyrenees. The tectonic evolution of both the original rift and the subsequent orogeny are
442 investigated in order to obtain new insights into the role of extensional structural inheritance in
443 the development of collisional orogens. The proposed model shows how, following an
444 extensional stage characterized by a hyper-extended rift system that also led to mantle
445 exhumation, Pyrenean shortening included two main stages. A pre-collisional stage produced
446 inversion of the hyper-extended rift system, whereas a syn-collisional stage involved the
447 southern proximal rift domain, resulting in the formation of the Axial Zone and the Iberian pro-
448 foreland.

449 **Munteanu et al.** present results from state-of-the art lithospheric scale analogue modelling of
450 compressional systems. Based on a systematic series of analog models the authors investigate
451 the structural imprint of local crustal weaknesses on the deformation front in contractional
452 settings. The model setup features brittle (granular) and viscous (fluid) multilayers similar to a
453 continental lithosphere. Because the models rest on a dense low viscosity fluid (similar to the
454 asthenosphere) they are isostatically balanced and applied lateral forces are transmitted

455 throughout the model domain. In the model, the crustal weaknesses are effectively localizing
456 early stage deformation. Depending on the size, shape, number and location of the weakness
457 zones either laterally continuous but bent structures evolve nucleating at the weak spots and
458 growing laterally into stronger lithosphere or discontinuities (transfer zones) form separating
459 strong and weak lithospheric domains. This study highlights that crustal weaknesses inherited
460 for example from an earlier rifting phase needs to be considered as a source of along-strike
461 irregularity of orogenic structures.

462 **Martins-Ferreira** integrates seismic interpretation (calibrated by well logs) and field data to
463 investigate the relationships between fault reactivation and thrusting in the cratonic region of
464 central Brazil. The study area is located between the opposed-verging Brasília and Araçuaí
465 Neoproterozoic belts, which formed during the initial stages of West Gondwana amalgamation.
466 Seismic interpretation suggests a tight relationship between shallow folds and thrusts and rift
467 inversion. Reverse-slip reactivation of inherited normal faults exerts a clear control on thrust
468 ramp nucleation. Widespread thrusting occurs over buried rifts characterized by multiple
469 inverted faults, whereas sedimentary successions overlying non-rifted basement are not
470 significantly affected by thrusting and folding. Furthermore, the development of salients and
471 recesses characterizing the Brasiliano-Pan Africano orogen appears to be controlled by the
472 morphology of buried rifts.

473 **Tavani et al.** analyzed meso-scale fractures and faults exposed in the Triassic to Miocene
474 sedimentary succession of the Lurestan region (Zagros Belt). The authors document
475 development of syn-sedimentary extensional fractures formed during Early Jurassic rifting then
476 in response to foreland flexuring and along-foredeep stretching during Late Cretaceous-Eocene
477 and late Miocene-Pliocene pulses of convergence. These repeated extensional episodes
478 produced oblique-slip reactivation of inherited basement structures, above which differential
479 compaction and subsidence prevailed during tectonically quiescent periods. The fractures

480 related to differential compaction add to the complexity of the ‘tectonic’ pattern formed during
481 true tectonic (extensional and compressional) pulses, leading to a complex articulated fracture
482 network. This study emphasizes how meso-scale structures may be relevant to regional-scale
483 tectonics and highlights the role of tectonic inheritance in fold-and-thrust belts at all scales.

484 * Deformation of basement rocks in fold-and-thrust belts

485 The paper by **Searle et al.** provides a review of the geology of the Caledonian Moine Thrust
486 zone in the Loch Eriboll region, NW Scotland. In addition to already published maps, the
487 authors use new maps, balanced and restored cross sections from fieldwork at Loch Eriboll
488 together with a cross-section from the Moine thrust hinterland to infer geological processes
489 including sequence of thrusting, shortening estimates and regional tectonic implications for this
490 famous fold-and-thrust belt in the NW Scottish Highlands. The authors suggest in a rather
491 provocative way that two major crustal-scale thrusts that extend down into the upper mantle
492 imaged on seismic profiles across the foreland, the Outer Isles and Flannan thrusts, are unrelated
493 spatially or temporally to the Moine thrust sequence and speculate on the regional significance
494 of these thrusts. Finally, the authors reflect on the metamorphic sequence overlying the Moine
495 Thrust and draw parallels to Himalayan style orogenesis.

496 **Bellahsen et al.** use the case study of the Bielsa basement unit in the Axial Zone of the Pyrenees
497 to provide new insights into the modes and style of upper crustal shortening in orogens. In the
498 study area, distributed strain associated with widespread, minor shear zones appears to predate
499 strain localization along major crustal ramps (as constrained by zircon fission-track data). The
500 authors suggest that chemical weakening – in the form of feldspar sericitization – exerts a major
501 control on basement rheology during the early stage of distributed shortening. Sericitization is
502 widespread, and occurs not only in ultramylonites, ultra-cataclasites and phyllonites, but also
503 in un-deformed granodiorites. Based on these observations, the authors propose that the strength

504 of the upper crust was very low at the onset of shortening, due to a high thermal gradient and
505 fluid circulation that induced large-scale sericitization in greenschist facies conditions.

506 * Geomorphic signatures of tectonics and surface processes in fold-and-thrust belts

507 **Obaid & Allen** focus on the Zagros fold-thrust belt, which forms one of the best-exposed
508 examples of and active fold-thrust belts in the world. Therefore, it is perfectly suited to
509 investigate the role of climate on landscape evolution in an actively deforming system. The
510 authors use geomorphic indices such as normalized channel steepness index or integrated relief
511 and hypsometric index to test how geomorphology is a sensitive indicator of tectonic processes.
512 The geomorphic indices show differences between different regions, correlating with climatic
513 differences. A possible interpretation is that wetter conditions retard plateau growth, whereas
514 dry climate allows for plateau growth, as the river draining the area have lower stream power
515 (which is partly also a function of rock erodibility). This does not imply that climate drives
516 landscape evolution. Instead, a positive feedback exists, where tectonics forms topographic
517 barriers that control precipitation. Despite local differences, there is topographic similarity
518 along five relatively evenly distributed swath profiles across the Zagros. Similar strain rates or
519 similar overall shortening may explain this topographic similarity across the belt.

520 **Sanchez Nassif et al.** combine forward modelling and structural reconstruction to link tectonic
521 steps to erosion rates from the Argentinian Precordillera Jachal section. Their workflow
522 successfully articulates at each time steps forward kinematic modelling and alpha calculation
523 following the Coulomb wedge analysis to deduce the cumulative erosion budget. Their
524 methodology allows to obtain the sequence of deformation as well as the associated erosion
525 rates on the basis of geological field evidences such as tectonic features and preserved
526 sedimentary structures, without any additional t-T data. Applied to the example of the Jachal
527 section, the authors validate their analysis with the comparison to the published section and

528 thermochronological data, showing a two-stepped evolution of the erosion rates (1 and 1.3
529 km/Myr) associated to the activity of the Niquivil fault.

530 *Active tectonics and seismicity in fold-and-thrust belts

531 The Hengchun Peninsula at the southern tip of Taiwan is of global importance, as it is one of
532 the few locations in the world where an accretionary prism is exposed on land. Furthermore,
533 the area is seismically very active and constraining the geometries of the structures is important
534 for geohazard assessment. **Deffontaines et al.** address this using a new high-resolution digital
535 terrain model and SAR images. The authors conclude that the peninsula consists of two ramp
536 structures at depth, and that the Hengchun Fault, interpreted as the major tectonic lineament of
537 the peninsula, extends towards the offshore. This study stimulates the debate on the geometry
538 of the structures of the Hengchun Peninsula and in other systems, showing the high
539 uncertainties at depth despite extensive data coverage. The study particularly highlights the
540 need for further research, as a nuclear power plant is located close to the Hengchun Fault.

541 **Mescua et al.** integrate field and wellbore data to discuss the stress field in the frontal sector of
542 the Malargüe fold-and-thrust belt (Andes of Argentina). Surface observations indicate N-S
543 thrusts and active NW to WNW and ESE strike-slip faults in the study area. Inversion of fault
544 kinematic indicators, combined with borehole breakout data and a mini-frac test, constrain the
545 Quaternary to recent stress state, which is characterized by a subhorizontal, E-W oriented
546 maximum principal stress, and by intermediate and minimum stresses with similar magnitudes
547 that are locally interchanged, producing a setting in which reverse and strike-slip faults are
548 alternatively active. The implications of the recognized structures for earthquake hazard are
549 examined.

550 **Rivas et al.** analysed seismic activity in the Precordillera in the Andean backarc region of
551 Argentina where a long record of large and damaging earthquakes in the last century exists. In

552 the northern part, which is poorly known either seismically or in terms of style of deformation,
553 the authors determine seismic locations, seismic moments, moment magnitudes, focal depths
554 and focal mechanisms for local earthquakes over the periods 2000–2002 and 2007–2010.
555 Overall the results agree with the E-W compression and shortening in the Andean northern
556 Precordilleran backarc region. The results further provide new constraints on the patterns of
557 earthquake distribution and on the Andean backarc crustal deformation.

558

559 * Exhumation and sediment routing in fold-and-thrust belts

560 **Odlum et al.** investigate the geological history of the South-Eastern Pyrenees, from the syn-
561 rift period to the syn-orogenic period. With the combination of several detrital proxies such as
562 U/Pb and (U-Th)/He on Zircons as well as U/Pb on rutiles and with an extensive mapping of
563 the bedrock ages dated with the same techniques, the authors retrace the source of the sediments
564 deposited from early Cretaceous to Oligocene times. By doing so, they provide a geological
565 scenario of hinterland exhumation, foreland basin evolution, and sediment routing system of
566 the south Eastern Pyrenees.

567 **Buford Parks & McQuarrie** present a study on the Central Andes highlighting the
568 importance of thermal, flexural, and kinematic models for understanding the evolution of fold-
569 and-thrust belts and orogens. The authors present a compilation of thermochronological data
570 and use them as independent tests for different structural models. The authors show that with
571 this method it is possible to gain insights on the sequence and rates of deformation and the
572 plausibility of different geometric models can be tested. Coupling with flexural models shows
573 that an additional geodynamic driver for uplift may be present, such as mantle delamination,
574 isostatic attainment, or lower crustal flow. Apart from insights on the geologic evolution of the
575 Central Andes, this contribution provides a tool for assessing kinematic models in fold-and-
576 thrust belts in general.

577 Based on the integration of field geology, seismic interpretation, apatite fission track and (U-
578 Th)/He (AHe) dating, as well analogue modelling, **Chang et al.** investigate the tectonic
579 evolution of the Kalpin fold-and-thrust belt. As this mountain belt accommodates crustal
580 shortening between the Tianshan and the Tarim Basin, unravelling its architecture, modes and
581 timing of development may significantly improve our understanding of the tectonic evolution
582 of central Asia. The timing of thrusting, constrained by low-temperature thermochronometry,
583 implies a southward propagation of Cenozoic deformation. Combined inverse and forward
584 thermal modelling is used to obtain information on the cooling and deformation of specific
585 thrust sheets, while sandbox models are used to simulate and confirm the thrust sequence. The
586 results are effectively used to provide new insights into the timing of deformation at the
587 northern margin of the Tarim Basin.

588 * Salt processes in fold-and-thrust belts

589 Based on a new regional balanced cross section, **Espurt et al.** present and discuss a new
590 interpretation of the Provence fold-and-thrust belt. The system is described as a Mesozoic
591 halokinetic salt province above a basement with strong structural inheritance, which has been
592 subsequently shortened during the Pyrenean and Alpine orogenies. They analysed the geometry
593 and timing of deformation and provide estimates of the amount of pre-orogenic contraction.
594 This paper highlights the major role of halokinetic processes and shows that a significant
595 amount of folding in fold-thrust belts can result from an early halokinetic fold system developed
596 during the pre-contractional passive margin evolution.

597

598 On the basis of field data, **Snidero et al.** describe and interpret the stratigraphic and structural
599 relationships between the Hormuz salt and its overburden around the Darmadan anticline in the
600 eastern Fars region of the Zagros fold-and-thrust belt (Iran). Their model describes for the first
601 time Late Jurassic-Early Cretaceous halokinetic sequences indicative of passive diapirism,

602 followed by early squeezing and tilting of the diapir's flanks during the Campanian-
603 Maastrichtian. They furthermore show that second order structural features indicate secondary
604 welding during upper Miocene times. Their work supports that the present structural trends of
605 the Zagros fold-and-thrust belt in the eastern Fars region are the result of the reactivation of
606 pre-existing salt structures.

607 * Case studies

608 Andean fold-and-thrust belts and basins

609 **Barrionuevo et al.** report on the structural kinematics of the Malargüe fold-and-thrust belt, and
610 the control imposed by the local stress field on magmatic and hydrocarbon fluid migration.
611 They propose that the structural framework controlling the magmatic activity corresponds to
612 inverted Mesozoic normal faults and Cenozoic thrusts, with oblique structures showing strike-
613 slip kinematics. Miocene dykes and sills were emplaced in relation to strike-slip and reverse
614 faults, respectively. Structural analysis suggests local switches from compressional to strike-
615 slip/compressional likely related to the similar values of the minimum (σ_3) and intermediate
616 (σ_2) principal stress with an E-W oriented maximum principal stress (σ_1), favouring the
617 emplacement of igneous intrusions and hydrocarbon migration through both thrusts and sub-
618 vertical strike-slip faults.

619 **Ronda et al.** present a series of balanced and sequentially restored geological sections across a
620 segment of the Southern Patagonian Andes between 46 and 48 °S. Opening of the Austral-
621 Magallanes basin and the Rocas Verdes back-arc oceanic basin as a result of Jurassic extension
622 were followed by Cretaceous to Cenozoic shortening. A regional westward-dipping listric
623 detachment, interpreted as formed during Jurassic extension, was reactivated during Andean
624 shortening as the main thrust detachment for the basement structures and the fold and thrust
625 belt. Based on the occurrence of angular unconformities and growth strata, three main

626 shortening stages have been recognized by the authors: (i) an early (Late Cretaceous) stage
627 associated with positive inversion and closure of the northernmost Rocas Verdes basin; (ii) a
628 Miocene stage (between 18 Ma and 10 Ma), which produced most of the shortening; and (iii) a
629 late (younger than 7 Ma) stage involving out-of-sequence thrusting in interior of the belt,
630 possibly associated with the onset of glaciations and glacial erosion.

631 Peri-Mediterranean fold-and-thrust belts and basins

632 **Masrouhi et al.** provide an extensive tectono-sedimentary study of the Southern Atlas front in
633 Tunisia. By a combination of field observations, well correlation and seismic profiles
634 interpretation, they propose a series of restored cross-sections across the particular area of the
635 Chott basin. The authors conclude on a mixed thick- and thin-skinned structural style with an
636 efficient décollement level defined by the Triassic evaporites. From their reconstruction, they
637 also find a decreasing amount of shortening from West to East, of a relatively small amount
638 (from 7 % to 1 %) that is explained by the inversion pattern. Masroushi et al documents that
639 this area, and the Chott basin in particular, is a large Mesozoic roll-over structure that was
640 reactivated during Alpine orogeny.

641 In their paper, **Khomsi et al.** provide a review of the known structural architecture of the Atlas
642 and Tell fold-and-thrust belts. Whereas the major tectonic steps accounting for the development
643 of the eastern Maghreb structures are rather well constrained by seismic data and structural
644 analyses, the overall configuration at depth, as well as the deep architecture of the underlying
645 basement remain poorly understood because only few wells yet penetrated deeper than the
646 Triassic series in the Atlas and adjacent foreland basin. The paper is intended at giving an
647 overview on the deep structural features affecting the Pan-African basement in the eastern
648 Maghreb by means of tentative regional cross-sections, in order to stimulate exploration of deep
649 hydrocarbon plays with particular focus on the pre-Triassic series.

650 Based on seismic profiles and field observations, **Balestra et al.** propose a 3D numerical model
651 of the Apennine-Maghrebian chain in the Mt. Kumeta and Mt. Rocca Busambra in Sicily. Their
652 model builds up on cross-section restoration based on interpreted seismic profiles. They are
653 incorporated to the 3D geomechanical model to provide a full and comprehensive view of the
654 studied area, especially useful when seismic data is of poor quality. By applying that technique,
655 the authors revisit the geological cross-sections by testing two-end member scenarios with a
656 one- or two-stepped structural sequence. They find that the along-strike variation in structural
657 style observed in the Trapanese unit is controlled by structural inheritance rather than by a
658 polyphase history.

659 **Vitale et al.** present a review on the stratigraphy, petrology, deformation and metamorphism of
660 the oceanic, sedimentary, magmatic and metamorphic successions from northern Calabria to
661 the Campania region of southern Italy (i.e., the Southern Ligurian Domain). These rock units
662 underwent subduction (Eocene) to be subsequently exhumed and exposed (Tortonian). Deep-
663 basin successions in the easternmost sector (close to the continental margin of Adria) were
664 obducted and frontally accreted (Aquitainian-Burdigalian). The tectonic transport for the
665 obducted successions was dominantly to SE, however new data provides a mean eastward-
666 directed transport during tectonic exhumation. Petro-chemical comparison between the
667 Southern Ligurian Domain mafic rocks with the corresponding rocks in the Alps, Corsica and
668 northern Apennines suggests an ocean continental transition setting for the Ligurian Domain.
669 Early orogenic stages of the southern Apennines-northern Calabria system were characterized
670 by a complex kinematic evolution of the subduction system, including the migration of the basal
671 and roof decollements within the subduction channel.

672 **Oliva-Urcia et al.** present a new magnetostratigraphic section from the Southern Pyrenees to
673 date the Oligocene-Miocene Pyrenean succession. They detail a 5-km section of continental
674 syn-tectonic sediments and link their magneto-stratigraphy results to the activity of the

675 Gavarnie and Guargua thrust sheets, active from 31 to 24 and from 24 to 21 Ma, respectively.
676 They also refine the dating on the latest deformation phase, of ~5 Myr younger than previously
677 published and the dating the Upper Riglos thrust system at 21 Ma. Finally, they also derive the
678 sediment accumulation rates through time that they correlate to first- and second order tectonic
679 activities, as well as to the signal of Ebro basin closure at 36 Ma.

680 Fold-and-thrust belts and basins of eastern Europe

681 **Tomek et al.** investigate the tectonics of the northeastern Variscan belt, in particular the
682 tectonic history of the Moravosilesian Culm Basin. Using structural, paleomagnetic and
683 magnetic anisotropy data, the authors argue that the basin underwent a change in deformation
684 history from early compression to late strike-slip dominated tectonics. Such a switch in
685 kinematics is important, as it provides the opportunity to test how earlier structures are
686 influencing later deformation phases. Similarly, this study shows that from studying foreland
687 fold-thrust belts at the regional scale, inferences can be drawn for the entire orogen and its plate
688 tectonic context. The authors conclude that, as opposed to some previous studies, the strike-slip
689 deformation did not play a major role during continent-continent collision in the area.

690 In order to analyse the timing of late-stage Neo-Tethys subduction and subsequent continent-
691 continent collision in central Turkey, **Gülyüz et al.** integrate field mapping, low-temperature
692 thermochronometry and structural analyses – including inversion of fault slip data – on the
693 Upper Cretaceous to Eocene infill of the Haymana basin. The results of paleostress analysis
694 point out that this basin, located at the junction between the Izmir-Ankara-Erzincan and the
695 Intra-Tauride suture zones, underwent initial N–S to NNE-SSW extension until the middle
696 Paleocene, followed by N–S synsedimentary shortening and coeval E-W directed extension
697 (possibly reaching the middle Miocene). Apatite (U-Th)/He cooling ages indicate that
698 exhumation of the southeast portion of the basin started in the early Oligocene, while the
699 northwest part was exhumed during the early Miocene. The differential uplift and unroofing of

700 the basin fill is tentatively related to the progressive NW-ward activation of a major fault
701 bounding the basin to the north, within the framework of the evolution of the whole basin from
702 an extensional forearc depocentre (Late Cretaceous to early Paleocene) to a foreland basin
703 during the subsequent collision between Taurides and Pontides.

704

705

706 Acknowledgements.

707 The Guest Editor team would like to acknowledge the work of the reviewers who have played
708 an extremely important part in maintaining a high level of rigor to the contributions. Journal
709 Editor-in-Chief Rob Govers is thanked for his support.

710

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